Interannual Variability in contributions of the Equatorial Undercurrent (EUC) to Peruvian Upwelling source water

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Abstract. Time-varying sources of upwelling waters off the coast of northern Peruvian are analysed in a Lagrangian framework, tracking virtual particles backwards in time for 12 months. Particle trajectories are calculated with temperature, salinity and velocity fields from a hindcast spanning 1988-2007, obtained with an eddy-resolving (1/12º) global configuration of the NEMO ocean model. At ~~30 and 100 m, where late-December~~ 30-m and 100-m, where coastal upwelling rates exceed 50 m per month, particles are seeded at monthly intervals in proportion to the upwelling rate. Ensemble maps of particle concentration, age, depth, temperature, salinity and density reveal that a substantial but variable fraction of the particles upwelling off Peru arrive via the Equatorial Undercurrent (EUC). Particles follow the EUC core ~~at around 250 m~~within the depth range 125-175-m, characterised by temperatures ~~of around 15-17ºC~~<17ºC, salinities in the range 34.9-35.2, and densities of ~~σ = 25.5-26.5~~ $\sigma_\theta = 25.5-26.5$. Additional inflows are via two slightly deeper branches further south from the main system, at around ≈3ºS and ≈8ºS. ~~The annual percentage of particles recruited by the EUC~~ Averaged across the hindcast, annual-mean percentages of particles upwelling at 30-m (~~17.5-47%and 16.5-54.6% , from 30 and 100 m respectively)~~ 100-m) associated with the EUC vary from 57.4% (52%) at 92°W to 19.2% (17.9%) at 165°W. Considerable interannual variability in these percentages reveal that more of the Peruvian upwelling can be ~~tracked~~ traced back to the EUC during warm events, such as El Niño~~and weak La Niña events~~. In contrast, upwelling waters are of more local origin during ~~a strong~~ cold events such as La Niña. Despite weaker EUC transport during El Niño, relative flattening of the equatorial thermocline brings the EUC upwelling waters much closer to the Peruvian coast than under neutral or La Niña conditions. Annually averaging EUC transport at specific longitudes, a notable negative-to-positive transition is evident during the major El Niño/La Niña events of ~~1997-99.~~ 1997/99. On short timescales, a degree of longitudinal coherence is evident in EUC transport, with transport anomalies at 160ºW evident at the Galapagos Islands (92ºW) around 30-35 days later. It is concluded that the Peruvian upwelling system is subject to a variable EUC influence, on a wide range of timescales, most notably the interannual timescale of El Niño Southern Oscillation (ENSO). Identifying this variability as a driver of shifts in population and catch data for several key species, during the study period, these new findings ~~may~~ might inform sustainable management of commercially-important fisheries off northern Peru.

1 Introduction

A key feature of the tropical Pacific circulation is the Equatorial Undercurrent (EUC) (Cromwell et al., 1954; Knauss, 1959; Lukas, 1986). The EUC originates in the western equatorial Pacific, just north of Papua New Guinea, as an eastward thermocline flow in the depth range 180-280 m at 147ºE. The EUC strengthens eastward (Bryden, 1985; Tsuchiya et al., 1989; Johnson et al., 2002 , reaching peak velocity and transport at around 140ºW ~~(Knauss, 1959)~~(Knauss, 1958, 1959), in a core located within ≈ 3°N and 3°S of the Equator (Blanke and Raynaud, 1997; Johnson et al., 2002; Brown et al., 2007). Near the Galapagos Islands, the EUC core shifts to around 0.5°S ~~(Kessler, 2006; Karnauskas et al., 2010)~~(Kessler, 2006; Karnauskas et al., 2010, 2020), before strengthening again towards the eastern boundary ~~(Johnson et al., 2002)~~(Lukas, 1986; Johnson et al., 2002), where it feeds major currents along the western South American coastlines of Ecuador, Peru and Chile (Lukas, 1986; Karnauskas et al., 2010; Montes et a . The EUC is located at the depth of the equatorial thermocline, which varies seasonally across the Pacific (Johnson et al., 2002) , leading to seasonality in EUC core velocity. This sub-surface flow plays a crucial role in regional climate and biogeochemistry, through substantial transport of ~~nutrient-~~ nutrient and carbon-rich cold water to the surface, feeding the so-called "cold tongue" upwelling region (Chavez et al., 1998; Pennington et al., 2006; Chavez and Messié, 2009; Qin et al., 2015; Wang et al., 2019).

Around the global coastline, wind-driven upwelling results in high nutrient supply to the surface layer, enhancing primary production where light and nutrient levels are optimal. At the eastern boundary of the Atlantic and Pacific basins in each hemisphere, the "Big Four" upwelling systems - Benguela, California, Iberia/Canary and Chile/Peru - are the most active in the world, accounting for approximately 12 out of 17 million metric tons of marine fish catch-year between 2000 to 2007 (representing 20% of the global taken over an area of less than 1% of the global ocean) according to FAO (on Chavez and Messié (2009)). Of the Big Four, the Peruvian part of the Peru/Chile system presents the highest average volume of upwelled waters (1.6 Sv), even though upwelling-favorable winds are the weakest in average (5.7 m s^{-1}) (Chavez and Messié, 2009; Kämpf and Chapman, 2016).

Peruvian upwelling is dynamically linked to the EUC, itself a sub-surface consequence of equatorial surface flows that are driven by easterly trade winds, pushing surface water to the west along the Equator in the South Equatorial Current and creating a pressure head in the western Pacific. Beneath this wind-dominated surface layer, where the Coriolis effect disappears at the Equator, an eastward pressure gradient drives water back to the east within the EUC, which shoals across the basin from around 250 m in the west to reach the surface in the east ~~(Knauss, 1959).~~ (Knauss, 1959; Johnson et al., 2002).

This equilibrium is interrupted as the trade winds weaken or reverse ~~under during~~ (specially west to 155ºW) under El Niño ~~, with which are associated~~ along with considerable flattening of the thermocline (McPhaden, 1999), seen in historical model simulations (Terada et al., 2020) and moored buoys (Kessler and McPhaden, 1995). Associated with El Niño are changes in the quantity and properties of upwelled waters off Peru.

~~El Niño affects not only the upwelling system, but also brings socioeconomic and cultural impacts to Peru.~~ For instance, in 1982-83 and 1997-98 El Niño events, principal fisheries ~~in Peru~~ such the Peruvian Anchovy, Peruvian hake ~~, giant squid~~ and others collapsed~~completely~~, due to migration and dispersion of the biomass in the region (Ñiquen and Bouchon, 2004; Tam et al., 2006; Wolff et al., 2007). Total pelagic fish landings decreased from 3.3 million tonnes in 1982 to 1.4 millions in 1983 and to almost zero in 1984, as catch species transitioned to Sardine (Arntz and Tarazona, 1989) and others such Jack Mackerel and Pacific Mackerel (Ñiquen and Bouchon, 2004). ~~During the 1997-98 event, the contribution of fisheries to Gross Domestic product (GDP) fell to less than 1%, impacting the employment of approximately 80,000 people (Kämpf and Chapman, 2016). Moreover, under these anomalously warm conditions, Peru experiences severe flooding and extensive climate-related damage to infrastructure, transport, health, employment, throughout the country, making Peru and other countries in the region highly vulnerable to El Niño.~~

In the present study, we address the origin and nature of the upwelled waters of one of the most productive upwelling systems in the world. We specifically address the role of the EUC in Peruvian ~~coastal upwelling. We~~ upwelling, relative to local wind-driven coastal upwelling in the zone 5-10ºS. We aim to quantify absolute and relative changes in the provenance of waters upwelling off northern Peru, to establish the extent of interannual variability and links to ENSO events. We investigate how El Niño and La Niña conditions thus modulate the Peruvian ~~Coast~~ upwelling system. ~~To proceed, we apply a Lagrangian method for tracking upwelling particles backwards through time, using~~ Central to our analysis is an ocean model hindcast spanning 1988-2007. We use a combination of Eulerian and Lagrangian diagnostics, the latter to track upwelling water backwards through time on annual timescales.

In the following, Section 2 describes the methodology used along with brief ~~details~~ specifications of the numerical model and hindcast, and details of our diagnostics. In Section 3, we present a range of results, to illustrate variability in the source of upwelling waters, and in the EUC itself. In Section 4, we ~~reach conclusions regarding~~ summarize and discuss our findings in relation to previous studies. In conclusion, Section 5, we emphasize the consequences of El Niño and La Niña for the EUC~~and~~, Peruvian upwelling, and ~~discuss our findings in relation to previous studies. Finally, Section 5 presents the Code~~ associated marine ecosystems. Section 6 provides details of code and data availability.

2 ~~METHODOLOGY~~Methodology

We first describe the model that provides the hindcast data needed for the Lagrangian analysis~~and~~, EUC transport calculations, and diagnostics of coastal upwelling, which are subsequently outlined.

2.1 Model Description

We sample 5-day averages of temperature, salinity and velocity ~~data~~ from a hindcast spanning 1988-2007, previously obtained with the Nucleus for European Modelling of the Ocean (NEMO) ocean model (Madec, 2008) in eddy-resolving global configuration (ORCA12), henceforth NEMO-ORCA12. For details of parameterization, initialisation and forcing of this hindcast, see Blaker et al. (2015). The model has a horizontal resolution at the equator of 1/12º (9.277 ~~km~~km), with 75 vertical levels from the surface up to 5902 meters depth, alongside finer grid spacing near the surface (38 levels from 0 to 411 ~~meters~~m). The region of focus in this study is the ~~South Eastern~~ eastern tropical Pacific, extracted as the region from 170ºW to 75ºW and from 5ºN to 15ºS, for the purposes of ~~Ariane~~ particle trajectory calculations. An advantage of using fields from a fully global model is that remote influences on the region of interest are fully represented, rather than prescribed at the boundaries in a regional model, which can be problematic. The high resolution NEMO-ORCA12 hindcast is evaluated with observations from three moored NOAA buoys, from the surface up to 400 m depth along the Equatorial Pacific.

2.2 Lagrangian Analysis

To efficiently analyze the provenance of water upwelling off the Peruvian Coast of NEMO-ORCA12, we use the ARIANE Lagrangian code, based on the original method of Blanke and Raynaud (1997). This mass-preserving numerical Lagrangian approach has proved to be an appropriate method for studying the origin and fate of water masses in a wide range of studies (Doos, 1995; Blanke et al., 2002). Ensembles of particles are "seeded" in the coastal upwelling zone off Peru~~were~~, defined here east of 82.5ºW, from 10ºS to the Equator. These particles are then tracked in "backward" mode, reversing in time the analytical calculation of particle progress through grid cells, to reveal ~~exact~~ a range of pathways.

~~Allocating~~ We allocate particles in proportion to ~~the upwelling rate~~ upward vertical velocity where this exceeds a threshold of 50 m per month (1.9x10^{-5} m.s^{-1}). We specifically allocate 1 particle per 10 m per month of upwelling in excess of the threshold, per grid cell. For example, in the simple case that upwelling is 90 m per month, the excess upwelling rate is 40 m per month, so we allocate 4 particles to this grid cell. Initial locations were evenly located within grid cells on the 5-day ORCA12 mesh.

A variable number of particles are released at three "release depths"~~(~~, model levels closest to ~~30 m, 50 m and 100 m), we release a variable number of particles per annual, from 669 (e.g.,1991 for 100 m, see Table 2, Figure 15 in annexes~~30-m, 50-m and 100-m, at monthly intervals throughout 1989-2007. This number ranges from 38,292 (at 30-m in 1998) up to ~~10829 (e.g.,1997 for 30 m, see Table 2, Figure 15 in annexes). Initial locations were evenly located within grid cells on the ORCA12 mesh, where the 5-day averaged upwelling rate exceeded 50 m per month.~~ 98,951 (at 30-m in 2005). Calculations based on initial upwelling ~~from the 50-m~~ at 50-m release depth were relatively similar to those ~~from the 30-m~~ at 30-m release depth, so we ~~will~~ show results for ~~30-m and 100-m~~ 30-m and 100-m release depths only. For examples of the initial positions of particles at both release depths, in Fig. 1 we illustrate seeded particles for the cases of 1997 and 1998.

Particles were released on ~~31 December of each~~ the last day of each month (December-January) for each year and followed back in time~~for 12 months~~, to sample inter-annual changes in 3D pathways and water properties (temperature, salinity, potential density) of consequence for Peruvian upwelling. Particle data were statistically analyzed on a grid of resolution 0.5º x 0.5º to quantify particle "concentration" as a fractional particle presence in Log10 scale, dividing the number of particle occurrences passing through each grid ~~square~~ cell by the total number of particles occurrences during the ~~1-year tracking period.~~ course of the year.

For an average representation of main pathways in the study region, we computed a "grand ensemble" for ~~1988-2007~~ the entire data 1989-2007 (for the ~~30-m and 100-m release depth~~30-m and 100-m release depths). Alongside particle concentration, we also average particle age, depth, salinity, temperature and potential density on the 0.5º x 0.5º grid, providing further context for interannual variability of inflow to the Peruvian upwelling system. Then, given the initial (total) number of upwelled particles off Peru (where upwelling exceeds 50 m per month) and the total number of particles crossing specific locations along the equatorial Pacific, we calculate a percentage recruitment from the EUC (3ºN -3ºS) west to east for further yearly comparisons.

2.3 Pathways and Transport

Informed by the Lagrangian analysis, we also compute EUC volume transport by integrating eastward flow between 3ºN to 3ºS, encompassing most particle trajectories, and in agreement with previous studies (Blanke and Raynaud, 1997; Johnson et al., 2002; Brown et al., 2007), at selected longitudes from 160ºW to 85ºW, every 10º from ~~160ºW-100ºW~~160ºW-110ºW, and at 95ºW, 92ºW and 85ºW respectively. Naturally accommodating EUC flow along shoaling isopycnals, we specifically compute transports binned in potential density, ~~σ~~σ_θ, in the range ~~$\sigma = 23.0 - 27.0$~~$\sigma_\theta = 23.0 - 27.0$, at intervals of ~~$\sigma = 0.1$~~$\sigma_\theta = 0.1$, using the ~~CDFTOOLS diagnostic package (~~NEMO “CDFTOOLS” diagnostic routine “cdfsigtrp” (see https://github.com/meom-group/CDFTOOLS). We thus assemble time series of 5-day averaged EUC transport in density space at the selected longitudes to get time series of monthly averages and anomalies.

3 ~~RESULT AND DISCUSSION~~

~~We~~

2.1 Diagnostics of coastal upwelling

Finally, to evaluate the contribution of local winds to the coastal upwelling, Ekman flux and Upwelling flux were calculated. We sample 5-day averages of vertical velocity east of 82.5ºW longitude and between 5ºS to 10ºS (avoiding the near-Equator zone) at 30 m, where this exceeds 50m per month. Area-integrated, we thus obtain the Upwelling flux in units of Sv. Then, to estimate the wind-driven component of coastal upwelling, we calculate zonal Ekman transport (M_x) associated with the meridional component of wind stress (τ_y) at each coastal gridpoint, per 1/12º of latitude, as $M_x = \tau_y / f\rho$, where f is the Coriolis parameter and ρ is a representative density for seawater (1026 $\mathrm{kg.m^{-3}}$). Multiplied by meridional gridcell length (1/12 x 111 km), we obtain offshore transport in units of Sv that is proportional to coastal upwelling, satisfying continuity of volume. Summed in the same latitudinal range (avoiding low values of f equatorward of 5ºS), we thus obtain a monthly index of wind-driven coastal upwelling by averaging the 5-daily index.

3 Results

We begin with a brief evaluation of the NEMO-ORCA12 hindcast in our study region. We then provide an overview of the Lagrangian calculations, before we focus on the pathways followed by water upwelling off Peru, and variation in the source and transport of these waters. We then consider variability in the wind-driven component of upwelling, concluding with analysis of EUC transport anomalies along the Equator.

3.1 Evaluation of NEMO-ORCA12 hindcast in the equatorial Pacific

To establish whether the NEMO-ORCA12 hindcast realistically simulates the equatorial Pacific circulation, we extract vertical profiles of the monthly mean velocities from the model at longitudes along the Equator (170ºW, 140ºW and 110ºW) corresponding with observational data from NOAA-ADCPs, as shown in Fig. ~~??~~2. The principal eastward subsurface (e.g. EUC) and westward surface currents (e.g. South Equatorial Current - SEC), as described in previous studies (Cromwell et al., 1954; Knauss, 1959; Lukas, 1986; Johnson et al., 2002), are clearly reproduced by NEMO/ORCA12. Shoaling of eastward-flowing core of the EUC towards the Galapagos Islands, can be well identified between 50 and 200 m depth with velocities higher than 0.5 ~~m s~~m.s^{-1} in most of the sections shown in ~~Figure ??.~~ Fig. 2. At 140ºW and 110ºW, the EUC core is seen at around 50 to 150 ~~m~~m, with values notably higher than 1 ~~m s~~m.s^{-1}. The highest core velocity along the equator, is found at 140ºW longitude in this comparison.

3.2 Upwelling rates and backtracking ~~on monthly~~ at annual timescales

We define the upwelling off the northern Peruvian coast where upward advection exceeds 1.9x10^{-5} ~~m s~~m.s^{-1} (50 ~~m~~ m per month), for each release depth in the water column, chosen to sample upwelling over a depth range typical of the Peruvian upwelling system (Kämpf and Chapman, 2016). To illustrate vertical velocities at two of our chosen depths, monthly NEMO-ORCA12 climatologies (1988-2007) at 33 ~~m~~ m and 100 ~~m~~ m are presented in ~~Figures ?? and ?? respectively .~~ Figs. 3 and 4 respectively (we have used here 33m, since is the closest depth available to 30m).

Monthly upward velocities are elevated along both the equatorial upwelling region and off the Peruvian coast, with clear seasonal variability. During austral summer (December-January-February), upwelling is stronger and more widespread north and south of the Equator, and slightly weaker off Peru. During austral autumn (March-April-May), equatorial upwelling is characterised as a narrower zone (restricted equatorward of 3º), with higher values towards the east; off Peru, stronger upwelling is evident, compared to summer. During austral winter (June-July-August), strong upwelling extends across the easternmost equatorial region (140ºW to Galapagos) and off Peru, when south-east trade winds are seasonally stronger.

A notable difference is evident in the ~~100-m~~ 100-m analysis, with downwelling stronger around the Galapagos Islands and near the Ecuadorian coast. This is also consistent with the climatology rates calculated for the Upwelling and Ekman coastal fluxes, peaking during austral autumn and winter seasons and reduced during austral summer season (when also winds are weakest), as forced by the meridional component of the wind stress along the north-central Peruvian coast.

~~Returning to particle release off Peru,~~ Returning to particles released off Peru, for most years (notably 1991, 1997 and ~~to schematically visualize how these particles interact when released at 30 m, we illustrate in Fig. ?? the first 30 days of back-tracking for each year of the hindcast. For most of this December-only ensemble, we observe an eastward~~ 2004 - not shown, we observe a westward origin for some particles, from the Galapagos Islands via the Equator~~(most clearly represented in 1991,1997, 2004)~~. This is consistent with eastward shoaling of the EUC (~~to depths below 40 m~~around 75 m). However, in other years, particles originate from a northern shallower region, in the vicinity of the Ecuadorian coast (e.g. 1988, 1989 ~~,~~ and 1990 - not shown). For particles initialised at 100 ~~m, Fig. ??,~~ m, we observe a similar pattern of inflow at greater ~~depth~~depths (around 125 m), although with a clearer origin from the equatorial band and southern hemisphere, this can be also observed in the grand ensemble.

3.3 Mean pathways at annual timescale

As outlined in Section 2.2, particle concentration and mean age maps are obtained on a 0.5º x 0.5º mesh for each year from ~~1988 to 2007. The 20 ensembles are then~~ 1989 to 2007, further averaged to obtain the “grand ensemble” ~~result for 30-m and 100-m release depths as~~ results for 30-m and 100-m release depths that are shown in Fig. ~~??~~ 5, right and left panels respectively. Tracking backwards, ~~Fig. ?? show the “grand mean ensemble” for the~~ Figs. 5 a-f show the corresponding particle concentration, age, depth, temperature, salinity, and potential density (~~σ) distributions~~σ_θ) distributions, for flows feeding the upwelling ~~off Peru, including particle age as contours, every 60 days (see contours~~ source water off Peru.

We use a logarithmic scale for particle concentration, to emphasize a particularly wide range of this diagnostic (in Fig. ~~??b, left and right panels), for particles released on 31 December. Ages are expressed as days since 31 December, so~~ 5a, red represents high concentration while blue colors low concentration in average). Mean particle age (5b) is expressed in days back-tracked, so for end-December releases, day 0 to day 365 runs from ~~31 December back to 1 January . We use a logarithmic scale for particle concentration, to emphasize a particularly wide range of this diagnostic.~~ December 31st back to January 1st.

In Fig. ~~??a~~5a,b (left and right), highest particle concentration and youngest age (transit time range between 0-120 days) are naturally located near the release ~~section (contour lines)~~ points in both experiments. ~~As particles spread~~ After around 2 months, particles are back-tracked to the Galapagos Islands with relative higher concentrations before further back-conveyance along the equatorial band. As particles are traced westward, high ~~concentrations~~ concentration along ≈ 3ºN to 3ºS are associated with sub-surface eastward flow of the EUC, widening slightly west of the Galapagos Islands. ~~For the 30-m release depth, particle~~ Particle concentration rapidly declines further south and north of 3º to the west of 110ºW, ~~while for 100-m release depth~~ although, we note patches of relatively higher ~~particle concentration (compared to Fig. ??a)~~ concentrations forming two branches in the southeast of ~~the~~ our region.

Inflows ~~remain at~~ from 30-m release depth remain at a relatively shallow depth (above ≈ ~~120 m~~75 m) during the first ~~120 days (spanning December to September in Fig. ??b, left and right panels) with temperature around 17ºC and salinities~~ 150-180 days (Fig. 5c, left panel) described by the Isotherms of 17ºC-20ºC and Isohalines between 34.9 and 35.1 along ~~$\sigma = 25.5 - 26.0$~~$\sigma_\theta = 25.5 - 26.0$ isopycnals (~~for particles released at 30 m, Fig. ??dleft panel). For the same period of time, particles released at 100 m are~~ Fig. 5d,e left panels). Particles released at 100-m more notably deepen after the first 60-90 days (October-September) of release, remaining above 125 m around Galapagos Island and the eastern coast. In this region, particles were characterised with temperatures in the range ~~15-17ºC~~ 15ºC-17ºC and salinities around ~~34.9 along isopycnals at $\sigma < 26.5$~~35-35.1 along isopyncnals $\sigma_\theta = 26.0 - 26.5$ (Fig. ~~??~~5f right panel). ~~After 240-300 days~~

West of Galapagos, after 210 days of back-tracking ~~(spanning May to March)~~in both experiments, particles lie at greater ~~depths (below 240 m) for both release depths~~depth (below 125 m) where particle concentration is the highest. Along the EUC core, between ≈ ~~1ºN~~ 2ºN to ≈ ~~1ºS~~2ºS, temperatures fall below ~~17ºC~~ 17ºC (Fig. ~~??~~5d, left and right panels) and salinities lie ~~between~~ in the range 34.9 – 35.2 (Fig. ~~??~~5e, left and right panels), combining for ~~$\sigma > 26.5$~~$\sigma_\theta > 26.0$ isopycnal (Fig. ~~??~~5f, left and right panels) and populating the isopycnal surface around $\sigma_\theta = 26.5$ for the deeper release depth.

~~Two~~ As previously mentioned, two relatively deep branches (Fig. ~~??~~5a, left and right panels), in agreement with previous studies (Lukas, 1986; Johnson and Moore, 1997; Donohue et al., 2002; Montes et al., 2010; Kuntz and Schrag, 2018), are identified further south of the main equatorial system~~, along~~ . Along ≈3ºS and ≈8ºS (Fig. ~~??~~5a, left and right panels)~~in agreement with previous studies (Lukas, 1986; Johnson and Moore, 1997; Donohue et al., 2002; Montes et al., 2010)~~, more clearly ~~identified~~ distinguished with potential density (~~i. e. Figure ??fleft panel)~~ Fig. 5f) for the deeper release depth experiment. From both release depths, these southern branches are associated with higher density (e.g. > ~~σ~~σ_θ ≈ 26.5 in the case of 100-m release depth), with the southern most branch (≈8ºS) carrying the densest inflow and the northern branch (≈3ºS) at somewhat lower density (i.e. Figure ~~??~~5f left panel)~~, in agreement with previous studies (Montes et al., 2010; Kuntz and Schrag, 2018) .~~ . ~~For those particles released at 30 m, the~~ Our results suggest that these branches are likely associated with a bifurcation of the EUC, to the west of the Galapagos Islands as previous studies have highlighted (Johnson et al., 2002; Karnauskas et al., 2010; Jakoboski et al., 2020). This source of Peruvian water masses in the upwelling region is thus directly modulated by the EUC, when the latter is at its strongest; this may be particularly the case from April to June, when EUC transports reach maxima (Table 1). We will return to this point in the following section.

Regarding the denser southern branch (≈8ºS, left panels for Fig. ~~??~~5d, Fig. ~~??~~5e and Fig. ~~??f) , is related to~~ 5f) traced by those particles released at 30-m, our experiments indicate that the high density is coincident with high salinity rather than low temperature (>35.1, ~~17ºC~~ 17ºC or less, and ~~σ~~$\sigma_\theta = 25.5 - 26.5$ ~~= 25.5-26.5~~). Somewhat different characteristics were found for particles released at ~~100-m~~ 100-m (right panels for Fig. ~~??~~5c, Fig. ~~??~~5d, Fig. ~~??~~5e and Fig. ~~??~~5f), for which temperature is a more dominant influence on density. ~~The branch at ≈3ºS is likely associated with a bifurcation of the EUC, to the west of the Galapagos Islands (Johnson et al., 2002; Karnauskas et al., 2010; Jakoboski et al., 2020). This source of Peruvian upwelling is thus indirectly fed by the EUC; this may be particularly the case from April to June, when EUC transport is at its strongest.~~

3.4 EUC transport across the eastern Pacific and the variable EUC contribution to Peruvian upwelling

~~Particles released off Peru in late December originate from depths in the range 240-260 m (Fig. ??c, left and right panels), mostly within 3º of the Equator (Fig. ??a, left and right panels), indicating~~ Back-tracked trajectories from the released particles off Peru reveled that a substantial proportion of ~~the Peruvian upwelling is~~ water masses are recruited from the main EUC system~~. With this perspective, we compute EUC transport as a function of density in the range 23.0 ≤ σ ≤ 27.0,at 160ºW~~, ~~150ºW, 140ºW, 130ºW, 120ºW, 110ºW, 95ºW, 92ºW and 85ºW - see Fig. ??. In these monthly transport climatologies (averaged over~~ traveling along the equatorial Pacific for around 10,000 km (from 170ºW to 80ºW) before reaching the Peruvian upwelling region. Integrated positive transports across density classes show a longitudinal-seasonal EUC flow variability when averaging over 20 years from 1988-2007 ~~), EUC flow is associated with highest eastward transport at thermocline densities, in the σ-range 25.8-26.7. The overall structure of~~ (Table 1 and Fig. 6). In the first half of the study region between 160ºW and 135ºW, EUC transport ~~in longitude-density space is consistent with observations and other models (Johnson et al., 2002; Qin et al., 2015)~~

~~. From September to December, eastward transport of warm waters by the intensified EUC leads to a flattening of the thermocline. Focussing here on the thermocline density range, westward flows are only evident, for part of the year, at high and low density in Fig. ??.~~ peaks from March to June with average values > 36.2 Sv. Maximum values during April-May are almost double the annual average of 36.3 Sv. East of 135ºW, EUC transport tends to peak during April-June (see Table 1), consistent with observational data (Johnson et al., 2002). Climatological EUC transport tends to decrease from August to December across the whole region.

~~Integrating positive transports across density classes, for total EUC transport,~~ This variability underpins transports of energy and other properties from the western Pacific to the ~~monthly climatology in Table ?? shows a degree of seasonality, with peak transport in April from the mid-Pacific~~ Peruvian Upwelling region. For instance, using Lagrangian trajectories, Qin et al. (2016) found that an elevated iron injection associated with the 1997/98 El Niño showed 30% higher total dissolved iron concentrations in the eastern equatorial Pacific around 13 months later, consistent with suppressed productivity responses in the central, eastern Pacific and off Peru, from models and observations (Ryan et al., 2006; Slemons et al., 2009; Chever et al., 2015). Moreover, eastward transports along the EUC are higher in the western part of the study region than close to Galapagos (~~57.1 Sv at~~ e.g. 36.2 Sv at 160ºW ~~) to Galapagos Island (23.7 Sv at~~ and 12.5 Sv at 92ºW~~). This seasonality~~ , on average), consistent with loss of transport to progressive upwelling along the Equator.

Seasonality is evident for most years of the hindcast, ~~shown in Fig.??a. The monthly transport anomalies shown in Fig. ??b) indicate transport anomalies that can exceed ± 20 Sv during the April peak, notably during the first half of 1990, 1992 and 1993, when negative EUC transport anomalies exceed -20 Sv at 160ºW. Of particular interest here is the development of negative anomalies during El Niño, followed by positive anomalies during La Niña.~~

~~During the historic~~ as shown in the monthly average and anomalies of Fig. 6. To emphasize seasonal cycles, we add contour lines for transports of 25 and 50 Sv in Fig. 6a. We note that the lowest values of <5 Sv coincide with major El Niño ~~of~~ events, such 1997/98 ~~,~~ and 1991/92, and other warm event such 2002/03 in the western region (≈ 160ºW-140ºW). At these times, the ~~EUC disappeared from the central Pacific from December 1997 to January 1998, when the~~ West Pacific Warm Pool ~~migrated eastward with the collapse of the trade winds and the equatorial cold tongue (McPhaden, 1999). Transport anomalies remarkably exceeded -20 Sv for much of 1997 (Fig.??b). Other notable periods of sustained negative EUC transport anomalies span the end of 2002 and beginning of 2003, and late 2006. All of these events were succeeded by strongly positive anomalies, from the most western region (160ºW) to the east, most strikingly over 1997-99 (Fig. ??b) , when the El Niño transitioned to a sustained La Niña.~~ expands across the central Pacific, the thermocline largely flattens, and the EUC quasi-disappears for a short period (McPhaden, 1999).

~~According to previous classifications, major El Ni~~During strong La Niñ~~o~~as 1988-1989 and 1998/~~La Ni~~99, positive anomalies in transport coincide with restoration of an eastward-shoaling thermocline after strong El Niñ~~a events are defined as an intense warming (cooling) over most of the equatorial Pacific, with~~ os the year before, associated with "peaks" of the ~~strongest oceanic signature located in the eastern/central region (McPhaden, 1999; Takahashi and Dewitte, 2015; Bertrand et al., 2020). Notable recent events are the 1982/83 and 1997/98~~ >50 Sv isoline in Fig. 6a, displaced westward from 140ºW to 160ºW. West-East displacement can be also seen following isolines for anomalies (Fig. 6b) during El Niñ~~os, the 1998~~os. For instance, in late 1996 and early 1997, the warm pool with anomalies >+10 Sv was rapidly transported from 160ºW to 92ºW, next to the Galapagos Islands - in less than a year - before reaching the Peruvian coast. A similar transient is observed for smaller events (e.g., 2002/~~99 La Niña, and the 2015~~03 and 2006/~~16 El Ni~~07), although less intense than during strong El Niñ~~o. Moderate Eastern (80-90ºW, 5ºN-5ºS) or Central (160ºE-150ºW, 5ºN-5ºS) El Niños are defined as modest equatorial warming with strongest~~os. We identify other strong displacements, such during 2000/~~weak oceanic signature off the South American coast (Yu et al., 2011). Moderate Eastern~~ 01 and 2004/05, when ENSO variability is not so clearly implicated.

This EUC transport variability is consistent with our Lagrangian trajectory analysis, most clearly during major events, when the signal is clear. Fig. 7 shows yearly variations of the number of particles (a and b for 30-m and 100-m respectively) and its percentage (d and c for 30-m and 100-m respectively) after a year of tracking. In general, our results suggest that particles originating in the western basin cross the Pacific mainly within the EUC, at different intensities varying between 19.2% - 52.0% for 30-m (17.9% - 57.4% for 100-m) between 165ºW and 92ºW (see dashed lines in Fig. 7c and Fig. 7d).

We found that during strong El Niñ~~os are apparent in 1991/92 and 2006/07, whilst the 2000/01~~ o and La Niña ~~may be categorised as a weak Eastern event (Yu et al., 2011).~~

~~Given the initial (total) number of upwelled particles off the north Peruvian coast (where upwelling exceeds 50 m per month) and the total number of particles crossing specific locations along the equatorial Pacific, we calculate percentage recruitment from the EUC (3ºN -3ºS) , shown in Fig. **??**. During the major El~~ events, the percentage of particles feeding the eastern basin changes dramatically. Highest and lowest percentages of particles are observed during the major ENSO of 1997/98. Particles recruited by the EUC in 1997 rapidly crossed the Pacific, showing in the easternmost region (92ºW) the highest record of 71.1% (80.7% for 100-m), consistent with a enhancement of volume transport in Qin et al. (2016). However, during a following cold La Niñ~~o of late 1997, a substantial fraction of particles tracked from the Peruvian upwelling zone crossed 160ºW in the EUC within the year. Evidently, the EUC contribution~~ a phase, in the same longitude range, smaller percentages arrived via the EUC (3.7% for 30-m and 1.7% for 100-m in Fig. 7c and Fig. 7d). This reduced contribution of the EUC to Peruvian upwelling ~~is most dominant during the major El~~ coincides with a western displacement of transport in mid-1998, which persists into 1999 (see anomalies in Fig. 6). Following the major La Niñ~~o of 1997-98 (Fig. **??**, Fig. **??**). Full details of the percentages of upwelling particles recorded earlier in the EUC at selected longitudes are provided in Tables **??** and **??**.~~

~~For the 30-m (100-m) release depth, the highest percentage of~~ a of 1998-99, notably into 2000, positive transport anomalies (>10Sv) prevailed across the Pacific (Fig. 6), consistent with a high percentage of particles ~~found in 1997 varied from 39.5% (34.4%) at 160ºW to 86.6% (93.2%) at 92ºW. This is consistent with flattening of the thermocline with the onset of~~ coming from the EUC (with 68.4%-76.7% to 39.2%-38.7% at 160ºW and 92ºW for 30-100-m respectively, see Fig. 7).

Another major event categorised as a “Moderate Eastern Event” (Yu et al., 2011), unfolded over 1991/92 in the Pacific. For the 1991-92 warm event, our trajectories in percentage showed that above-average numbers of particles are recruited by the EUC close to Galapagos (up to 53.5% for 30-m and 43.7% for 100-m, in 1991 at 165ºW). In the case of a 2002/03 event, we observed the same tendency for warm and cold phases of the event, although less intense compared to that of 1991/92. Conversely, for the 2006/07 “Moderate Eastern” El Niño, we can differentiate between the warm and cold phases, with higher particles recruited by the EUC in ~~1997 (McPhaden, 1999), which limited the upwelling of EUC waters in the central Pacific and~~

~~allowed more of these waters to progress all the way to the eastern boundary, where upwelling continued along the Peruvian coast. Conversely, during~~ 2006 at 165ºW (52.7% for 30-m and 74.9% for 100-m) than during 2007 La Niña (43.1% for 30-m and 53.6% for 100-m). Evidently, the EUC contribution to Peruvian upwelling is most dominant before-during major El Niño or anomalously warm events, than cold or major La Niña ~~, smaller percentages of upwellingparticles arrived via the EUC. Lowest values in 1998 varied from 0.0% (2.3%)at 160ºW to 27.7% (46.4%) at 92ºW, for 30-m (100-m) release depth (see Fig. ??, Tables ?? and ??). This reduced contribution of the EUC to Peruvian upwelling paradoxically coincides with anomalously strong transports at 160ºW during late 1998, which persist into 1999 (transport and anomalies >50 Sv and >20 Sv respectively, see Fig. ??). However, the lowest percentages of upwelling particles in the longitude range 120-160ºW during~~ conditions.

3.5 The coastal wind-driven contribution to Peruvian upwelling

Upwelling off Peru is augmented by equatorward winds that induce offshore transport in a surface layer, supplied by "rapid upwelling" in a relatively narrow coastal zone. Further offshore, "slow upwelling" is induced by negative wind stress curl that induces large-scale Ekman suction (Rykaczewski and Checkley, 2008). Variability in wind-driven upwelling has been often linked to the biological decadal variability in upwelling system such as off Peru (Rykaczewski and Checkley, 2008; Espinoza-Morriberón et we focus on variability in coastal upwelling associated with the meridional component of wind stress (Fig. 8a), used to compute Ekman offshore transport which we term the Ekman flux (Fig. 8b), compared to the full Upwelling flux, obtained from the vertical velocity component at 30 m (Fig. 8c). The three time series in Fig. 8 all trend towards positive anomalies after 2000, most obviously for the Ekman flux, that is a direct consequence of strengthening in the equatorward wind stress. This may be part of a wider strengthening trend in the trade wind system of the Pacific, that has been identified as a driver of the early-2000s hiatus in global warming (England et al., 2014). We further observe that, during strong and Moderate La Niña ~~supports a hypothesis that only particles reaching an easternmost extension of the EUC are likely to influence Peruvian upwelling, consistent with evidence from 1997 and other~~ events such 1988-1999, 1998-1999 and 1993-1994, equatorward wind stress decreases drastically during austral summer with values =< 0.01 N.m^{-2}, and a further minimum during the relatively warm event of summer 2000. From the turn of the millennium, anomalies are generally positive (exceeding climatology), peaking in late 2005 at ≈ 0.07N.m^{-2}.

During the 1997-1998 El Niño~~events (Firing et al., 1983; Halpern, 1987; Mcphaden et al., 1990; Johnson et al., 2000).~~

~~A second highest peak, for the number of upwelling particles and percentage associated with~~, wind stress was not completely shut down in the northern part off Peru, although presenting negative anomalies (reduction) in the first quarter or the year. This is consistent with the high percentage of particles appearing off Peru during extreme El Niños that arrived via the EUC, ~~happened during the Eastern Weak La~~ rather than being locally sourced via coastal Ekman dynamics. An alternative source for upwelled waters is further evidenced in the Upwelling flux (Fig. 8), where positive anomalies coincide with a developing El Niñ~~a of 2000/01 (Fig.??).In 2000, the percentage of particles arriving~~o (e.g., second half of 1997).

3.6 Longitudinal coherence of EUC transport anomalies

To the extent that EUC transport anomalies are coherent across the equatorial Pacific, variability of upwelling via the EUC ~~varied between 35.2% (26.5%)~~may be predictable. To examine the longitudinal coherence of EUC transport anomalies, we compute correlation between EUC transport at each selected longitude (Table 2). Despite generally low values, we find significant (p-value< 0.01, 99% confidence interval) positive correlations at zero-lag. This is consistent with rapid and near-simultaneous variations in EUC transport across the eastern Pacific, although weakening of correlation with separation (in longitude) indicates that transport anomalies may advect eastward, associated with inertial terms in the momentum balance and eastward propagation of Kelvin waves.

The development of an anomalous EUC flow to the east may thus be predictable for an appropriate time lag. Experimenting with lagged cross correlations of transport at 160ºW and ~~76.6% (73.5%) at 160ºW and~~ 92ºW~~for 30-m (100-m) released depth. By 2001, recruited from the west was substantially reduced, with only 5.9% (12.8%) present~~, we obtain highly significant $R = +0.31$ (p-value < 0.01) between transport at 160ºW and transport 30-35 days later at 92ºW. With further investigation, beyond scope here, more skilful EUC predictability may provide useful advance warning of substantial changes in the Peruvian upwelling system, and subsequent impacts on important marine resources. A further step would be assess the variability and predictability of year-round inflows to the Peruvian upwelling system, also beyond the scope of this investigation.

4 Summary and Discussion

With a range of diagnostics, we have evaluated variability of the Peruvian upwelling system in an eddy-resolving model hindcast spanning 1988-2007. Wind-driven coastal upwelling is associated with alongshore winds, complemented by shoaling of the Equatorial Undercurrent. Local and remote drivers of upwelling are highly variable on seasonal to interannual timescales. Contrasting scenarios are shown in Fig. 9. In Fig. 9a, we highlight upwelling dominated by local winds, when the EUC shoals well to the west of the Galapagos Island. In Fig. 9b, the EUC shoals more gradually to the east, reaching the coastal upwelling zone, where weakened winds otherwise limit local upwelling. ~~Enhanced EUC influences on Peruvian upwelling late in 1997 and 2000 are a consequence of very different variability. During the strong El Niño of 1997, the EUC was severely weakened but conveyed particles all the way to the eastern boundary. In contrast, during the Eastern Weak La Niña of 2000, EUC strengthening extended at least to 92ºW (see Fig. ??), conveying more particles eastward. Rather than upwelling along the Equator, as during a strong La Niña, more EUC water reached the eastern boundary during this weak La Niña.~~

4.1 ~~Impacts on Peruvian upwelling linked to variable EUC contributions~~

~~These episodes of enhanced recruitment from the western basin likely impact Peruvian upwelling~~ We now consider impacts on Peruvian upwelling that are linked to variable EUC contributions. Both the 1997/98 and 2000/01 events had important effects in fisheries. May-June 1997 fieldwork conducted by the Peruvian Sea Institute (IMARPE), revealed an unusual southern and deeper (> 150-200 m) migration of Peruvian Hake (Merluccius gayi peruanus; generally distributed between 0ºN to 8ºS), easily reaching 12ºS and even more southern areas along the Peruvian coast (Castillo et al., 1997). Conversely, Hake tended to migrate northward and sometimes closer to the surface ~~,~~ (pelagic behaviour, considering that hake is a demersal specie), due bottom oxygen deficiency, during the La Niña of 1998/99. This is consistent with the ~~biogeochemical~~ physical-biogeochemical character of the EUC. Where a southward branch of the EUC reaches the northern Peruvian coast (Fig. ~~??~~5), it refreshes the sub-surface layer with high concentrations of dissolved oxygen (Echevin et al., 2020) and other vital nutrients ~~.~~ (Qin et al., 2016; Espinoza-Morriberón et al., 2019).

A similar impact afflicts the second most abundant demersal species in Peruvian coastal waters, the Conger eel (Ophichthus remiger). During ~~recent~~ El Niño events (December 1997 to May 1998; March-July 1992) and subsequent La Niña events (November 1993 to March 1994; April-August 1998), significant ~~enhanced~~ increased and reduced Catch per Unit Effort (CPUE) was reported, respectively (Castillo et al., 2000; Martina, 2004). The oceanographic and fisheries evidence together suggest crucial links between demersal species and the EUC. More recently, Martina (2018) also highlighted dramatic reduction in biomass during 2004/05, followed by significant increase in 2007. The latter increase coincides with positive EUC transport anomalies taking place during early 2007 (see Fig. ~~??~~6b).

During the moderate Eastern Pacific El Niño of 1991/92, an alternative transition is observed. A relatively high percentage of upwelling particles is traced back to the EUC, ~~with 31.4% (9.4~~amounting to 48.2% (36.5%) and ~~53.2% (57.4~~22.6% (8.7%) at 160ºW and 92ºW ~~respectively~~ during 1991, for ~~30-m (100-m~~30-m (100-m) release depths respectively (Fig. ~~??b~~ 7c and Fig. ~~??~~7d). These enhanced percentages of EUC origin, at 160ºW from 1991 to 1992, coincide with relatively ~~small anomalies in eastward transport~~ high anomalies in transports (Fig. ~~??~~7b). ~~Through~~ From the second half of 1991, east of 140ºW, positive anomalies >+10Sv were found, while at the end of 1991 to the first half of 1992, ~~particles from 30-m release depth were present in much lower percentages along the EUC, while particles back-tracked from 100-m appeared in higher percentages (see Tables ?? and ??), consistent with a subsequently deeper thermocline and EUC during 1992. Related to the~~ high negative anomalies were observed from 160ºW to 110ºW. This is consistent with eastward transport of warm pool waters and the relatively high percentage of EUC particles found in 1992 near Galapagos (92ºW), associated with deepening of the thermocline followed by relaxation-shoaling, in the eastern Pacific (Kessler and Mcphaden, 1995).

Related to 1991/92 El Niño~~was~~ , a dramatic size-age and biomass reduction ~~in 1992, along with spatial migration (Castillo, 1996), as during the major El Niño of 1997. These changes were previously attributed to a combination of overfishing by foreign fishing vessels during 1989-90 and~~ of the Peruvian Hake happened. This was first attributed to an overfishing (Castillo, 1996), although observations indicated an enhanced EUC at this time. This likely led to drastic relocation of species farther south, causing high mortality for adults and leaving young species in the northern region of Peru. Plus, sub-optimal oceanographic conditions ~~during the 1991/92 El Niño,~~ when the EUC ~~disappeared from most of the eastern Pacific. While species~~ weakened after its peaks passed. Species are often slow to recover in size-age-abundance indicators~~, even~~ . Even after the major El Niño of 1997/98, ~~recovery is observed~~ Hake has shown signs of recovery in CPUE data as wider size structure and higher catch numbers (Ballón et al., 2008; Lassen et al., 2009a).

~~In summary, the EUC contribution to upwelling along the northern Peruvian coast is sometimes substantial, altering environmental conditions of consequence for many marine species. It~~ Furthermore, it has further been suggested that some species also use the EUC as a migration conveyor, such as the Pacific Giant squid (Dosidicus gigas). Largely absent from east Pacific coastal waters over 1950-90, catches of this species increased from the beginning of 1991 through 1992. During 1995-96, overfishing impacted this population (in the northern American coast), primary due to the high catch effort of foreign fishing vessels (Contreras Paya, 2017). Subsequent to the El Niño, of ~~1997-98~~1997/98, a denser nearshore migration of the squid, more accessible for fishing, was ~~observed~~ reported off the North American coast (CALCOFI, 2000). Changes in coastal squid populations may be linked to the equatorially and then coastally trapped Kelvin waves (e.g., Shaffer et al. (1997)), which may favour southward (northward) migration when the EUC is stronger (weaker) around the Galapagos Island, as the squid seek out for optimal conditions (Cheung W. L. et al., 2018).

We identify less intense events at other times, such late 2002~~to early 2003~~ /03 (Fig. ~~??b~~ 6 and Fig. ~~??d~~7), when Peruvian Hake almost ~~disappeared from a fishery off northern Peruvian~~collapsed off northern Peru, leading to a moratorium on fishing ~~this species~~ until March 2004 (Lassen et al., 2009b; Benites and Barriga, 2011). ~~Following~~ This followed the moderate Eastern El Niño of ~~2006-07, a substantial percentage of particles upwelling at 100-m in late~~ 2006/07, with positive transport anomalies (>+10 Sv) during the first half of 2006, transporting water masses to the east. Then, from the end of 2006 to early 2007~~are traced back to the EUC, in contrast to very few particle back-traced from 30-m, suggesting that the EUC remained anomalously deep following the El Niño.~~

4.1 ~~Longitudinal coherence of EUC transport anomalies~~

~~To examine the longitudinal coherence of EUC transport anomalies , we compute correlation between EUC transport at each selected longitude (Table ??). Despite generally low values, we find significant (p-value< 0.01, 99% confidence interval) positive correlations at zero-lag.~~ , negative anomalies were observed in the west, indicating that water masses reached the east. This is consistent with ~~rapid and near-simultaneous variations in EUC transport across the eastern Pacific, although weakening of correlation with separation (in longitude) indicates that transport anomalies may advect eastward, associated with inertial terms in the momentum balance and eastward propagation of Kelvin waves. The development of anomalous flow to~~ higher percentage of particles in 2006 and a reduction of those in 2007 (Fig. 7) from both experiments.

In summary, the ~~east may thus be predictable for an appropriate time lag. Experimenting with lagged cross correlations of transport at 160ºW and 92ºW, we obtain highly significant R = +0.31 (p-value < 0.01)between transport at 92ºW and transport 30-35 days earlier at 160ºW. With further investigation, beyond scope here, more skilful EUC predictability may provide useful advance warning of substantial changes in the Peruvian upwelling system, and subsequent impacts on important marine resources. A further step would be assess the variability and predictability of year-round inflows to the Peruvian upwelling system, also beyond the scope of this preliminary investigation~~EUC contribution to upwelling region along the northern Peruvian coast is sometimes substantial, depending on the type of event (strong, moderate or weak), altering environmental conditions of consequence for many marine species. Changes in the EUC in the Pacific, as a first stage before El Niño happens off Peru, can have positive and negative effects in fisheries regarding the species (Icochea et al., 1989; Castillo et al., 1997; Contreras Paya, Identifying this variability as a driver of shifts in population for several key species, might efficiently inform sustainable management of fisheries in Peru.

5 Conclusions

We have systematically quantified the origin of waters upwelling in one of the most productive regions in the global ocean, ~~the Peruvian Upwelling System. Analyses are~~ based on a high resolution NEMO-ORCA12 eddy-resolving model hindcast spanning 1988-2007, which includes major warm and cold events associated with El Niño and La Niña respectively. Through Lagrangian analysis of virtual particle ensembles, we ~~established that~~ identified an interannually variable fraction of ~~the particles~~ upwelled water that is recruited via the Equatorial Undercurrent (EUC), flowing between ≈ 3ºN to 3ºS at ~~around 250 m~~depths ranging 125 m to 175 m. A key ~~inference~~ finding is that the northern Peruvian upwelling system is sensitive to highly variable EUC inflow.

Particle back-trajectories - sampling the upwelling at depths of 30-m and 100-m off Peru - trace the EUC as far west as 170ºW on an annual timescale, moving at depths, temperatures, salinities and densities that are consistent with observations (Johnson et al., 2002). Back-trajectories further identified two relatively deep branches south of the main equatorial system, along ≈3ºS and ≈8ºS, in agreement with previous studies (Lukas, 1986; Johnson and Moore, 1997; Donohue et al., 2002; Montes et al., 2010). Only a small percentage of particles otherwise originate from latitudes poleward of 3º, as far west as 170ºW.

Over ~~the 1988-2007 hindcast period~~1989-2007, we quantified the variable contribution of EUC waters to Peruvian upwelling~~in late December~~. At 92ºW, we identify highest and lowest percentages of EUC-sourced particles during the strong El Niño of 1997/98 and the subsequent strong La Niña of 1998/99, respectively. The EUC is thus most influential – in relative terms – when the equatorial thermocline ~~is flattened during~~ deepens across the eastern Pacific, during early El Niño, allowing ~~flow all the way to~~ the EUC to extend as far as the eastern boundary and upwelling to the east of Galapagos. During a strong La Niña, EUC waters conversely upwell with a strongly shoaling thermocline to reach the surface layer in the central Pacific, ~~far to the west of the eastern boundary~~with weaker EUC upwelling further to the east. In this scenario, the waters upwelling off Peru are of local provenance ~~.~~ (see Fig. 9).

Variable provenance of oxygenated, nutrient-rich EUC waters in the 1990s can be linked to substantial changes in Peruvian coastal fisheries. Further variability in the 2000s is associated with a range of El Niño and La Niña events. On sub-annual timescales, particles most rapidly cross the eastern Pacific with the EUC during peak transport around March and April, consistent with previous studies (Flores et al., 2009). Correlations between 5-day averaged EUC transport at selected longitudes in the range 92-160ºW indicate a high degree of longitudinal coherence, with evidence of a time lag of 30-35 days for transport anomalies at 160ºW to reach the Galapagos Islands at 92ºW.

In highlighting the impact of variable EUC influences on regional biogeochemistry, ecosystems and fisheries, this study provides the basis for informed analysis and prediction of an unfolding El Niño or La Niña, for a sustainable approach to management of marine resources in the Peruvian upwelling system. A next step would be to include biogeochemical analyses, using models and observations, to better understand the consequences of variable nutrient supply for primary productivity at the base of the food chain.

6 Code and data availability

The NEMO-ORCA12 data analysed here is archived at the National Oceanography Centre, Southampton. The original version of ARIANE software used here, available from http://stockage.univ-brest.fr/ grima/Ariane/, was adapted at the National Oceanography Centre and the University of Southampton. Specific trajectory data and NEMO-ORCA12 diagnostics presented here are available from the authors, on request. Equatorial current data from the TAO/TRITON Array are available from NOAA via https://www.pmel.noaa.gov/gtmba/pmel-theme/pacific-ocean-tao.

Author contributions. .

All authors have contributed equally to this paper.

Competing interests. .

The authors declare that they have no conflict of interest.

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References

Arntz, W. E. and Tarazona, J.: Summary for Policymakers, in: Climate Change 2013 - The Physical Science Basis, edited by Intergovernmental Panel on Climate Change, vol. 52, pp. 1–30, Cambridge University Press, Cambridge, https://doi.org/10.1017/CBO9781107415324.004, https://www.cambridge.org/core/product/identifier/CBO9781107415324A009/type/book_part, 1989.

Ballón, M., Wosnitza-Mendo, C., Guevara-Carrasco, R., and Bertrand, A.: The impact of overfishing and El Niño on the condition factor and reproductive success of Peruvian hake, Merluccius gayi peruanus, Progress in Oceanography, 79, 300–307, https://doi.org/10.1016/j.pocean.2008.10.016, http://dx.doi.org/10.1016/j.pocean.2008.10.016, 2008.

Benites, C. and Barriga, E.: La Poblacion de la merluza durante el verano 2004. Crucero BIC Olaya 0401-02, Tech. rep., Instituto del Mar del Peru, Callao, https://doi.org/ISSN 0378-7702, 2011.

Bertrand, A., Lengaigne, M., Takahashi, K., Avadí, A., Florence, P., and Chris, H.: El Niño Southern Oscillation (ENSO) effects on fisheries and aquaculture, FAO Fisheries and Aquaculture Technical Paper No. 660., Rome, https://doi.org/10.4060/ca8348en, 2020.

Blaker, A. T., Hirschi, J. J., McCarthy, G., Sinha, B., Taws, S., Marsh, R., Coward, A., and de Cuevas, B.: Historical analogues of the recent extreme minima observed in the Atlantic meridional overturning circulation at 26°N, Climate Dynamics, 44, 457–473, https://doi.org/10.1007/s00382-014-2274-6, 2015.

Blanke, B. and Raynaud, S.: Kinematics of the Pacific Equatorial Undercurrent: An Eulerian and Lagrangian approach from GCM results, Journal of Physical Oceanography, 27, 1038–1053, https://doi.org/10.1175/1520-0485(1997)027<1038:KOTPEU>2.0.CO;2, 1997.

Blanke, B., Arhan, M., Lazar, A., and Prévost, G.: A Lagrangian numerical investigation of the origins and fates of the salinity maximum water in the Atlantic, Journal of Geophysical Research C: Oceans, 107, 27–1, https://doi.org/10.1029/2002jc001318, 2002.

Brown, J. N., Godfrey, J. S., and Fiedler, R.: A zonal momentum balance on density layers for the central and eastern equatorial Pacific, Journal of Physical Oceanography, 37, 1939–1955, https://doi.org/10.1175/JPO3090.1, 2007.

Bryden, H. L. B. E. C.: Dignostic Model of the Three-Dimensional Circulation in the Upper Equatorial Pacific Ocean, Journal of Physical Oceanography, 15, 1255–1273, 1985.

CALCOFI: California Cooperative Oceanic Fisheries Investigations, Tech. Rep. 5, National Oceanic and Atmospheric administration, National Marine Fisheries service, La Jolla CA, https://doi.org/10.1163/156851801511873, http://www.calcofi.org/publications/calcofireports/v41/CalCOFI_Rpt_Vol_41_2000.pdf, 2000.

Castillo, R.: Informe tecnico sobre la situacion de la merluza 1990-1995, Tech. rep., Instituto del Mar del Peru, Callao, http://biblioimarpe.imarpe.gob.pe/bitstream/123456789/899/1/IP17.pdf, 1996.

Castillo, R., Ganoza, F., Aliaga, A., Gutiérrez, M., and Guevara-Carrasco, R.: Distribución, concentración y biomasa de la merluza peruana en otoño 1997 con el método hidroacústico. Crucero BIC Humboldt 9705-06, Tech. rep., Instituto del Mar del Peru, Callao, http://biblioimarpe.imarpe.gob.pe/handle/123456789/1486, 1997.

Castillo, R., Gomez, E., and Paredes, F.: Pesqueria y biologia de la Anguila comun Ophichthus pacifici (Gunther)* en el Peru, Tech. rep., Instituto del Mar del Peru, Callao, http://biblioimarpe.imarpe.gob.pe/bitstream/123456789/1168/1/IP134.pdf, 2000.

Chavez, F. P. and Messié, M.: A comparison of Eastern Boundary Upwelling Ecosystems, Progress in Oceanography, 83, 80–96, https://doi.org/10.1016/j.pocean.2009.07.032, http://dx.doi.org/10.1016/j.pocean.2009.07.032, 2009.

Chavez, F. P., Strutton, P. G., and McPhaden, M. J.: Biological-physical coupling in the central equatorial Pacific during the onset of the 1997-98 El Niño, Geophysical Research Letters, 25, 3543–3546, https://doi.org/10.1029/98GL02729, 1998.

Cheung W. L., W., Bruggeman, J., and Butenschön, M.: Chapter 4: Projected changes in global and national potential marine fisheries catch under climate change scenarios in the twenty-first century, in: Impacts of climate change on fisheries and aquaculture. Synthesis of current knowledge, adaptation and mitigation options., edited by Barange, M., Bahri, T., Beveridge C. M., M., Cochrane L., K., Funge-Smith, S., and Poulain, F., vol. 627, chap. 4, pp. 113–138, Food and Agriculture Organization of the United Nations, Rome, 2018.

Chever, F., Rouxel, O. J., Croot, P. L., Ponzevera, E., Wuttig, K., and Auro, M.: Total dissolvable and dissolved iron isotopes in the water column of the Peru upwelling regime, Geochimica et Cosmochimica Acta, 162, 66–82, https://doi.org/10.1016/j.gca.2015.04.031, http://dx.doi.org/10.1016/j.gca.2015.04.031, 2015.

Contreras Paya, I.: DOCUMENTO CONSOLIDADO, Convenio de desempeño 2016. Estatus y posibilidades de explotacion biologicamente sustentable de los principales recursos nacionales del 2017: Jibia, 2017, Tech. rep., Instituto de Fomento Pesquero, https://doi.org/10.13140/RG.2.2.20164.63365, http://www.ifop.cl/wp-content/contenidos/uploads/RepositorioIfop/InformeFinal/P-483253_jibia.pdf, 2017.

Cromwell, T., Montgomery, R., and Stroup, E.: Equatorial Undercurrent in Pacific Ocean Revealed by New Methods, Science, 119, 648–649, https://doi.org/10.1126/science.119.3097.648, 1954.

Donohue, K. A., Firing, E., Rowe, G. D., Ishida, A., and Mitsudera, H.: Equatorial pacific subsurface countercurrents: A model-data comparison in stream coordinates, Journal of Physical Oceanography, 32, 1252–1264, https://doi.org/10.1175/1520-0485(2002)032<1252:EPSCAM>2.0.CO;2, 2002.

Doos, K.: Interocean exchange of water masses, Journal of Geophysical Research, 100, 499–514, 1995.

Echevin, V., Gévaudan, M., Espinoza-morriberon, D., Tam, J., Aumont, O., Gutierrez, D., and Colas, F.: Physical and biogeochemical impacts of RCP8 . 5 scenario in the Peru upwelling system, Biogeosciences, pp. 1–42, https://doi.org/doi.org/10.5194/bg-2020-4, 2020.

England, M. H., Mcgregor, S., Spence, P., Meehl, G. A., Timmermann, A., Cai, W., Gupta, A. S., Mcphaden, M. J., Purich, A., and Santoso, A.: Recent intensification of wind-driven circulation in the Pacific and the ongoing warming hiatus, Nature Climate Change, 4, 222–227, https://doi.org/10.1038/nclimate2106, 2014.

Espinoza-Morriberón, D., Echevin, V., Colas, F., Tam, J., Gutierrez, D., Graco, M., Ledesma, J., and Quispe-Ccalluari, C.: Oxygen variability during ENSO in the Tropical South Eastern Pacific, Frontiers in Marine Science, 5, 1–20, https://doi.org/10.3389/fmars.2018.00526, 2019.

Firing, E., Lukas, R., Sadler, J., and Wyrtki, K.: Equatorial undercurrent disappears during 1982-1983 El Niño, Science, 222, 1121–1123, https://doi.org/10.1126/science.222.4628.1121, 1983.

Flores, R., Tenorio, J., and Dominguez, N.: Variaciones de la Extension Sur de la Corriente Cromwell frente al Peru entre los 3 y 14°S, Tech. rep., Instituto del Mar del Peru, Callao, http://biblioimarpe.imarpe.gob.pe/bitstream/123456789/1091/1/BOL24%281-2%29-7.pdf, 2009.

Halpern, D.: Observations of Annual and El Niño Thermal and Flow Variations at 0º110ºW and 0º95ºW during 1980-1985, Journal of Geophysical Research, 92, 8197–8212, https://doi.org/http://dx.doi.org/10.1029/JC092iC08p08197, 1987.

Icochea, L., Chipollini, A., and Ñiquen, M.: Analisis de la pesqueria de arrastre pelagica en la costa peruana durante 1983-1987 y su relacion con el medio ambiente, in: Memorias del Simposio Internacional de los Recursos Vivos y las Pesquerias en el Pacifico Sudeste, pp. 455–465, Instituto del Mar del Peru, Viña del Mar, https://biblioteca.imarpe.gob.pe/opac_css/index.php?lvl=notice_display&id=13778, 1989.

Jakoboski, J., Todd, R. E., Brechner Owens, W., Karnauskas, K. B., and Rudnick, D. L.: Bifurcation and upwelling of the equatorial undercurrent west of the galápagos archipelago, Journal of Physical Oceanography, 50, 887–905, https://doi.org/10.1175/JPO-D-19-0110.1, 2020.

Johnson, G. C. and Moore, D. W.: The pacific subsurface countercurrents and an inertial model, Journal of Physical Oceanography, 27, 2448–2459, https://doi.org/10.1175/1520-0485(1997)027<2448:TPSCAA>2.0.CO;2, 1997.

Johnson, G. C., McPhaden, M. J., Rowe, G. D., and McTaggart, K. E.: Upper equatorial Pacific Ocean current and salinity variability during the 1996-1998 El Niño-La Niña cycle, Journal of Geophysical Research: Oceans, 105, 1037–1053, https://doi.org/10.1029/1999jc900280, 2000.

Johnson, G. C., Sloyan, B. M., Kessler, W. S., and McTaggart, K. E.: Direct measurements of upper ocean currents and water properties across the tropical Pacific during the 1990s, Progress in Oceanography, 52, 31–61, https://doi.org/10.1016/S0079-6611(02)00021-6, 2002.

Kämpf, J. and Chapman, P.: Upwelling Systems of the World. A Scientific Journey to the Most Productive Marine Ecosystems, Springer Nature, Switzerland, https://doi.org/10.1007/978-3-319-42524-5, 2016.

Karnauskas, K. B., Murtugudde, R., and Busalacchi, A. J.: Observing the Galápagos-EUC interaction: Insights and challenges, Journal of Physical Oceanography, 40, 2768–2777, https://doi.org/10.1175/2010JPO4461.1, 2010.

Karnauskas, K. B., Jakoboski, J., Johnston, T. M., Owens, W. B., Rudnick, D. L., and Todd, R. E.: The Pacific Equatorial Undercurrent in Three Generations of Global Climate Models and Glider Observations, Journal of Geophysical Research: Oceans, 125, 1–12, https://doi.org/10.1029/2020JC016609, 2020.

Kessler, S. and Mcphaden, J.: The 1991-1993 El Nifio in the central Pacific, Deep Sea Research II, 42, 295–333, https://doi.org/https://doi.org/10.1016/0967-0645(95)00041-N, 1995.

Kessler, W. S.: The circulation of the eastern tropical Pacific: A review, Progress in Oceanography, 69, 181–217, https://doi.org/10.1016/j.pocean.2006.03.009, 2006.

Kessler, W. S. and McPhaden, M. J.: The 1991-1993 El Niño in the central Pacific, Deep-Sea Research Part II, 42, 295–333, https://doi.org/10.1016/0967-0645(95)00041-N, 1995.

Knauss, J. A.: Observations of the Pacific Equatorial Undercurrent, Nature, pp. 601–602, https://doi.org/https://doi.org/10.1038/182601a0, https://www.nature.com/articles/182601a0, 1958.

Knauss, J. A.: Measurements of the Cromwell current, Deep Sea Research (1953), 6, 265–286, https://doi.org/10.1016/0146-6313(59)90086-3, 1959.

Kuntz, L. B. and Schrag, D. P.: Hemispheric asymmetry in the ventilated thermocline of the tropical pacific, Journal of Climate, 31, 1281–1288, https://doi.org/10.1175/JCLI-D-17-0686.1, 2018.

Lassen, H., Barriga, E., Palacios, J., Vargas, N., Díaz, E., and Argüelles, J.: Evaluación del estado del stock de merluza (Merluccius gayi peruanus Ginsburg) en el mar peruano. 2008, Tech. Rep. 1-2, Instituto del Mar del Peru, Callao, https://hdl.handle.net/20.500.12958/1016, 2009a.

Lassen, H., Barriga, E., Palacios, J., Vargas, N., Díaz, E., and Argüelles, J.: Evaluación del estado del stock de merluza (Merluccius gayi peruanus Ginsburg) en el mar peruano. 2008, Tech. Rep. 1-2, Instituto del Mar del Peru, Callao, 2009b.

Lukas, R.: The termination of the Equatorial Undercurrent in the eastern Pacific, Progress in Oceanography, 16, 63–90, https://doi.org/10.1016/0079-6611(86)90007-8, 1986.

Madec, G.: NEMO ocean engine, Note du P^ole de mod´elisation de l'Institut Pierre-Simon Laplace No 27, France, 2008.

Martina, J.: Analisis de la captura por unidad de esfuerzo (CPUE) de la anguila comun (Ophichthus pacifici) y su relacion con la Corriente Cromwell, Ph.D. thesis, Universidad Nacional Agraria La Molina, http://ban.lamolina.edu.pe/search~S1*spi?/cM11+M37+-+T/cm++++11+m37+t/-3%2C-1%2C0%2CB/frameset&FF=cm++++11+m37+t&1%2C1%2C, 2004.

Martina, J.: Estimación del coeficiente de capturabilidad (q) y de la biomasa de anguila común (Ophichthus remiger) del norte de Perú mediante análisis geoestadístico de las densidades de captura using likelihood-based geostatistical method on fish density, Anales Cientificos, 79, 168–177, https://doi.org/http://dx.doi.org/10.21704/ac.v79i1.1159, ISSN2519-7398(Versiónelectrónica), 2018.

McPhaden, M. J.: Genesis and Evolution of the 1997-98 El Niño, Science, 283, 950–954, https://doi.org/https://doi.org/10.1126/science.283.5404.950, 1999.

Mcphaden, M. J., Hayes, S. P., and Mangum, L. J.: Variability in the Western Equatorial Pacific Ocean during the 1986-87 El Niño/Southern Oscillation Event*, Journal of Physical Oceanography, 20, 190–208, https://doi.org/https://doi.org/10.1175/1520-0485(1990)020<0190:VITWEP>2.0.CO;2, 1990.

Montes, I., Colas, F., Capet, X., and Schneider, W.: On the pathways of the equatorial subsurface currents in the eastern equatorial Pacific and their contributions to the Peru-Chile Undercurrent, Journal of Geophysical Research: Oceans, 115, 1–16, https://doi.org/10.1029/2009JC005710, 2010.

Ñiquen, M. and Bouchon, M.: Impact of El Niño events on pelagic fisheries in Peruvian waters, Deep-Sea Research Part II: Topical Studies in Oceanography, 51, 563–574, https://doi.org/10.1016/j.dsr2.2004.03.001, 2004.

Pennington, J. T., Mahoney, K. L., Kuwahara, V. S., Kolber, D. D., Calienes, R., and Chavez, F. P.: Primary production in the eastern tropical Pacific: A review, Progress in Oceanography, 69, 285–317, https://doi.org/10.1016/j.pocean.2006.03.012, 2006.

Qin, X., Gupta, A. S., and van Sebille, E.: Variability in the origins and pathways of Pacific Equatorial Undercurrent, Journal of Geophysical Research: Oceans, 120, 3113–3128, https://doi.org/10.1002/ 2014JC010549., 2015.

Qin, X., Menviel, L., Sen Gupta, A., and van Sebille, E.: Iron sources and pathways into the Pacific Equatorial Undercurrent, Geophysical Research Letters, 43, 9843–9851, https://doi.org/10.1002/2016GL070501, 2016.

Ryan, J. P., Ueki, I., Chao, Y., Zhang, H., Polito, P. S., and Chavez, F. P.: Western Pacific modulation of large phytoplankton blooms in the central and eastern equatorial Pacific, Journal of Geophysical Research: Biogeosciences, 111, 1–14, https://doi.org/10.1029/2005JG000084, 2006.

Rykaczewski, R. R. and Checkley, D. M.: Influence of ocean winds on the pelagic ecosystem in upwelling regions, Proceedings of the National Academy of Sciences of the United States of America, 105, 1965–1970, https://doi.org/10.1073/pnas.0711777105, 2008.

Shaffer, G., Pizarro, O., Djurfeldt, L., Salinas, S., and Rutllant, J.: Circulation and low-frequency variability near the Chilean coast: Remotely forced fluctuations during the 1991-92 El Niño, Journal of Physical Oceanography, 27, 217–235, https://doi.org/10.1175/1520-0485(1997)027<0217:CALFVN>2.0.CO;2, 1997.

Slemons, L., Gorgues, T., Aumont, O., Menkes, C., and Murray, J. W.: Biogeochemical impact of a model western iron source in the Pacific Equatorial Undercurrent, Deep-Sea Research Part I: Oceanographic Research Papers, 56, 2115–2128, https://doi.org/10.1016/j.dsr.2009.08.005, http://dx.doi.org/10.1016/j.dsr.2009.08.005, 2009.

Takahashi, K. and Dewitte, B.: Strong and moderate nonlinear El Niño regimes, Climate Dynamics, 46, 1627–1645, https://doi.org/10.1007/s00382-015-2665-3, 2015.

Tam, J., Purca, S., Duarte, L. O., Blaskovic, V., and Espinoza, P.: Changes in the diet of hake associated with El Niño 1997-1998 in the northern Humboldt Current ecosystem, Advances in Geosciences, 6, 63–67, https://doi.org/10.5194/adgeo-6-63-2006, 2006.

Tam, J., Taylor, M. H., Blaskovic, V., Espinoza, P., Michael Ballón, R., Díaz, E., Wosnitza-Mendo, C., Argüelles, J., Purca, S., Ayón, P., Quipuzcoa, L., Gutiérrez, D., Goya, E., Ochoa, N., and Wolff, M.: Trophic modeling of the Northern Humboldt Current Ecosystem, Part I: Comparing trophic linkages under La Niña and El Niño conditions, Progress in Oceanography, 79, 352–365, https://doi.org/10.1016/j.pocean.2008.10.007, http://dx.doi.org/10.1016/j.pocean.2008.10.007, 2008.

Taylor, M. H., Tam, J., Blaskovic, V., Espinoza, P., Michael Ballón, R., Wosnitza-Mendo, C., Argüelles, J., Díaz, E., Purca, S., Ochoa, N., Ayón, P., Goya, E., Gutiérrez, D., Quipuzcoa, L., and Wolff, M.: Trophic modeling of the Northern Humboldt Current Ecosystem, Part II: Elucidating ecosystem dynamics from 1995 to 2004 with a focus on the impact of ENSO, Progress in Oceanography, 79, 366–378, https://doi.org/10.1016/j.pocean.2008.10.008, http://dx.doi.org/10.1016/j.pocean.2008.10.008, 2008.

Terada, M., Minobe, S., and Deutsch, C.: Mechanisms of future changes in equatorial upwelling: CMIP5 intermodel analysis, Journal of Climate, 33, 497–510, https://doi.org/10.1175/JCLI-D-19-0128.1, 2020.

Tsuchiya, M., Lukas, R., Fine, R. A., Firing, E., and Lindstrom, E.: Source waters of the Pacific Equatorial Undercurrent, Progress in Oceanography, 23, 101–147, https://doi.org/10.1016/0079-6611(89)90012-8, 1989.

Wang, Q., Wang, F., Feng, J., Hu, S., Zhang, L., Jia, F., and Hu, D.: The Equatorial Undercurrent and Its Origin in the Region Between Mindanao and New Guinea, Journal of Geophysical Research: Oceans, 124, 2313–2330, https://doi.org/10.1029/2018JC014842, 2019.

Wolff, M., Taylor, M., Mendo, J., and Yamashiro, C.: A catch forecast model for the Peruvian scallop (Argopecten purpuratus) based on estimators of spawning stock and settlement rate, Ecological Modelling, 209, 333–341, https://doi.org/10.1016/j.ecolmodel.2007.07.013, 2007.

Yu, J. Y., Kao, H. Y., Lee, T., and Kim, S. T.: Subsurface ocean temperature indices for Central-Pacific and Eastern-Pacific types of El Niño and La Niña events, Theoretical and Applied Climatology, 103, 337–344, https://doi.org/10.1007/s00704-010-0307-6, 2011.

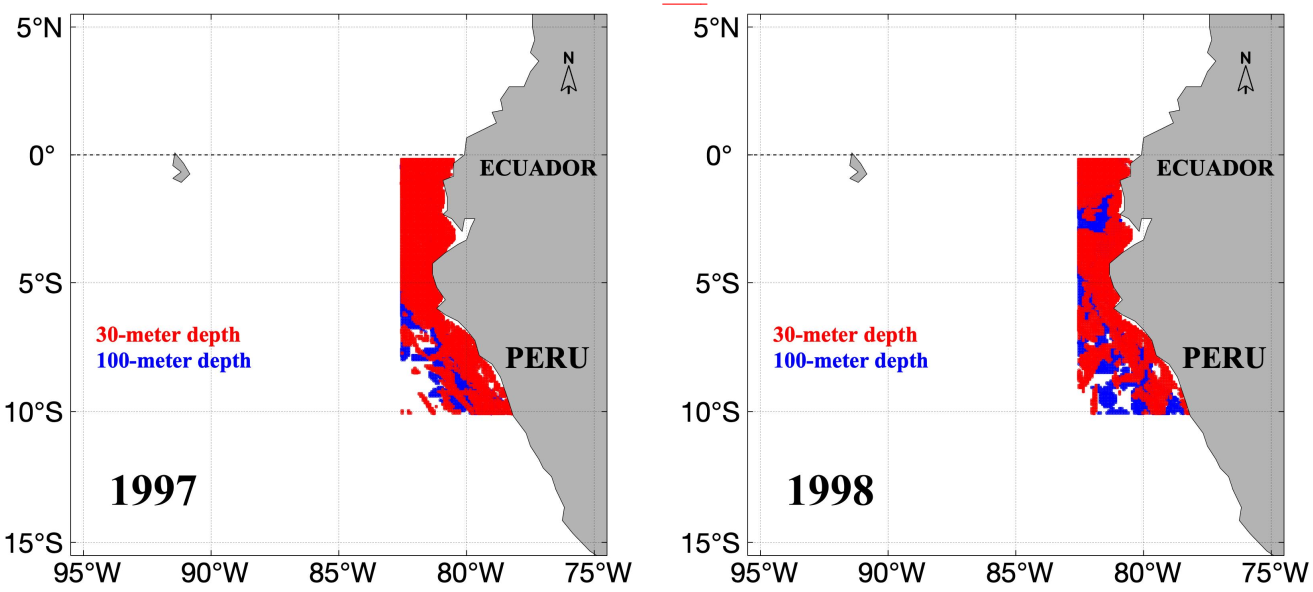

Figure 1. ~~Vertical profiles~~ Representation of ~~zonal~~ particles “seeded” (initial positions) off Peru and ~~meridional components of current~~ Ecuador at 30-m depth (red dots) and 100-m depth (blue ~~respectively~~dots) ~~averaged over 1988-2007,~~ for upwelling rates exceeding 1.9×10^{-5}m s^{-1} (50 m per month) in the ~~NEMO-ORCA12 hindcast and from NOAA mooorings~~region bounded by 82.5ºW, ~~on~~ 10ºS and the Equator, in this particular case for 1997 and 1998 years. Particles were released at ~~170ºW~~the end of each month December-January, ~~140ºW~~ for each year, over the period 1988 to 2007. Consider that initial particles here represented as blue and ~~110ºW~~red dots are overlapping each-other, although, 30-m released depth is in the front while 100-m released depth is in the background respectively.

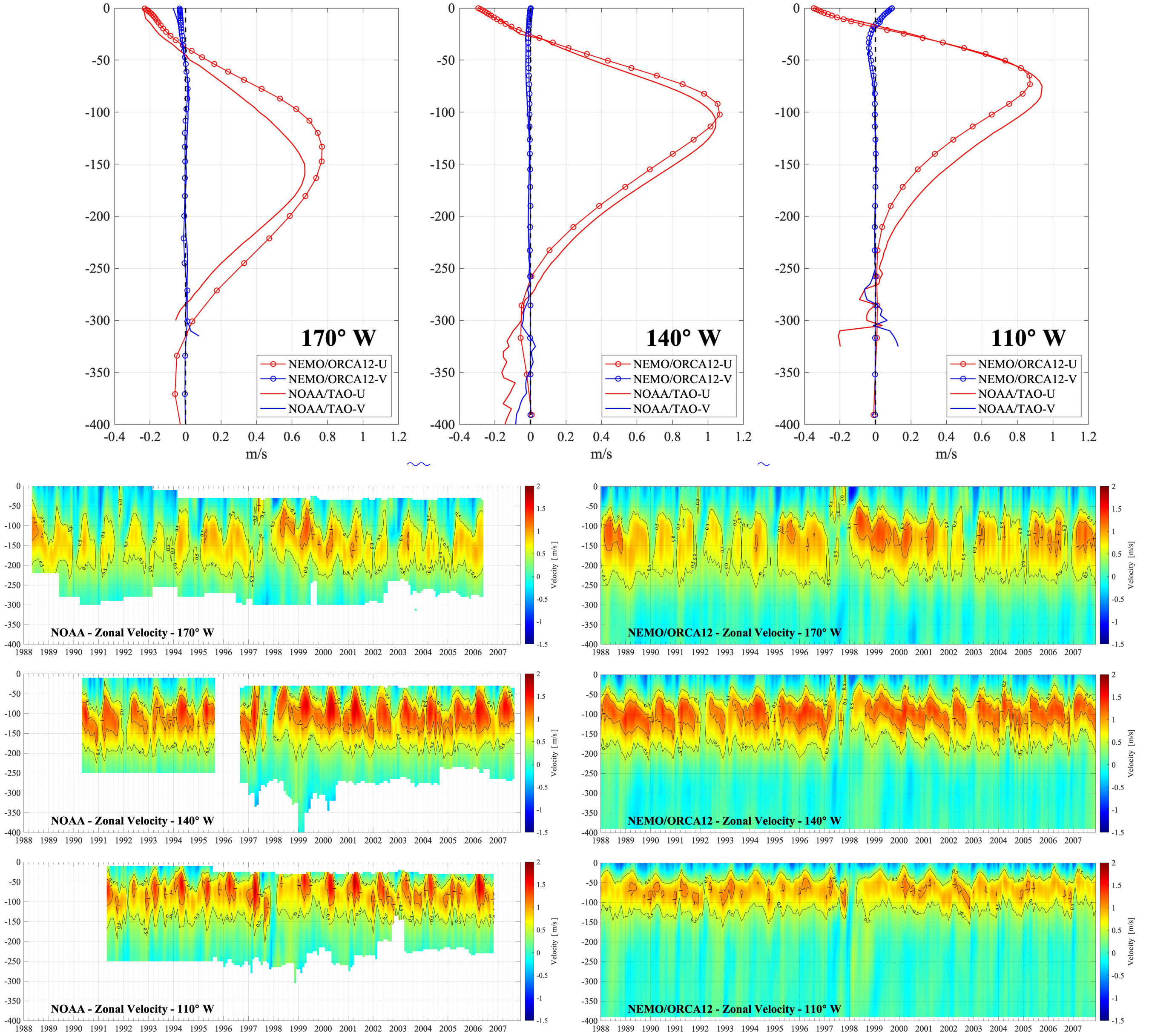

Figure 2. ~~Monthly~~ Vertical profiles for zonal ~~component of current along the equatorial Pacific at 170ºW, 140ºW~~ and ~~110ºW~~meridional velocity components (red and blue respectively) averaged over 1988-2007 (first top panel), ~~from the NEMO-ORCA12 hindcast~~ and ~~NOAA/TAO in-situ data. Contour~~ monthly vertical sections for zonal velocity component (bottom panels, contour lines highlight 0.5 ~~m s~~m.s^{-1} and ~~1 m s~~1.0 m.s^{-1}) in the NEMO-ORCA12 hindcast and NOAA mooring at 170ºW, 140ºW and 110ºW longitudes respectively.

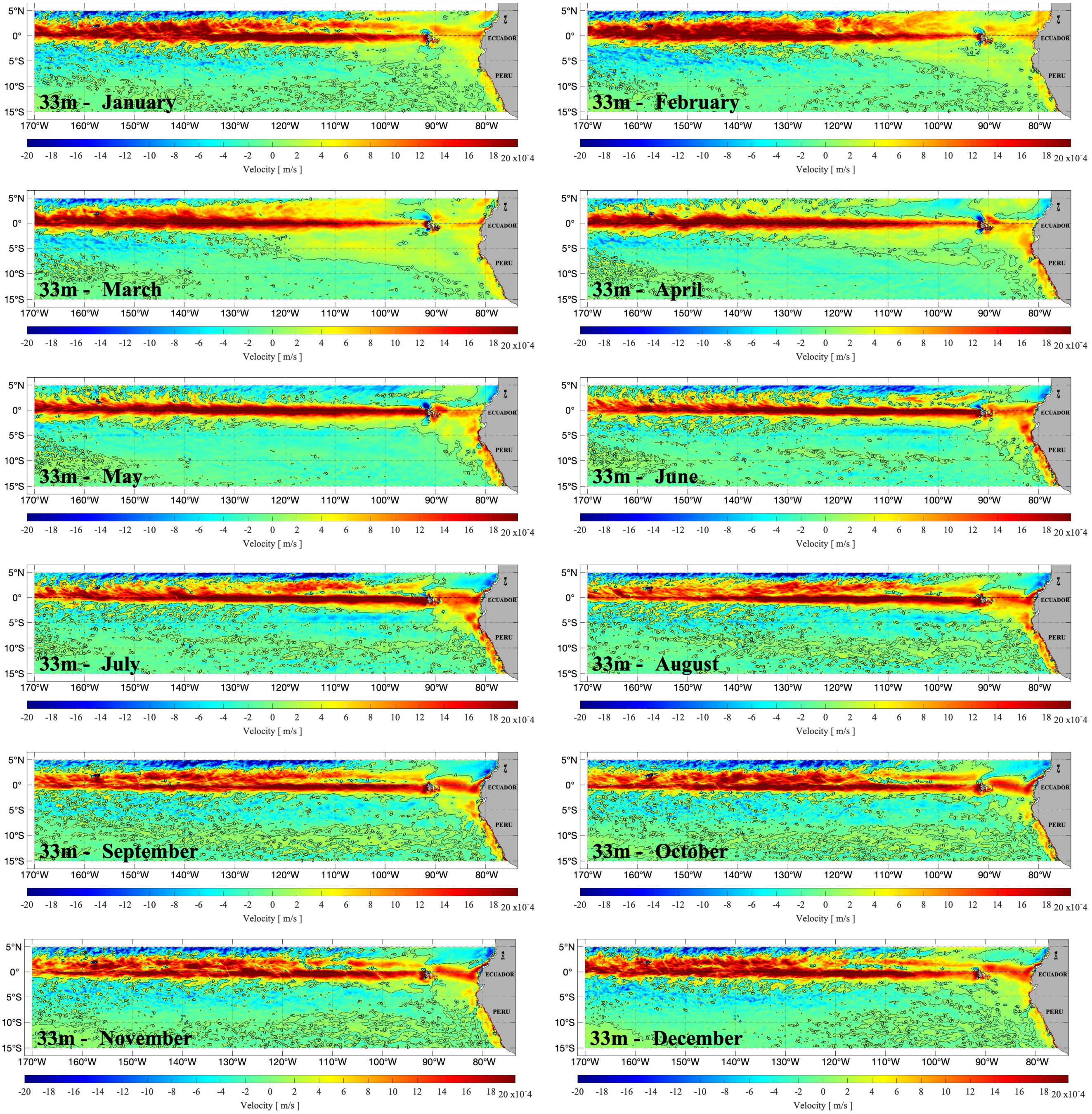

Figure 3. Monthly climatology (averaged over 1988-2007) of vertical velocity at a depth of 33 ~~m~~ m in the eastern Pacific, ~~in the~~ from NEMO-ORCA12 hindcast.

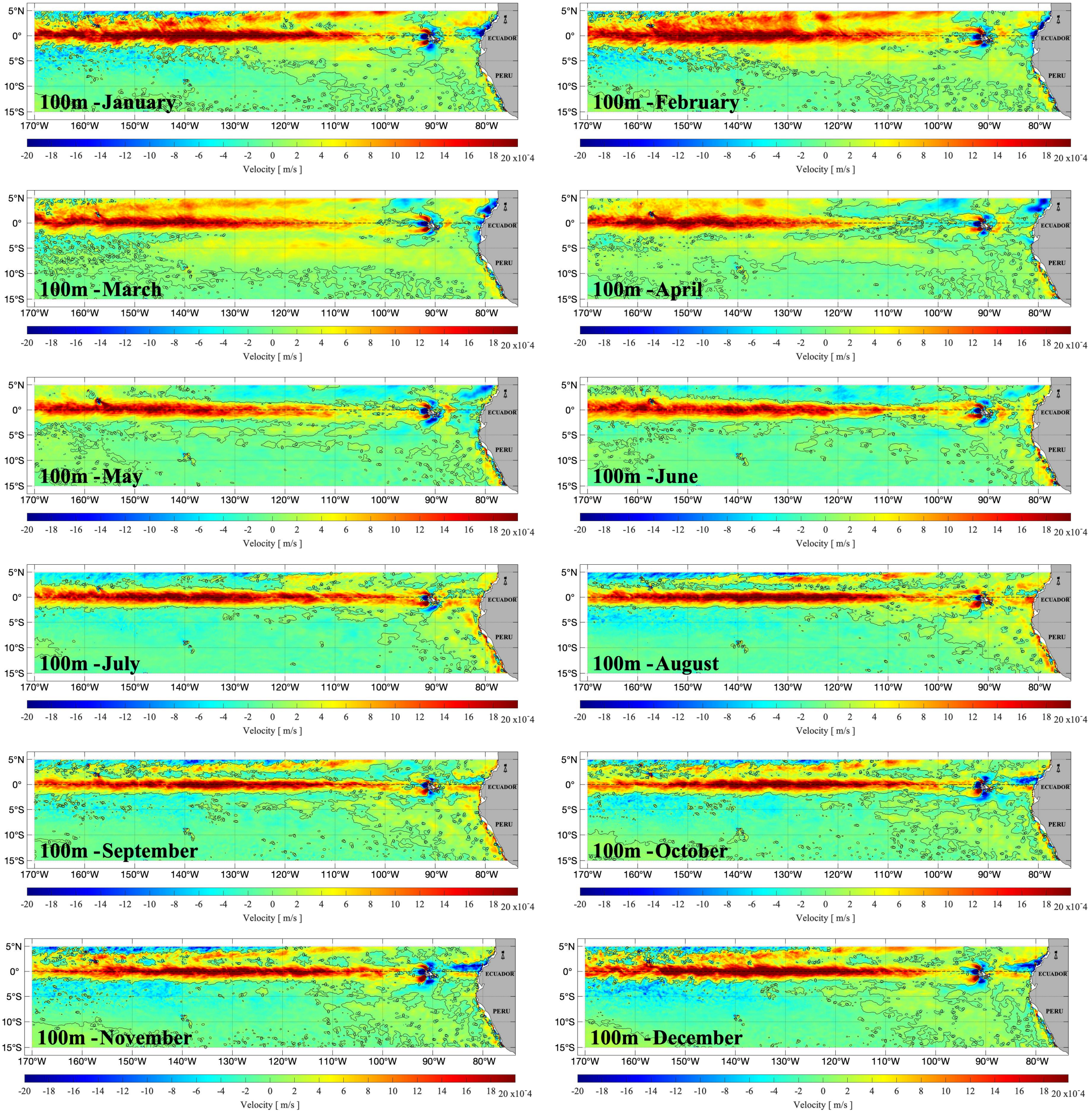

Figure 4. Monthly climatologies (averaged over 1988-2007) of vertical velocity at a depth of 100 ~~m~~ m in the eastern Pacific, from NEMO-ORCA12 hindcast.

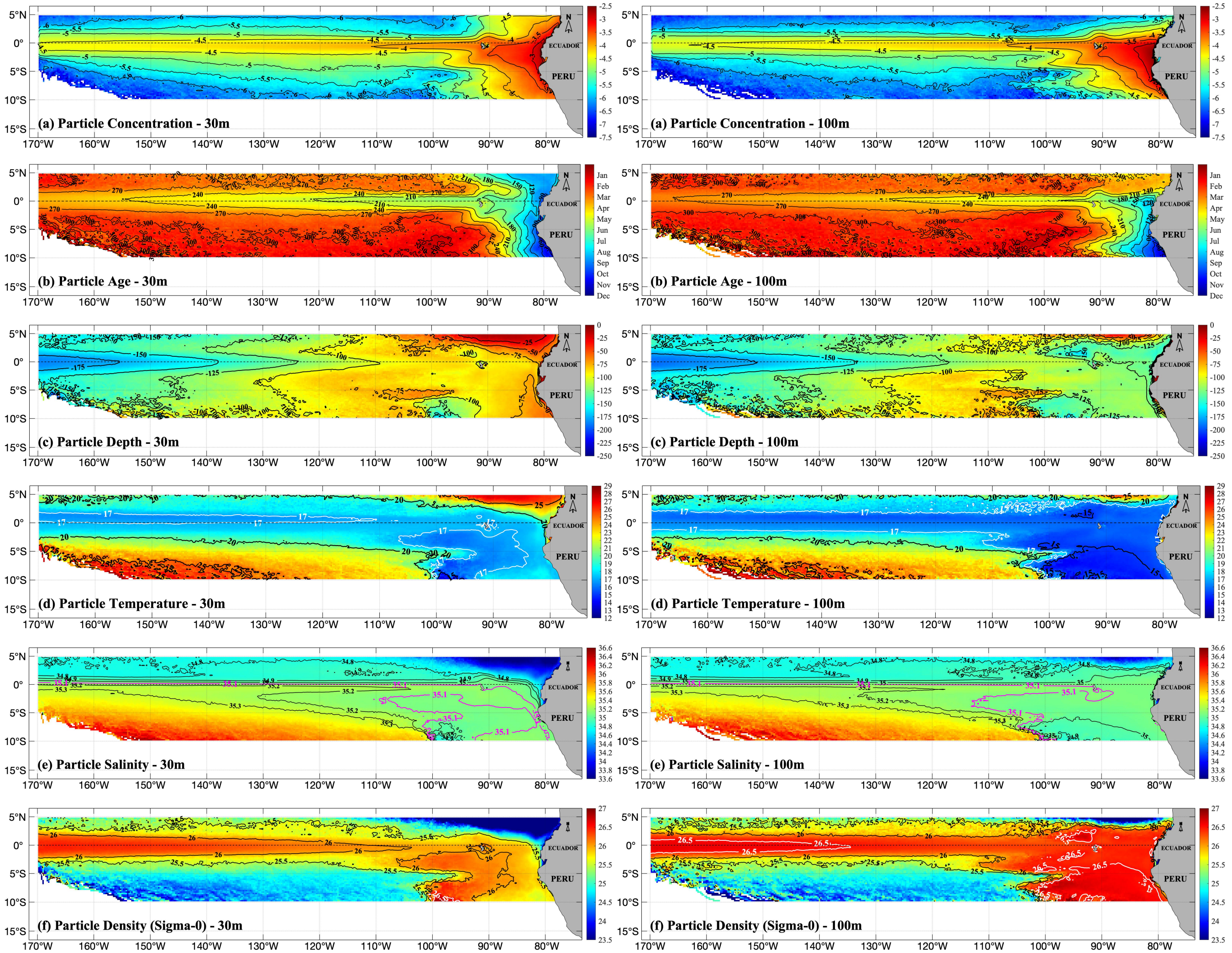

Figure 5. ~~Integrated EUC transport based on calculations at 160ºW~~Grand ensemble of back-trajectory data (averaging the 1989-2007 ensembles), ~~150ºW, 140ºW, 130ºW, 120ºW, 110ºW~~ for particles back-tracked at the end of each month (December-January) at 30-m depth (left panel) and ~~92ºW~~ 100-m depth (~~Galapagos Islands~~right panel) respectively, ~~in the latitude range 3ºN-3ºS~~binning at 0.5º x 0.5º resolution: (a) ~~monthly averages~~particle concentration (calculated as a fractional particle presence in Log10 scale, by dividing the number of particle occurrence passing through each grid cell by the total number of particles occurrences during the course of the year); (b) ~~monthly anomalies~~particle age (days, backward in time); (c) particle depth (m); (d) particle temperature (ºC), (e) particle salinity (psu); (f) particle potential density (σ_θ, kg.m^{-3}).

Monthly Climatology	160ºW	150ºW	140ºW	130ºW	120ºW	110ºW	95ºW	92ºW
January	~~26.0~~ 25.3	~~30.9~~ 28.9	~~33.6~~ 30.7	~~33.8~~ 30.1	~~25.6~~ 24.9	~~15.9~~ 18.4	~~12.4~~13.5	12.1
February	~~30.3~~ 28.3	~~27.2~~ 27.4	~~21.3~~ 25.3	~~20.2~~ 23.1	~~19.4~~ 21.3	~~15.0~~ 17.2	~~6.8~~13.2	12.2
March	~~47.0~~ 44.2	~~47.6~~ 45.0	~~43.9~~ 39.1	~~36.3~~ 30.8	~~24.0~~ 22.1	~~19.0~~ 17.1	~~10.7~~12.6	11.7
April	~~57.1~~ 53.5	~~58.6~~ 54.0	~~55.2~~ 49.9	~~44.1~~ 43.8	~~34.4~~ 33.6	~~30.1~~ 26.9	~~23.7~~22.2	18.7
May	~~58.6~~ 56.1	~~55.8~~ 51.3	~~43.2~~ 44.0	~~36.8~~ 37.6	~~30.2~~ 31.0	~~27.1~~ 26.1	~~17.8~~18.3	14.6
June	~~45.0~~ 45.2	~~42.0~~ 42.5	~~36.2~~ 37.0	~~31.6~~ 32.5	~~27.5~~ 28.0	~~23.5~~ 21.9	~~16.3~~15.1	13.0
July	~~39.8~~ 39.2	~~35.8~~ 34.7	~~30.2~~ 30.0	~~23.9~~ 25.1	~~20.7~~ 20.6	~~19.1~~ 18.5	~~5.9~~11.3	9.5
August	~~31.4~~ 31.9	~~29.1~~ 29.7	~~25.6~~ 27.2	~~23.0~~ 25.1	~~21.9~~ 21.6	~~18.2~~ 17.1	~~12.6~~11.6	10.3
September	~~26.2~~ 26.1	~~25.8~~ 27.4	~~27.2~~ 29.1	~~27.4~~ 28.3	~~31.8~~ 27.1	~~27.6~~ 23.5	~~17.3~~ 14.0	12.4
October	~~25.5~~ 26.5	~~27.6~~ 27.8	29.3	~~30.4~~ 29.2	~~26.9~~ 27.1	~~18.6~~23.6	16.5	13.8
November	~~29.0~~ 28.5	~~35.4~~ 32.6	~~38.1~~ 34.2	~~37.6~~ 33.0	~~34.0~~ 30.5	~~26.2~~ 22.9	~~7.7~~11.8	10.4
December	~~32.6~~ 29.1	~~38.5~~ 33.4	~~43.0~~ 35.8	~~40.6~~ 34.1	~~36.8~~ 31.1	~~30.2~~ 23.4	~~18.4~~13.0	11.2
~~AVERAGE~~ Annual average	~~36.3~~ 36.2	~~36.4~~ 36.2	~~34.6~~ 34.3	~~31.5~~ 31.0	~~26.8~~ 26.6	~~21.5~~ 21.4	~~12.6~~14.4	12.5

Table 1. Monthly Climatology (from 5-day ~~averaging between~~ averaged data from 1988 to 2007) and annual average for the EUC transports (~~Sv~~Sv) at 160ºW, 150ºW, 140ºW, 130ºW, 120ºW, 110ºW, and 92ºW.

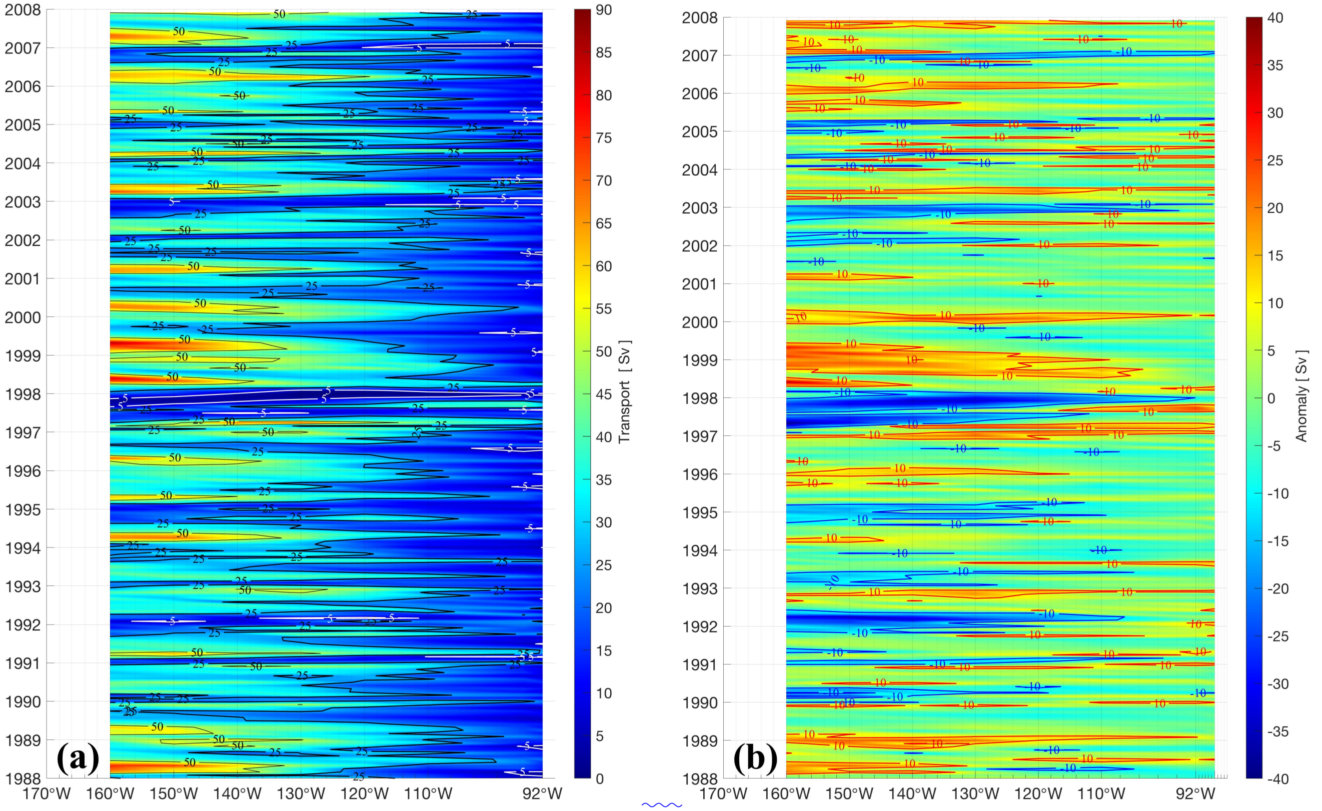

Figure 6. Integrated EUC transport based on calculations at 160ºW, 150ºW, 140ºW, 130ºW, 120ºW, 110ºW, 95ºW and 92ºW (Galapagos Islands), in the latitude range 3ºN-3ºS where the concentration of particles backtracked is the highest: (a) monthly averages; (b) monthly anomalies.

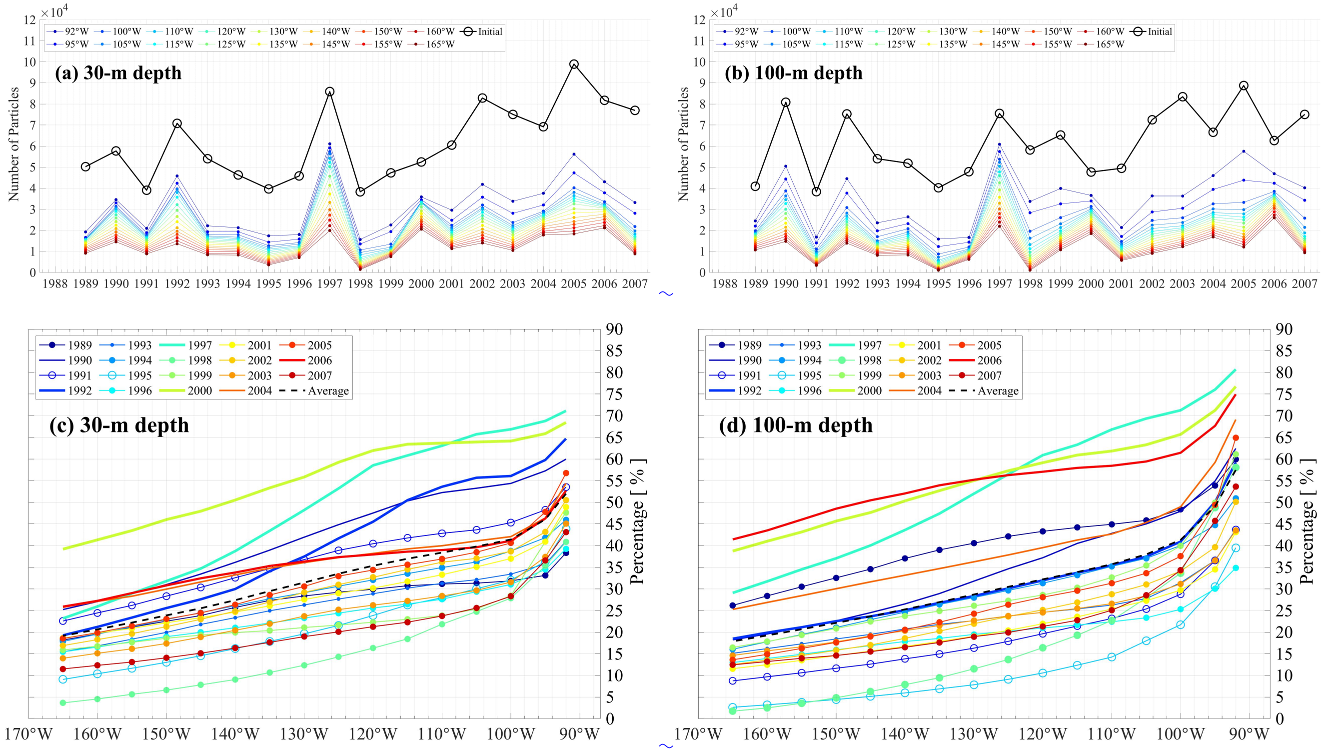

Figure 7. Annual number (from December-January monthly releases) of particles and percentage associated to the EUC (3ºN-3ºS) at 160ºW, 150ºW, 140ºW, 130ºW, 120ºW, 110ºW, 100ºW and 92ºW, for 30-m releases in (a) and (b), and for 100-m releases in (c) and (d).

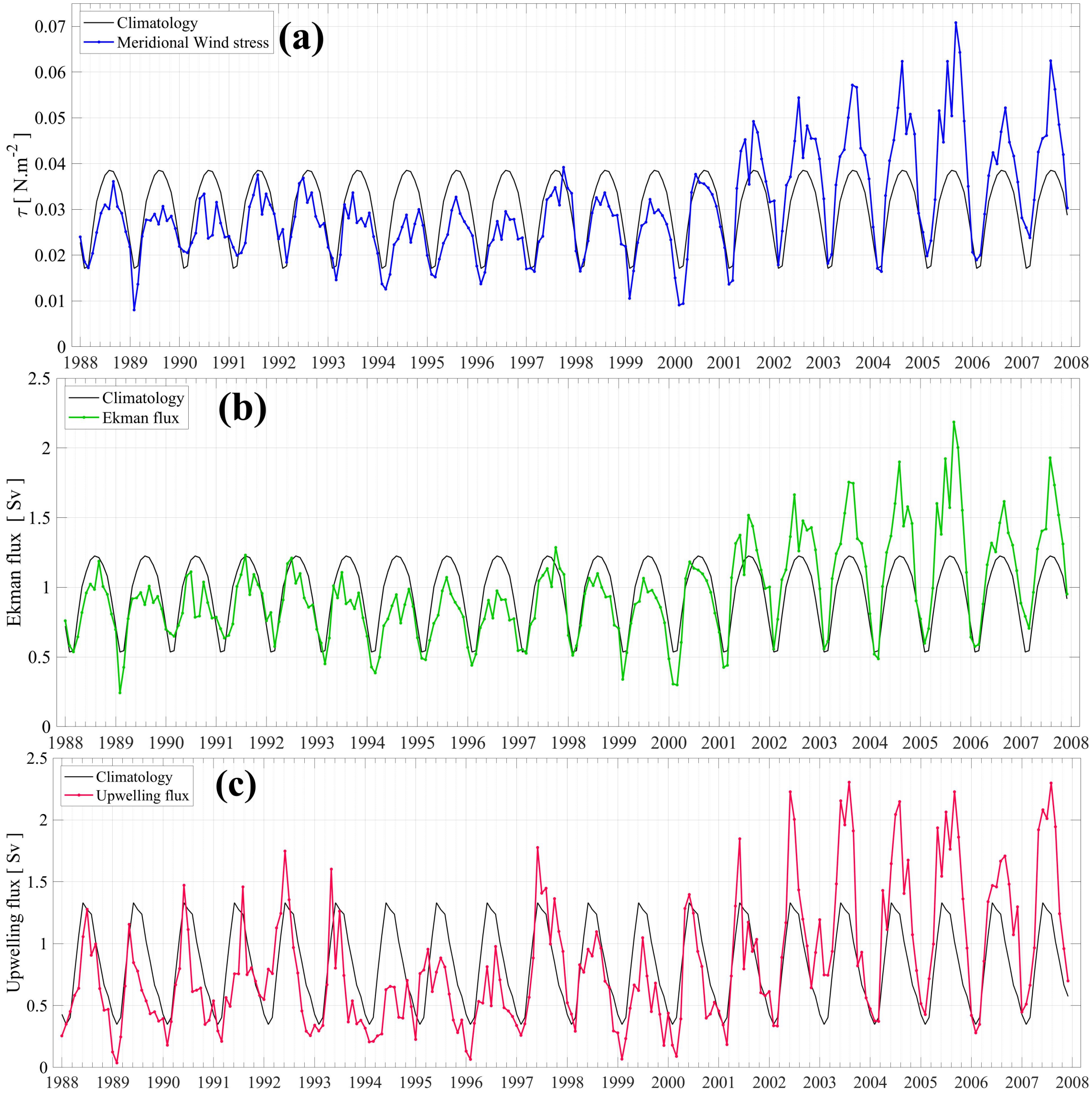

Figure 8. Diagnostics of upwelling in the north Peruvian Upwelling System, spanning the latitude range 5-10ºS: (a) equatorward wind stress $(\mathrm{N.m^{-2}})$: (b) Ekman flux (Sv); (c) Upwelling flux (Sv). Climatologies are obtained by averaging across the 1988-2007 time series.

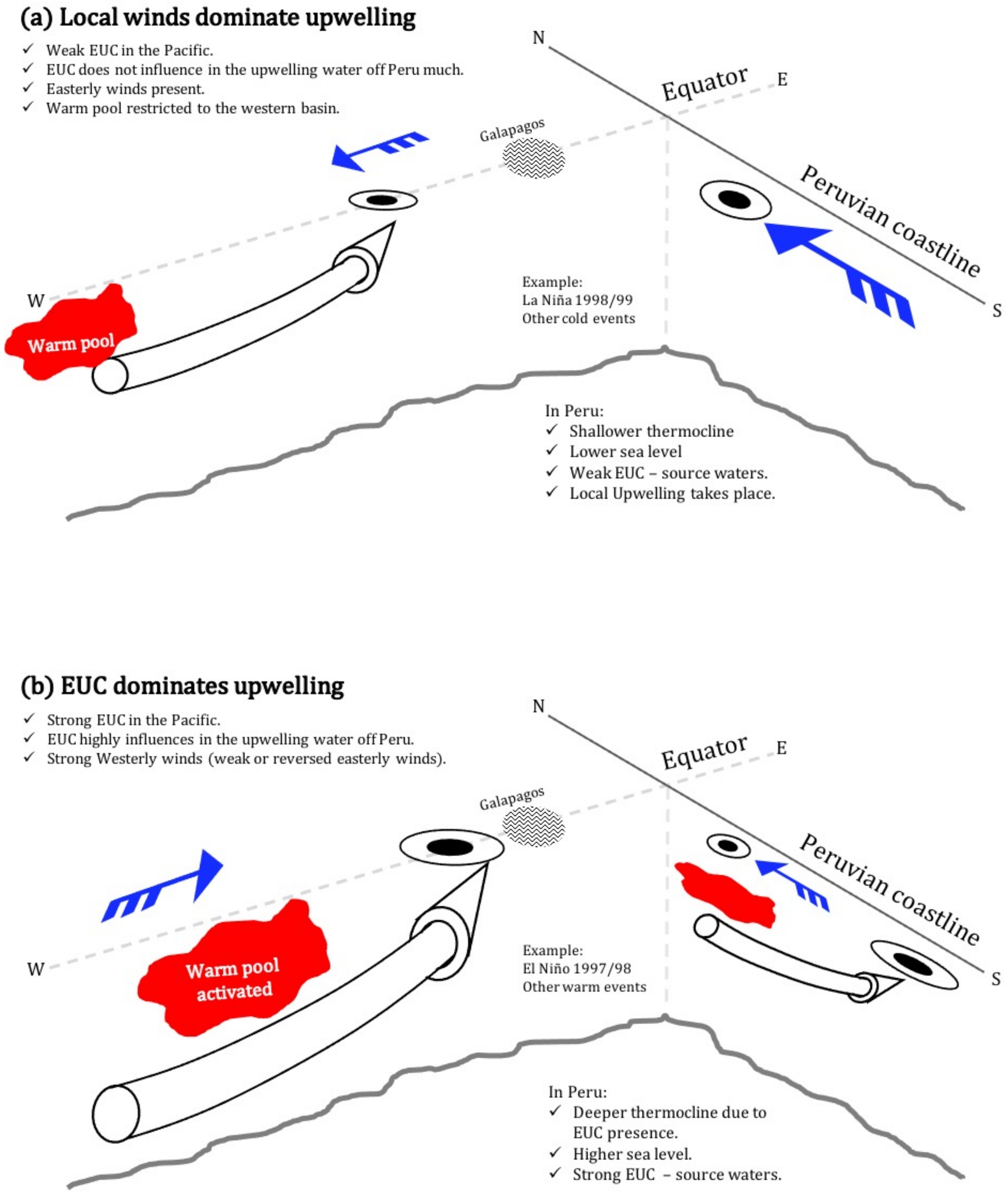

Figure 9. Schematic drivers and processes associated with the Peruvian upwelling system: (a) dominated by local winds; (b) dominated by EUC source waters.

R	160ºW	150ºW	140ºW	130ºW	120ºW	110ºW	92ºW
160ºW	1.00	0.79	0.53	0.27	N/S	N/S	-0.17
150ºW	0.79	1.00	0.71	0.45	0.23	N/S	-0.14
140ºW	0.53	0.71	1.00	0.68	0.42	0.19	-0.11
130ºW	0.27	0.45	0.68	1.00	0.68	0.39	N/S
120ºW	N/S	0.23	0.42	0.68	1.00	0.70	0.20
110ºW	N/S	N/S	0.19	0.39	0.70	1.00	0.43
92ºW	-0.17	-0.14	-0.11	N/S	0.20	0.43	1.00

Table 2. Correlation coefficient (R) matrix at 0-lag for the EUC transport at the selected longitudes, for R values with p-values <0.01 (99 % confidence interval). Not significant R values are indicated as N/S.