Ocean Sci. Discuss., https://doi.org/10.5194/os-2020-72-AC2, 2020 © Author(s) 2020. This work is distributed under the Creative Commons Attribution 4.0 License.



OSD

Interactive comment

Interactive comment on "Norwegian Sea net community production estimated from O_2 and prototype CO_2 optode measurements on a Seaglider" by Luca Possenti et al.

Luca Possenti et al.

l.possenti@uea.ac.uk

Received and published: 9 December 2020

We would like to thank Reviewer 1 for providing constructive and insightful comments. We will incorporate their suggestions into a revised manuscript. Reviewer 1's comments have been reproduced below in black, with the authors' response in blue.

General comments

Parameterization for deriving phosphate and silicate concentrations along the glider track from 'spot' samples collected during four cruises over the deployment (March, May, June, and October). Sampling restricted to the southern half of the transect. And





yet, the uncertainties were only 1.3 and 0.13 umol kg-1 for silicate and phosphate? I hope this parameterization is discussed in detail (in the text or an appendix). I also hope that some sensitivity analysis was completed regarding the impact of differing nutrient concentrations (within a reasonable range for the region & study period) on CO2SYS calculations. I'm also concerned about the use of chloroform to preserve nutrient samples.

The reviewer's comment highlighted that our method description has been too brief on nutrient analysis. In fact, we collected 58 discrete samples along the glider transect in March, May, June and October, which were analysed for total alkalinity (AT), dissolved inorganic carbon (CT) as well as nutrient concentrations. In addition, we used 52 CT and AT discrete samples collected at a weather station (OWSM) located 270 km northeast of the glider transect (Figure 1 of the discussion paper) in March, May, June, August and November. For the latter samples, no nutrient measurements were undertaken. Instead, we filled in these gaps with the nutrient concentration nearest in time and depth from the 58 along-transect samples, which we expected to be sufficiently close to the actual concentration at OWSM. We then derived the CO2 concentration (c(CO2)) using the MATLAB toolbox CO2SYS (Van Heuven et al., 2011), using the phosphate and silicate concentrations to account for their contributions to AT. This c(CO2) was used as reference to calibrate the glider output (Figure 1, in red the samples collected along the transect and in black at OWSM). The nutrient concentrations were not used in any further calculations. To assess the uncertainty of the final silicate and phosphate concentration we calculated the uncertainty in the calculation of the OWSM c(CO2) using the interpolation uncertainty 1.1 and 0.12 μ mol kg-1 for the silicate and phosphate concentration, respectively. The uncertainty was calculated as the root mean square error of the interpolation, which estimates the standard deviation of the error distribution. See the interpolation for phosphate and silicate in the supplement PDF. The derived nutrient concentration led to a mean error of 0.04 μ mol kg-1 in the calculation of c(CO2). We followed the procedure described by Hagebo and Rey (1984) to collect and preserve the nutrients. We are not aware of any problems

OSD

Interactive comment

Printer-friendly version



with the use of chloroform to preserve nutrient samples. The 2019 GO-SHIP Repeat Hydrography Nutrient Manual only discourages the use of acid (which would require neutralisation before analysis) and mercuric chloride (a long-term environmental hazard) (Becker et al., 2020).

For the lag correction to the CO2 optode, data from the glider ascents are compared against those from descents. However, there is significant horizontal distance between a glider ascent and descent, unlike what one might expect for a CTD cast from a ship. By minimizing the differences observed between glider ascents & descents, you are loosing information and I'm not sure the lag correction is necessarily reliable. I would suggest comparing potential temperature and salinity in glider ascents & descents. Do they match? If so, then perhaps this method is OK. If not, the authors may need to revise the lag correction method.

The CO2 optode lag was corrected using the algorithm of Miloshevich, (2004) that uses the sensor response time (τ). The τ was calculated minimising the difference between each glider ascent and descent (see Figure 5 of the discussion paper). In the equation, we used the median of all the τ values (1384 s). The correction decreased the difference between the glider and descent of the raw p(CO2) from (71±30) to (21±26) μ atm. Using the median τ is a robust indicator for the lag time, even there were short-term variations (e.g. due to internal tides or waves) between the descent and ascent. To assess if the glider was in the same water mass between ascent and descent we looked at potential temperature and salinity. The mean difference between the descent and ascent was (0.13±0.33) °C for potential temperature and 0.02±0.04 for salinity. These small mean differences for potential temperature and salinity show that the method is sufficiently robust.

Why is the correlation between the discrete samples and optode output CO2 partial pressure (Figure 6) so much better when using CO2 concentration vs. partial pressure (from the discrete samples)? The authors should at least offer some educated guesses or speculation.

OSD

Interactive comment

Printer-friendly version



The better correlation with c(CO2) was probably related due to an inadequate temperature-parameterisation of the sensor calibration function. The sensor output depends on the changes in pH that are directly related to the changes of c(CO2) in the membrane and – indirectly – p(CO2), via Henry's Law. Sensor and external water p(CO2) should be in equilibrium. The calibration is supposed to correct for the temperature-dependence of the sensor output (Atamanchuk et al, 2014). The observation that the sensor output correlated better with c(CO2) than p(CO2) is perhaps due to a fortuitous cancellation of an inadequate temperature-parameterisation and the temperature-dependence of the Henry's Law relationship between c(CO2) than p(CO2).

I am concerned about the potential impact of advection on the NCP calculation. The study focuses on a SE-NW transect, in a region where waters are transported in a meridional direction along well known currents (NwAC and NCC, as shown in Fig. 1). Can the authors be certain that the time rate of change in O2 and DIC does not reflect advection of water through the transect? What steps did the authors take to ensure that changes in O2 and DIC were truly a function of time and not space? Differentiating temporal vs. spatial changes in measured variables from gliders is not a trivial task and prior studies have typically used repeating spatial patterns to form a 'box' in order to compute O2 and/or carbon budgets for the estimation of NCP. In this study, the glider did not survey a box but a transect in a region of potentially meandering currents and a frontal region separating two water mass regimes. The authors need to do a better job justifying their methods and eliminating (or at least minimizing) doubt that spatial variations and/or advection contribute significantly to the observed changes in oxygen over the study period.

We have assumed that the main processes controlling the surface dissolved inorganic carbon and oxygen concentrations are biological production and respiration as well as air-sea gas exchange and vertical transport. Even though there are well-known currents, horizontal gradients are reduced due to constant stirring from winds and tides

OSD

Interactive comment

Printer-friendly version



and therefore net advective fluxes are likely to be small (Gislefoss et al., 1998; Falck and Gade, 1999). Previous estimates of net community production in the Norwegian Sea have also neglected advective fluxes (Falck and Gade, 1999; Skjelvan, Falck, Leif G. Anderson, et al., 2001; Falck and Anderson, 2005; Kivimäe, 2007). For example, Gislefoss et al, (1998) considered minimal the effect of horizontal advection on N(CT) during the summer because CT changes were largely controlled by biology and airsea interactions. For NCP estimates on shorter timescales (days to a few weeks). advective fluxes and water-mass movement would have to be taken into account, but this would require a different survey design, involving multiple platforms (Alkire et al., 2014), beyond the scope of the present study. However, the query from the reviewer prompted us to revisit our NCP calculation, which showed that at the glider turn-around points, inventory changes were calculated over relatively short time-scales of a few days. Therefore, to minimise the effect of horizontal advection in the new version of the manuscript we will extend the time interval used to calculate the inventory changes from less than a week to an average of 50 days. This was achieved by calculating the concentration difference between two transects when the glider moved in the same direction (e.g. transects 1-3, 2-4 and 3-5 all in N-S direction) instead of two consecutive transects.

The authors indicate a separation of the NCP calculation based on water masses with a cutoff at S = 35 that distinguishes between the two primary water masses influencing the study area: Norwegian Atlantic Current (NwAC) water and the Norwegian Coastal Current (NCC) water. It is also stated that salinities between 32 and 34 were encountered in the top 50 m, signifying influence of NCC water. I'm curious whether the authors took mixing into account between the two water masses in the region where NCC was encountered. How might this impact the NCP calculations? Also, I would have appreciated more information regarding the separation. Was NCP calculated separately for each of the two regions? Were they then averaged together to present a single NCP number for O2 and DIC?

OSD

Interactive comment

Printer-friendly version



The daily value of the net community production (N) was calculated separately for NCC and NwAC. The annual N was then calculated combining the two water masses to be consistent with the previous studies (Falck and Gade, 1999; Skjelvan, Falck, Leif G Anderson, et al., 2001; Falck and Anderson, 2005; Kivimäe, 2007). Here again, the reviewer's comment inspired a revision of our calculation method and for the new version of the paper, we will calculate daily and annual N without separating NCC from NwAC. We change the methodology to be consistent with previous studies, to extend and homogenise the time difference used to calculate inventory and entrainment and to minimise the impact of horizontal advection.

Integration of oxygen & DIC over a specific depth range for the calculation of NCP may be subject to vertical heaving of isopycnals. What steps did the authors take to ensure that such vertical displacement did not impact the calculations? What about vertical mixing from the bottom up? The authors calculate an entrainment flux that focuses on periods when the mixed layer depth exceeded the limit of integration (45 m), but do not discuss the possibility of mixing across the bottom boundary. Admittedly, this is probably minimal, unless there were periods of isopycnal heaving (which looks probable, from the temperature distribution shown in Fig. 8), but the possibility should have been investigated and (at least briefly) mentioned in the manuscript.

We thank the reviewer for the interesting suggestion. Vertical heaving has an effect entraining the water from below the integration. We do already consider the effect of entrainment in the calculation of the net community production, in the form of terms E(O2) (Eq. 10) and E(CT) (Eq. 16). In response to the reviewer's comment, we have also estimated the diapycnal mixing flux and will incorporate this into the new version of the manuscript. The diapycnal mixing flux (Fv) was calculated from the vertical oxygen concentration gradient. In the calculation, we used a vertical eddy diffusivity (Kz) of 10–5 m s-2 derived for the Nordic Seas by Naveira Garabato et al. (2004). The effect of Fv for O2 and CT was calculated at zmix when it was deeper than the integration depth zlim and at zlim when zmix was shallower than zlim. See the

OSD

Interactive comment

Printer-friendly version



equations in the supplement PDF. A positive sign of Fv(CT) means a decrease of the dissolved inorganic concentration in the layer of interest between surface and zlim; a negative sign corresponds to an increase. In the new version of the manuscript, we will add a new Figure 2 that shows Fv as a function of time during the glider deployment. The results show that Fv is negligibly small: $Fv(CT) = (0.05\pm0.3)$ mmol m-2 d-1 and (-0.02 ± 0.33) mmol m-2 d-1 for O2. For that reason, diapycnal mixing will not be used to calculate N.

Specific comments

Please clarify units of N(CT) and N(O2). Are they both expressed as mmol C m-2 d-1 or do they differ (e.g., mmol C m-2 d-1 vs. mmol O2 m-2 d-1)? After getting to section 3.6, it's clear they were reported in different units, but readers shouldn't have to wait that long to be sure.

Both N(O2) and N(CT) are expressed in mmol m-2 d-1. In the case of N(O2), this is a flux of O2, in the case of N(CT), a flux of inorganic carbon.

Preservation of nutrient samples with chloroform is not a recommended procedure. . .

See our answer on page 1.

Figure 2 indeed shows that, on average, the oxygen concentration at higher latitudes was greater (by 10-15 umol kg-1) than those measured at lower latitudes. However, the oxygen concentration decreases fairly linearly with time in both regions (lower and higher latitudes). Why is this the case? I wouldn't think it was short-term drift as such drift should be minimal in oxygen optodes. Does this results, perhaps, from a longitudinal gradient in oxygen concentrations? Figure 3 shows a similar 'drift', or time rate of change, in the gain factor computed to correct the optode oxygen. I am surprised there is such an apparent, continued drift in the optode sensor response. I would have expected a large, initial drift ('storage' drift') but then would have thought the optode response to be relatively stable over a deployment period of around 8 months. Can the

OSD

Interactive comment

Printer-friendly version



authors show the individual, median oxygen concentrations and standard deviations from the discrete data? I'm curious how stable the oxygen concentrations are in this density/depth range (around 427 to 1000 m).

The oxygen concentrations for $\sigma 0 > 1028$ kg m–3 decreased linearly in both regions because the oxygen optode drifted continuously during the deployment (Figure 2 of the paper and Figure 3 where in red is the uncorrected oxygen, in blue the corrected oxygen and in yellow the discrete sample used as reference). In the new version of the manuscript, we will add a figure with all the samples collected and the glider data before and after the correction showing how the corrected glider oxygen is within the variability of the discrete samples and how stable the O2 concentration is in this depth range. It was possible to use waters of these potential densities because were always well below the mixed layer depth and therefore subject to limited seasonal and interannual variability. The salinity of the discrete samples varied from 34.88 to 34.96, with a mean of (34.90 ± 0.01) and the temperature varied from 0.45 to -0.76 °C with a mean of (-0.15±0.36) °C. Variations are due to differences in deep-water masses. Therefore, we only used the glider and discrete samples collected at latitudes north of 64° N because this reflects the largest part of the transect. Also, the region south of 64° N contained just 5 days of archived samples. See also reply 2 to Reviewer 2. We added in the appendices a plot with all the discrete samples and the glider oxygen before and after the correction.

Line 269: "The thermal lag of the glider conductivity sensor was corrected for. . ." What?

The correct phrase should be "The thermal lag of the glider conductivity sensor was corrected using the method of Gourcuff (2014)."

Can the authors please define cN(Chl a)? Is this the computed chlorophyll concentration, using factory-defined coefficients?

Yes, in the conversion from the raw chlorophyll to the chlorophyll concentration, we

OSD

Interactive comment

Printer-friendly version



used the factory-based coefficients.

Line 363: ". . . because after this dive, the CO2 optode stopped sampling. . ."

We meant "For the subsequent dives, the CO2 optode stopped sampling in the first 150 m (Figure 2.8d)."

Line 364: ". . .raw c(O2) data was calibrated and drift-corrected and c(CO2) was driftand lag-corrected and recalibrated, then used to. . ." I'm not going to focus my review on grammar corrections, so I suggest the authors carefully re-read the manuscript to avoid any additional grammar or spelling mistakes that should be addressed prior to publication.

Apologies if the sentence structure was unclear. We meant to say that "The raw c(O2) data were drift-corrected and calibrated. The CO2 output was drift and lag-corrected and then calibrated against cC(CO2) from nearby discrete samples. The calibrated glider cG(O2) and cG(CO2) were used to calculate inventory changes and air-sea exchange fluxes (Φ) to evaluate the net community production changes."

Plot isopycnals on panels of Fig. 8. I'd also recommend plotting the mixed layer depth and highlighting zlim (dotted line?).

We changed figure 8 adding the mixed layer depth, zlim and the isopycnals (Figure 4).

Line 375: What is "against year-day"? Please re-word this sentence.

Year-day means day of the year and varies from 1 to 365. We will change all occurrences of year-day in the manuscript to "day of the year".

Lines 456-457: Can the authors please expand on how NCP was calculated? It is stated that, "The two Ns were calculated as the difference in inventory changes between two transects when the glider was in the same water mass." Two transects? So, is one transect equivalent to the glider moving over the entire transect in one direction and the second transect is the glider moving back over the transect in the opposite di-

Interactive comment

Printer-friendly version



rection? Is the NCP calculated only for the NwAC water mass? So any changes within the NCC water mass are removed from the analysis?

Yes, it is correct, that we used one transect with the glider moving in one direction and the following transect with the glider moving in the opposite direction. To calculate the net community production (N) the data were binned into 0.1° latitude intervals and the inventory changes were calculated as the difference of the integrated c(O2) and CT every time the glider was in the latitude bin. The air-sea flux was the instantaneous flux when the glider was in the bin and the entrainment was considered as the concentration changes when the mixed layer deepened between two transects in the same latitude bin. The daily N was calculated separating the two water masses (NCC and NwAC) and the annual N was calculated as the mean of the daily N considering the two water masses together. Following the prompt for the reviewer on the possible influence of horizontal advection (see above), the revised version of the manuscript will use an amended methodology to calculate net community production will. We will use the difference of CT and O2 between two transects when the glider moved in the same direction (e.g. southeast to northwest). We will not use two consecutive transects. This means that inventory changes will be calculated based on a similar time difference between the two samples. For two consecutive transects, the time difference between the two samples would be smaller at the beginning and the end of the transect and larger in the middle. Also, to correct for the variability of the wind speed, we will use flux-weighted gas transfer velocities for O2 and CO2 (Reuer et al, 2007), rather than instantaneous fluxes (as before). kw(O2) and kw(CO2) will be normalised using the daily wind speed in the latitude bin in the time interval used to calculate the inventory changes. The time interval is the time between two samples used to calculate the inventory changes and entrainment. The air-sea flux is based on the concentration measured at the time of the second transect used to calculate the inventory changes and the entrainment flux.

It is important to compare NCP estimates with those of previous studies; however, it

OSD

Interactive comment

Printer-friendly version



is difficult to know how comparable the numbers are in Table 3 because it is not clear where in the Norwegian Sea these various studies took place. It is also difficult because zlim varies largely among the studies. The fact that three of the four compared studies used zlim >= 100 m also calls into question why exactly the current study decided on zlim = 45 m, particularly since the mixed layer depth varied so largely and often exceeded zlim.

We used zlim = 45 m because this corresponds to the average depth of the euphotic zone, which is the region of interest for net community production from a biogeochemical and ecological point of view. Previous studies may have used zlim = 100 m for operational reasons (e.g. constrained by discrete sampling depths). To show the influence of zlim on N, we calculated N for zlim = 30 m and 100 m. N(CT; 30 m) was 0.6 and N(CT; 100 m) was -0.04 mol m-2 a-1. N(O2; 30 m) was 4.6 and N(O2; 100 m) was 4.3 mol m-2 a-1. In the case of N(CT), the derived two values are lower to the previous studies where N(CT) varied from 8.6 to 2.0 mol m-2 a-1. N(CT; 100 m) was negative because the deep integration depth included water below the euphotic zone where the remineralisation of organic matter can increase CT. This signal is not present in N(O2) because the changes were largely controlled by $\Phi(O2)$ that was always positive. The calculated N(O2) is in agreement with previous studies, which gave results between 2.6 and 11 mol m-2 a-1. In the discussion, we will add a section where we explain the location and the period when the previous studies took place. Falck and Anderson (2005) used historical data from 1960 to 2000 collected in the area from 62 to 70° N and from 1991 to 1994 collected at OWSM. Skjelvan et al (2001) used data collected from 67.5° N 9° E to 71.5° N 1° E and along 74.5° N from 7 to 15° E from 1957 to 1970 and 1991 to 1998. Kivimäe (2007) used the oxygen measured at OWSM from 1955 to 2005 and Falck and Gade (1999) used data collected in all the Norwegian Sea from 1955 to 1988. The glider in the transect moved from to 66.3 °N 4 °W to 63 °N 4 °E (Figure 5 where the black dots are the glider dives, the green box the region used by Falck and Gade (1999), the yellow lines the transects used by Skjelvan et al. (2001), the azure box the region used by Falck and Anderson (2005) and the red dot the loInteractive comment

Printer-friendly version



cation that corresponds to the Ocean Weather Station M (OWSM) used by Kivimäe (2007)).

Bibliography

Alkire, M. B. et al. (2014) 'Net community production and export from Seaglider measurements in the North Atlantic after the spring bloom', Journal of Geophysical Research: Oceans. Wiley Online Library, 119(9), pp. 6121–6139.

Atamanchuk, D. et al. (2014) 'Performance of a lifetime-based optode for measuring partial pressure of carbon dioxide in natural waters', Limnology and Oceanography: Methods, 12(2), pp. 63–73. doi: 10.4319/lom.2014.12.63.

Becker, S. et al. (2020) 'GO-SHIP repeat hydrography nutrient manual: the precise and accurate determination of dissolved inorganic nutrients in seawater, using continuous flow analysis methods', Frontiers in Marine Science. Frontiers, 7, p. 908.

Falck, E. and Anderson, L. G. (2005) 'The dynamics of the carbon cycle in the surface water of the Norwegian Sea', 94, pp. 43–53. doi: 10.1016/j.marchem.2004.08.009.

Falck, E. and Gade, G. (1999) 'Net community production and oxygen fluxes in the Nordic Seas based on O2 budget calculations', 13(4), pp. 1117–1126.

Foltz, G. R. et al. (2003) 'Seasonal mixed layer heat budget of the tropical Atlantic Ocean', Journal of Geophysical Research: Oceans. Wiley Online Library, 108(C5).

Gislefoss, J. S. et al. (1998) 'Carbon time series in the Norwegian sea', 45, pp. 433–460.

Gourcuff, C. (2014) 'ANFOG Slocum CTD data correction', (March).

Hagebo, M. and Rey, F. (1984) 'Storage of seawater for nutrients analysis', Fisken Hav., 4, 1, 12. Van Heuven, S. et al. (2011) 'MATLAB program developed for CO2 system calculations', ORNL/CDIAC-105b. Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, US Department of Energy, Oak Ridge, Tennessee, 530.

Interactive comment

Printer-friendly version



Kivimäe, C. (2007) 'Carbon and oxygen fluxes in the Barents and Norwegian Seas: production, air-sea exchange and budget calculations'. The University of Bergen.

Miloshevich, L. (2004) 'Development and Validation of a Time-Lag Correction for Vaisala Radiosonde Humidity Measurements', pp. 1305–1328.

Naveira Garabato, A. C. et al. (2004) 'Turbulent diapycnal mixing in the Nordic seas', Journal of Geophysical Research: Oceans. Wiley Online Library, 109(C12).

Obata, A., Ishizaka, J. and Endoh, M. (1996) 'Global verification of critical depth theory for phytoplankton bloom with climatological in situ temperature and satellite ocean color data', Journal of Geophysical Research: Oceans. Wiley Online Library, 101(C9), pp. 20657–20667.

Reuer, M. K. et al. (2007) 'New estimates of Southern Ocean biological production rates from O2/Ar ratios and the triple isotope composition of O2', Deep Sea Research Part I: Oceanographic Research Papers. Elsevier, 54(6), pp. 951–974.

Skjelvan, I., Falck, E., Anderson, Leif G, et al. (2001) 'Oxygen fluxes in the Norwegian Atlantic current', Marine chemistry. Elsevier, 73(3–4), pp. 291–303.

Skjelvan, I., Falck, E., Anderson, Leif G., et al. (2001) 'Oxygen fluxes in the Norwegian Atlantic Current', Marine Chemistry, 73(3–4), pp. 291–303. doi: 10.1016/S0304-4203(00)00112-2.

Please also note the supplement to this comment: https://os.copernicus.org/preprints/os-2020-72/os-2020-72-AC2-supplement.pdf OSD

Interactive comment

Printer-friendly version



Interactive comment on Ocean Sci. Discuss., https://doi.org/10.5194/os-2020-72, 2020.

Interactive comment



Fig. 1.

Printer-friendly version









Printer-friendly version







Fig. 3.

Printer-friendly version



OSD



Printer-friendly version







Interactive comment





Printer-friendly version

