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6 **Technical Note: Estimation of global loss of freshwater based on sea level**  
7 **changes over geological time**

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28 **Abstract**

29 Water vapour at the upper layer of the atmosphere undergoes light-dependent photolysis generating  
30 reactive hydrogen ions *in statu nascendit* that escape to space by different mechanisms. Besides  
31 hydrogen, other volatile gases such as methane molecules and helium atoms also escape to space in  
32 smaller quantities or traces such as oxygen. The escape of hydrogen through the planetary air leak  
33 cannot be reliably judged. Our estimation of global freshwater loss used another approach based on the  
34 sea level changes that continuously fluctuated over geological time. The most reliable evidence for  
35 eustatic sea level changes was provided by geologists estimating the shifts of shorelines generating  
36 sedimentary deposits. The sea level changes turned to volumetric data of a) radii of the Earth ( $r_1$ ) to  
37 calculate the volume of the geoid Earth ( $V_1$ ) comparing and validating them with available  
38 estimations, b) average sea depth ( $r_2$ ) comparing volumetric values of the sea with best-fitting values  
39 ( $V_2$ ). c) showing the correlation between geological changes ( $r_3$ ) and corresponding sea volumes ( $V_3$ ).  
40 These data, along with the sea volume of the infant Earth, allowed to plot a calibration curve to  
41 calculate the sea level belonging to the volume and *vice versa*. Geologic data indicate the  
42 shrinkage of freshwater pools during interglacial dilution periods and the remarkable long-term  
43 salination of the ocean.

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45 **Keywords**

46 water vapour, photolysis of water, H escape, freshwater reserves, sea level rises, salination of  
47 sea

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54 **1. Introduction**

55 Water vapour at the upper layer of the atmosphere undergoes light-dependent photoionisation  
56 generating reactive hydrogen ions *in statu nascendi* that escape to space by different mechanisms. The  
57 escape velocity of hydrogen to the outer space is the highest among light gases followed by helium,  
58 ammonia and water vapour and is dependent on gravity, which in the inner planets of our solar system  
59 is lowest on Venus followed by Earth and Mars. The escape of gases, among others, is dependent on  
60 the temperature, rate of cooling, greenhouse effect. It was estimated that the Sun would be brighter and  
61 hotter in about 1 billion years and the global surface temperature could rise to about 47 °C. The Earth  
62 will be likely to lose its water (Schröder et al., 2008) similarly to what already happened to Venus.

63 Losses of light gases are unlikely to be inherent parts of the evolution, at least in the inner planets of  
64 our solar system. Vast losses of gases also came from the activity of the young Sun that blew light  
65 gases, primarily hydrogen and helium to the outer planets known as gas giants. An important factor  
66 related to the global loss of water is the polar wind that drives hydrogen and other ions to space seen  
67 as faint yellow area near the Earth's poles.

68 The shrinkage of freshwater reservoirs including atmospheric moisture (snow, rain, clouds) and ice  
69 (glaciers, polar ice, ice sheets, permanent snow) is contributed by the freshwater loss of global  
70 warming. The atmospheric escape of hydrogen on Earth is assumed to result in approximately 3kg/s  
71 loss of hydrogen and about 50 g/s loss of helium (Zahnle, 2006; Catling and Zahnle, 2009). The  
72 abundant CH<sub>4</sub> originating from methanogenesis could have supported the escape of hydrogen to space  
73 by orders of magnitude faster than today (Catling et al., 2001). Our own repeated calculations showed  
74 that under the conditions that exist today, the H escape could have resulted in only about 0.02% loss of  
75 the recent ocean volume. The explanation to this negligible loss could be that the escape of water: a)  
76 was much faster earlier, b) came from different sources that could sum up or c) was not significant  
77 during the evolution of Earth.

78 Regarding oxygen, only small quantities were found to escape to the Moon from the Earth (Terada  
79 et al., 2017). O<sub>2</sub> produced by photosynthesis absorbed in oceans and seabed rocks started to gas out  
80 about 1,850 Mya but was absorbed mainly by land surfaces. When the ocean saturated, oxygen began  
81 to accumulate in the atmosphere (Holland, 2006).



82 An important warning signal of freshwater loss is the increasing gap between saltwater and  
83 freshwater, which is now 97 *versus* 3 %. The question is how this ratio could have changed from a  
84 dilute too salty and concentrate sea during evolution. The highest known sea level some 3,700-3,800  
85 million years ago, suggested that the Infant Sea could have contained 26% more water than today  
86 (Pope et al., 2012; Rosing et al., 2010). Ice ages generated low sea levels particularly during the  
87 Proterozoic Snowball Earth period when Earth's surface became almost entirely frozen and covered by  
88 slush, snow and ice (Kirschwink, 1992; Allen and Etienne, 2008; Pu et al., 2016). The opposite  
89 tendency, namely the accumulation of substantial freshwater reservoirs was assumed to take place in  
90 the late Proterozoic aeon before the Cambrian period (541 Mya) and between the Ordovician and Silur  
91 epochs (Pu et al., 2016; Haq et al., 1988; Holland et al., 1986). Severe glaciations with low sea levels  
92 were followed by extreme high sea levels, *e.g.* during the Paleozoic era at the end of the Ordovician  
93 period (~450 Mya). Although the Ordovician lasted for only 45 million years and represented only  
94 ~1% of the age of the Earth, the life on Earth during this time diversified rapidly. The unprecedented  
95 radiation of species in Ordovician is accounted for by the dilute, yet optimal osmolarity (0.2 – 0.4  
96 Osm) of the oxygenated ocean and the abundance of freshwater supply.

97 Gas escape theories did not provide reliable means to judge the global loss of water on Earth.  
98 Thus geological sea levels were turned into volumetric changes to give a reasonable explanation to  
99 the shrinkage of freshwater.

## 100 **2. Assessing freshwater loss**

### 101 **2.1. Estimation of sea volume and turning sea levels to volumetric data**

102 We used calculations of interconversions to turn sea levels and volumes into each other. The validity  
103 of the data obtained was tested by using:

- 104 *i)* Geometric radii of the geoid Earth and average radius
- 105 *ii)* Reliable estimates of average sea depth
- 106 *iii)* Comparison of volumetric estimates and selection among best-fitting values
- 107 *iv)* Testing whether volumetric changes belonging to sea levels can be used to construct a  
108 calibration curve

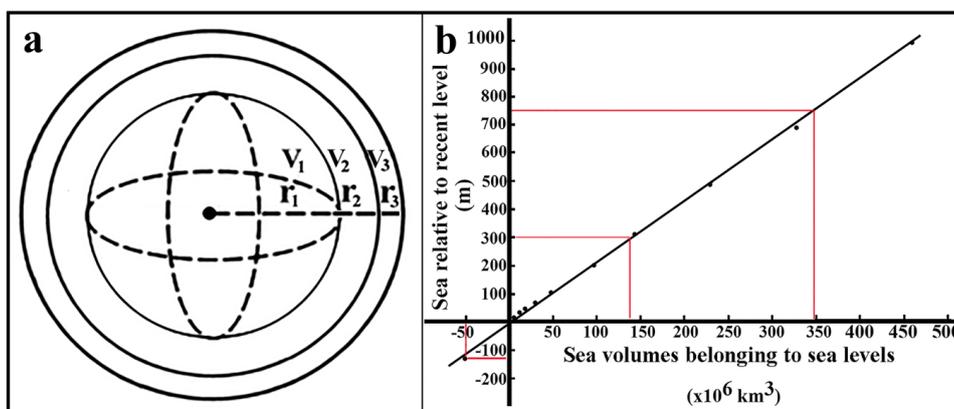


109 v) Plotting the calibration curve to show the relationship between sea level rises and volumetric  
110 changes in sea volumes (Banfalvi, 2017).

## 111 2.2 Sea volumes

112 The extension of the calibration curve included the highest known sea level of the sea of the Infant  
113 Earth (~ 750 m) relative to the recent sea level (6,371,008 m). Compared to the average radius of the  
114 Globe (6367.3 km), the average depth of the sea (3682.2 m) is negligible thus the linearity of the  
115 calibration curve is not questionable (Fig. 1). Most of the data of the sea volumes were 30-40 years  
116 old, differed from each other and were significantly higher than the recent satellite measurements  
117 (<https://www.livescience.com/6470-ocean-depth-volume-revealed.html>). One could not explain the  
118 higher estimates of current sea volumes by recent changes. Rather, the amounts of undersea  
119 mountains, ocean ridges and other geographical features were not subtracted from the bulk of the  
120 ocean resulting in higher estimates. These structures under the sea level include *i*) the Globe itself, *ii*)  
121 continental shelves, slopes, rises, insular shelves that surround continents and islands, and *iii*) zones of  
122 the ocean floor (continental margins, deep-ocean basins, mid-ocean ridges, sediments) that are merged  
123 and referred to as undersea features.

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126 Figure 1. Demonstration of sea levels and volumes (a), plotting them into a calibration curve (b).

127 a) Data used for the calculation of underwater volumes of the continental shelf and other underwater  
128 features include:  $r_1$  = average of equatorial and polar radii of geoid Earth (6367.3 km) (Banfalvi,



129 2017).  $r_2$  = radius of Earth + depth of sea (6,371,008 m) (Williams, 2012).  $V_1$  = volume of Earth  
 130 calculated from  $r_1$  ( $1,081.320 \times 10^9 \text{ km}^3$ ).  $V_2$  = volume of sea ( $1332.9 \times 10^6 \text{ km}^3$ ) (Banfalvi, 2017).  
 131  $V_3$  = actual volume arising from sea level rise ( $0.09 \text{ m} \times 361.84 \text{ km}^3 = 32.5 \text{ km}^3$ ). b) Relationship  
 132 between sea level elevations and sea volumes relative to recent data. The data of abscissa represent  
 133 eustatic sea level elevations and those of the ordinate the volumetric changes relative to recent data.  
 134 Lowest sea level (-130 m) (~20,000 years ago), higher (~300 m) (~500 million years ago) and the  
 135 highest sea level of the Infant Earth (3,700-3,800 million years ago). Modified with permission  
 136 (Banfalvi, 2017).

137 Table 1 shows the recent volumetric data of water on Earth (Table 1). The continental margin is  
 138 constituting about 28% of the oceanic area (Cook and Carleton, 2000). All the underwater features  
 139 represent an estimated 29.5 % of sea volume and will be more precisely determined as the  
 140 measurements will become even more accurate. The most reliable data obtained by satellite  
 141 measurements of sea surface face some light reflection problems (Charette and Smith, 2010). The  
 142 resolution of the sea depth needs further "fine-tuning" by ship-based echo sonar and other methods.  
 143 Satellite measurements show that the average ocean depth is 3,682.2 m. This value multiplied by the  
 144 area of sea ( $361.84 \times 10^6 \text{ km}^2$ ) results in the volume of the ocean of  $1,332.4 \times 10^6 \text{ km}^3$  (Charette and  
 145 Smith, 2010). These measurements show a good match ( $\pm 0.1\%$ ) with other values published within  
 146 the last ten years (Table 1).

147 Table 1

148 Volumetric data of water on Earth

149 #	150 Volume 151 of	152 Volume 153 $\times 10^6 \text{ km}^3$	154 Volume 155 (%)	156 References
157 1	*Sea	1332.4	96.9	Charette and Smith, 2010
158		1332.9	96.9	Banfalvi, 2017
159		1335	97.1	Eakins and Sharman, 2010
160		1335	97.1	Durack, 2015
161	2 **Sea (average)	1333.8	97.0	
162	3 † Freshwater in reservoirs	25.54	1.86	Durack, 2015
163	4 ‡ Freshwater in land	15.66	1.14	Durack, 2015
164	5 Freshwater total (#3 + #4)	41.20	3.00	Durack, 2015
165	6 Water on Earth (#2+#5)	1376.2	100	Durack, 2015
	7 Sea (Infant Earth rel. to #2)	1681.1	126	Banfalvi, 2017
	8 Global loss of water (#7-#6)	305	-22	This estimate



166 \*Sea volume, average of four estimates rounded up to  $1335 \times 10^6 \text{ km}^3$  (70.015%).  
167 † Freshwater reservoirs: ice, glaciers, permanent snow.  
168 ‡ Freshwater inland: lakes, artificial lakes, ponds, streams, wetland and groundwater.  
169 Estimation of the global loss of water was obtained by subtracting the recent volume of the sea from  
170 the volume belonging to the highest sea level.  
171 The largest volume represented by the volume of Earth + volume of the sea. Calculated from  $r_2$ :  
172  $6,371,008 \text{ m} \rightarrow 1,083.210 \times 10^9 \text{ km}^3$ .  
173 Volume of underwater features including continental shelf. Subtraction of volumes calculated from  $r_2$   
174 and  $r_1$ :  $1,083.210 \times 10^9 \text{ km}^3 - 1,081.320 \times 10^9 \text{ km}^3 \rightarrow 557 \times 10^6 \text{ km}^3$  (29.5%)  
175 Summary of volumes of sea  $1332.9 \times 10^6 \text{ km}^3$  (70.5%) plus underwater features and continental shelf:  
176  $557 \times 10^6 \text{ km}^3$  (29.5%), total  $1889.9 \times 10^9 \text{ km}^3$  (100 %).  
177 Calculation of continental shelf: land surface area of the World  $\times$  average depth of ocean:  $148.42 \times 10^6$   
178  $\text{km}^2 \times 3,6822 \text{ km} = 548.42 \times 10^6 \text{ km}^3$ .  
179 Volume of underwater features without continental shelf:  $557 - 548.42 = 8.57 \times 10^6 \text{ km}^3 = 0.45\%$ .

180 Our calculations showed that the volume of the underwater features excluding the continental shelf  
181 represents only 0.45 % of the size of the ocean, confirming that the average seabed (ocean floor) is a  
182 relatively flat surface. John Murray (1888) utilised a simple model and measured the depth of the  
183 ocean at several locations then calculated the ocean volume by only 1.2% higher than the current  
184 estimates of the amount of the world's oceans (Charette and Smith, 2010). Our data showed a close  
185 relationship between sea level rises and volumetric changes in the sea. These data were useful to  
186 create a calibration curve (Figure 1) (Banfalvi, 2017). The calibration curve was then extended to  
187 apply it to the highest sea levels. More importantly, the calibration curve served to estimate global  
188 water loss.

189 The elevations of sea level rises originate mainly from the melting of freshwater reservoirs and  
190 thermal expansion during interglacial periods. Predictions forecasted that the recent interglacial  
191 period could melt about 80% of the Earth's ice and snow reserves of  $\sim 50 \times 10^6 \text{ km}^3$  (Berger and  
192 Loutre, 2002). Others judged that the freshwater reserve is only about half ( $24 \times 10^6 \text{ km}^3$ ) of the  
193 earlier estimation (Shiklomanov, 1995). The newest estimate of freshwater is somewhat higher ( $25.54$   
194  $\times 10^6 \text{ km}^3$ ) (Durack, 2015) but only about half of the maximum of the latest ice age some 120,000  
195 years ago. By fitting sea level rises and volumes into the calibration curve (Fig. 2), sea levels can be  
196 turned into volumetric values and *vice versa*.



197 The highest reported sea level was at the period of the Infant Earth (Pope et al., 2012; Rosing et  
198 al., 2010). To estimate the global loss of freshwater, the highest sea level rise of the Infant Earth  
199 served as a basis. This sea level could have been by 750 m higher than it is today (Fig. 1b) with a sea  
200 volume of  $\sim 1,681 \times 10^6 \text{ km}^3$ . Volumetric data of sea and freshwater are summarised in Table 1. The  
201 calibration curve (Figure 1b) takes into consideration that in the presence of continental crust, the sea  
202 level is about 29.5 % higher than in its absence. Due to the lack of landmasses, the early Earth was  
203 assumed to be completely covered with water (Rosing et al., 2010). Thus the sea level of the Infant  
204 Earth could have been roughly 530 m higher than it is today.

### 205 **3. Geochemical stability *versus* dynamic nature of global osmolyte system**

206 Despite large sea level fluctuations, only small changes can be traced if at all, leading to the  
207 conclusion that in general, the rate of input and output in the sea was nearly equal in agreement with  
208 the long-term mean of its salt concentration (Pope et al., 2012). The idea of a general geochemical  
209 balance of the sea is related to the limited foreseeable future of man that is not longer than 100 years  
210 but provided a model to make such constancy plausible (Rubey, 1951; Railsbeck et al., 1989). The  
211 view of a dynamic osmolyte rather than a steady-state ocean system is gaining ground, by measuring  
212 short-term volumetric decreases during ice ages; sea level rises during interglacial periods and long-  
213 term salination of the ocean (Banfalvi, 1991; 2016; 2017).

214 The Unified Sea level Rise Projection ([http://southeastfloridaclimatecompact.org/wp-](http://southeastfloridaclimatecompact.org/wp-content/uploads/2015/10/2015-Compact-Unified-Sea-level-Rise-Projection.pdf)  
215 [content/uploads/2015/10/2015-Compact-Unified-Sea-level-Rise-Projection.pdf](http://southeastfloridaclimatecompact.org/wp-content/uploads/2015/10/2015-Compact-Unified-Sea-level-Rise-Projection.pdf)) serves as a reasonable  
216 source of information to predict the sea level rise in the 21st-century relative to the 1992 mean sea  
217 level. The short-term rise envisaged 15-25 m by 2030. The medium-term sea level rise projected 25-66  
218 m with a less likely possibility of reaching 86 m by 2060. The long-term projection predicted nearly  
219 79-155 m sea level rise and a less likely extension to 206 m (Southeast Florida Regional Climate  
220 Change Compact, 2015). The melting of  $20 \times 10^6 \text{ km}^3$  ice and snow would correspond to about 80%  
221 loss of the available freshwater reservoir and would reduce the global freshwater volume to  $\sim 1.51\%$   
222 from 3% and cause a sea level rise of about 50 m (Figure 1). A higher sea level rise is less likely to  
223 occur than that predicted by 2060 in the Unified Sea level Rise Projection. Nevertheless, it may  
224 correspond to the forecast of the Intergovernmental Panel on Climate Change (IPCC) 5th Assessment



225 Report (ARS5) model projection (Nerem et al., 2018). It is unrealistic to assume the complete melting  
226 and loss of the freshwater reservoir during the recent interglacial period. A higher than 50 m sea level  
227 rise is unlikely to be reached simply because of the shortage of the freshwater reserve. The complete  
228 melting of the freshwater reservoir would be catastrophic, especially to the terrestrial vertebrates that  
229 are fully dependent on freshwater supply. Forced sea level rise patterns were predicted and could  
230 continue in the coming decades with elevated rates of rises (Fasullo and Nerem, 2018), but the rise  
231 may not be linear with time. The initial faster rate of sea level rise could slow down due to the  
232 shortage of freshwater reserve.

#### 233 **4. Conclusions**

234 One explanation for the global freshwater loss is that water vapour at the upper atmosphere  
235 photohydrolysed to H and oxygen. H escaped to the outer space; oxygen formed the protective ozone  
236 layer. Calculations showed that the recent escape of hydrogen to space is low, but it could have been  
237 much higher at the earlier periods of global history. Calculations related to the hydrogen escape theory  
238 did not allow far-reaching conclusions regarding a significant water loss. Thus evolutionary changes  
239 of sea levels and their converted data to sea volumes were used to calculate the loss of water on Earth.

240 The consequence of gradual loss of water on Earth is that the salt concentration of sea increases.  
241 Based on Raoult's law applied to the sea as a global osmolyte system, the evaporation of water vapour  
242 decreases and results in less precipitation and freshwater.

243 Characteristic biological phenomena accompany the loss of freshwater. Although the recent sea  
244 level rise will not significantly reduce the salinity of the sea, due to the limited freshwater reserves, its  
245 effect will severely impact flat seashores impacting many large cities. The life of sea animals will be  
246 hardly affected, unlike land vertebrates, including man that are entirely dependent on freshwater. The  
247 migration of people driven by the shrinkage of available freshwater and the spread of deserts are  
248 continuing. Due to the freshwater shortage, the habitat of species is decreasing at an alarming rate,  
249 threatening with extinction many endangered freshwater species.

250 It is concluded that different sea levels with decreasing heights during evolution can be used to  
251 estimate the global loss of water. The sea level of the Infant Earth could have been by 26% more  
252 voluminous than it is today. The second-largest sea levelrise (~300 m) took place some 500 million



253 years ago. Fluctuations indicate record-low sea level (-130 m) about 20,000 years ago. The basis of  
254 estimation of the global loss of freshwater is the subtraction of the recent volume of the sea from the  
255 amount belonging to the highest sea level. The substantial freshwater shrinkage contributed by man  
256 and the salination of sea demand counteractions to be taken or already in effect to protect life on Earth.

#### 258 **Data availability**

259 The MS will be deposited into the institutional repository of DEA (Educational and Research  
260 Support) of the University of Debrecen and National Library.

#### 261 **Author contribution**

262 All activities related to the preparation and publication of the manuscript were carried out by  
263 the single author.

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266 **Conflict Of Interest Statement** The author declares to have no conflict of interest.

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269 **Ethical consideration** The author has no ethical issues to be considered.

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