## Reply to Editor's comments of 5 July 2020 on the paper "Predicting tidal heights for extreme environments: From 25 h observations to accurate predictions at Jang Bogo Antarctic Research Station, Ross Sea, Antarctica"

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9 Format We are grateful for the Editor's final comments on our paper. The review process has been useful in improving this 10 paper. Below we reply to the Editor's review, with each individual comment copied in blue, a response written below it, and 11 then the final modified text or figure copied below the response.

Topic Editor Decision: Publish subject to minor revisions (review by editor) (05 Jul 2020) by Philip Woodworth Comments
 to the Author: 5 July 2020

Comments on revised version of "Predicting tidal heights for extreme environments: From 25 hr observations to accurate
 predictions at Jang Bogo Antarctic Research Station, Ross Sea, Antarctica" by Byun and Hart

19 I did not look again in detail at the responses to the reviewers as I had seen them already. I just read the new version afresh 20 and below are some remarks. This draft is certainly much improved on the first one but there are some remaining things that 21 need attending to, some trivial, including obvious problems with a couple of figures. I believe when these issues are fixed the 22 next version should be fine.

*I must admit that there are aspects of the method used here I don't understand (other than where I can relate to them as a response method). I guess one has to use it to understand it properly, although I am sure it will interest other people.*

#### 27 8 - in the regional

- **Response:** This change has been made as suggested.
- **This text now reads:** "Accurate tidal height data for the seas around Antarctica are much needed, given the crucial role of these tides in the regional and global ocean, marine cryosphere, and climate processes".

#### 32 *12 - using a record from*

33 *13 - regime*34 • **Res**

- **Response:** These two changes have been made as suggested.
- **The text now reads:** "This study evaluates the ability of a relatively new, tidal species based approach, the Complete Tidal Species Modulation with Tidal Constant Corrections (CTSM+TCC) method, to accurately predict tides for a temporary observation station in the Ross Sea, Antarctica, using a record from a neighbouring reference station characterised by a similar tidal regime".

#### 40 53-67 - I must say, as I think I said before, that some of this attempted justification

- 41 *is a bit over the top. But does no harm I guess.* 
  - **Response:** This text has been shortened in response to this comment.
- The text now reads: "Floating ice shelves occupy around 75% of Antarctica's perimeter (Padman et al., 2018). Tidal oscillations at the ice-ocean interface influence the location and extent of grounding zones (Padman et al., 2002), and control heat transfer and ocean mixing in cavities beneath the marine cryosphere (Padman et al., 2018) and the calving and drift of icebergs (Rignot et al. 2000). Tides also affect variability in polynyas; seasonal sea ice patterns; and thus the functioning of marine ecosystems. And tides affect the dynamics of landfast sea ice, which provides aircraft landing zones (Han and Lee, 2018).
- Accurate Antarctic region tide data are needed for models examining changes in global climate and ocean circulation
   (Han and Lee, 2018) while coastal tide data are needed for ice mass balance and motion studies (Padman et al., 2008;
   Rignot et al. 2000; Rosier and Gudmundsson, 2018). Ice thickness is typically measured by subtracting tidal heights
   from highly accurate but relatively low resolution (temporally or spatially) satellite or in situ observations of ice
   surface elevation (Padman et al., 2008). Where ice shelves and glacier tongues occur, grounding zone and ice flexure
   mechanics make ice thickness and motion determination challenging, so that accurate tidal height inputs are crucial
   (Wild et al. 2019)".

#### 57 68 - the applicability

- **Response:** This change has been made as suggested.
- This text now reads: "In this study, we tested the applicability of Byun and Hart's (2015) CTSM+TCC method".
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61 80 and 96 - could you please make it clear what the JBARS and CR instruments are delivering i.e. real sea level or subsurface pressure? In the JBARS case it must be SSP as it is a bottom-mounted instrument, unless there is some processing to 63 remove air pressure that is not mentioned. The CR instrument is a bubbler gauge I believe, although your text does not say 64 so, so that would be delivering sea level I guess - you can check with Glen Rowe. This does not matter much for diurnal and 65 semidiurnal tides but it certainly does for the longer period ones. See below.

- Response: We agree that this was not fully clear, so the different data types have been clarified in this section (see text below) as well as attention drawn to these differences in Sect. 5.1, where they become significant in the long period tides discussion.
- Concerning the JBARS data, this text now reads: "This mission collected the first, 19 day sea level related record for JBARS: 10 min interval subsurface pressure observations were recorded between 28 Jan. and 16 Feb. 2017 using a bottom-mounted pressure sensor (WTG-256S AAT, Korea), and the data were converted to sea level heights using the hydrostatic equation. High-frequency sea level oscillations (<3 hr) were removed from the observation record using a fifth-order low-pass Butterworth filter".</li>
- Concerning the ROBT data, this text now reads: "Five minute interval seawater pressure data have been collected at ROBT since November 2011 using GEOKON 4500 series standard piezometers, vented to the atmosphere, with this data converted to sea level heights using the hydrostatic equation".

#### 78 102 - I think you mean you infer P1 from K1 and K2 from S2? It would be best to express it that way.

- **Response:** This change has been made as suggested.
- This text now reads: "Also the inference method was used to infer the P<sub>1</sub> constituent from the K<sub>1</sub>, and the K<sub>2</sub> constituent from the S<sub>2</sub>, with their amplitude ratios and phase lag differences obtained from harmonic analysis of the long-term ROBT reference station records".

105 - it would be worth adding 'at the two stations' again at the end of the sentence. When I read it first time it looked like
you were saying all the amplitudes were the same at each station.

- **Response:** This change has been made as suggested.
- This text now reads: "Analyses revealed that the two main diurnal (O<sub>1</sub> and K<sub>1</sub>) and semidiurnal (M<sub>2</sub> and S<sub>2</sub>) tides
   had similar amplitudes at the two stations, with the diurnal (semidiurnal) amplitudes being slightly larger (smaller)
   at ROBT than at JBARS, and the phase lags of all four tides having only slightly different values at the two
   stations".

#### 92 125 - from anytime --> throughout

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- **Response:** We have rewritten this sentence, dividing it sentence into two clauses, and reversing the order of the two clauses to make the meaning clearer.
- **This text now reads:** "LH<sub>r</sub> can come from any time period, but must comprise high quality (e.g. few missing data) tidal height observations throughout".

#### 98 143 - between the complete JBARS

- **Response:** This change has been made as suggested.
- This text now reads: "Comparisons were made between the complete JBARS observations and the 17 prediction data sets generated for each campaign to identify which 25 hr short-term data window produced optimal  $\eta_0(\tau)$  results".

104 157 - 'spring tide' --> 'diurnal tide'. 'Springs' and 'neaps' are usually reserved for semidiurnals, although some people like to
 105 refer to diurnal springs (not recommended).

- **Response:** Agreed we removed mention of spring tide for the diurnal record here.
- **This text now reads:** "The maximum tidal range occurred on 9 Feb., with step (ii) data slices from this date producing predictions with a low (but not the lowest) RMSE (4.81 cm)".

110 163-166 - please can you make this much clearer? What I believe you are doing is first relating 2017 JBARS to the 2017 CR
111 predictions through the method and seeing when it works best (Fig 5). Then you use the constants from the best section of
112 2017 to produce predictions for 2019. Right? (Fig 6 in principle). There is no CR data used directly in the latter. Again
113 right? This could be worded better.

- **Response:** This text of Sect. 4.1 has been improved to clarify the point about the data used, and to point out that this comment refers to the effect of using certain data in step 2 of the method on the final prediction results.
- This text now reads: "Interestingly, RMSEs and R<sup>2</sup> values between the 2019 CTSM+TCC tidal predictions and observations were almost identical to those of the 2017 comparisons, revealing that our approach performed consistently across different prediction years.
- As in the 2017 experiments, the 2019 prediction dataset made using the 8 Feb. 2017 data slices (i.e., in step (ii) of the method) produced the lowest RMSE (5.3 cm) and highest  $R^2$  (0.913) values of the 2019 experiments (Fig. 5b)".

Anyway there is a problem with figures. Fig 6 has 2019 in the header but the figure itself is the same as Fig 5a. And then the end of this para refers to Fig 7 but that is for lunar declination and would be better including in the next section.

124 **Response:** Fig. 6 illustrated the same sort of results as in Fig. 5a, with very slight differences in the RMSEs (that were too small to pick up visually) and the same  $R^2$  values – so we have deleted the original Fig. 6 as we now see it 125 was not necessary to explain our key points here. We also swapped around order of Fig. 5. and 7. So that the 126 127 example prediction versus observation results figure now comes first (now Fig 5., formerly Fig.7) and the RMSE 128 evaluation of the predictions versus observations comes second (Fig.6 now, formerly Fig. 5). Fig. 8, which 129 focusses on lunar declination, is introduced in the next section - Sect. 4.2. 130 131 Section 4.2 - again the figures do not seem to be assigned to the text as well as they might. 132 Response: This comment made us reflect carefully on the results figure placement and order. This was also required given the recommended (and accepted) deletion of several figures. As such, we have carefully improved 133 134 the placement of references to figures in Sect 4.1 and 4.2. 135 The placement in Sect. 4.2 of Fig. 7 is now as follows: "Similarly, in a diurnal tide regime or a mixed, mainly 136 diurnal tide regime, preferred temporary station observation days can be estimated based on the lunar declination 137 (Fig. 7), which varies at a period of 13.66 days. That is, maximum tidal range days can be estimated for JBARS 138 based on the day of the Moon's greatest northern (GN) and southern (GS) declinations. The time between the 139 Moon's semi-monthly GN and GS declinations and their effects on tidal range, called the age of diurnal inequality 140 (ADI), is commonly 1 to 2 days. The GN and GS lunar declinations during our temporary station summertime 141 observation periods occurred on 8 Feb. 2017 (GN) and on 6 Jan. 2019 (GS) respectively (Fig. 7), with the 142 maximum diurnal tides at JBARS expected around 1 day after each lunar declination peak". 143 144 145 Also it occurred to me at this point - why is 25 hr so important? It is obviously the minimum for looking at one tidal cycle, 146 although you don't actually say that anywhere. But how well would say short records of 50 or 100 hours do? 147 **Response:** Yes – you are right that we omitted to explain this point explicitly. Please see below new text inserted 148 into Sect. 2, where we discuss the data sets used. This text now reads: "Note that short-term records >25 hr may be used in CTSM+TCC but, as demonstrated in 149 150 Byun and Hart (2015), large tidal range (range being twice amplitude) and high data quality have a much greater 151 positive impact on prediction results than any increase in the length of the concurrent short-term records employed". 152 153 198 - it would be worth saying somewhere that this negative AT is very unusual. You do say somewhere that AT is 1 or 2 154 days in most parts of the world so this aspect of this area is worth noting. Response: A small additional note has been added as suggested. 155 This text in Sect. 4.3 now reads: "Note that the negative AT values in Table 1 are an unusual feature of the Ross 156 157 Sea tides, given that elsewhere spring tides commonly occur a day or two after the full and new moon". 158 159 215 - better to say: Srun excluded all long-period tides (i.e. the 6 listed in Table 2) 160 216 drop 'the' 217 ditto 161 and I think you mean 'And the CSTM+TCC repeated in each case'. 162 **Response:** These changes have been made as suggested. 163 This text now reads: 164 "To investigate the main cause of the apparent fortnightly prediction biases in our results, we examined the effects of 165 two fortnightly tidal constituents (Mf, and MSf) at ROBT using T\_TIDE. Three 2019 tidal prediction experiments were 166 conducted: 167 Srun excluded all long-period tides (see list of exclusions in Table 2); • 168 Run1 was based on Srun but also incorporated M<sub>f</sub>; and • Run2 was based on Srun but also incorporated Mf and MSf; 169 with T\_TIDE predictions made for each case. Comparisons between Run1 and Srun predictions revealed that exclusion 170 171 of the  $M_{\rm f}$  tide (2.7 cm amplitude) can produce prediction biases during periods of lunar declination change, with 172 comparisons between Run2 and Run1 results revealing that the additional exclusion of the MS<sub>f</sub> tide (1.2 cm amplitude) 173 intensifies the biases. While these results elucidate an issue with predicting Ross Sea tides based on the diurnal and semidiurnal species alone, the aforementioned differences in gauge and record types in themselves can also result in 174 175 different harmonic analysis results and, in turn, different prediction results.". 176 177 But please also see above my question at line 80 about SSP and sea level. If you have sea level at CR and SSP at JBARS then 178 you cannot assume the long period tides (especially Sa and Ssa) are the same. 179 **Response:** Yes – thank you for raising this valid point. There is a difference in the JBARS and ROBT 180 measurements and therefore their optimal usage. To fully address this issue we have now: added a note to acknowledge the different measurement types in the text describing the observations 181 0 182 collection (please see earlier in this file); 183 added the note regards the Srun/ Run1/ Run2 experiment findings (please see response just above); and 0 in addition, we have now added a note about the ROBT record used and its limitations in terms of 184 0 185 atmospheric 'noise' in the sea level record. This text of Sect. 5.1 now reads: "Table 2 summarises the characteristics of 6 long-period tides (Sa, Ssa, MSm, Mm, 186 187 M<sub>f</sub>, MS<sub>f</sub>) at the ROBT station, derived from tidal harmonic analysis of year-long (2013) in situ observation records. 188 Note that since the ROBT observation record was derived from seawater pressure measurement, and thus includes proportionately large non-tidal (atmospheric) sea level variations, caution should be exercised in comparing the 189 190 harmonic analysis results of the non-astronomical constituents, which are affected by seawater density and 191 atmospheric forcing (i.e.  $S_a$  and  $S_{sa}$ )".

221 - why does an ice sheet respond only to MSf? Surely it responds to them all? 193

- 194 **Response:** Ice shelves respond at several tidal frequencies, including the MS<sub>f</sub>, as highlighted by Rosier and Gudmundsson (2008, p. 1709) in their paper: "The non-linear rheology of ice means that, as an ice shelf bends to 195 accommodate vertical tidal motion, stresses generated in the grounding zone reduce the effective viscosity of ice. 196 197 This leads to modulation of ice shelf velocity at a number of frequencies, including the MS<sub>f</sub> frequency, which is 198 readily observed on many Antarctic ice shelves (King et al., 2011; Minchew et al., 2016; Gudmundsson et al., 2017; 199 Rosier et al., 2017a)".
- Our text reads: "Rosier and Gudmundsson (2018) found that ice flows are modulated at various tidal frequencies, 200 201 including that of the MS<sub>f</sub> tide".
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203 228-238 - surely the main thing is the amplitudes of K1 and O1 are almost the same, so they will double or cancel over a 204 fortnight, and when they are cancelling then the small semidiurnals will manifest themselves. Obvious really. 205

Response: Yes agreed. We have modified the text to emphasis this main (simpler) explanation.

This text now reads: "The combination of these out-of-phase tidal species generates double peaks (or double troughs) around low and high tide (Fig. 8b) for periods when the diurnal tide amplitude is low, due to the similar amplitude  $K_1$  and  $O_1$  tides cancelling each other out across a fortnight, allowing the combined  $M_2$  and  $S_2$  amplitudes to temporarily approach or exceed that of the combined K1 and O1 tides (Fig. 8c)".

#### 211 243 - please can you put underflow and overflow arrows on the colour scale of Fig 11? As it stands there is no area allowed below 0.25 or above 3 in the plot. Also there is a red (?) blob half way down the east coast of the peninsula that draws the 212 eye and you don't comment on. 213

- **Response:** These figure changes have been made as suggested, and the red blob you mention has now been explicitly mentioned and explained.
- This text now reads: "The Weddell Sea is dominated by mixed, mainly semidiurnal tides, excepting the 216 217 semidiurnal area mentioned and another small area exhibiting diurnal tides (F>3) at around 76.5°S, where 218 amphidromic points (i.e. zero amplitudes) occur for both the M<sub>2</sub> and S<sub>2</sub> tides".
- 219 Figure 9 (formerly Fig. 11) now looks like this:



221 Figure 9. Distribution of tidal form factor (F) values around Antarctica. Note the magenta area (72°S) on the Antarctic Peninsula's 222 Weddell Sea coast denotes the only area with a properly semidiurnal tide regime (F < 0.25) in the Antarctic region.

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224 251-256 - these lines are just repeating yourself.

- **Response:** We agree that this was repetitive, considering what is stated earlier in the introduction and in the subsequent conclusion section, so have deleted the original lines 253 to 256.
- 227 The ending of Section 5.2 now reads: "As the nodal amplitude factor variations of the diurnal and semidiurnal 228 tides are out of phase, this leads to differing tidal responses around Antarctica over 18.61 years, particularly between the Ross and Weddell Seas (see details for ROBT in Byun and Hart, 2019). Given that CTSM+TCC is 229 230 based on modulated tidal amplitude and phase lag corrections for each diurnal and semidiurnal species, this

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233 402 - records --> record (three times)

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#### 404 - you mean P1 from K1 and K2 from S2. Your use of brackets is confusing.

- **Response:** These changes and clarifications have been made as suggested.
- This caption text now reads: "Table 1. Major tidal harmonic results for diurnal and semidiurnal constituents from harmonic analyses of sea level observations: the year-long (2013) record from Cape Roberts (ROBT), and 17.04 day record (29 Jan. to 15 Feb. 2017) and 20.54 day record (29 Dec. 2018 to 18 Jan. 2019) from Jang Bogo Antarctic Research Station (JBARS) in the Ross Sea (see source details in Sect. 2). For the JBARS tidal harmonic analyses, the inference method was used to infer the P<sub>1</sub> constituent from the K<sub>1</sub>, and the K<sub>2</sub> constituent from the S<sub>2</sub>, with their amplitude ratios and phase lag differences obtained from harmonic analysis of the long-term ROBT 2013 reference station record".

approach is applicable in studying a continent with such a diversity of tidal regime types".

244 408 - from a harmonic analysis of one year-long

- **Response:** This change has been made as suggested.
- **This text now reads:** "Table 2. Harmonic constants for 6 long-period tidal constituents, derived from harmonic analysis of one year-long observations (2013) measured at the Cape Roberts sea level gauge (ROBT), using T\_Tide (Pawlowicz et al., 2002). Note that this gauge is a vented piezometer so caution should be exercised in interpreting the results (particularly those for S<sub>a</sub> and S<sub>sa</sub>) given the inclusion of proportionately large non-tidal (atmospheric) variations in this kind of sea level record".

### Please see my question at line 80. Sa and Ssa will be very different if SSP or sea level are being used, although I believe if CR is a bubbler gauge then it should be sea level ok - check with Glen.

• **Response:** Thank you for raising this point – please see paper adjustments detailed earlier in this file.

#### 256 Fig 2b - the lons and lats are upside down

- **Response:** Yes the orientation of these have been corrected.
- Figure 2 now looks like:



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261 *Fig 4f - degrees is missing from y-title* 

261 11g 4J acgrees is missing from y title 262 430 - of the entire 369 ... record ... entire 17 ... record

- **Response:** These figure and caption corrections have been made as suggested.
- Figure 4 and its caption now look like:
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Figure 4. Daily amplitudes (a, c); phase lags (b, d); amplitude ratios (e); and phase lag differences (f) of the  $K_1$  and  $M_2$  tides (representative diurnal and semidiurnal tide species) at ROBT (a, b) and JBARS (c, d), and between JBARS and ROBT (e, f), calculated from 'daily' slices of the 29 Jan. to 14 Feb. 2017 ROBT tidal predictions and JBARS sea level observations. In addition, thick blue ( $K_1$ ) and thin pink ( $M_2$ ) horizontal lines in the panels indicate the amplitudes and phase lags derived from harmonic analyses of the entire 369 day 2013 ROBT sea level record (a, b) and of the entire 17 day 2017 JBARS sea level record (c, d), along with their amplitude ratios and phase lag differences (e, f).

#### 273 274 *Fig 5 - drop*

• **Response:** The top half (5a) of this figure illustrated the same idea (and was based on very similar, but not identical, RMSE data) to that in Fig. 6. Since Fig. 5a was explained first in the paper, we have kept Fig. 5a, and instead deleted Fig. 6 (please also see above notes on swapping order of the old Fig. 5 and Fig. 7).

#### 279 440 - record

- **Response:** This change has been made as suggested
- The Fig. 6 (formerly Fig. 5) caption now reads: "Figure 6. (a) Time series (29 Jan. to 14 Feb. 2017) of Root 281 Mean Square Errors (RMSE, thick blue line with •) and coefficients of determination (R2, thin black line with •) 282 283 between JBARS 10 min interval sea level observations and the CTSM+TCC prediction datasets, generated for this 284 site using harmonic analysis results from the JBARS daily (25 hr) sea level data slices and concurrent daily (25 hr) 2017 tidal prediction data slices and harmonic analysis results from ROBT station's year-long (2017) tidal 285 predictions. (b) Time series of predicted 2017 tidal heights (thin blue line) and daily tidal ranges (thick black line 286with •) for ROBT, based on harmonic analysis of this station's 2013, 5 min interval sea level record, plus an 287 288 indication of the moon's phase and declination."

#### 290 Fig 6 - is the same as Fig 5a

• **Response:** Please see above where we explain why we deleted the former Fig. 6. We also adjusted the figure

numbering throughout according to this deletion (and also due to the recommended deletion below of Fig. 9).

294 452/3 - drop the outer set of (). Too many of them.

- **Response:** This change has been made as suggested.
- This text now reads: "The JBARS predictions were generated via the CSTM+TCC method using a daily (25 hr) slice of local sea level observations from 8 Feb. 2017 (dashed box in (a)), along with concurrent (to time periods a and b) ROBT predictions; and year-long (2017) 5 min interval ROBT tidal predictions".

Fig 9 - this is trivial stuff, isn't it? (a) and (b) are anyway near identical and when reduced for print you won't see any
 difference. I think I would drop this figure but see what you think

- Response: Yes fair point. This is better explained with just the existing 3 lines. We have deleted Fig. 9 as
   suggested. We swapped the word 'showing' for 'revealing' since we don't *show* any figure now and just report on what the test results revealed.
- The remaining (now unillustrated) text now reads: "Comparisons between *Run1* and *Srun* predictions revealed that exclusion of the M<sub>f</sub> tide (2.7 cm amplitude) can produce prediction biases during periods of lunar declination change, with comparisons between *Run2* and *Run1* results revealing that the additional exclusion of the MS<sub>f</sub> tide (1.2 cm amplitude) intensifies the biases".

309310 *Fig 11 - see comments above* 

• **Response:** Thank you – these have all been attended to above, improving this figure (now numbered Fig. 9).

# 313 Predicting tidal heights for extreme environments: From 25 hr 314 observations to accurate predictions at Jang Bogo Antarctic Research 315 Station, Ross Sea, Antarctica

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320 Abstract. Accurate tidal height data for the seas around Antarctica are much needed, given the crucial role of these tides in 321 the regional and global ocean, marine cryosphere, and climate processes. However obtaining long term sea level records for 322 traditional tidal predictions is extremely difficult around ice affected coasts. This study evaluates the ability of a relatively 323 new, tidal species based approach, the Complete Tidal Species Modulation with Tidal Constant Corrections (CTSM+TCC) 324 method, to accurately predict tides for a temporary observation station in the Ross Sea, Antarctica, using a records from a 325 neighbouring reference station characterised by a similar tidal regime. Predictions for the 'mixed, mainly diurnal' regimes of Jang Bogo Antarctic Research Station (JBARS) were made and evaluated based on summertime (2017; and 2018 to 2019) 326 327 short-term (25 hr) observations at this temporary station, along with tidal prediction data derived from year-long observations 328 (2013) from the neighbouring 'diurnal' regime of Cape Roberts (ROBT). Results reveal the CTSM+TCC method can produce accurate (to within ~5 cm Root Mean Square Errors) tidal predictions for JBARS when using short-term (25 hr) tidal data from 329 330 periods with higher than average tidal ranges (i.e. those at high lunar declinations). We demonstrate how to determine optimal 331 short-term data collection periods based on the Moon's declination and/or the modulated amplitude ratio and phase lag difference between the diurnal and semidiurnal species predicted from CTSM at ROBT (i.e. the reference tidal station). The 332 333 importance of using long period tides to improve tidal prediction accuracy is also considered and, finally, the unique tidal 334 regimes of the Ross Sea examined in this paper are situated within a wider Antarctic tidal context using FES2014 model data.

#### 335 Copyright statement (will be included by Copernicus)

#### 336 1 Introduction

Conventionally, year-long sea level records are used to generate accurate tidal height predictions via harmonic methods (e.g. 337 338 Codiga, 2011; Foreman, 1977; Pawlowicz et al., 2002). Obtaining long term records for such tidal analyses is extremely 339 difficult for sea ice affected coasts like that surrounding Antarctica. As a compliment to in situ tidal records, recent work has significantly advanced our understanding of tide models for the shallow seas around Antarctica and Greenland via the 340 341 assimilation of laser altimeter data and use of Differential Interferometric Synthetic Aperture Radar (DInSAR) imagery, amongst other methods (Padman et al., 2008; 2018; King et al., 2011; Wild et al., 2019). However, Byun and Hart (2015) 342 343 developed a new approach to successfully predict tidal heights based on as little as 25 hr of sea level records when combined with neighbouring reference site records, using their Complete Tidal Species Modulation with Tidal Constant Corrections 344 (CTSM+TCC) method, on the coasts of Korea and New Zealand. Demonstrating the usefulness of this method for generating 345 accurate tidal predictions for new sites on sea ice affected coasts is the motivation for this study. We focus on the Ross Sea, 346 347 Antarctica, as our case study area.

Long-term, quality sea level records in the Ross Sea are few and far between, and include observations from gauges operated by New Zealand at Cape Roberts (ROBT); by the United States in McMurdo Sound (see reference to data in Padman et al., 2003); and by Italy at Mario Zucchelli Station (Gandolfi, 1996), all in the eastern Ross Sea. Permanent sea level gauge

351 installations in this extreme environment must accommodate or somehow avoid surface vents freezing over with sea ice, and 352 damage to subsurface instruments from icebergs. There is also the challenge of securing and preventing damage to the cables 353 that join the subsurface instruments to their onshore data loggers and power supplies, across the seasonally dynamic and harsh 354 coastal and subaerial environments of Antarctic shorelines. At ROBT, these issues have been avoided by sheltering the sea 355 level sensor towards the bottom of a 10 m long hole, drilled through a large, shoreline boulder, from its surface  $\sim 2$  m above 356 the sea and sea ice level, to ~6 m below sea level, below the base of the sea ice (Glen Rowe, Technical Leader Sea Level Data, New Zealand Hydrographic Authority, pers. comm. 13 Dec. 2019). In the absence of a suitable permanent gauge site, 357 358 hydrographic surveys have been conducted at the Korean Jang Bogo Antarctic Research Station (JBARS). Such surveys are 359 best conducted during the summertime predominantly sea ice free window around mid-January to mid-February. Even then, 360 mobile ice (Fig. 1) and severe weather events frequently hinder surveys via instrument damage or loss, not to mention the logistical difficulties of instrument deployment and recovery (Rignot et al. 2000). Accurate tidal records from the Ross Sea 361 362 and other areas around Antarctica are thus scarce compared to those available from other regions, although these data are much 363 needed given the crucial role of tidal processes around this continent (Han et al., 2005; Jourdain et al., 2018; Padman et al., 2003; 2018). 364

Floating ice shelves occupy around 75% of Antarctica's perimeter (Padman et al., 2018). Tidal oscillations at the ice-ocean interface influence the location and extent of grounding zones (Padman et al., 2002; Rosier and Gudmundsson, 2018), and control heat transfer and ocean mixing in cavities beneath the marine cryosphere (Padman et al., 2018; Wild et al., 2019) and the calving and subsequent drift of icebergs (Rignot et al. 2000). Tides also affect variability in polynyas; patterns of seasonal sea ice\_patterns; and thus the functioning of marine ecosystems. And tides affect the dynamics of landfast sea ice, which provides aircraft landing zones\_for Antarctic science operations (Han and Lee, 2018).

371 Accurate Antarctic region tideal input data are needed for models examining changes in global climate and ocean circulation, 372 including for the generation of Antarctic bottom water (Han and Lee, 2018; Wild et al., 2019) while c. Data on coastal tide 373 datas are also essential needed for studies of ice mass balance and motion studies (Han and Lee, 2018; Padman et al., 2008; 374 2018; Rignot et al. 2000; Rosier and Gudmundsson, 2018; Wild et al., 2019). Ice thickness is typically measured byvia the 375 subtractingion of tidal heights oscillations from highly accurate, but relatively low frequencyresolution (temporally or 376 spatially) satellite or, satellite in situ based observations of ice surface elevation and/or from in situ Global Navigation Satellite 377 System (GNSS) instrument observations (Padman et al., 2008). For floating ice, this procedure is relatively straightforward 378 butW-where ice shelves and glacier tongues occur, the mechanics of grounding zones and ice flexure mechanics render-make 379 the determination of ice thickness and motion determination challenging-(Padman et al. 2018; Rosier and Gudmundsson, 380 2018), making so that the accuratecy of tidal height inputs are crucial for effective ice modelling (Wild et al. 2019).

In this study, we tested <u>the</u> applicability of Byun and Hart's (2015) CTSM+TCC method in an extreme observation environment using 25 hr short-term records from JBARS, our temporary tidal observation station, and year-long data from ROBT, the neighbouring reference station. Sect. 2 of this paper details the JBARS and ROBT observation data sets used to generate harmonic tidal analysis results and CTSM+TCC tidal predictions. Sect. 3 explains how the CTSM+TCC method was applied and adapted in this case study (with Appendix 1 detailing the calculations), while Sect. 4 demonstrates the CTSM+TCC tidal prediction capability. Sect. 5 discusses the generation of fortnightly tide effects and double tidal peaks; and situates the Ross Sea tides examined in this paper within the wider context of Antarctic tidal regimes.

#### 388 2 Antarctica's major tides: Observations and background

#### 389 2.1 Study sites and data records

- 390 The Korea Hydrographic and Oceanographic Agency (KHOA) survey team went to JBARS in Northern Victoria Land's Terra
- 391 Nova Bay, Ross Sea, Antarctica, in the austral summertime of 2017 (Fig. 2) for a preliminary fieldtrip to conduct hydrographic

- 392 surveys and produce a nautical chart. This mission collected the first, 19 day sea level related records for JBARS: 10 min 393 interval subsurface pressure observations were-data, recorded between 28 Jan. and 16 Feb. 2017 using a bottom-mounted 394 pressure sensor (WTG-256S AAT, Korea) with the data converted to sea level heights using the hydrostatic equation. High-395 frequency sea level oscillations (<3 hr) were removed from the observation record using a fifth-order low-pass Butterworth 396 filter. Note that the first and last days of this campaign comprised partial day records, so we excluded these end days from our 397 tidal prediction experiments, since our method requires continuous 25 hr input data (i.e. covering one tidal cycle minimum 398 and, for convenience, starting at midnight). That left 17 days and 1 hour of useable tidal observation data as the basis of the 399 primary JBARS observation record. Note that short-term records >25 hr may be used in CTSM+TCC but, as demonstrated in 400 Byun and Hart (2015), large tidal range (range being twice amplitude) and high data quality have a much greater positive 401 impact on prediction results than any increase in the length of the short-term observation records employed.
- 402 For the purposes of a full-scale survey, three additional, discontinuous sea level observation records were measured by KHOA 403 at JBARS between 29 Dec. 2018 and 11 Mar. 2019, all at 10 min intervals using the same instrument. Of these, the 20.54 day 404 record produced between 29 Dec. 2018 and 18 Jan. 2019 comprised relatively high quality data with small residuals (i.e. 405 observations minus predictions). We used this additional dataset (hereafter referred to as the JBARS 2019 observations) to 406 verify CTSM+TCC method tidal predictions generated from input parameters derived from 'daily' (25 hr) slices of the 2017 407 sea level records. Due to the short duration of the KHOA survey team's forays into the Ross Sea, and in the absence of a 408 permanent tide station at JBARS, it was not possible to collect the year-long sea level records that are commonly employed to 409 obtain reliable tidal harmonic constants for tidal prediction.
- Approximately 269 km south of JBARS, there is a permanent tidal observation station named after its location on Cape Roberts
  (ROBT), operated by Land Information New Zealand (LINZ) and recording at intervals since November 1990 (Fig. 2). Five
  minute interval sea<u>water level-pressure</u> data have been collected at ROBT since November 2011 using Standard Piezometers
- 413 (GEOKON Model 4500 series Standard Piezometers, GEOKON), vented to the atmosphere, with this data converted to sea
- 414 level heights using the hydrostatic equation. Part of the 2017 record from this site was unavailable online at the time of starting
- 415 this research, so instead we chose as our reference records the 2013 ROBT sea level data, a quality year-long dataset with few 416 missing points.

#### 417 **2.2 Tidal characteristic analyses and descriptions**

418 Using the T\_TIDE toolbox (Pawlowicz et al., 2002), we obtained the tidal harmonic constants of the 8 and 6 major tidal 419 constituents for ROBT and JBARS, respectively (Table 1). Also the inference method was used to infer the P1 constituent from 420 the K<sub>1</sub>, and the K<sub>2</sub> constituent from the S<sub>2</sub><del>separate out neighbouring diurnal (K<sub>1</sub> and P<sub>1</sub>) and semidiurnal (S<sub>2</sub> and K<sub>2</sub>) tide</del> 421 constituents, with their amplitude ratios and phase lag differences obtained from harmonic analysis of the long-term ROBT 422 reference station records. Analyses revealed that the two main diurnal ( $O_1$  and  $K_1$ ) and semidiurnal ( $M_2$  and  $S_2$ ) tides had 423 similar amplitudes at the two stations, with the diurnal (semidiurnal) amplitudes being slightly larger (smaller) at ROBT than 424 at JBARS, and the phase lags of all four tides having only slightly different values at the two stations. The amplitude differences 425 result in slightly different tidal form factors at the two sites (e.g., F in Table 1).

#### 426 3 Using the CTSM+TCC tidal prediction methodology in the Ross Sea

427 Having analysed the tidal harmonic constants at the two stations, we then employed the CTSM+TCC method (Byun and Hart,

428 2015) to generate tidal height predictions for JBARS, our 'temporary' tidal observation station (subscript o), using ROBT as

429 the 'reference' station (subscript *r*). This prediction approach (see Appendix 1 for the detailed calculations, and Byun and Hart

430 (2015) for explanation of procedure development) is based on:

- 431 (i) using long-term (1 year, in our case) reference station records  $(LH_r)$  and CTSM calculations to make an initial
- 432 anytime ( $\tau$ ) tidal prediction ( $\eta_r(\tau)$ ), which involves summing tidal species' heights for the reference station (Fig.3); 433 (ii) comparing the tidal harmonic constants (amplitude ratios and phase lag differences) of representative tidal 434 constituents (e.g., M<sub>2</sub> and K<sub>1</sub>) for each tidal species between the temporary and reference stations (Fig. 4), calculated 435 using T\_TIDE and concurrent short-term records ( $\geq$ 25 hr duration, starting at midnight) from the temporary (SH<sub>o</sub>)

436 and reference (SH<sub>r</sub>) stations; and

437 (iii) using the step (ii) comparative data and the TCC calculations for each tidal species to adjust the  $\eta_r(\tau)$  tidal species' 438 heights in order to generate accurate, anytime tidal height predictions for the temporary tidal station ( $\eta_o(\tau)$ ).

439 In this Ross Sea case study we used the 2017 JBARS tidal observation records (i.e. 17.04 days from 00:00 29 Jan. to 01:00 15 440 Feb.) as a source of  $SH_o$ , keeping the second JBARS 2019 observation record for evaluation purposes.

Importantly, this method assumes that the reference and temporary tidal stations are situated in neighbouring regimes with similar dominant tidal constituent and tidal species characteristics, and that the tidal properties between the two stations remain similar through time. As explained above, both JBARS and ROBT have tidal regimes that are primarily dominated by diurnal tides. LH<sub>r</sub> can come from any time period, but must comprise high quality (e.g. few missing data) tidal height observations throughoutfrom anytime.

446 Byun and Hart (2015) recommended the use of short-term records gathered during periods of calm weather, to minimise errors 447 due to atmospheric influences. They employed observational data for both  $SH_{o}$  and  $SH_{r}$  but as demonstrated in this paper the method can also be applied using tidal predictions as a source of  $SH_r$ . This adjustment in approach arose since for the 2017 448 449 JBARS observation time period, the concurrent 2017 ROBT records available online (LINZ, 2019) had multiple missing data. 450 We solved this issue by producing a year-long synthetic 2017 record for ROBT using T\_TIDE (Pawlowicz et al., 2002) and the 2013 (i.e. LHr) observational record as input data. The 17.04 days of predicted tides that were concurrent with the 2017 451 452 JBARS observation record were then used as our  $SH_r$  source. While this CTSM+TCC method adjustment was procedurally small, it represents an important adaptation in the context of generating tidal predictions for stations situated in extreme 453 454 environments, since concurrent temporary and reference station observations might be rare in such contexts.

455 When using CTSM+TCC, if the available temporary tidal station observation record covers multiple days, it is best practice 456 to experiment by generating multiple  $\eta_o(\tau)$ , each using different concurrent pairs of SH<sub>o</sub> and SH<sub>r</sub> daily data slices in step (ii) 457 above, to produce daily amplitude ratios and phase lag differences between the two stations for the diurnal  $K_1$  and semidiurnal 458  $M_2$  tidal constituents. Comparisons are then made between the different  $\eta_0(\tau)$  data sets produced and the original temporary 459 station observations, to determine the optimal 25 hr window to use: once selected, tidal height predictions can be generated for the temporary observation station for any time period. Thus, 17 individual 25 hr duration data slices were clipped from the 460 461 2017 JBARS observation records and from the concurrent ROBT predictions, forming 17 pairs of  $SH_{e}$  and  $SH_{r}$  'daily' slices. Each paired data set was then used with LHr to generate tidal height predictions for JBARS covering both the 2017 and 2019 462 463 KHOA observation campaign time periods. Comparisons were made between the complete JBARS observations and the 17 464 prediction data sets generated for each campaign to identify which 25 hr short-term data window produced optimal  $\eta_o(\tau)$ 465 results.

#### 466 4 Results

#### 467 4.1 Tidal prediction evaluation

468 CTSM+TCC was used to produce 17 different JBARS tidal prediction data\_sets for the period 29 Jan. to 14 Feb. 2017, based 469 on harmonic analysis results of the 'daily' (25 hr)  $K_1$  and  $M_2$  amplitudes and phase lags at our two tidal observation stations 470 (Fig. 4). Figure 5a illustrates one such tidal height prediction data set, in comparison to the observed tides. In order to evaluate

- 471 the <u>17 different prediction results</u>, each predictioned tidal height data set data set was compared with the concurrent
- 472 JBARS field observations via Root Mean Square Error (RMSE) and coefficient of determination (R<sup>2</sup>) statistics.
- 473 As illustrated in Fig. 5, the RMSE and R<sup>2</sup> results varied in relation to the JBARS tidal range (range being twice amplitude),
  474 with greater accuracy evident in predictions made using data derived from periods with above average tidal ranges.
- 475 In the JBARS area of the Ross Sea during our 2017 observation period, above average tidal ranges corresponded to the period
- 476 when the moon was near its greatest northern declination. RMSEs between the 2017 observations and predictions ranged from
- 477 4.26 cm to 20.56 cm, while R<sup>2</sup> varied from 0 to 0.94, across the 17 'daily' experiments (Fig. 6). Eleven of the experiments
  - 478 produced accurate results (i.e. excluding those derived from 31 Jan.; and 1 to 4 and 14 Feb. data slices). Daily datasets from
    479 periods with relatively high tidal ranges (>83.5 cm) produced predictions with RMSEs <5 cm and R<sup>2</sup> values >0.92. The
    480 maximum <u>tidal spring tidal range</u> occurred on 9 Feb., <u>with: the step (ii)</u> data slices from this date producinged predictions with
    481 a low (but not the lowest) RMSE (4.81 cm). The predictions with the lowest RMSE (4.259 cm) and highest R<sup>2</sup> value (0.941)
    482 were produced using data slices from one day earlier, 8 Feb. 2017 (Fig. 5a and Fig. 6). In contrast to the majority of successful
    483 experimentsprediction datasets, that e experiment based on using the data derived from the 2 Feb. 2017 data slices in step (ii)
  - $\frac{484}{485} = \frac{11.95}{485} \text{ cm} + \frac{11.9$
  - 486 Interestingly, RMSEs and R<sup>2</sup> values between the 2019 CTSM+TCC tidal predictions and observations were almost identical
  - 487 to those of the 2017 comparisons, revealing that our approach performed consistently across different prediction years.
  - <u>As in the 2017 experiments, the 2019 prediction dataset made using the 8 Feb. 2017 data slices (i.e., in step (ii) of the method)</u>
    produced the lowest RMSE (5.3 cm) and highest R<sup>2</sup> (0.913) values of the 2019 experiments (Fig. 5b).
  - Across both the 2017 and 2019 prediction time periods, the RMSE and R<sup>2</sup> results varied in relation to the JBARS tidal range,
    with greater accuracy evident in predictions made using step (ii) 2017 data slices from periods with above average tidal ranges.
    In the JBARS area of the Ross Sea during the 2017 short-term observation period, above average tidal ranges corresponded to
    the period when the moon was near its greatest northern declination (Fig. 6). As with the 2017 predictions, RMSEs between
    the 2019 predictions and observations were lower when generated using data slices from 2017 periods at high lunar declination
    (Fig.6). For example, 2019 predictions made using input data derived from the 8 Feb. 2017 data slices produced the lowest
    RMSE (5.3 cm) and highest R<sup>2</sup> (0.913) values of the 2019 experiments (Fig. 7).
  - 497

498 <u>Collectively t</u>These results show that the CTSM+TCC method can be used successfully to predict tidal heights for JBARS, 499 when using <u>short-term</u> observation records gathered from periods at high lunar declination, (sometimes called tropic tides and 500 <u>thus above average tidal ranges</u>), with relatively calm weather, together with observation or prediction records from the 501 neighbouring reference station ROBT.

#### 502 4.2 Determining the ideal short-term sea level observation period when using CTSM+TCC

The previous section verified that the CTSM+TCC method can be used to generate accurate tidal predictions based on 25 hr sea level records, from periods with above average tidal ranges, for a temporary station in a mixed, mainly diurnal regime and a reference station in a diurnal regime. The question arises as to how to determine optimal observation days in such settings to produce the most accurate tidal predictions.

507 For semidiurnal or mixed, mainly semidiurnal tidal regimes, we can estimate preferred temporary station observation days,

508 those with the largest tidal ranges, based on the moon's phase, without reference to tide tables. That is, spring tides commonly

509 occur just a day or two after the full and new moon, which reoccurs at a period of 14.76 days. The time lag between the full or

510 new moon and the spring tide is called the age of the tide (AT).

- 511 Similarly, in a diurnal tide regime or a mixed, mainly diurnal tide regime (Fig. 5), preferred temporary station observation
- 512 days can be estimated based on the lunar declination (Fig. 7), which varies at a period of 13.66 days. That is, maximum tidal

- 513 range days can be estimated for JBARS based on the day of the Moon's greatest northern (GN) and southern (GS) declinations.
- 514 The time between the Moon's semi-monthly GN and GS declinations and their effects on tidal range, called the age of diurnal
- 515 inequality (ADI), is commonly 1 to 2 days. As shown in Fig. 8, T-the GN and GS lunar declinations during our temporary
- 516 station summertime observation periods occurred on 8 Feb. 2017 (GN) and on 6 Jan. 2019 (GS) respectively (Fig. 7), with the
- 517 maximum diurnal tides at JBARS expected around 1 day after each lunar declination peak.
- 518 Thus, when planning to use the CTSM+TCC tidal prediction method for places characterised by diurnal or mixed, 519 predominantly diurnal tidal regimes, we can use knowledge of the moon's declination to select potential sea level observation 520 days.

#### 521 4.3 Comparison of ROBT and JBARS tidal species characteristics

The CTSM+TCC tidal prediction method is based on the assumption that the tidal harmonic characteristics of each tidal species are very similar between the temporary and reference stations. This is because the reference station tidal species' CTSMs form the basis of the tidal predictions for the temporary observation station. To test the validity of this assumption, we examined the phase lag (*G*) differences of the two major diurnal and semidiurnal tidal constituents using *ADI* and *AT*, calculated as:

526 
$$ADI(day) = \left(\frac{G_{K_1}-G_{O_1}}{\omega_{K_1}-\omega_{O_1}}\right)/24$$
, and (1)

527 
$$AT(day) = \left(\frac{G_{S_2} - G_{M_2}}{\omega_{S_2} - \omega_{M_2}}\right)/24$$
, (2)

where  $\omega_{K_1}$  (= 15.0410686° hr<sup>-1</sup>),  $\omega_{O_1}$  (= 13.9430356° hr<sup>-1</sup>),  $\omega_{S_2}$  (= 30.0000000° hr<sup>-1</sup>), and  $\omega_{M_2}$  (= 28.9841042° hr<sup>-1</sup>) are the angular speeds of the K<sub>1</sub>, O<sub>1</sub>, S<sub>2</sub> and M<sub>2</sub> tides, respectively. Results revealed that the *ADI* are very similar, and there is <1 day *AT* difference, between ROBT and JBARS respectively (Table 1), indicating that the tidal characteristics of the representative tidal constituents for each species between the two stations are very similar, in particular the dominant diurnal species. Note that the negative *AT* values in Table 1 are an unusual feature of the Ross Sea tides, given that elsewhere spring tides commonly occur a day or two after the full and new moon. The *ADI* and *AT* is similarities between our two stations y-explains why we found the CTSM+TCC method successful in generating our the Ross Sea tidal predictions.

#### 535 5 Discussion

#### 536 5.1 Explaining fortnightly tide effects and double tide peaks in the Ross Sea tidal predictions

- 537 We have demonstrated that the CTSM+TCC approach can produce reasonably accurate tidal predictions (RMSE <5 cm, R<sup>2</sup> 538 >0.92) for a new site in the Ross Sea, Antarctica, based on 25 hr temporary station observation records from periods with 539 above average tidal ranges, plus neighbouring reference station records. Our results compare favourably with those of Han et 540 al. (2013), who reviewed the tidal height prediction accuracy of 4 models for Terra Nova Bay, Ross Sea: these models 541 generated similar quality results to our CTSM+TCC results, with R<sup>2</sup> values between 0.876 and 0.907, and RMSEs ranging 542 from 3.6 to 4.1 cm. However, as shown in Fig. 57, our results contain a changing fortnightly timescale bias in estimates. This 543 error pattern likely resulted from our application of CTSM+TCC considering only 2 major tidal species (diurnal and 544 semidiurnal) whilst ignoring several long period and small amplitude short period tides.
- 545 Table 2 summarises the characteristics of 6 long-period tides (S<sub>a</sub>, S<sub>sa</sub>, MS<sub>m</sub>, M<sub>m</sub>, M<sub>f</sub>, MS<sub>f</sub>) at the ROBT station, derived from
- 546 tidal harmonic analysis of year-long (2013) in situ observation records. Note that since the ROBT observation record was
- 547 derived from seawater pressure measurement, and thus includes proportionately large non-tidal (atmospheric) sea level
- 548 variations, caution should be exercised in comparing the harmonic analysis results of the non-astronomical constituents, which
- 549 are affected by seawater density and atmospheric forcing (i.e.  $S_a$  and  $S_{sa}$ ).

- 550 To investigate the main cause of the apparent fortnightly prediction biases in our JBARS results, in particular that in the 2019
- 551 predictions (Fig. 7b), we examined the effects of two fortnightly tidal constituents ( $M_f$ , and  $MS_f$ ) at ROBT using T TIDE.
- 552 Three 2019 tidal prediction experiments were conducted:

555

- *Srun* excluded all long-period tides (see list of exclusions in Table 2);
  - *Run1* was based on *Srun* but also incorporated the M<sub>f</sub>; and
  - *Run2* was based on *Srun* but also incorporated the M<sub>f</sub> and MS<sub>f</sub>.
- with T TIDE predictions made for each case. Comparisons between Run1 and Srun predictions revealed show that exclusion of the M<sub>f</sub> tide (2.7 cm amplitude) can produce prediction biases during periods of lunar declination change (Fig. 9a), with comparisons between Run2 and Run1 results showing revealing that the additional exclusion of the MS<sub>f</sub> tide (1.2 cm amplitude) intensifies the biases (Fig. 9b). While these results elucidate an issue with predicting Ross Sea tides based on the diurnal and semidiurnal species alone, the aforementioned differences in gauge and record types in themselves can also result in different
- 561 harmonic analysis results and, in turn, different prediction results.
- Rosier and Gudmundsson (2018) found that ice flows are modulated at various tidal frequencies, including that of the  $MS_f$ tide. However, because these tides' amplitudes have small signal-to-noise ratios (<1) with large standard errors (Table 2), caution should be exercised when elucidating fortnightly tide effects using these constituents. Nevertheless, studies indicate that incorporating major and minor tidal constituents, including long period tides, into tidal predictions may be advantageous for their use in ice flow and ice-ocean front modelling specifically (e.g. Rignot et al., 2000; Rosier and Gudmundsson, 2018). Consideration of additional, long period tides in predictions is one recommendation we have for future work on improving tidal predictions for Ross Sea coasts.
- 569 Another characteristic of our results needing explanation is the double tidal peaks evident in both the tidal observations and 570 predictions at JBARS. These peaks occur, for example, in Fig. 7b-5b between Jan. 11<sup>th</sup> and 17<sup>th</sup>, 2019. To explore why these 571 double peaks occur, we generated JBARS tidal height predictions using Eq. (A1) and the 2019 tidal constants listed in Table 572 1 for the two major diurnal and semidiurnal tides. Fig. 810 a shows separately the resulting diurnal (with their period of 13.66 573 days) and semidiurnal (with their period of 14.77 days) species' tide predictions. The combination of these out-of-phase tidal 574 species generates double peaks (or double troughs) around low and high tide (Fig. 810b) for periods when the diurnal tide 575 amplitude is are-low, due to the similar amplitude  $K_1$  and  $O_1$  tides cancelling each other out across a fortnight, allowing the 576 combined  $M_2$  and  $S_2$  amplitude to temporarily approach or exceed that of the combined  $K_1$  and  $O_1$  tides (Fig. 8c).-and the 577 amplitude ratio of the semidiurnal to diurnal tide species is >0.5 (Fig. 10c). Double peaks also occur around high tide during 578 periods of low lunar declination (Fig. 8b), when the semidiurnal to diurnal species amplitude ratio is >1, and the phase lag 579 difference between the diurnal and semidiurnal species is between 78° and 46° (Fig. 10). Since the semidiurnal tides are 580 slightly stronger, and the diurnal tides are slightly weaker, at JBARS compared to at ROBT (Table 1), these double tide peaks 581 occur more commonly at JBARS (e.g., compare Fig. 5b and Fig. 7).

#### 582 5.2 Understanding the contrasting tidal environments around Antarctica

- Figure <u>911</u> illustrates the form factors of tidal regimes in the seas surrounding Antarctica, according to FES2014 model data. There are large areas characterised by diurnal (F>3); mixed, mainly diurnal (1.5 < F < 3); and mixed, mainly semidiurnal (0.25 < F < 1.5) forms. Only in a small area half-way along the Weddell Sea coast of the Antarctic Peninsula (at 72°S) do tides exhibit a semidiurnal form (F < 0.25). The Weddell Sea is dominated by mixed, mainly semidiurnal tides, excepting the semidiurnal area mentioned and another small area exhibiting diurnal tides (F>3) at around 76.5°S, where amphidromic points (i.e. zero amplitudes) occur for both the M<sub>2</sub> and S<sub>2</sub> tides. Strong diurnal tides predominate in the Ross Sea area of West
- 589 Antarctica, around to the Amundsen Sea. In addition, a small area near Prydz Bay (Fig. 2) in East Antarctica exhibits diurnal

- 590 and mixed mainly diurnal tides. The rest of the seas surrounding Antarctica\_, including the Weddell Sea, are predominantly
- 591 characterised by mixed, mainly semidiurnal tides.
- 592 Since diurnal tides have larger nodal amplitude factor and nodal angle variations than semidiurnal tides (Pugh and Woodworth,
- 593 2014), areas like the Ross Sea will have larger variations in tidal height across the 18.61 year lunar nodal cycle compared to
- <sup>594</sup> areas like the Weddell Sea. As the nodal amplitude factor ngle-variations of the diurnal and semidiurnal tides are out of phase,
- 595 this leads to differing tidal responses around Antarctica over 18.61 years, particularly between the Ross and Weddell Seas (see
- 596 details for ROBT in Byun and Hart, 2019). Given that CTSM+TCC is based on modulated tidal amplitude and phase lag
- 597 corrections for each diurnal and semidiurnal species, this approachit is applicable in studying a continent with such a diversity
- 598 of tidal regime types. Accurate (cm scale) quantification of the contrasting tidal behaviours and environments around
- 599 Antarctica's margins are not only of use for polar station maritime operations, they are essential for estimating ice flows to the
- 600 sea. This paper has shown how the CTSM+TCC approach may be used to complement existing efforts to quantify variations
- 601 in tidal processes around Antarctica, in particular for places with sparse in situ tidal monitoring, such as the Ross Sea.

#### 602 6 Conclusions

This paper has demonstrated the usefulness of the CTSM+TCC method for tidal prediction in extreme environments, where long-term tidal station installations are difficult, using the Ross Sea in Antarctica for our case study. Here CTSM+TCC methods can be employed for accurate tidal height predictions for a temporary tidal observation station using short-term ( $\geq$ 25 hr) sea level records from this site, plus long-term (1 year) tidal records from a neighbouring reference tidal station. Essentially the temporary and reference station sites must share similarities in their main tidal constituent and tidal species characteristics for CTSM+TCC to produce acceptable results.

- 609 Using this approach, an initial tidal prediction time series is generated for the temporary station using CTSM and the reference 610 station long-term records. The temporary station predicted time series can then be adjusted via TCC of each tidal species, 611 based on harmonic comparisons between the short-term temporary station observation record and its corresponding modelled 612 predictions, leading to improved accuracy in the tidal predictions. The modulated amplitude ratio and phase lag difference
- between diurnal and semidiurnal species predicted from CTSM at the reference station can be used as an indicator for selecting
- optimal short term observation dates at a temporary tidal station.
- This paper has further demonstrated that the CTSM+TCC approach can be employed successfully in the absence of concurrent short-term (25 hr) records from the reference station, since a tidal harmonic prediction program can be used to produce a synthetic short-term record for the reference station, based on a quality long-term (1 year) record from that site.
- 618 The proper consideration of long-period tides in the CTSM+TCC approach remains a challenge, as outlined in this study, with
- 619 the solutions to this issue likely to improve tidal predictions further. However, this study demonstrates that the CTSM+TCC
- 620 method can already produce tidal predictions of sufficient accuracy to aid local polar station maritime operations, as well as
- 621 starting to help resolve gaps in the spatial coverage of tidal height predictions for scientists studying important issues, such as
- 622 the rate and role of ice loss along polar coastlines.

#### 623 Code Availability

The T\_TIDE based CTSM code is available from <u>https://au.mathworks.com/matlabcentral/fileexchange/73764-ctsm\_t\_tide</u>.

#### 625 Data Availability

626 The sea level data used in this paper are available from LINZ (2019) for selected ROBT records, with the remaining ROBT

627 records available by email application (customersupport@linz.govt.nz); and the JBARS records used are available on request

628 from KHOA (infokhoa@korea.kr). Details of the FES2014 tide model are found in Carrère et al. (2016) and via

 $629 \quad \underline{https://www.aviso.altimetry.fr/en/data/products/auxiliary-products/global-tide-fes.html.}$ 

#### 631 Appendix 1

632 This appendix describes the calculations involved in using the CTSM+TCC approach as employed in this Ross Sea, Antarctica,

case study. For a fuller description of the development of this approach and its application in semidiurnal and mixed, mainly
 semidiurnal tidal regime settings, see Byun and Hart (2015).

635 As explained in the main body of this paper, we used 25 hr slices of the 2017 short-term observations from JBARS (SH<sub>o</sub>), our

636 temporary tidal observation station (subscript *o*), and 2013 year-long observations (LH<sub>r</sub>) and 2017 short-term tidal predictions

- $(SH_r, concurrent with SH_o)$  from ROBT, our reference tidal station (subscript r), as the basis of JBARS tidal prediction
- 638 calculations. We then employed the full 17.04 day 2017 JBARS tidal observation data set, and an additional 21.54 day 2019

639 JBARS tidal observation dataset, to evaluate the success of the CTSM+TCC tidal prediction calculations for this site.

640 The CTSM+TCC, expressed as the summation of each tidal species cosine function, includes three key steps:

641 (i) calculating each tidal species' modulation at the reference tidal station;

- 642 (ii) comparing the tidal harmonic constants between the temporary observation and reference stations (e.g., the tidal
  643 amplitude ratios and phase lag differences of each representative tidal constituent for each tidal species calculated
  644 from concurrent observation records between two stations); and
- 645 (iii) adjusting the tidal species modulations calculated in the first step using the correction factors calculated in the
   646 second step to produce predictions for the temporary tidal station.

As a first step, tidal height predictions for the temporary station  $(\eta_o(\tau))$  were initially derived from reference station predictions ( $\eta_r(\tau)$ ) on the assumption that the tidal properties between the two stations remain similar through time. Using the modulated amplitude  $(A_r^{(s)})$  and the modulated phase lag  $(\varphi_r^{(s)})$  for each tidal species, this step is expressed as:

650 
$$\eta_r(\tau) = \sum_{s=1}^k A_r^{(s)}(\tau) \cos\left(\omega_R^{(s)} t - \varphi_r^{(s)}(\tau)\right)$$
 (A1)

651 with

$$652 \qquad A_r^{(s)}(\tau) = \sqrt{\sum_{i=1}^m \left[ f_i^{(s)}(\tau) \ a_i^{(s)} \right]^2 + 2\sum_{i=1}^{m-2} \sum_{j=i+1}^m \left[ f_i^{(s)}(\tau) \ a_i^{(s)} \right] \left[ f_j^{(s)}(\tau) \ a_j^{(s)} \right] \cos\left\{ \left( \omega_i^{(s)} - \omega_j^{(s)} \right) t + \left[ V_i^{(s)}(t_0) + u_i^{(s)}(\tau) - G_i^{(s)} \right] - \left[ V_j^{(s)}(t_0) + u_j^{(s)}(\tau) - G_j^{(s)} \right] \right\}$$

- 653 (A2)
- 654 and

$$655 \quad \varphi_r^{(s)}(\tau) = tan^{-1} \left( \frac{\sum_{i=1}^m f_i^{(s)}(\tau) a_i^{(s)} \sin[(\omega_i^{(s)} - \omega_R^{(s)})t + V_i^{(s)}(\tau_0) + u_i^{(s)}(\tau) - G_i^{(s)}]}{\sum_{i=1}^m f_i^{(s)}(\tau) a_i^{(s)} \cos[(\omega_i^{(s)} - \omega_R^{(s)})t + V_i^{(s)}(\tau_0) + u_i^{(s)}(\tau) - G_i^{(s)}]} \right)$$
(A3)

where superscript s denotes the type of tidal species (e.g., 1 for diurnal species and 2 for semidiurnal species); m is the number 656 of tidal constituents;  $t_0$  is the reference time; t is the time elapsed since  $t_0$ ; and  $\tau = t_0 + t$ ;  $\omega_i^{(s)}$  are the angular frequencies 657 of each tidal constituent (subscripts i and j);  $\omega_R^{(s)}$  are the angular frequencies of each tidal constituent representing a tidal 658 species (subscript R); with the dominant tidal constituent of each tidal species used as the representative for that species (e.g., 659  $K_1$  and  $M_2$  are used as representative of the diurnal and semidiurnal species, respectively). For each tidal constituent,  $a_i^{(s)}$  and 660  $G_i^{(s)}$  are the tidal harmonic amplitudes and phase lags (referenced to Greenwich);  $f_i^{(s)}(\tau)$  is the nodal amplitude factor of each 661 tidal constituent;  $u_i^{(s)}(\tau)$  is the nodal angle; and  $V_i^{(s)}(t_0)$  is the astronomical argument. T\_TIDE was used for tidal harmonic 662 analysis as well as for calculation of the nodal amplitude factors; nodal angles; and astronomical arguments; for the 663 664 representative tidal constituents.

As the second step, under the 'credo of smoothness' assumption that the admittance or 'ratio of output to input' does not change significantly between constituents of the same species (Munk and Cartwright, 1966; Pugh and Woodworth, 2014), the amplitude ratio and phase lag difference of each representative tidal constituent for each tidal species between the temporary and reference stations were calculated from the results of tidal harmonic analyses of concurrent 25 hr data slices (starting at

- 669 00.00) from the temporary observation and reference tidal stations (i.e. from SH<sub>o</sub> and SH<sub>r</sub>). The process of selecting the optimal
- 670 25 hr window for the concurrent data slices from amongst the 17.04 days of available records is explained in Sect. 3.
- 671 Once this 2017 window was selected, the third step involved adjusting the tidal predictions at the reference station calculated 672 from Eq. (A1), to represent those for the temporary station ( $\eta_o(\tau)$ ), by substituting the daily (i.e. SH<sub>o</sub> and SH<sub>r</sub>) amplitude ratios 673  $\left(\frac{a_o^{(s)}}{a_r^{(s)}}\right)$  and phase lag differences  $\left(G_o^{(s)} - G_r^{(s)}\right)$  for the tidal constituents (K<sub>1</sub> and M<sub>2</sub>) representing the diurnal and semidiurnal
- tidal species between the temporary and reference stations into Eq. (A1) as follows (Byun and Hart, 2015):

675 
$$\eta_o(\tau) = \sum_{s=1}^k A_o^{(s)}(\tau) \cos\left(\omega_R^{(s)} t - \varphi_o^{(s)}(\tau)\right)$$
 (A4)

676 with 
$$A_o^{(s)}(\tau) = A_r^{(s)}(\tau) \left(\frac{a_o^{(s)}}{a_r^{(s)}}\right)$$
, and (A5)

677 
$$\varphi_o^{(s)}(\tau) = \varphi_r^{(s)}(\tau) + G_o^{(s)} - G_r^{(s)}$$
 (A6)

678 Substituting Eqs. (A5) and (A6) into Eq. (A4),  $\eta_o(\tau)$  can be expressed as:

679 
$$\eta_o(\tau) = \sum_{s=1}^k A_r^{(s)}(\tau) \left(\frac{a_o^{(s)}}{a_r^{(s)}}\right) \cos\left[\omega_R^{(s)}t - \left(\varphi_r^{(s)}(\tau) + G_o^{(s)} - G_r^{(s)}\right)\right]$$
 (A7)

680

681 The T\_TIDE based CTSM code is available from https://au.mathworks.com/matlabcentral/fileexchange/73764-ctsm t tide.

#### 682 Author contribution

D-SB conceived of the tidal prediction idea behind this paper, and drafted initial results sections. Both authors worked on

684 initial and final versions of the full manuscript.

#### 685 Competing interests

- The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be
- 687 construed as a potential conflict of interest.

#### 688 Special issue statement (will be included by Copernicus)

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- 745 ice-shelf grounding zones, The Cryosphere, 13(12), 3171-3191, doi.org/10.5194/tc-13-3171-2019, 2019.

746Table 1. Major tidal harmonic results for diurnal and semidiurnal constituents from harmonic analyses of sea level observations:747the year-long (2013) records from Cape Roberts (ROBT), and 17.04 day records (29 Jan. to 15 Feb. 2017) and 20.54 day records (29748Dec. 2018 to 18 Jan. 2019) from Jang Bogo Antarctic Research Station (JBARS) in the Ross Sea (see source details in Sect. 2). For749the JBARS tidal harmonic analyses, the inference method was used to infer the P1 constituent from the K1, and the K2 constituent750from the S2, with their amplitude ratios and phase lag differences obtained from harmonic analysis of the long-term ROBT 2013751reference station record the inference method was applied to separate out the K1 (S2) and P1 (K2) tidal constituents, using inference752parameters estimated from the ROBT 2013 harmonic analysis.

Tidal constituents & characteristics		ROBT (2013) 369 days		JBARS (2017) 17.04 days		JBARS (2019) 20.54 days	
		Amp. (cm)	Pha. (°)	Amp. (cm)	Pha. (°)	Amp. (cm)	Pha. (°)
Diurnal	O1	21.1	202	19.6	208	16.0	208
	$\mathbf{K}_1$	20.5	217	16.3	214	14.9	216
	$\mathbf{P}_1$	6.6	215	5.2	213	4.8	214
	$Q_1$	4.4	190	-	-	-	-
Semidiurnal	<b>M</b> <sub>2</sub>	5.3	5	6.7	4	6.3	34
	$S_2$	4.9	309	6.4	329	5.7	320
	$N_2$	3.8	255	-	-	-	-
	<b>K</b> <sub>2</sub>	1.8	315	2.4	333	2.4	328
F		4.1		2.7		2.6	
		(diurnal form)		(mixed, mainly diurnal)		(mixed, mainly diurnal)	
ADI (day)		0.57		0.23		0.30	
AT (day)		-2.30		-1.44		-2.87	

753 Note: Amp. denotes amplitude; Pha. denotes phase lag, referenced to  $0^{\circ}$  Greenwich; F is the amplitude ratio of the  $(K_1 + O_1)/(M_2 + O_2)$ 

754 S<sub>2</sub>) tides; and *ADI* and *AT* denote the age of diurnal inequality and the age of the tide.

#### Table 2. Harmonic constants for 6 long-period tidal constituents, derived from harmonic analysies of <u>one</u> year-long observations (2013) measured at the Cape Roberts sea level gauge (ROBT), using T\_Tide (Pawlowicz et al., 2002). Note that this gauge is a vented piezometer so caution should be exercised in interpreting the results (particularly for S<sub>a</sub> and S<sub>sa</sub>) given the inclusion of proportionately large non-tidal (atmospheric) variations in this kind of sea level record

Constituent		Amplitude (cm)	Amplitude standard error (cm)	Phase lag (°)	Phase lag standard error (°)	SNR
Solar annual	Sa	5.8	4.8	75	50	1.5
Solar semi-annual	Ssa	0.1	3.3	352	194	0.06
Lunar monthly	MS <sub>m</sub>	0.4	3.5	57	254	0.02
Luna monuny	Mm	2.9	3.8	139	102	0.59
Lunar fortnightly	MS <sub>f</sub>	1.2	3.0	281	189	0.14
Eunar forunghury	Mf	2.7	3.9	153	101	0.47

759 Phase lags are referenced to 0° Greenwich, and SNR denotes the signal-to-noise ratios.



Figure 1. Drifting ice, including icebergs and mobile sea ice, around the Jang Bogo Antarctic Research Station (JBARS),
 photographed on 29 Jan. 2017.



Figure 2. Maps showing (a) the locations of the two tidal observation stations employed in this study within a wider Antarctic context:
 Jang Bogo Antarctic Research Station (JBARS, ▲) and Cape Roberts (ROBT, ●); and (b) the case study station locations relative

to two other (previous) temporary tidal observations stations, McMurdo Station (•), and Mario Zucchelli Station (•), in the Ross
 Sea.



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773 Figure 3. Modulated tidal (a) species amplitudes and (b) phase lags for the diurnal and semidiurnal tidal species, calculated from Cape Roberts (ROBT) tidal prediction data (29 Jan. to 14 Feb. 2017), using Appendix 1 Eqs. (A1) and (A3). 774

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Figure 4. Daily amplitudes (a, c); phase lags (b, d); amplitude ratios (e); and phase lag differences (f) of the  $K_1$  and  $M_2$  tides (representative diurnal and semidiurnal tide species) at ROBT (a, b) and JBARS (c, d), and between JBARS and ROBT (e, f), calculated from 'daily' slices of the 29 Jan. to 14 Feb. 2017 ROBT tidal predictions and JBARS sea level observations. In addition, thick blue ( $K_1$ ) and thin pink ( $M_2$ ) horizontal lines in the panels indicate the amplitudes and phase lags derived from harmonic analyses of the <u>entire</u> 369 day 2013 ROBT sea level records (a, b) and of the <u>entire</u> 17 day 2017 JBARS sea level records (c, d), along with their amplitude ratios and phase lag differences (e, f).





786 Figure 5. (a) Time series (29 Jan. to 14 Feb. 2017) of Root Mean Square Errors (RMSE, thick blue line with •) and coefficients of 787 determination (R<sup>2</sup>, thin black line with ○) between JBARS 10 min interval sea level observations and the CTSM+TCC prediction datasets, generated for this site using harmonic analysis results from the JBARS daily (25 hr) sea level data slices and concurrent 788 789 daily (25 hr) 2017 tidal prediction data slices and harmonic analysis results from ROBT station's year-long (2017) tidal predictions. 790 (b) Time series of predicted 2017 tidal heights (thin blue line) and daily tidal ranges (thick black line with  $\clubsuit$ ) for ROBT, based on 791 harmonic analysis of this station's 2013, 5 min interval sea level records, plus an indication of the moon's phase and declination.





Figure 6. Time series of Root Mean Square Errors (RMSE, thick blue line with ●) and coefficients of determination (R<sup>2</sup>, thin black
 line with ○) between JBARS 10 min interval sea level observations (29 Dec. 2018 to 18 Jan. 2019) and the CTSM+TCC prediction
 data sets generated for this site (using harmonic analysis results from daily (25 hr) summertime 2017 sea level data slices from
 JBARS along with concurrent daily (25 hr) tidal prediction slices and harmonic analysis results from ROBT station's year-long
 (2017) tidal predictions).



801 Figure 75. Time series of JBARS sea level observations (Obs.), predicted tidal heights (Pred.), and sea level residuals (Diff.) from (a) 29 Jan. to 14 Feb. 2017; and (b) 29 Dec. 2018 to 18 Jan. 2019. The JBARS predictions were generated via the CSTM+TCC method 802 803 (using a daily (25 hr) slice of local sea level observations from 8 Feb. 2017 (dashed box in (a)), along with concurrent (to time periods 804 a and b) ROBT predictions; and year-long (2017) 5 min interval ROBT tidal predictions). RMSE and R<sup>2</sup> denote the comparison 805 Root Mean Square Errors and coefficients of determination, respectively.



<sup>Figure 6. (a) Time series (29 Jan. to 14 Feb. 2017) of Root Mean Square Errors (RMSE, thick blue line with ●) and coefficients of
determination (R<sub>2</sub>, thin black line with ○) between JBARS 10 min interval sea level observations and the CTSM+TCC prediction
datasets, generated for this site using harmonic analysis results from the JBARS daily (25 hr) sea level data slices and concurrent
daily (25 hr) 2017 tidal prediction data slices and harmonic analysis results from ROBT station's year-long (2017) tidal predictions.
(b) Time series of predicted 2017 tidal heights (thin blue line) and daily tidal ranges (thick black line with ♦) for ROBT, based on
harmonic analysis of this station's 2013, 5 min interval sea level record, plus an indication of the moon's phase and declination.</sup> 



Figure 87. Time series of the Moon's declination, calculated at daily intervals for two observation periods: (a) 1 Jan. to 15 Feb. 2017; and (b) 16 Dec. 2018 to 30 Jan. 2019. Dashed boxes indicate the sea level observation windows examined in this study.



Figure 9. Time series of ROBT tidal predictions (a) made without long-period constituents (SRun, i.e. excluding the constituents
 listed in Table 2) versus with the Mr tide (Exp1); and (b) time series of ROBT tidal predictions made (SRun) without the long-period
 constituents versus (Exp2) with the MSr and Mr tides. All predictions were generated based on tidal harmonic analysis results from
 the year-long (2013) ROBT sea level records.



Figure <u>810</u>. Time series (29 Dec. 2018 to 18 Jan. 2019) of (a) predictions of the diurnal (K<sub>1</sub>+O<sub>1</sub>) tides (blue line) and the semidiurnal (M<sub>2</sub>+S<sub>2</sub>) tides (magenta line) for JBARS; (b) their combined JBARS predictions (red line) and observations (black dashed line); (c) the ROBT diurnal (blue line) and semidiurnal (magenta line) species amplitudes and their ratio (green line); and (d) the ROBT diurnal (blue line) and semidiurnal line) species phase lags and their difference (diurnal – semidiurnal) (green line).

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Figure <u>911</u>. Distribution of tidal form factor (F) values around Antarctica. Note the magenta area (72°S) on the Antarctic Peninsula's
Weddell Sea coast denotes the only area with a properly semidiurnal tide regime (F<0.25) in the Antarctic region.</li>