RESPONSE TO REVIEWER 1

Reviewer’s comments are inserted in italics and blue, and responses in regular font.

Many thanks for these comments.

GENERAL COMMENTS

The manuscript deals with an interesting, regionally relevant topic. The presented results and conclusions will certainly serve as a basis for future oceanographic, hydrological and geophysical investigations in the Caspian Sea environment. Given the particular sensitivity of the Caspian Sea and its mean water level to the changing climate, and in view of the current focus of the international research community on quantifying the variability of ocean tides over climatic time scales, especially the presented numerical experiments that reveal the dependence of both the tidal pattern and local seiches frequencies on the mean sea level is timely and highly relevant. Methods, results and conclusions are well presented. The paper is well structured; the amount and choice of figures, tables and equations is appropriate; the wording is generally adequate and clear (some suggestions for rewording are given below). There are relatively few bibliographic references, but this might indicate in fact a shortage of previous work on this topic.

There are a few papers devoted to the problem of tides in the Caspian Sea. We tried to include in the review all the main papers on this topic.

SPECIFIC COMMENTS

The scientific approach and the applied methods are valid and do, in my opinion, not require any correction.

Having myself a primarily observational background, I would have welcomed more information about the confrontation of the presented model with available sea-level data (tide gauges, altimetry). This is certainly a subjective preference, and I realize that this has been dealt in previous work (Medvedev et al. 2017, 2019: tide gauges) or may be the subject of upcoming publications (altimetry). Nevertheless, the reader could be provided with some additional interesting information, perhaps even in a qualitative way, adding a few sentences to the Introduction or Discussion. For example, what is the proportion of the relatively small astronomic tides in the total observable sea-level variation (i.e., how efficient is the presented model to predict real sea-level changes)?

We have added in the current article a comparison of modeling results with the results of harmonic analysis according to observations on coastal tide gauges. In particular, we added one additional comparison figure for harmonics M$_2$ and K$_1$ and a short section with text.
We added a few paragraphs to the discussion reflecting the assessment of the contribution of tides to the variance of total sea level fluctuations with periods from 6 hours to 2 days. In the current research, we estimated the contribution of gravitational tides to the sea level variance based on the numerical modelling results. We made two numerical experiments: 1) with the tidal input; 2) with meteorological forcing produced by the fields of wind and air pressure variations over the Caspian Sea for 1979 from NCEP/CFSR reanalysis (Saha et al., 2010). We calculated the variance of tidal sea level variability (excluding long-period constituents) and the variance of the meteorological sea level variations in the first frequency band from 0.1 to 6 cpd and the second frequency band from 0.5 to 6 cpd. Then we estimated the relative contribution (in percent) of tides to the total sea level variance in the Caspian Sea.

The maximal contribution of tides to the total sea level variance has been located in the east part of the Middle Caspian: up to 29% for the first frequency band and up to 53% for the second frequency band. In the western part of the Southern Caspian and in Turkmen Bay the tidal contribution of total variance for the second frequency band from 0.5 to 6 cpd is up to 40%. The minimum contribution has been observed in the Northern Caspian, where strong storm surges occur; and near the Absheron Peninsula, where the amphidromic points of the diurnal and semidiurnal tides are located.

The last paragraph of the Conclusions (p19, l6-11) is interesting and deserves more space. If the presented model indeed qualifies as an appropriate complement of global ocean tide models, this would be a mayor outcome of the paper and increase significantly its value. This question should be discussed in more detail. Purely empiric ocean tide models based on satellite altimetry should not be "distorted" by any assumption on the Caspian MSL. Here, again, rises the question about how well agree the presented model and altimetry, which could be offered as an outlook to future work or posed as an open research question. An invalid assumption on the Caspian MSL could indeed affect dynamical and assimilation models, but then it should be demonstrated which particular global ocean tide model assumes a wrong Caspian sea level and to which extent the tidal signal is distorted.

We rewrote this paragraph a bit and tried to make it clearer. Most of the models presented by Stammer et al. (2014) don’t include the Caspian Sea: FES14, EOT11a, TPXO9, GOT4.10, OSU12, DTU10, HAMTide. The TPXO9 included the Caspian Sea, but the MSL of the sea was 0 m with respect to the BHS. This invalid assumption shifted the coastline and significantly increased the sea area and as a result distorted the tide in this sea.
TECHNICAL CORRECTIONS

p1,l14: I understand that the splitting into two amphidromies occurs only in the diurnal case. If so, make this explicit: "For the diurnal constituents, the Absheron Peninsula splits this system into two separate amphidromies..." or so.

p1,l20-22: rephrase, e.g.: Numerical experiments with tidal simulation were made using different mean sea levels of the Caspian Sea (within a range of 5 m). The results indicate that the spatial features of the tides are strongly sensitive to changes of the mean sea level.

p1,l25: I prefer "one of the major drivers of ocean water motion" to "one of the major types".

p2,l1: "unique object for THE analysis"

p2,l5: "7.7 cm based on AN analysis of 30-day"

p2,l6: "performed A spectral analysis"

p2,l9: I prefer "Analyzing annual series..." to "Having analyzed annual series..."

p2,l14: "for different parts of the Caspian Sea" instead of "for different sea parts"

p2,l14: "... tide gauges. A maximum tidal range..."

p2,l16: "performed A high-resolution spectral analysis"

p2,l17: I prefer: "Southern (or Northern) Caspian" to "South (or North) Caspian" throughout the text. Also "Central Caspian" to "Middle Caspian".

p2,l18: "of THE diurnal radiational constituent S1"

p2,l19: "than those of THE gravitational constituents"

p2,l22: "examination OF specific tidal features"

p2,l23: "in THE deep-water areas"

p2,l25: check reference format: "Caspian Sea (Medvedev et al. 2019)." or "Caspian Sea in Medvedev et al. (2019)."

p3,l2: "we used A 2D version"

p3,l3: "in THE two-dimensional shallow water equations"

p3,l12: check if this complies with the journal’s reference format, or if there is another reference to this model to cite.

p3,l18: "THE energy dissipation of THE generated flows is caused by THE vertical turbulent viscosity. THE friction..."

p3,l22: "is THE flow velocity above..."

p4,l2: I prefer "to avoid a vanishing bottom drag in very deep waters"

p4,l4: "THE numerical simulations"
p4l6: "In section 3.1, a mean sea level (MSL) of the Caspian of -28 m with respect to the Baltic Height System (BHS, relative to the zero of the Kronstadt gauge) was adopted in the numerical modelling."

p4,l8: "from -25 m to -30 m with respect to the BHS. THE boundary conditions..." Consider replacing "** m of the BHS" by "** m with respect to the BHS" throughout the text.

p4,l11: reference format (see p2,l25); same in the following sentence

p5,Fig1: Kizlyar Bay is not indicated; also the tide gauges stations listed in Table 2 would be helpful to display. If necessary, the isobath annotation could be thinned out, or even omitted, if a color scale would be provided. In the caption, include a reference to the bathymetry model:

"Figure 1: The bathymetry of the Caspian Sea according to..."

p5,l8: "A numerical model with A MSL of -28 m with respect to the BHS..."

p6,l1: perhaps better "examine" instead of "consider"

p6,l2: "taking THE major constituents..."

p6,l3: "THE diurnal pattern includes..."

p6,l5: "have [or: feature] A counterclockwise rotation."

p6,l9: reference format (see p2,l25); I suggest rephrasing, e.g.: "Medvedev et al. (2019) showed that the results of numerical modelling are not really reliable in the Northern Caspian due to the very shallow depths in this area with about 20% of this part of the Caspian being less than 1 m deep (Baydin and Kosarev, 1986)."

p6,l12: "... have a spatial distribution similar to that of K1."

p6,l19: "The areas ... are: 1) the western part of the Southern Caspian..." or, better:

"Maximum M2 amplitudes are found in 1) the western part of the Southern Caspian..."

p6,l24: "of THE major tidal constituents at selected cities [or: towns] around the Caspian Sea"

p7,Fig2: Add panel identifier "a)" and constituent identifier "K1" to the left panel.

p8,Tab2: After a quick glance at Fig. 2 it seems somewhat surprising that Fort Shevchenko features the largest tidal range. Perhaps this deserves a short comment.

p8,l7: "form factor AS determined by the AMPLITUDE ratio of the MAJOR diurnal and semidiurnal constituents"

p8,l11: "In general, A mixed mainly semidiurnal..."

p8,l12: "Only in THE western and eastern parts..."

p9,l3: "Based on the results of THE numerical modeling..."

p9,l6: I prefer "features a pattern similar to the M2 amplitude distribution"

p9,l9: I prefer here "are included" to "are presented"
"THE spatial structure..." I am used to the expression "semi-major axis" instead of "major semi-axis" - revise throughout the text, same for semi-minor axis. "The largest M2 current velocities are observed [or: found] in," or: "The areas of the largest M2 current velocities are:"

"up to 12.5 cm/s and 11.7 cm/s"

I prefer "depending on local topographic features."

"and THE tidal currents are nearly rectilinear."

"The spatial pattern of THE S2 tidal currents"

Perhaps "only the S2 semi-major axis is half of that of M2."

"repeats THE pattern of M2, too."

I prefer "by a factor of 1.8." to "by 1.8 times."

Semi-major axis velocity magnitudes (cm/s) for M2 tidal currents. Blue ellipses indicate clockwise circulation, red ellipses counterclockwise circulation.

I prefer: "3.4 Numerical experiments with VARYING MSL"

"THE interannual MSL variability...

"and 20% of this area has a depth less than 1 m" - this has already been stated above, consider dropping this statement here.

"As a result, changes of the Caspian MSL by 2-3 m (as observed, e.g., between 1974 and 1994) lead to significant changes in the hydrodynamics of the Northern Caspian as well as in coastal waters of the Central and Southern Caspian."

"THE spatial characteristics of natural resonant oscillations in the basin (seiches)"

"experiments with tidal simulationS using..."

"This corresponds to the natural range of MSL changes..."

"THE results of these experiments allow to estimate [or: identify] the changes..."

"repeats THE pattern of M2, too."

"by a factor of 1.8." to "by 1.8 times."

I prefer "depending on local topographic features."

"and THE tidal currents are nearly rectilinear."

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Perhaps "only the S2 semi-major axis is half of that of M2."

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Semi-major axis velocity magnitudes (cm/s) for M2 tidal currents. Blue ellipses indicate clockwise circulation, red ellipses counterclockwise circulation.

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"As a result, changes of the Caspian MSL by 2-3 m (as observed, e.g., between 1974 and 1994) lead to significant changes in the hydrodynamics of the Northern Caspian as well as in coastal waters of the Central and Southern Caspian."

"THE spatial characteristics of natural resonant oscillations in the basin (seiches)"

"experiments with tidal simulationS using..."

"This corresponds to the natural range of MSL changes..."

"THE results of these experiments allow to estimate [or: identify] the changes..."
The amplitude of the diurnal tide

This is caused... AT low MSL.

Strong modifications of the diurnal tidal pattern due to MSL changes occur along the transition between Northern and Central Caspian. At a MSL of -25 m the largest amplitudes are located...

With decreasing MSL, large amplitudes extent farther west.

At a MSL of..., revise throughout the text.

These changes are probably caused by a strong modification... and, as a result, of the frequency [or: resonant] properties of this part of the sea [or: subbasin].

The change in the spatial structure

The maximum tidal range of 22 cm is found in the Mangyshlak Bay for a MSL of -25 m. At this MSL the tidal range in the Turkmen Bay amounts to 13 cm and in the Türkmenson, sy Gulf to 15.5 cm.

The changes in tidal characteristics...

tidal vector diagrams... for different MSL of the Caspian Sea.

However, the M2 phase lag... 100_ and the M2 amplitude doubles: from 2.5 cm at a MSL of ...

THE results of THE numerical tidal modelling... of A harmonic analysis...

... spectra FROM the western coast" (or: at)

simulation of THE K1 amplitude

eastern part of THE Southern Caspian

..., the area and length of the island increase significantly, and the islands becomes an effective boundary to the west, reflecting THE tidal waves...

reference format: Badyukova

... with a maximum elevation of the island of 2 m at a MSL of -28 m.

Thus, in the experiments assuming a MSL of -25 m, the island was completely submerged.

A comparison of the island’s...

I prefer "has gradually moved eastward and has changed its geometrical configuration due to the redistribution of deposits and erosion."

The greatest contribution to this process originates from eolian redistribution." reference format: Nikiforov

THE magnitude and direction of THE generated wind fields

A spectral analysis...
When the MSL of the Caspian Sea decreases, the $Q$-factor of seiches with a period of about 12 hours significantly decreases in the Türkmenbaşy Gulf and at a MSL of -29 m it does not exceed any more the spectral noise level.

"In the Turkmen Bay a decrease in MSL from -26 m to -29 m causes the spectral peak of the main seiches mode to migrate towards lower frequencies..."

Consider dropping "Apparently,". "This is due to the progressive elongation of Ogurja Ada Island which represents the western boundary of the bay."

"In this study THE tidal dynamics of the Caspian Sea HAVE been numerically investigated."

"THE results of THE numerical simulation"

I prefer "hydrodynamics" to "water dynamic"; "... might have been underestimated so far."

"OUR numerical experiments indicate... sensitive to changes in the MSL. A modification..."

"including the Northern Caspian, which results in significant changes in the frequency response of the basin. This is also confirmed by..."

"in THE improvement of"

"Stammer et al. (2014) present a detailed comparison of the main modern global barotropic tide models."

I prefer "We believe that our findings on the tidal dynamics can help to better understand the diurnal and semidiurnal variability in the sea level and currents in the Caspian Sea."

(Author contribution): It think that "IP" should be replaced by "IM" throughout section 7.

We agree with all comments in technical corrections section and clarified these sentences.

Some comments:

The reviewer correctly noted that Fort Shevchenko features the largest tidal range among other cities in Table 2. The maximum tidal range has been observed in the Turkmen Bay (21 cm), but there are no major cities in this area. We added short comment about it in text of manuscript.
RESPONSE TO REVIEWER 2

Reviewer’s comments are inserted in italics and blue, and responses in regular font.

Many thanks for these comments.

5 SPECIFIC COMMENTS

I generally find the paper and topic interesting but in many ways the paper is positioning itself between two chairs. One describing the result of the numerical modelling (not the numerical model) and the second the effect of lake level change. So given the title of the paper this was a bit of surprise to me.

So I would very much like to see the title reflecting this better like “ocean tides under changing lake level”. In general I find the scientific approach and the applied methods valid though I have the same problem as the first reviewer that no information of the presented model is given as this is given in previous work. (Medvedev et al. 2017, 2019: tide gauges). I find the investigation of love numbers misplaced in this context as this is likely dealt with in the reference work, and I suggest this is substituted with more quantitative discussion on the quality of the model.

Our paper describes the result of the numerical modelling. We added in new version of manuscript information about the confrontation of the presented model with tide gauge data.

We believe that the main results of presented paper tidal charts for amplitudes and phase lags of the major tidal harmonics, form factor, tidal range and velocity of tidal currents. The numerical results with the changes of the mean sea level are secondary. Therefore, we believe that the current title of the article reflects well the results presented in it. We didn’t do the investigation the Love numbers, but simply describe the model parameters.

TECHNICAL

Figure 1 is nice but identical to another publication by the leading author. As the evaluation of the ocean tide model in Table 2 is done for a number of cities surrounding the Caspian it would be much more appropriate if Figure 1 was changes to represent the location of these cities and I personally have no clue to where the cities are located. This would make reading easier.

We corrected Figure 1 and added some names of bays and cities from Table 2.

Of interest I am very puzzled about the >21 cm tides in the TB described in Figure 3 because it does not relate very well to the amplitudes of the two major constituents in Figure 2 and the 4 major constituents in Table 2. the major semi-diurnal constituents explains a maximum of 7 cm or 1/3 of the tidal range in TB and the Maximum tidal range (R) for the 4 major explains less than 1/2 of the signal. Consequently there must be other major constituents not mentioned in this paper that is responsible and likely dominating?. Again the fact that I do not know the location of the cities in Table 2 makes it hard to determine the location of maximum amplitudes. The paper deserves an detailed explanation of this phenomena (is it astronomical constituents, overtides???).

We think that there is no surprise in the estimates of the tidal range. The tidal range was calculated as the maximum range of tidal sea level oscillations during one lunar day (~25 hours). This is approximately equal to twice the sum of the four major constituents (M2, S2, K1, and O1). For example in Table 2 for Fort Shevchenko we have $H(M2)= 2.47$ cm, $H(S2)= 0.92$ cm,
$H(K1)=0.56 \text{ cm}, H(O1)=0.30 \text{ cm}$. The twice sum of these constituents is 8.5 cm, that is relate well to the tidal range in Table 2 for this city (8.9 cm). In Turkmen Bay (see Fig. 3 in new version of paper) we have $H(M2)=6 \text{ cm}, H(S2)=2.6 \text{ cm}, H(K1)=0.73 \text{ cm}, H(O1)=0.47 \text{ cm}$. The twice sum of these constituents is 19.6 cm, that is relate well to the tidal range = 21 cm in Fig. 4. The differences in the magnitude of the tidal range in paper and presented twice sums are caused by the contribution of semidiurnal constituents N2 and K2 with amplitudes of about 0.5-1 cm.

The paper briefly mentions the form factor $F$ in Table 2 and later in the paper gives one sentence about it. The form factor is detailed in previous publications by Medevedev, and I would leave it out of describe it much more detailed in this publication.

In current paper, we show for the first time a map of form factor for the Caspian Sea. Consequently, we will keep the short description of it.

When discussing numerical experiments with different MSL more information on the accuracy of the bathymetry used must be provided. The discussion on Page 13 following Figure 6 is interesting but again I question on the Turkmen Bay.

Figure 4 could benefit from names on the regional features.

Figure 5 6 and 7 should be reconsidered and redrawn for consistency.

Figure 5 used 26 28 and 29 meters, Figure 6 25, 27 and 29 meters and Figure 7 25-30 meters. so they all are consistent. Figure 5 also needs a bit of "regional" explanation for the reader. How can two cities 300 km apart. Exhibit sea level changes differing by 0.5 meters from 1900 until now.

Since 1980 the sea level curve matches but before it differs up to 0.5 meters?-

We added some information on the accuracy of the bathymetry.

We added names on the regional geographical features.

We redrawn Figure 5. We done new bathymetry maps for MSL = -25, -27, -29 m. The difference in the mean sea level of 0.3 m in Baku and Makhachkala caused us questions too. We wanted to show the mean sea level of the whole Caspian Sea, but since it is different depending on the station, we decided to show two stations in the figure. We checked several sources of data on the Caspian level (http://www.caspcom.com/ и http://caspi.ru/). In both databases the data shown in Fig. 5 differences in the average level at 0.3 m in the first half of the 20th century. We believe that this feature in interannual sea level variability is caused by the local conditions of these stations (for example, tectonic movements), which led to a relative change in the absolute height of the tidal pole. But since these questions are not the purpose of this study, we decided to show in this figure only one station (Makhachkala).

Figure 8 is interesting in attempting to explain the spectral density at different MSL regimes. I guess this is the key to the large tides in the Turkmen Bay, and the key to which constituents are responsible for the large tides. This deserved more attention and investigation and explanation in my opinion.

We have added a few more words to this section.
In the discussion there is a bit of uncertainty to the discussion of the large tides in the Turkmen Bay. The height of the island is in the paper claimed to be 3-5 meters by the author and 5-8 meters from the SRTM. SRTM was measured in the Early 2000’s where sea level was -27.5 meters, so there is inconsistency here.

We agree with the reviewer that there is inconsistency in the height of the Ogurja Ada Island. The SRTM was in February 2000 where sea level was -27 m. When we took the SRTM data we expected to see the height of the island about 1-3 m, but it turned out to be higher. Because of this difference between the results of Badyukova, 2015, GEBCO database and SRTM data we put this paragraph in the discussion. We have not yet found more reliable information about the height of the island and now we can’t say who is right. We will try to study this in more detail in the subject of upcoming future researches.

Many thanks.
Numerical modelling of the tides in the Caspian Sea
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³Institute of Ocean Sciences, Sidney, B.C., Canada
Correspondence to: Igor Medvedev (patamates@gmail.com)

Abstract. The Caspian Sea is the largest enclosed basin on the Earth and a unique subject for analysis of tidal dynamics. The Caspian Sea has independent tides only, which are generated directly by tide-generating forces. Using the Princeton Ocean Model (POM), we examine the spatial and temporal features of tidal dynamics in the Caspian Sea in detail. We present tidal charts for amplitudes and phase lags of the major tidal harmonics, form factor, tidal range and velocity of tidal currents. Semidiurnal tides in the Caspian Sea are determined by a Taylor amphidromic system with counterclockwise rotation. The largest M₂ amplitude is 6 cm and is located in the Turkmens Bay. For the diurnal constituents, the Absheron Peninsula splits this system into two separate amphidromes with counterclockwise rotation to the north and to the south of it. The maximum K₁ amplitudes (up to 0.7–0.8 cm) are located in: 1) the southeastern part of the Caspian Sea, 2) the Türkmenbaşy Gulf, 3) the Mangyshlak Bay, and 4) the Kizlyar Bay. The semidiurnal tides prevail over diurnal tides in the Caspian Sea. The maximum tidal range has been observed in the Turkmen Bay, up to 21 cm. The highest velocity of the total tidal currents is observed in the straits to the north and south of Ogurja Ada, up to 22 cm/s and 19 cm/s, respectively. We made numerical experiments with tidal simulations using different mean sea level MSL levels of the Caspian Sea (from -25 within a range of 5 m to -30 m). The results indicate that the spatial features of the tides are strongly sensitive to the MSL changes of the mean sea level.

1 Introduction

Tides, one of the major types of ocean water motion, are formed under the influence of tide-generating forces of the Moon and the Sun and the rotation of the Earth. Tides can be represented as the sum of two types of oscillations: (1) the co-oscillating tide caused by the tidal influx from an adjacent basin, and (2) the independent tide, which is generated directly by the tide-generating forces (Defant, 1961). Co-oscillating tides dominate in marginal seas, generated by tidal waves penetrating from the adjoining ocean or seas. In isolated inland seas (e.g., the Black Sea and the Baltic Sea), independent tides strongly prevail as tidal waves from adjacent basins cannot significantly penetrate the sea (Medvedev et al., 2013; Medvedev et al., 2016; Medvedev, 2019).
The Caspian Sea is a unique object for the analysis of independent tide formation as it is the largest enclosed basin on Earth.

Tides in the Caspian Sea have been studied for a long time, though not on a regular basis. Malinovsky (1926) showed that the semidiurnal tides dominate in the Caspian Sea and the spring tidal range was 7.7 cm based on an analysis of 30-day hourly records at three tide gauges. German (1970) performed a spectral analysis of three-month observational series at eight tide gauges and distinguished the diurnal and semidiurnal constituents through different generation mechanisms: semidiurnal tides had a gravitational origin while diurnal tides were formed by sea-breezes. Kosarev and Tsyganov (1972) and found the maximum tidal range 12 cm at Ogurja Ada. Spidchenko (1973) estimated the amplitudes and phase lags at seven sites. Having analyzed annual series of hourly observations at six tide gauges, Levyant et al. (1994) hypothesized that a semidiurnal tidal wave is represented by a counterclockwise amphidromic system (like a Kelvin wave) with a center in the Apscheron Threshold’s area.

Medvedev et al. (2017) estimated the amplitudes and phase lags of major tidal constituents for different sea parts of the Caspian Sea based on an analysis of long-term hourly data from 12 tide gauges. Maximum tidal range of 21 cm was found at the Aladga (eastern part of the Southern Caspian). Medvedev et al. (2017) also performed a high-resolution spectral analysis and determined that the diurnal sea level oscillations in the Middle Caspian have a gravitational origin, while those in the Southern Caspian are mainly caused by radiational effects: the amplitude of the diurnal radiational constituent $S_1$ is much higher than those of the gravitational constituents $O_1$, $P_1$, and $K_1$. In the Northern Caspian, there are no gravitational tides and only weak radiational tides are observed. A semidiurnal tide is predominant in the Middle Caspian and in the Southern Caspian.

An analysis of tide gauge data allows for the examination of specific tidal features at different sites, but not for the estimation of the spatial structure of tides in the deep-water areas of the Caspian Sea. Therefore, in order to capture these spatial structures we adapted the numerical Princeton Ocean Model (POM) to the Caspian Sea in (Medvedev et al., 2019). The developed POM reproduces the tides and meteorological sea level variability with periods ranging from several hours to a month and tides in the Caspian Sea, which, In present we can use this model to describe in detail the spatial and temporal peculiarities of tidal dynamics of the whole the Caspian Sea.
2 Data and methods

2.1 Numerical model description

In this study we used a 2D version of the Princeton Ocean Model (POM) (Mellor, 2004). The forcing term in the two-dimensional shallow water equations was specified through the gradients of tidal potential over the Caspian Sea:
\[
\bar{F}_T = -(1 + k - h) \nabla \bar{\Omega},
\]  
where \(k\) and \(h\) are the Love numbers and \(\bar{\Omega}\) is the tidal potential. Love numbers \(k\) and \(h\) relate the body Earth tide (and associated perturbations) to the potential. We used frequency-dependent values of \(h\) and \(k\) calculated by Wahr (1981) (Table 1). The tidal potential was calculated for spherical harmonics via formulas provided by Munk and Cartwright (1966) and included all the main tidal components (> 80), including major diurnal, semi-diurnal, shallow water and long-period constituents. Additionally, our numerical model includes the ocean tidal loading potential obtained from FES2014 (Finite Element Solution tidal model) produced by NOVELTIS, LEGOS and CLS Space Oceanography Division and distributed by AVISO, with support from CNES (http://www.aviso.altimetry.fr/).

Table 1. Love numbers and the elasticity factors for major tidal constituents (Wahr, 1981; Kantha and Clayson, 2000).

<table>
<thead>
<tr>
<th>Constituent</th>
<th>Frequency (cpd)</th>
<th>( h )</th>
<th>( k )</th>
</tr>
</thead>
<tbody>
<tr>
<td>long-period</td>
<td>0.606</td>
<td>0.604</td>
<td>0.299</td>
</tr>
<tr>
<td>Q(_1)</td>
<td>0.8932</td>
<td>0.603</td>
<td>0.298</td>
</tr>
<tr>
<td>O(_1)</td>
<td>0.9295</td>
<td>0.603</td>
<td>0.298</td>
</tr>
<tr>
<td>P(_1)</td>
<td>0.9973</td>
<td>0.581</td>
<td>0.287</td>
</tr>
<tr>
<td>K(_1)</td>
<td>1.0027</td>
<td>0.520</td>
<td>0.256</td>
</tr>
<tr>
<td>J(_1)</td>
<td>1.0390</td>
<td>0.611</td>
<td>0.302</td>
</tr>
<tr>
<td>semidiurnal</td>
<td>0.609</td>
<td>0.609</td>
<td>0.302</td>
</tr>
<tr>
<td>shallow</td>
<td>0.609</td>
<td>0.609</td>
<td>0.302</td>
</tr>
</tbody>
</table>

Energy The energy dissipation of the generated flows is caused by the vertical turbulent viscosity. Friction The frictional force in the momentum equations is determined by the speed of the bottom flow and the friction coefficient:
\[
(t_{bx}, t_{by}) = (C_b u_b |\bar{u}_b|, C_b v_b |\bar{u}_b|),
\]  
where \(\bar{u}_b = (u_b, v_b)\) is the flow velocity above the bottom boundary layer (which is assumed to be equal to the barotropic velocity \(\bar{u}_b\) for the 2D model), \(C_b\) is the bottom friction coefficient which has the following form:
\[ C_b = \max\left[\frac{\kappa^2}{(\ln(0.5H/z_0))}, 0.0025\right], \]  

where \( \kappa = 0.4 \) is the von Kármán constant, \( z_0 \) is the bed roughness length. A minimum value for the bottom friction coefficient, \( C_b = 0.0025 \), was applied in order to avoid having the vanishing bottom drag effect vanish when the water depth is in very large deep waters.

Numerical The numerical simulations were performed on a grid of 507 by 659 nodes with a constant step of 1’ in latitude and longitude, created from GEBCO bathymetry data of the Caspian Sea with a resolution of 30 arcseconds. For the Caspian Sea GEBCO used the gridded data set provided by Hall (2002). This dataset based on over 280,000 bathymetric soundings and points digitized from bathymetric contours, taken from 107 Russian navigational charts. In section 3.1, for numerical modelling the mean sea level (MSL) of the Caspian was set at -28 m with respect to the Baltic Height System (BHS, relative to the zero of the Kronstadt gauge) was adopted in the numerical modelling. In the numerical experiments in section 3.2, the MSL of the Caspian was varied from -25 to -30 m with respect to the BHS. Boundary The boundary conditions for the tidal model are zero flow normal to the coast (at the 2 m depth contour).
Figure 1. The bathymetry of the Caspian Sea according to the GEBCO database. The MSL is -28 m with respect to BHS. Black points are tide gauges used for validation of the numerical model. Other designations: TI is the Tyuleniy Island, TA is the Tyuleniy Archipelago, MB is Mangyshlak Bay, KB1 is Kizlyar Bay, KB2 is Kazakh Bay, AP is the Absheron Peninsula, AT is the Absheron Threshold, CI is the Chilov (Zhiloy) Island, TG is Türkmenbaşy Gulf, and TB is Turkmen Bay.
2.2 Tidal model validation

In (Medvedev et al., 2019), the model was validated by hourly sea level observations at eight tidal gauges in the Caspian Sea (Fig. 1). In (Medvedev et al., 2019), several experiments with different values of the bed roughness length were performed. The best tide reproduction accuracy at the eight sites was obtained at $z_0 = 0.01$ m, which is used here to determine the bottom friction coefficient $C_b$. All presented results of the tidal analysis are given in (3) to determine the bottom friction coefficient $C_b$. The comparing of the amplitudes ($H$) and phase lags ($G$) of tidal components calculated from the results of numerical modeling and based on observations is presented in Fig. 2. The error in the calculations of the amplitude of harmonic $M_2$ at Baku, Svinoi Island, Fort Shevchenko, Bektash, and Ogurchinsky Island did not exceed 0.1–0.2 cm. This error for Kara-Bogaz-Gol and Krasnovodsk was 0.3–0.4 cm. The phase lag error for six tide gauges varied from 0° to 6°, for Ogurchinsky Island was 36°, and for Krasnovodsk was 26°. The amplitude error of harmonic $K_1$ at seven tide gauges was 0.1–0.2 cm, and for Baku was 0.4 cm. The phase lag errors varied from 1° to 50°. All presented results for phase lags in the tidal analysis are relative to Greenwich Mean Time (GMT).
Figure 2. The comparison of (a) amplitudes and (b) phase lags of harmonics $M_2$ (red) and $K_1$ (blue) estimated by results of the numerical modelling ($H_{\text{model}}$ and $G_{\text{model}}$) and tide gauge observations ($H_{\text{obs}}$ and $G_{\text{obs}}$).

Figure 1. The bathymetry of the Caspian Sea. Black points are tide gauges used for validation of the numerical model. Other designations: TI is the Tyuleny Island, TA is the Tyuleniy Archipelago, MB is the Mangyshlak Bay, KB is the Kazakh Bay, CI is Chilov (Zhiloy) Island, TG is Türkmenbaşy Gulf, TB is Turkmen Bay.

3 Results

3.1 Numerical modelling of tides

A numerical model with an MSL of -28 m of with respect to the BHS was used in order to reproduce the tides in 1978. For this year we had the maximum number of tide gauge’s data. Amplitudes and phase lags of major tidal constituents were calculated using classical harmonic analysis (Pugh and Woodworth, 2014). In this section, we consider the spatial pattern of diurnal and semi-diurnal tides taking the major constituents $K_1$ and $M_2$ as an example.

Diurnal tidal pattern includes a complicated amphidromic system in the Middle Caspian (Fig. 2a). The Absheron Peninsula splits this system into two separate amphidromes to the north and south of it. Both amphidromic systems have a counterclockwise rotation. Near the Absheron Peninsula, the $K_1$ amplitude is less than 0.15 cm. The maximum $K_1$ amplitudes (up to 0.7–0.8 cm) are located in: 1) the southeastern part of the Caspian Sea, 2) the Türkmenbaşy Gulf, 3) the Mangyshlak Bay, and 4) the Kizlyar Bay. Another amphidrome, with counterclockwise rotation, is formed in the Northern Caspian. In (Medvedev et al., 2019) we showed that the results of numerical modelling are not really reliable in the Northern Caspian and are not really trustworthy. For the reason of due to the very shallow depths of this area with about 20% of this part of the North-Caspian having depths being less than 1 m deep (Baydin and Kosarev, 1986). Other diurnal tidal constituents have similar spatial distribution similar to that of $K_1$. The amplitudes of these constituents are up to 0.5 cm for $O_1$, and up to 0.25 cm for $P_1$. The amplitude of the other diurnal tidal constituents in the Caspian Sea does not exceed 0.1 cm.

Semidiurnal tides in the Caspian Sea are determined by a Taylor amphidromic system with counterclockwise rotation. This system is the result of the superposition of two Kelvin waves propagating in opposite directions (Fig. 2b). The amphidromic point of this system is located...
80 km east of the Absheron Peninsula. The minimum M₂ amplitudes are located in: 1) east of the Absheron Peninsula, 2) western and 3) eastern parts of the Northern Caspian. The areas with the maximum M₂ amplitudes are observed in: 1) western part of the Southern Caspian, up to 2.4 cm; 2) the Kazakh Bay, up to 3.2 cm; 3) the Mangyshlak Bay, up to 3.2 cm; and 4) the Türkmenbasy Gulf, 3.9 cm. The largest M₂ amplitude is 6 cm and is located in the Turkmen Bay. Other semidiurnal tidal constituents have a similar spatial distribution to M₂. The S₂ amplitude in the Turkmen Bay is 2.6 cm, N₂ is 1.1 cm, and K₂ is 0.7 cm. The amplitudes and phase lags of the major tidal constituents at main cities in selected towns around the Caspian Sea are presented in Table 2.

![Amplitude Map](image)

Table 2: Tidal Amplitudes and Phase Lags for Main Cities in Selected Towns around the Caspian Sea.
Figure 23. Tidal maps of co-amplitudes (cm) (shaded) and co-phase lags (degrees, GMT) (solid lines) for (a) $K_1$ and (b) $M_2$. 
Table 2. Amplitudes ($H$) and Greenwich phase lags ($G$) of major tidal constituents, the form factor ($F$), and maximum tidal range ($R$) at main cities in the Caspian Sea.

<table>
<thead>
<tr>
<th>Station</th>
<th>Country</th>
<th>$M_2$</th>
<th>$S_2$</th>
<th>$K_1$</th>
<th>$O_1$</th>
<th>$F$</th>
<th>$R$, cm</th>
</tr>
</thead>
<tbody>
<tr>
<td>Bandar-e Anzali</td>
<td>Iran</td>
<td>2.12</td>
<td>354</td>
<td>0.92</td>
<td>354</td>
<td>0.24</td>
<td>112</td>
</tr>
<tr>
<td>Rudsar</td>
<td>Iran</td>
<td>1.83</td>
<td>353</td>
<td>0.77</td>
<td>353</td>
<td>0.26</td>
<td>149</td>
</tr>
<tr>
<td>Tonekabon</td>
<td>Iran</td>
<td>1.64</td>
<td>352</td>
<td>0.68</td>
<td>352</td>
<td>0.33</td>
<td>160</td>
</tr>
<tr>
<td>Chalus</td>
<td>Iran</td>
<td>1.43</td>
<td>351</td>
<td>0.59</td>
<td>351</td>
<td>0.41</td>
<td>171</td>
</tr>
<tr>
<td>Babolsar</td>
<td>Iran</td>
<td>1.07</td>
<td>357</td>
<td>0.41</td>
<td>359</td>
<td>0.56</td>
<td>187</td>
</tr>
<tr>
<td>Bandar</td>
<td>Kazakhstan</td>
<td>1.10</td>
<td>360</td>
<td>0.42</td>
<td>10</td>
<td>0.78</td>
<td>197</td>
</tr>
<tr>
<td>Torkaman</td>
<td>Kazakhstan</td>
<td>1.90</td>
<td>88</td>
<td>0.74</td>
<td>98</td>
<td>0.54</td>
<td>215</td>
</tr>
<tr>
<td>Hazar</td>
<td>Turkmenistan</td>
<td>2.27</td>
<td>129</td>
<td>0.85</td>
<td>140</td>
<td>0.60</td>
<td>234</td>
</tr>
<tr>
<td>Türkmenbaşy</td>
<td>Turkmenistan</td>
<td>2.43</td>
<td>167</td>
<td>0.99</td>
<td>169</td>
<td>0.40</td>
<td>253</td>
</tr>
<tr>
<td>Garabogaz</td>
<td>Turkmenistan</td>
<td>2.30</td>
<td>186</td>
<td>0.89</td>
<td>187</td>
<td>0.29</td>
<td>303</td>
</tr>
<tr>
<td>Aktau</td>
<td>Kazakhstan</td>
<td>2.47</td>
<td>210</td>
<td>0.92</td>
<td>210</td>
<td>0.56</td>
<td>326</td>
</tr>
<tr>
<td>Fort</td>
<td>Kazakhstan</td>
<td>1.16</td>
<td>48</td>
<td>0.40</td>
<td>67</td>
<td>0.77</td>
<td>75</td>
</tr>
<tr>
<td>Shevchenko</td>
<td>Kazakhstan</td>
<td>1.22</td>
<td>228</td>
<td>0.48</td>
<td>244</td>
<td>0.53</td>
<td>25</td>
</tr>
<tr>
<td>Lagan</td>
<td>Russia</td>
<td>1.22</td>
<td>229</td>
<td>0.48</td>
<td>244</td>
<td>0.52</td>
<td>24</td>
</tr>
<tr>
<td>Makhachkala</td>
<td>Russia</td>
<td>1.23</td>
<td>227</td>
<td>0.48</td>
<td>241</td>
<td>0.51</td>
<td>23</td>
</tr>
<tr>
<td>Kaspysk</td>
<td>Russia</td>
<td>1.27</td>
<td>223</td>
<td>0.49</td>
<td>235</td>
<td>0.48</td>
<td>25</td>
</tr>
<tr>
<td>Izberbash</td>
<td>Russia</td>
<td>1.80</td>
<td>231</td>
<td>0.74</td>
<td>239</td>
<td>0.25</td>
<td>28</td>
</tr>
<tr>
<td>Sumqayit</td>
<td>Azerbaijan</td>
<td>2.18</td>
<td>36</td>
<td>0.96</td>
<td>8</td>
<td>0.05</td>
<td>358</td>
</tr>
<tr>
<td>Baku</td>
<td>Azerbaijan</td>
<td>2.32</td>
<td>7</td>
<td>1.03</td>
<td>8</td>
<td>0.10</td>
<td>30</td>
</tr>
<tr>
<td>Gobustan</td>
<td>Azerbaijan</td>
<td>2.24</td>
<td>2</td>
<td>1.00</td>
<td>3</td>
<td>0.13</td>
<td>71</td>
</tr>
<tr>
<td>Neftçala</td>
<td>Azerbaijan</td>
<td>2.40</td>
<td>1</td>
<td>1.06</td>
<td>2</td>
<td>0.21</td>
<td>71</td>
</tr>
<tr>
<td>Lankaran</td>
<td>Azerbaijan</td>
<td>2.29</td>
<td>359</td>
<td>1.00</td>
<td>359</td>
<td>0.21</td>
<td>79</td>
</tr>
</tbody>
</table>

3.2 Form factor and tidal range, and role of tidal oscillations in the sea level variability

The results of our analysis indicate that semidiurnal tides prevail over diurnal tides in the Caspian Sea. We estimated the form factor as determined by the amplitude ratio of the major diurnal and semidiurnal constituents (Pugh and Woodworth, 2014):

$$F = \frac{H_{M_2} + H_{O_1}}{H_{M_2} + H_{S_2}}.$$  \hspace{1cm} (4)

Tides have a semidiurnal form in the eastern part of the Middle Caspian ($F < 0.25$), in the western part of the South Caspian ($F < 0.25$), and in the Turkmen Bay ($F \sim 0.14-0.17$) (Fig. 3a-4a). In general, a mixed mainly semidiurnal tide ($0.25 < F < 1.5$) is observed in other areas of the Caspian Sea. Only in the western and eastern parts of the North Caspian and at the semidiurnal
amphidromic point (80 km east of the Absheron Peninsula) the tide has a mixed mainly diurnal form ($F > 1.5$).

Based on the results of the numerical modelling of diurnal, semidiurnal and shallow tidal constituents at each grid node the 18.6-year tidal time series have been predicted at each grid node. The tidal range was calculated as the maximum range of tidal sea level oscillations during one lunar day (~25 hours). The tidal co-range picture has features a pattern similar with the $M_2$ amplitude distribution (Fig. 3b). The maximum tidal ranges have been observed in 1) the Kazakh Bay, up to 12 cm; 2) the-Mangyshlak Bay, up to 12 cm; 3) the-Türkmenbaşy Gulf, 13 cm; and 4) the-Turkmen Bay, up to 21 cm. The form factor and tidal range at main cities in the Caspian Sea are presented included in Table 2. Fort Shevchenko features the largest tidal range among other cities in Table 2. The maximum tidal range has been observed in Turkmen Bay (21 cm), but there are no major cities in this area.

---

**Table 2:**

<table>
<thead>
<tr>
<th>City</th>
<th>Tidal Range (cm)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kazakh Bay</td>
<td>up to 12 cm</td>
</tr>
<tr>
<td>Mangyshlak Bay</td>
<td>up to 12 cm</td>
</tr>
<tr>
<td>Türkmenbaşy Gulf</td>
<td>13 cm</td>
</tr>
<tr>
<td>Turkmen Bay</td>
<td>up to 21 cm</td>
</tr>
</tbody>
</table>

---

**Diagram:**

- **Form factor**
- **Tidal range (cm)**

---

**Figure 3a:**

- **Figure 3b:**
In Medvedev et al. (2017) we estimated the role of tidal oscillations in the sea level variability in the Caspian Sea. We calculated the relative contribution of tides (gravitational and radiational) to the total sea level variance in the frequency band from 0.5 to 6 cpd for eleven tide gauges. The maximum contribution is observed at Bektash (27%). At Aladga, which has the tidal range of 21 cm, the tidal contribution to the sea level variance was 22.5%. The least relative tidal role was on the western coast: 7.6% at Makhachkala and 11.7% at Baku.

In the current research, we estimated the contribution of gravitational tides to the sea level variance based on the numerical modelling results. We made two numerical experiments: 1) with the tidal input; 2) with meteorological forcing produced by the fields of wind and air pressure variations over the Caspian Sea for 1979 from NCEP/CFSR reanalysis (Saha et al., 2010). We calculated the variance of tidal sea level variability (excluding long-period constituents) and the variance of the meteorological sea level variations in the first frequency band from 0.1 to 6 cpd and the second frequency band from 0.5 to 6 cpd. Then we estimated the relative contribution (in percent) of tides to the total sea level variance in the Caspian Sea.

The maximum contribution of tides to the total sea level variance has been located in the east part of the Middle Caspian: up to 29% for the first frequency band and up to 53% for the second frequency band. In the western part of the Southern Caspian and in Turkmen Bay the tidal contribution of total variance for the second frequency band from 0.5 to 6 cpd is up to 40%. The minimum contribution has been observed in the Northern Caspian, where strong storm surges occur; and near the Absheron Peninsula, where the amphidromic points of the diurnal and semidiurnal tides are located.
Figure 5. The relative contribution (in percent) of tides to the total sea level variance in the Caspian Sea in (a) the first frequency band from 0.1 to 6 cpd and (b) the second frequency band from 0.5 to 6 cpd.

3.3 Tidal currents

Tidal dynamics are characterized not only by sea level oscillations but also by periodic currents. Spatially, the spatial structure of the amplitudes of the semi-major semi-axis (amplitude) of tidal currents (Fig. 6) differs from the pattern of the tidal sea level amplitude distribution. Areas of the largest $M_2$ currents velocity (current speeds (semi-major semi-axis) have been observed (Fig. 4): are found in: 1) in the Mangyshlak Bay near the Tyuleniy Archipelago, up to 6.5 cm/s; 2) Absheron Strait which separates the Absheron Peninsula from the Chilov Island, up to 7.5 cm/s;
and 3) in the straits to the north and south of Ogurja Ada (the Ogurchinsky Island), up to 12.5 cm/s and 11.7 cm/s, respectively. The $M_2$ ellipse parameters (semi-major and semi-minor semi-axes, angle amplitudes, the direction of inclination, maximum current speed, phase lag) change depending on local topographic features of the water area. At the highest velocities, the rotation of the ellipse occurs in a clockwise direction. In straits and in shallow waters (for example, in the Turkmen Bay), the semi-minor semi-axis approaches zero and the tidal currents are nearly rectilinear. The spatial pattern of the $S_2$ tidal currents in the Caspian Sea has the same structure as $M_2$: the amplification areas and the ellipse parameters remain; only the $S_2$ semi-major semi-axis is 2 times weaker than the half of that of $M_2$ major semi-axis. Since $M_2$ and $S_2$ have the largest current velocities in the Caspian Sea, the spatial pattern of the maximum total tidal currents, calculated from time series computed for 18.6 years, repeats the pattern of $M_2$ again, too. The maximum total tidal current velocity in the Caspian Sea exceeds the $M_2$ velocity on average by a factor of 1.8 times. The highest velocity of the total tidal currents is observed mainly in the following straits: 1) the Mangyshlak Bay near the Tyuleniy Archipelago, up to 11.5 cm/s; 2) Absheron Strait which separates the Absheron Peninsula from the Chilov Island, up to 13 cm/s; and 3) in the straits to the north and south of Ogurja Ada, up to 22 cm/s and 19 cm/s, respectively.
Figure 4. Major semi-major axis speed magnitudes (cm/s) for M2 tidal currents. Blue ellipses indicate clockwise circulation, red ellipses are counterclockwise circulation. Other designations: CI is the Chilov (Zhiloy) Island, TG is Türkmenbaşy Gulf, TB is Turkmen Bay.

3.4 Numerical experiments with different varying MSL

Interannual The interannual MSL variability is one of the main features of the hydrological regime of the Caspian Sea (Bolgov et al., 2007). MSL variations lead to changes in the area and volume of the sea and result in changes in the frequency-selective properties of both the entire Caspian Sea and its individual parts (Fig. 5). The mean depth of the Northern Caspian is about 5–6 m and 20% of this area has a depth less than 1 m. As a result, the MSL changes of the Caspian Sea by 2–3 m (for example, from observed, e.g., between 1974 and 1994) lead to major changes in the sea dynamics of both North the Northern and as well as in coastal waters of the Middle and South the Southern Caspian. Due to long-term MSL changes, the spatial characteristics of natural oscillations of the basin (seiches) and the tidal pattern should also change. In the present study, we made numerical experiments with tidal simulations using different MSL of the Caspian Sea: from -25 m to -30 m with respect to the BHS. It is This corresponds to the natural range of MSL changes of the Caspian Sea under climatic conditions typical for the Sub-Atlantic climatic interval of the Holocene epoch ("risk zone", Bolgov et al.,
Results: The results of these experiments allowed us to estimate the changes in tidal patterns of the Caspian Sea throughout the 19th–20th centuries. Numerical results showed that MSL changes in the course of those centuries led to a significant restructuring of the spatial structure of natural sea level oscillations of the whole sea and its individual parts (Middle and Southern Caspian).
Figure 57. Changes of the mean sea level (MSL) of the Caspian Sea at Makhachkala (blue line) and Baku (red line) and the bathymetry of the sea with the MSL -26, -28, -27, and -29 m of with respect to the BHS. The orange areas are land territory created as a result of the assumed MSL changes below -25 m.

Spatial The spatial structure of the semidiurnal and diurnal tides is modified with the MSL changes of the Caspian Sea (Fig. 68). The $M_2$ amphidromic point shifts eastward by about 10 km with a decrease in the MSL from -25 to -29 m, it leads to a general displacement of the area with amplitudes of 1.5–2 cm and also to the east. As a result, the $M_2$ amplitude decreases by 0.2–0.3
cm (up to 10–20% of amplitude) along almost the entire eastern coast of the Middle Caspian. In the Southern Caspian, the tidal amphidromy also shifted to the east and amplitude increases at the amplitudes increase along the western coast of the sea. An area of amplification of semidiurnal tides with amplitudes up to 6.5 cm is formed in the Mangyshlak Bay (Northern Caspian) with the MSL of -25 m. When the MSL drops to -28 m, the amplitude in this bay decreases to 3.2 cm. An area of large amplitudes is again formed with a maximum of 5.5 cm in the Mangyshlak Bay (near the Tyuleniy Archipelago) with the MSL of -30 m.

The most interesting and complex modification of the tidal pattern occurs at the east coast of the sea. In the Türkmenbaşy Gulf, the amplitude decreases from 4.4 cm at the MSL of -25 m to 3.1 at the MSL of -29 m. The reverse picture is observed in the Turkmen Bay: the amplitude increases from 3.5 cm to 6.5 cm. The Turkmen Bay is a shallow semi-enclosed bay, with the Oguja Ada Island situated on its western border. This island is a narrow sandy spit approximately 42 km long and 1–1.5 km wide. The island’s height currently does not exceed 3–5 m (Badyukova, 2015). Thus, when the MSL of the Caspian Sea is -25 m, a significant part of the island is submerged. Results of our numerical experiments show that the presence of the island creates a western boundary in the Turkmen Bay. It leads to a change in frequency properties of the bay and as a consequence in an increase in the amplitude of semidiurnal tides.

More pronounced modification occurs in the diurnal tide pattern with the MSL changes. With the MSL of -25 m with respect to the BHS, there is a more noticeable separation of amphidromy near the Absheron Peninsula into two separate systems: to the northeast and south of the peninsula. The amplitude of the diurnal tide on the western coast of the Southern Caspian is 0.1–0.15 cm higher (up to 50% of amplitude) with the MSL of -29 m than for the MSL of -25 m. On the eastern coast of the Southern Caspian, the K1 amplitude varies weakly with the MSL changes (by 10%). However, the K1 phase lags are modified. This is caused by the influence of the Oguja Ada Island with a low MSL.

Strong modifications of the diurnal tidal pattern due to MSL changes in the mean sea level occur in the water area on the border of the North along the transition between Northern and Middle Caspian. At the MSL of -25 m, the largest amplitudes are located near the Tyuleniy Archipelago, up to 0.7–0.8 cm and in the Mangyshlak Bay, 1 cm. With decreasing MSL, large amplitudes area begins to expand to the west. At the MSL of -29 m, maximum amplitudes are already reached at the west coast of the Northern Caspian (near the Tyuleniy Island), up to 1.1 cm. These changes are apparently caused by a strong modification of the bottom topography of the shallow Northern Caspian and as a result, of the frequency (resonant) properties of this sea area subbasin.
The change in the spatial structure of the tidal range with the change in the MSL is similar to the $M_2$ amplitude pattern. The largest maximum tidal range of 22 cm is found in the Mangyshlak Bay, which is up to 22 cm (an for a MSL of -25 m). At this MSL, the tidal range in the Turkmen Bay is 13 cm; and in the Türkmenbaşy Gulf 15.5 cm. When the MSL decreases, the tidal range in Mangyshlak Bay decreases, and on the contrary, it increases in the Turkmen Bay. With the MSL of -29 m, the tidal range in the Turkmen Bay is 23 cm, whereas it is only 14 cm in the Mangyshlak Bay.
Figure 68. Tidal maps for the amplitude of harmonic $M_2$ (a, b, c), $K_1$ (d, e, f), and tidal range (g, h, i) with different MSL of the Caspian Sea: $-25\,\text{m}$ (a, d, g), $-27\,\text{m}$ (b, e, h), $-29\,\text{m}$ (c, f, i).
Changes in tidal characteristics can be very significant at individual sites. Figure 79 shows tidal vector diagrams that display the M2 model amplitude and phase lag for different sites at different MSL of the Caspian Sea. The amplitude and phase lag changes are relatively small at Makhachkala, Baku, and Bektash. However, the M2 phase lag for Ogurja Ada changes by about 100°. The M2 amplitude changes doubles: from 2.5 cm at a MSL of -25 m to 5 cm at a MSL of -30 m.

Figure 79. The changes in the amplitude (cm) and phase lag (degrees) at four sites with different MSL based on numerical modelling results.
4 Discussion

Results The results of the numerical tidal modelling in this study are in good agreement with the results of a harmonic analysis of tide gauge data of the Caspian Sea (Medvedev et al., 2017). Medvedev et al. (2017) demonstrated that the diurnal peak is absent in the sea level spectra for the western coast of the South Caspian (Baku, Svinoy Island), which is confirmed by the result of the numerical simulation of the K1 amplitude (Fig. 2a3a). Diurnal tides in the South Caspian are radiational and are formed under the influence of sea-breeze winds (Medvedev et al., 2017).

An unexpected result was obtained for the eastern part of the South Caspian. With a high MSL (for example, -25 m) a significant part of the territory of the Ogurja Ada Island is below the water level. As a result, it makes it easier for tidal waves to penetrate the Turkmen Bay. With a low MSL (for example, -29 m), the area increases significantly, and the island becomes an additional western border effective boundary to the west, reflecting the tidal waves which penetrate the Turkmen Bay. According to (Badyukova, 2015), currently, the island’s height (with the MSL of -27.5 m) does not exceed 3–5 m. According to elevation data derived from the Shuttle Radar Topography Mission (SRTM, Farr et al., 2007), the island’s maximum height is also 5-8 m. We used the GEBCO database to create our numerical grid for the model, with the island’s maximum height being the island of 2 m with a MSL of -28 m. Thus, in the experiments with the assumption of a MSL of -25 m, the island was completely submerged. According to historical records in 1835, when the MSL of the Caspian Sea was -25.5 m, the central elevated part of the island was not flooded by the sea (the maximum height being about 3.5 m). According to (Badyukova, 2015), that island actually represents preserved fragments of a coastal delta plain which built on transgressive coastal bars and subsequently merged into one island. Comparison A comparison of the island’s coordinates in 1850 with 2013 (Badyukova, 2015) shows that the island has gradually moved toward eastward and has changed its geometrical configuration due to the land at the expense of redistribution of its deposits after erosion and changes of length and configuration. The greatest role in contribution to this process belongs to eolian processes redistribution. According to (Nikiforov, 1964), from one meter of the beach every hour, 5 kg of sand is carried inland with a wind speed of 4.9 km/s.

Numerical experiments were conducted with forcing produced by synthetic wind fields in order to assess changes in natural oscillations (seiches) with a change in the MSL. Magnitude The magnitude and direction of the generated wind fields varied randomly every six hours.
Spectral analysis of the simulated wind sea level variability showed that a decrease in the MSL leads to change in the period and Q-factor of natural oscillations of the Türkmenbaşy Gulf and the Turkmen Bay. When the MSL of the Caspian Sea in the Türkmenbaşy Gulf decreases, the Q-factor of seiches in Türkmenbaşy Gulf with a period of about 12 hours significantly decreases. At the MSL of -29 m, it does not exceed any more the spectral noise level (Fig. 8a, 10a). The Q-factor of natural oscillations of the bay with a period of about 7 hours increases. In the Turkmen Bay decrease in the MSL from -26 m to -29 m, causes the spectral peak of the main seiche mode to migrate to the low-frequency band, thus the seiche period approaches the period of the harmonic M$_2$ (Fig. 8b). Apparently, this is due to elongation of the solid western boundary of the bay in the form of the Ogurja Ada Island. The closeness of the period of natural oscillations to the tidal period (12.42 h) affects the structure of the tidal oscillations, thus the “sensitivity” of the tides to changes in the MSL is determined by the proximity or distance from the natural period.

In Turkmen Bay a decrease in MSL from -26 m to -29 m causes the spectral peak of the main seiche mode to migrate towards lower frequencies. Thus the period of this seiche mode approaches the period of the harmonic M$_2$ (Fig. 10b). This is due to the progressive elongation of Ogurja Island which represents the western boundary of the bay. The closeness of the period of natural oscillations (seiches) to the tidal period (12.42 h) affects the structure of the tidal oscillations. The “sensitivity” of the tides to the changes in the MSL is determined by the proximity or distance from the natural period.

![Figure 8](image1.png)  ![Figure 10](image2.png)

Figure 810. Sea level spectra in (a) the Türkmenbaşy Gulf and (b) the Turkmen Bay at different MSL of the Caspian Sea.
5 Conclusions

In this study the tidal dynamics of the Caspian Sea have been numerically-investigated numerically. The numerical simulation was forced by the direct action of the equilibrium tidal potential. The main objective of the study was the mapping of tidal characteristics in the Caspian Sea. For the first time, it was possible to construct detailed co-tidal maps of the tidal sea level and tidal current ellipses for the major harmonics using a numerical hydrodynamic model taking into account the data of long-term sea level observations. Results The results of the numerical simulation indicate that maximum tidal amplitudes are located in the south-eastern part of the sea. We have shown that tidal currents can reach more than 20 cm/s in certain sea areas (for example, in straits), which is comparable to the magnitude of permanent sea currents. It means that the role of tides in the water dynamic hydrodynamics of isolated (non-tidal) seas seems to have been underestimated so far.

Numerical Our numerical experiments indicate that the spatial features of tides are sensitive to changes in the MSL changes. Modification A modification of the tidal pattern is caused by changes in the bathymetry and geometry of the coastline of shallow areas of the sea including the Northern Caspian, which results in significant changes in the frequency response of the basin changes significantly. It is also confirmed by the changes in the natural oscillation (seiche) structure of the Caspian Sea.

In recent decades significant progress has been achieved in the improvement of global barotropic tide models. This progress has been supported by available satellite altimetry. Stammer et al. (2014) presented a detailed comparison of the main modern global barotropic tide models. Some of these models (FES14, EOT11a, TPXO9, GOT4.10, OSU12, DTU10, HAM Tide) don’t include the Caspian Sea as well. But due to only the lower MSL of TPXO8 includes the Caspian Sea relative to the global MSL, the results for MSL of it was uncorrect, 0 m with respect to the BHS. This invalid assumption shifted the Caspian Sea in global tidal models have been seriously, coastline, significantly increased the sea area, and as a result distorted the tide in this sea.

We believe that our findings on the tidal dynamics may be quite helpful in understanding and currents in the Caspian Sea.
6 Data availability

Data and results in this article resulting from numerical simulations are available upon request from the corresponding author.

7 Author contributions

The concept of the study was jointly developed by IPIM and EK. IPIM did the numerical simulations, analysis, visualization and manuscript writing. EK prepared the numerical grids and participated in the analyses and the interpretation of the results. IF adapted the numerical Princeton Ocean Model (POM) to the Caspian Sea and participated in the verification stage. IPIM prepared the paper with contributions from EK and IF.

8 Competing interests

The authors declare that they have no conflict of interest.

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