

A NEW LAGRANGIAN BASED SHORT TERM PREDICTION METHODOLOGY FOR HF RADAR CURRENTS

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ABSTRACT

The use of High Frequency Radar (HFR) data is increasing worldwide for different applications in the field of operational oceanography and data assimilation, as it provides real-time coastal surface currents at high temporal and spatial resolution. In this work, a Lagrangian based empirical real-time, Short-Term Prediction (L-STP) system is presented in order to provide short term forecasts of up to 48 hours of ocean currents from HFR data. The method is based on the finding of historical analogues of Lagrangian trajectories obtained from HFR surface currents. Then, assuming that the present state will follow the same temporal evolution as the historical analogue did, we can obtain a short-term prediction of the surface currents. The method is applied to two HFR systems covering two areas with different dynamical characteristics: the southeast Bay of Biscay and the central Red Sea. The L-STP improves on previous prediction systems implemented for the SE Bay of Biscay and provides good results for the Red Sea study area. A comparison of the L-STP methodology with predictions based on persistence and reference fields has been performed in order to quantify the error introduced by this approach. Furthermore, a temporal sensitivity analysis has been addressed to determine the limit of applicability of the methodology regarding the temporal horizon of Lagrangian prediction. A real-time skill-score has been developed using the results of this analysis which allows to identify periods when the short-term prediction performance is more likely to be low and persistence can be used as a better predictor for the future currents.

1. INTRODUCTION

The coastal zone is under increasing human pressure. On the one hand, during recent decades coastal seas have been experiencing intensified activity for recreation, transport, fisheries and marine-related energy production. Simultaneously, continued growth of the global coastal population largely contributes to increase the problem of the wastewater discharge which, in many cases, results in serious damage to coastal marine ecosystems. A better understanding of the dynamical processes responsible for the surface oceanic transport, is a prerequisite for the efficient management of the coastal ocean. These processes are responsible of the transport and fate of pollutants, nutrients, jellyfish, harmful algal blooms, plastics, etc, and improving the capacity of monitoring and forecasting the coastal area, is necessary to identify regions of accumulation or dispersion of these harmful materials. This requirement is driving the set-up of a growing number of multi-platform operational observatories designed for the continuous monitoring of the coastal ocean (e.g., US IOOS, EU EOOS, SOCIB, Australian IMOS, etc.). In the need of providing a long-term framework for the development and improvement of the European Marine coastal observations, the JERICO Research infrastructure has been putting efforts (through JERICO, JERICO-NEXT and JERICO-S3 projects) to develop methods and tools for the production of high-quality marine data, and the sharing of expertise and infrastructures between the exiting observatories in Europe. Moreover, due to the need of forecasting applications for response to emergency situations such oil spills, or search and rescue operations, many of the existing operational observatories are linked with operational ocean forecasting models with or without data assimilation (e.g. MARACOOS, NOAA Global Real-Time Ocean Forecast System, COPENICUS Marine Environment Monitoring System). Typically, constituted with different in-situ point-wise observational platforms (such as moored

buoys, tidal gauges, wave buoys, etc.) a significant number of these observatories now employ land-based High Frequency Radars (HFR), that provide real-time coastal currents with unprecedented coverage and resolution (e.g. Paduan and Rosenfeld, 1996; Kohut and Glenn, 2003; Abascal et al., 2009; Solabarrieta et al., 2014, Rubio et al. 2017; Paduan and Washburn, 2013). Each HFR coastal site measures radial surface currents moving away or approaching its antenna, based in the shift of the first peak (Bragg peak) of the Doppler spectra (Crombie 1955, Barrick et al 1977). Combining the overlapping radial vectors from at least 2 antennas provides surface true vector currents (Barrick et al., 1977, Lipa and Barrick, 1983). Several studies have compared in-situ current measurements with HFR observations (e.g., Schott et al. 1985; Hammond et al. 1987; Paduan and Rosenfeld 1996, Emery et al. 2004; Paduan et al., 2006; Ohlmann et al. 2007; Liu et al., 2014; Solabarrieta et al, 2014, Bellomo et al., 2015; Lana et al., 2016; Hernandez-Carrasco et al., 2018b) and have repeatedly demonstrated the validity of this technology. Presently, more than 250 HFR antennas are installed being active worldwide (Roarty et al., 2019; <http://global-hfradar.org/>).

The range and the spatial resolution of the HFR current systems depend on their working frequency and the conductivity of the water over which the system is measuring. Ranges vary from 15 to 220 km range and spatial resolution from 250 m to 12 km. Typically, a 12 MHz radar has a range ~70 km with a spatial resolution of 2-5 km. HFR systems usually average current measurements for one hour, although some average currents for shorter periods, such as 30 minutes. Due to their high spatio-temporal resolution, HFR data are commonly used in real time for search and rescue (Ullman et al., 2006) or oil spill prediction/mitigation emergency response (Abascal et al., 2017).

94 The performance of HFR for measuring near-real time surface currents has resulted
95 in the development of assimilation strategies that incorporate the HFR measured
96 surface currents into ocean coastal models (Breivik and Sætra, 2001, Oke et al 2002,
97 Paduan and Shulman 2004, Stanev et al., 2011, Barth et al., 2011) some of which
98 have been tested for short periods of time (Chao et al., 2009). However, assimilation
99 of HFR data into models is still a computationally expensive and complex issue, not
100 to mention operational applications of such a procedure. Because of these constraints,
101 the availability of real-time high-resolution HFR current fields has led to alternative
102 solutions in order to obtain short term prediction (STP) of surface coastal currents,
103 through the direct use of HFR historical and nowcast observations using different
104 approaches (e.g. Zelenke 2005, Frolov et al. 2011, Barrick et al., 2012, Orfila et al.
105 2015, Solabarrieta et al. 2016, Vilibić et al, 2016, Ren et al., 2019, see Table 1).

106

107 The above-mentioned studies develop and implement different STP approaches
108 (harmonic analysis of the last hours, genetic algorithms, numerical models, ...)
109 which often require additional data, or long training periods of data without gaps
110 which can jeopardize the general utility of these methods in real time (Hardware
111 failures due to power issues, communications or environmental conditions often
112 result in spatio-temporal gaps within HFR datasets. Spatial gaps can be filled on a
113 real-time basis but the filling of long temporal gaps is not straightforward). Several
114 gap-filling methodologies have been developed for HFR data sets: Open Modal
115 Analysis, (OMA) (Kaplan and Lekien, 2007), Data Interpolating EOFs (DINEOF)
116 (Hernandez-Carrasco et al., 2018), and Self-Organizing Maps (SOM) (Hernandez-
117 Carrasco et al., 2018). The OMA method has been used for spatial gap filling in this
118 paper mainly because it's well functioning has been demonstrated (Kaplan and
119 Lekien, 2007, Hernández-Carrasco et al., 2018) and it is easily applicable in real time,

with available codes that will also be applied for trajectories' generation, later in this paper (HFR_progs MATLAB package: <https://github.com/rowg/hfrprogs>).

A widely used method in time series prediction, especially in early weather forecasting, is the method of analogues. It is based on the assumption that if the behavior of a system at a given time is similar to some other situation in the historical record, then the evolution in the future of state will be similar to the evolution observed in the same historical record. Simply stated, two analogue fields are two distinct fields that are close enough considering some metric, to be considered as equivalent. The finding of the best (nearest) analogue of a specific time does not require a historically continuous dataset, as long as it contains subsets of observations that extend longer than the testing period. These analogue events occur naturally in the environment and this methodology has been applied and tested in atmospheric forecasts (Lorenz, 1969, Jianping et al,1993, Prince and Goswami 2007, Shao and Li 2013).

Given the motivation described above, and developed partially in the framework of JERICO-NEXT project, we present a Lagrangian-based Short-Term Prediction (L-STP from now on) methodology using existing HFR datasets, to be applied to current real-time observations. The uniqueness of this approach is two-fold: first the historical Eulerian velocity fields are used to construct a catalogue of Lagrangian trajectories and second, using the trajectories obtained from present observations, analogues in the past dataset are searched in order to obtain the best predictive match. The method is based on Lagrangian computations since they have proven to be robust in identifying dynamical flow structures and they are direct measurements of transport of substances at sea.

146 Then, it is worth highlighting that this is the first time that the analogues technique
147 is applied to the HFR-derived ocean surface currents to obtain short-term forecast.
148 The L-STP is intended to be implemented operationally requiring low computational
149 cost (seconds to few minutes for each forecast, depending on the size of the historical
150 dataset) and it is easy to implement using existing HFR data processing tools.

2. DATA AND METHODS

2.1 Data

HFR data from two distinct oceanographic regions have been used for the evaluation, validation and testing of the developed methodology in this paper (Figure 1): Left: The Bay of Biscay (hereinafter BoB HFR) and Right: The central Red Sea region (hereinafter Red Sea HFR). These two study regions are used to evaluate the skill of the method with different dynamical conditions, and with a sufficient set of observations to provide a database suited to the efficient research of appropriate analogues. The BoB HFR system, located in the southeastern corner of the Bay of Biscay, in the Basque Country, is composed of two CODAR Seasonde sites, working since 2009 which transmit at 4.5MHz frequency covering up to 200km range and providing hourly surface velocity field at 5 km of spatial resolution. The dataset used in this study spans the period from 2012 to 2015. The Red Sea HFR system is located on the central western coast of Saudi Arabia and is also composed of two CODAR Seasonde sites, operational since June 2017, transmitting at 16.12MHz frequency, covering up to 120 km range and providing the hourly surface velocity field at 3 km spatial resolution. The dataset from June 2017 to October 2018 has been used in this study.

The BoB HFR has been chosen as the pilot system for testing the developed methodology because of our previous knowledge regarding the circulation and dynamical processes in the study area (Rubio et al 2013, Solabarrieta et al 2014, Solabarrieta et al., 2015, Rubio et al., 2018, Hernandez-Carrasco et al. 2018). The resulting methodology is then applied to the operational Red Sea HFR dataset, as a study case. Coastal dynamics in the BoB show a clear seasonality where cyclonic and anticyclonic eddies dominate in winter and summer, respectively in responding

to local winds and the mean coastal current (Iberian Poleward Current) (Esnaola et al., 2013, Solabarrieta et al., 2014). The circulation in the central Red Sea also demonstrates a clear seasonality (Sofianos and Johns, 2003; Yao et al., 2014a, 2014b; Zarokanellos et al., 2016, 2017) linked to the seasonal winds of the area (Abualnaja et al., 2014; Langodan et al., 2017). The region is dominated by eddy activity, with both cyclonic and anticyclonic eddies dominating the region (Zhan et al., 2014; Zarokanellos et al. 2016). Due to the only recently available dataset (since mid-June 2017 to present) the detailed small-scale surface circulation processes of this area is under characterization at the moment.

The primary difference between the two HFR systems is the operating frequency (5MHz for the BoB system and 16 MHz for the Red Sea system) resulting in a larger spatial coverage for the BoB HFR than for the Red Sea HFR (200km range vs. 120km, respectively), but with higher spatial resolution for the latter (3km and 5 km, respectively). This difference in the spatial resolution should result in better capturing the small-scale dynamical features in the Red Sea that could influence the selection of an analogue.

The data from both systems have been processed similarly. The spectra of the received backscattered signal are converted into radial velocities using the Multiple Signal Classification (MUSIC) algorithm (Schmidt 1986). HFR Progs MATLAB package is then used to combine radial currents and generate gap-filled total 2D currents, using the Open Modal Analysis (OMA) methodology of Kaplan and Lekien (2007).

2.2 Lagrangian analogues

The proposed methodology is based on the analogue finding approach, using a historical catalogue of maps of Lagrangian trajectories and finding the most similar one (detailed later in this section) to that of the last 48 hours (target field). Then, the next 48-hour time evolution of the closest (chosen) analogue provides the forecast for the target period. In other words, if we find a state in the historical database that is close enough to the target field (given a metric), the forecast for the current observations will evolve in the same way as did for the chosen analogue. Analogue finding has been applied in several geophysical variables in different regions (Zorita and von Storch, 1999; Fernandez-Ferrero et al., 2009, 2010; Ibarra-Berastegi et al., 2011; Martin et al., 2014; Seubert et al., 2014; Ibarra-Berastegi et al., 2015).

The analogue finding was first applied to eulerian surface velocity fields of the BoB HFR System (not shown), but the results did not improve the previously published STP results for the study area. The methodology was tested subsequently using a four-year dataset (2012-2015) of trajectory maps computed for the SE BoB, where the trajectory maps from the three first years was used as the search catalogue for analogues (2012-2014) (hereinafter “Lagrangian catalogue”), and the remaining year (2015) was used as a test case (hereinafter “test period”). Then the method was applied to the Red Sea dataset, for the period of July 2017-October 2018. As the period was short (1 year and 4 months), we have used the whole period to build the Lagrangian catalogue and act as a test period at the same time. In this case, for the analogues search the 5-days period around the date of the target field was removed from the catalogue at each iteration, to avoid temporal overlapping with the target field.

To build the Lagrangian catalogue we first generated hourly fields of 25 virtual particle trajectories on a regular grid, (blue dots of Figure 2), which were advected by the OMA HFR surface currents (without considering diffusion) during 48 hours. To this end we used the Lagrangian module included in the HFR_Progs MATLAB package, following the same procedure for the test period. Then, for each hour of the test period, the method searched the most similar Lagrangian patterns in the Lagrangian catalogue dataset. To increase the efficiency of the processes, the search was done in two steps. First, we looked for potential analogues with a similar main drift. To do that we computed and compared the position of the centroid of the 25 trajectories of each analogue to that of the target field, and discarded the analogues whose centroid was at a distance $> \delta_{cg}$. The value of the δ_{cg} needs to be small enough to minimize computational time but sufficiently large to as to not lose potential analogues. We explored different values of this distance threshold and we found that $\delta_{gc}=10\text{km}$ produces a good compromise between computational cost and number of potential analogues in both study areas. Then, in a second step, we computed the Lagrangian errors (\mathcal{E}) between the trajectories of the target field and the potential analogues, defined as:

$$\mathcal{E} = \sum ((\delta_{6h})^2 + (\delta_{12h})^2 + (\delta_{24h})^2 + (\delta_{36h})^2 + (\delta_{48h})^2) \quad \text{Eq. (1)}$$

Where δ_t is the mean separation distance [km] at time t between the trajectories belonging to the target field and each of the potential analogues (being $t=6, 12, 24, 36$ and 48 hours inside the trajectories lifetimes).

Finally, the potential analogue with the lowest \mathcal{E} was selected as the best analogue ($\mathcal{E}_{ANL} = \min(\mathcal{E})$) and the velocity fields during the next 48 h from that analogue provides STP currents for the target period (hereinafter “L-STP fields”). *Figure 3*

shows an example of the values of \mathcal{E} , through the potential analogues for a specific case.

Figure 2 provides an example of the selected analogue (Figure 2b) and corresponding L-STP fields (Figure 2d) for a given target field (Figure 2a) and the ‘truth’ trajectories for the following 48 hours from the date of the target field (Figure 2c). The associated temporal series of errors for the target field and the potential analogues are shown in *Figure 3*, where the value of ε_{ANL} is marked using a red dot. (corresponding to the error between the trajectories of the L-STP field in Figure 2d and the truth trajectories for the forecast period in –Figure 2c).

To assess the performance of the methodology, we computed forecasted trajectories based on persistence of currents (hereinafter ‘persistence fields’). To obtain simulated trajectories using persistence currents, the particles were advected during 48 hours using a constant velocity field (target field) during the 48 hours of simulation:

$$v(x,y,t_f+t_i)=v(x,y,t_s),$$

where t_f = study time and t_i =[t_f : t_f+48h].

The mean drift of the truth forecasted trajectories is also computed for each simulation period (the means drift is considered as the average of the distances moved by each particle during 48 hours).

The Lagrangian errors between the truth trajectories and the L-STP and between the truth trajectories and the persistence field were also computed as follows:

$$\epsilon_{STP} = \Sigma ((\delta_{_6h})^2 + (\delta_{_12h})^2 + (\delta_{_24h})^2 + (\delta_{_36h})^2 + (\delta_{_48h})^2) \quad \text{Eq. (2)}$$

where δ_t is the mean separation distance between truth field's and the L-STP field trajectories for $t = t : t+48$ (following 48 hours from the study time)

$$\epsilon_{PRS} = \Sigma ((\delta_{_6h})^2 + (\delta_{_12h})^2 + (\delta_{_24h})^2 + (\delta_{_36h})^2 + (\delta_{_48h})^2) \quad \text{Eq. (3)}$$

where δ_t is the mean separation distance between truth field's and Persistent field trajectories for $t = t : t+48$ (following 48 hours from the study time)

All the process for the selection and validation of the analogue with the different variables has been summarized in *Figure 4*.

The time series and spatial distribution of the ϵ_{STP} and ϵ_{PRS} errors have been analyzed for both study areas. Finally, ϵ_{STP} and ϵ_{PRS} time series have also been calculated and compared to the time series of the ϵ_{ANL} , in order to evaluate if the ϵ_{ANL} can be used as an indicator of the expected skill of the L-STP with respect to the persistence.

3. RESULTS

The performance assessment results for the BoB HFR system are described in section 3.1 and the temporal and spatial forecast for both study areas are shown in section 3.2.

3.1 Assessment of the L-STP skills

Figure 5 shows the ϵ_{ANL} through year 2015 for the BOB study area, together with the ϵ_{STP} and ϵ_{PRS} . The mean value of the ϵ_{PRS} is 73% higher than the ϵ_{STP} . Black dots over the timeline in Figure 5 show the times when ϵ_{STP} is higher than the ϵ_{PRS} , which occurs 12% of the time. Focusing on the times when the ϵ_{PRS} is lower than the ϵ_{STP} (black dots of the timeline in Figure 5), it can be seen that they mostly occur during winter months. Previous works in this area have shown that there are high persistent eastward currents that can last for several weeks during winter months (Solabarrieta et al., 2014), which can explain the better performance of the persistence fields in this period.

The correlation between ϵ_{ANL} and ϵ_{STP} is 0.46 while correlation between ϵ_{ANL} and ϵ_{PRS} is 0.05, for the whole test year (2015) (Figure 5).

The hourly values of ϵ_{STP} and ϵ_{PRS} have been plotted against their corresponding hourly ϵ_{ANL} values for the test year, ordered from minimum to maximum along the x-axis in Figure 6. We observe that, when ϵ_{ANL} is low (less than 853 km² for this data set), ϵ_{STP} is smaller than ϵ_{PRS} . However, as ϵ_{ANL} increases, ϵ_{STP} and ϵ_{PRS} converge until an inflection point beyond which ϵ_{STP} is slightly greater than ϵ_{PRS} . For the SE BoB experiment, the inflection point occurs at $\epsilon_{ANL} = 853$ km² and 88% of cumulative

ϵ_{ANL} . Results from the Red Sea HFR system indicates a similar pattern (not shown), when the inflection point occurs at $\epsilon_{ANL} = 821 \text{ km}^2$ and at 86.4% of cumulative ϵ_{ANL} .

Further analysis to elucidate the time periods that largely contribute to the errors, compared to persistence are presented hereinafter. ϵ_{ANL} has been plotted together with the mean separation distances of the trajectories using STP and persistent currents (hereinafter δ_{STP} δ_{PRS}), after 6, 12, 24, 36 and 48 hours for each target field (Figure 7). δ_{STP} is always higher than the δ_{PRS} for the 6 hours' simulation. But the values of δ_{STP} show better results for simulations at 12, 24, 36 and 48 hours. The values of the correlation coefficient (R^2) between the ϵ_{ANL} and δ_{STP} and between ϵ_{ANL} and δ_{PRS} after 6, 12, 24, 36 and 48 hours are summarized in Table 2. Values of R^2 for ϵ_{ANL} and δ_{PRS} are small (almost no correlation), varying between 0.01 and 0.11, while correlations between ϵ_{ANL} and δ_{STP} are higher, varying between 0.19 and 0.56, and showing higher correlation (>0.39) after 12 hours of simulations. The behavior of the Red Sea HFR system figures (not shown) is similar to the BoB HFR system.

3.2 L-STP performances in the selected study areas

Mean separation distances between truth and forecasted trajectories after different periods of integration times have been computed for both systems, for the best analogues, i.e., before the inflection point of $\epsilon_{STP} > \epsilon_{PRS}$ (Figure 6), in order to evaluate the temporal forecast capabilities of the methodology. Only analogues with $\epsilon_{ANL} < 853 \text{ km}^2$ (BoB system) have been used to generate this analysis, as those are the periods when the methodology produces good results. Separation distances computed for the whole test year 2015, are shown in Figure 8, for the BoB HFR observations.

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352 The separation distances between the measured trajectories and predicted persistent
353 and STP trajectories, have similar values during the first 6 hours (4km) of the
354 forecast period, with slightly better results for persistent trajectories. But after 6
355 hours, the separation distance for the forecast based on persistent currents increases
356 faster than using L-STP. At 24 hours, the separation distance is 11 km for persistence
357 forecasts and 8km for L-STP forecasts. The values are 12 and 18km, respectively,
358 after 48 hours of simulation. The mean drift values of the truth trajectories show that
359 the mean drift is similar to the L-STP separation distances, during the 48 hours.

360

361 Temporal mean separation distances between truth and forecasted trajectories for the
362 Central Red Sea HFR System, computed for the whole test time are shown in Figure
363 9. Only the best analogues with ε_{ANL} less than inflection point, i.e., $\varepsilon_{ANL} < 821\text{km}^2$,
364 have been used to generate this analysis. The separation distances for the STP
365 forecasts are higher than those forecasts with persistent currents during the first 15
366 hours. After 15 hours, quality of forecasts reversed where STP produced better
367 results than persistence.

368

369 Spatial distribution of the difference between δ_{PRS} and δ_{STP} at 6, 12, 24 and 48
370 hours, for the BoB and the Red Sea study areas, are shown in Figure 10 and Figure
371 11.

372

373 For the BoB HFR system, the differences are not appreciated during the first 6 hours.
374 But after 12 hours of simulation, the advantage of the L-STP is clear in most of the
375 study area, especially outside the continental shelf slope where persistent currents
376 dominate the circulation. The separation values between δ_{PRS} and δ_{STP} increase up
377 to 10km after 48hours of simulation.

378

379 For the Red Sea, the significant differences between STP and Persistence start after

380 24 hours of simulation, and continue until 48 hours.

4. DISCUSSION

In this work, a new methodology to forecast HFR currents has been described and the skill of the proposed STP methodology is analyzed. Different analyses are performed in order to check the spatial and temporal capabilities of the proposed methodology.

The methodology is based on the search of analogues in a trajectory (Lagrangian) space using a previously generated trajectory field catalogue. The values of the δ_{STP} , compared to previous works in the BoB area showed that the L-STP produces accurate predictions, which demonstrates the ability of the Lagrangian approach to capture key dynamical features needed to accurately predict the proper dynamical conditions.

Significant correlation values between ϵ_{ANL} and δ_{STP} , suggest that the ϵ_{ANL} can be considered as a real-time skill-score metric for the L-STP. Both BoB and the central Red Sea show a similar behavior; although the ϵ_{ANL} values are different, the accumulative % of the transition point is similar in both cases.

Figure 7 shows that after 12 hours of simulation, the L-STP provides a better prediction than the persistence field for more than 80% of the cases (reaching more than 90% of the cases for 36 and 48 hours of simulation). The minimum ϵ_{ANL} value for the δ_{STP} and δ_{PRS} cross point is 714km^2 . Figure 6, for the total ϵ_{ANL} shows the same behavior being 853km^2 the transition analogue error value between STP and Persistence.

For the BoB HFR System, temporal δ_{STP} shows values of 3.5km, 5.5km and 8km, after 6, 12, and 24 hours respectively. The δ_{STP} values are similar to the δ_{PRS} values during the first 6 hours of simulation but δ_{STP} are lower after that, with 3km and 5.5km of difference between them, after 24 and 48 hours of simulation, respectively (*Figure 8*). As stated in previous work, that the circulation over the BoB area is dominated by a stable, persistent current field during winter (Solabarrieta et al., 2014) which is reflected by these results where persistence has good or even slightly better forecasting skill during the first 6 forecast hours than the proposed methodology.

The δ_{STP} values for the BoB HFR system are similar to the ones obtained by Solabarrieta et al., 2016, for the whole year but δ_{STP} are better for summer months, for the same study area. They used the linear autoregressive model, described in Frolov et al., 2012, to forecast HFR current fields and the errors using that approach were 2.9 and 7.9km after 6 and 24 hours. Although the results obtained in this work improve only during certain periods the forecast presented in Solabarrieta et al. 2016, the presented methodology has three advantages over the previous method: it is easily run in real time; it does not require a continuous training period; and it is able to discriminate the times when the usage of the persistence is applicable. On the negative side, it requires the generation of a catalogue of past trajectories as the search space for analogues, but once it is ready, it is easily increasable in real time, without extra pre-analysis; just adding new trajectory fields to the previous catalogue.

The values of the δ_{STP} for the Red Sea HFR system follow a similar pattern to the BoB results, with higher separation distances. This may be related to the limited time

span of the available dataset, as a better closest analogue may be found in a longer dataset.

The spatial comparison of the δ_{STP} and δ_{PRS} for the BoB HFR system (Figure 10), shows that the L-STP has better skills for the entire study area after 12 hours of simulations. The skills of the L-STP with respect to the persistence increases with time, showing up to 10km of improvement relative to persistence at 48 hours in some parts of the study area. For the spatial distribution, after 12 hours, the smallest differences between δ_{STP} and δ_{PRS} occurred over the slope. This is explained by existence of persistent seasonal Iberian Poleward Current that flows along the continental slope toward the east along the Spanish coast and northward along the French coast (Solabarrieta et al. 2014). In other words: although the L-STP can be performant in periods of persistent currents, the persistence field can show a better forecast for a short temporal scale (48h). L-STP will improve those forecasts, as soon as spatiotemporal variability increases.

The results for the Red Sea HFR system are similar but the benefit of the L-STP methodology appears only after 12 hours of simulation. Spatially, the improvement is again lower where persistent currents occur, as it is the case of the Eastern Boundary Current that flows northward following the eastern Red Sea Coastline in the study area (Bower and Farrah, 2015; Sofianos and Johns, 2003; Zarokanellos et al., 2017). The dominance of the persistent currents is evident in the lower values of the difference between the STP forecasts and the Persistence forecasts as shown in Figure 11 and in comparison with Figure 10.

We have compared the capabilities of the L-STP forecast against the forecast based on the persistency of currents. The L-STP method requires long training periods but

461 performs better during non-persistent periods. Previous efforts to forecast surface
462 currents from HFR data have shown similar results compared with the methodology
463 presented in this paper. However, the advantage of the L-STP method is that it can
464 be used in near real time, with short and non-continuous datasets of around 2-3 years,
465 provided that a Lagrangian catalog representative for the study area can be built.

466

467 The HFR Progs MATLAB package ([https://](https://cencalarchive.org/~cocmpmb/COCMPwiki)
468 cencalarchive.org/~cocmpmb/COCMPwiki) has been used to generate total currents
469 from radial files to fill the spatial gaps of the surface current field using the OMA
470 method, and to generate Lagrangian trajectories. This methodology could be easily
471 included in this package so the final users could get forecast currents, in the same
472 time that they generate total currents.

5. CONCLUSION

A methodology to short term forecast of the surface currents in real-time has been proposed. This methodology provides accurate forecast of sea surface currents and its capability has been tested in terms of spatial and temporal distributions. The good functioning and confidence of this methodology has been demonstrated in the previous sections and also, its capability to be applied in real time. The methodology has been successfully applied to two distinct coastal regions to evaluate its capabilities in different hydrodynamic regimes, although further analysis using data from more areas is required to generalize the methodology.

Relationships between ϵ_{ANL} and $\epsilon_{STP}/\epsilon_{PRS}$ suggest that the ϵ_{ANL} can be considered as a reliable indicator of the method's performance. Taking in consideration all the analyses done in this work, we propose to use STP currents for trajectory or velocity field predictions from 12 hours forward, if the ϵ_{ANL} value is lower than 80% of the cumulative ϵ_{ANL} . If ϵ_{ANL} is higher, or the forecast is just for the next 6 hours, the use of the persistence field is suggested. We also suggest that the ϵ_{ANL} value and forecast transition time need to be carefully evaluated for each study region. This, of course, infers that a minimum data set is required before the L-STP method can be applied.

Further analysis of analogue finding approaches is required to improve the observed results, especially during periods when currents are persistent. The use of longer dataset as a training period may improve this aspect. Then, the next step would be to test the methodology for additional periods and other regions, to analyze the possibility to find analogues for different sub-regions and to evaluate its functionality in an operational mode.

499 The methods to find the minimum training period for each system should be
500 analyzed deeper in future works. The minimum training period will be directly
501 related to the variability of the local dynamics and those should be considered during
502 the analysis.

503 DATA AVAILABILITY

504

505 The Red Sea HF Radar data can be requested through:

- 506 • <https://lthdatalib.kaust.edu.sa>

507

508 Historical and NRT Bay of Biscay HF Radar data can be requested through:

- 509 • Euskoos portal: [https://www.euskoos.eus/en/data/basque-ocean-](https://www.euskoos.eus/en/data/basque-ocean-meteorological-network/high-frequency-coastal-radars/)
510 [meteorological-network/high-frequency-coastal-radars/](https://www.euskoos.eus/en/data/basque-ocean-meteorological-network/high-frequency-coastal-radars/)
- 511 • Emodnet Physics -
512 <http://www.emodnetphysics.eu/Map/platinfo/piradar.aspx?platformid=10>
513 [273](http://www.emodnetphysics.eu/Map/platinfo/piradar.aspx?platformid=10)
- 514 • CMEMS Instac - [http://marine.copernicus.eu/services-portfolio/access-to-](http://marine.copernicus.eu/services-portfolio/access-to-products/?option=com_csw&view=details&product_id=INSITU_GLO_UV_NRT_OBSERVATIONS_013_048)
515 [products/?option=com_csw&view=details&product_id=INSITU_GLO_UV_N](http://marine.copernicus.eu/services-portfolio/access-to-products/?option=com_csw&view=details&product_id=INSITU_GLO_UV_NRT_OBSERVATIONS_013_048)
516 [RT_OBSERVATIONS_013_048](http://marine.copernicus.eu/services-portfolio/access-to-products/?option=com_csw&view=details&product_id=INSITU_GLO_UV_NRT_OBSERVATIONS_013_048)

AUTHORS CONTRIBUTION

- Lohitzune Solabarrieta: She has worked on the set up of the methodology, data analysis, manuscript writing and final submission.
- Ismael Hernandez-Carrasco: He has worked on the set up of the methodology and the manuscript writing.
- Anna Rubio: She has worked on the set up of the methodology, data analysis, and manuscript writing.
- Alejandro Orfila: He has worked on the configuration of the methodology, data analysis and the manuscript writing.
- Michael Campbell: He has worked on the configuration of the methodology, especially in the pre-configuration that led us to rule out other data prediction methodologies. He has also contributed on the manuscript writing.
- Ganix Esnaola: He has worked on the configuration of the methodology, especially in the pre-configuration that moved us to the usage of analogues on this paper. He has also contributed on the manuscript writing.
- Julien Mader: He has contributed on the writing of the manuscript.
- Burton H. Jones: He has contributed on the writing of the manuscript

535 **COMPETING INTERESTS**

536 The authors declare that we have no conflict of interest

ACKNOWLEDGEMENTS

This work was funded by a Saudi Aramco-KAUST Center for Marine Environmental Observation (SAKMEO) Postdoc fellowship to Lohitzune Solabarrieta, and from the Integrated Ocean Processes (IOP) Group in KAUST. We acknowledge the support of the LIFE-LEMA project (LIFE15 ENV/ES/000252), the European Union's Horizon 2020 research and innovation program under grant agreement No. 654410 & 871153 (JERICO-NEXT and JERICO-S3 Projects), the Directorate of Emergency Attention and Meteorology of the Basque Government, the MINECO/FEDER Projects MUSA and MOCCA (CTM2015-66225-C2-2-P; 256RTI2018-093941-B-C31). and the Department of Environment, Regional Planning, Agriculture and Fisheries of the Basque Government (Marco Program). This work was partially performed while A. Orfila was a visiting scientist at the Earth, Environmental and Planetary Sciences Department at Brown University through a Ministerio de Ciencia, Innovación y Universidades fellowship (PRX18/00218). Ismael Hernandez-Carrasco acknowledges the Vicenç Mut contract funded by the Balearic Island Govern and the European Social Fund (ESF) Operational Programme.

The HF radar-processing toolbox HFR_Progs use to produce OMA was provided by D. Kaplan and M. Cook, Naval Postgraduate School, Monterey, CA, USA.

REFERENCES

- Abascal, A. J., Castanedo, S., Medina, R., Losada, I. J., Álvarez-Fanjul, E.: Application of HF radar currents to oil spill modelling. *Mar. Pollut. Bull.* 58 (2), 238–248, 2009
- Abascal A. J., Sanchez, J., Chiri, H., Ferrer, M. I., Cárdenas, M., Gallego, A., Castanedo, S., Medina, R., Alonso-Martirena, A., Berx, B., Turrell, W. R., Hughes, S. L.: Operational oil spill trajectory modelling using HF radar currents: A northwest European continental shelf case study. *Marine Pollution Bulletin*, Volume 119, Issue 1, Pages 336-350, ISSN 0025-326X, <https://doi.org/10.1016/j.marpolbul.2017.04.010>, 2017.
- Abualnaja, Y., Papadopoulos, V. P., Josey, S. A., Hoteit, I., Kontoyiannis, H., and Raitsos, D. E.: Impacts of climate modes on air–sea heat exchange in the Red Sea, *J. Clim.*, 28, 2665–2681, doi:10.1175/JCLI-D-14-00379.1, 2015.
- Barrick, D. E.: Extraction of wave parameters from measured HF radar sea-echo Doppler spectra. *Radio Sci.*, 12, 415–424, doi:10.1029/RS012i003p00415, 1977.
- Barrick D.E., Fernandez, V., Ferrer, M.I., Whelan, C., and Breivik, Ø.: “A short-term predictive system for surface currents from a rapidly deployed coastal HF-Radar network,” *Ocean Dyn.*, vol. 62, no. 5, pp. 725–740, 2012.
- Barth, A., Alvera-Azcárate, A., Beckers, JM., Staneva J., Stanev E.V., and Schulz-Stellenfleth J.: Correcting surface winds by assimilating high-frequency radar surface currents in the German Bight. *Ocean Dynamics*, 2011, vol 61: 599. <https://doi.org/10.1007/s10236-010-0369-0>, 2011.
- Bellomo, L., Griffa, A., Cosoli, S., Falco, P., Gerin, R., Iermano, I., Kalampokis, A., Kokkini, Z., Lana, A., Magaldi, M.G., Mamoutos, I., Mantovani, C., Marmain, J., Potiris, E., Sayol, J.M., Barbin, Y., Berta, M., Borghini, M., Bussani, A., Corgnati, L., Dagneaux, Q., Gaggelli, J., Guterman, P.,

- Mallarino, D., Mazzoldi, A., Molcard, A., Orfila, A., Poulain, P. M., Quentin, C., Tintoré, J., Uttieri, M., Vetrano, A, Zambianchi, E. and Zervakis, V.: Toward an integrated HF radar network in the Mediterranean Sea to improve search and rescue and oil spill response: the TOSCA project experience. Toward an integrated HF radar network in the Mediterranean Sea to improve search and rescue and oil spill response: the TOSCA project experience, Journal of Operational Oceanography, 8:2, 95-107, DOI: 10.1080/1755876X.2015.1087184, 2015.
- Bower, A. S., and Farrar, J. T.: Air–sea interaction and horizontal circulation in the Red Sea. In N. M. A. Rasul & I. C. F. Stewart, (Eds.), The Red Sea, Springer Earth System Sciences (pp. 329–342). Berlin, Germany: Springer. https://doi.org/10.1007/978-3-662-45201-1_19, 2015.
- Breivik, Ø, and Saetra, Ø.: Real time assimilation of HF radar currents into a coastal ocean model. Journal of Marine Systems, Volume 28, Issues 3–4, April 2001, Pages 161-182. [https://doi.org/10.1016/S0924-7963\(01\)00002-1](https://doi.org/10.1016/S0924-7963(01)00002-1), 2001.
- Chao, Y., Li Z., Farrara, K., McWilliams, J.C., Bellingham, J., Capet, X., Chavez, F., Choi, J., Davis, R., Doyle, J., Fratantoni, D. M., Li P., Marchesiello, P., Moline, M.A., Paduan, J., Ramp, S.: Development, implementation and evaluation of a data-assimilative ocean forecasting system off the central California coast. Deep Sea Research, Vol. 56, Issues 3-5, pp 100-126. <https://doi.org/10.1016/j.dsr2.2008.08.011>, 2009.
- Crombie, D. D.: Dopler Spectrum of Sea Echo at 13.56-Mc/s', Nature 175, 681-682, 1955.
- Emery, B. M., Washburn L., and Harlan, J. A.: Evaluating radial current measurements from CODAR high-frequency radars with moored current meters. J. Atmos. Oceanic Tech- nol., 21, 1259–1271, doi:10.1175/1520-0426(2004)021,1259: ERCMFC.2.0.CO;2, 2004.

611 Esnaola, G., Sáenz, J., Zorita, E., Fontán, A., Valencia, V., and Lazure, P.: Daily
 612 scale wintertime sea surface temperature and IPC-Navidad variability in the
 613 southern Bay of Biscay from 1981 to 2010, *Ocean Sci.*, 9, 655–679,
 614 <https://doi.org/10.5194/os-9-655-2013>, 2013.

615 Fernández-Ferrero, A., Sáenz, J., Ibarra-Berastegi, G., Fernández, J.: Evaluation of
 616 statistical downscaling in short range precipitation forecast. *Atmos. Res.* 94,
 617 448–461, 2009.

618 Fernández-Ferrero, A., Sáenz, J., Ibarra-Berastegi, G.: Comparison of the
 619 performance of different Analog-Based Bayesian probabilistic precipitation
 620 forecasts over Bilbao, Spain. *Mon. Weather Rev.* 38, 3107–3119, 2010.

621 Frolov, S., Paduan J., Cook M., and Bellingham J.: Improved statistical prediction
 622 of surface currents based on historic HF- radar observations. *Ocean Dyn.*, 62,
 623 1111–1122, doi:10.1007/s10236-012-0553-5, 2012.

624 Hammond, T.M., Pattiaratchi, C.B., Osborne, M.J., Nash, L.A., Collins, M.B.:
 625 Ocean surface current radar (OSCR) vector measurements on the inner
 626 continental shelf. *Continental Shelf Research*. Volume 7, Issue 4, Pages 411-
 627 431. [https://doi.org/10.1016/0278-4343\(87\)90108-7](https://doi.org/10.1016/0278-4343(87)90108-7), 1987.

628 Hernández-Carrasco, I., Solabarrieta, L., Rubio, A., Esnaola, G., Reyes, E., and
 629 Orfila, A.: Impact of HF radar current gap-filling methodologies on the
 630 Lagrangian assessment of coastal dynamics, *Ocean Sci.*, 14, 827-847,
 631 <https://doi.org/10.5194/os-14-827-2018>, 2018.

632 Hernández-Carrasco, I., Orfila, A., Rossi, V., and Garçon, V.: Effect of small-scale
 633 transport processes on phytoplankton distribution in coastal seas, *Scientific*
 634 *Reports*, 8:8613, <https://doi.org/10.1038/s41598-018-26857-9>, 2018b.

635 Ibarra-Berastegi, G., Saenz, J., Ezcurra, A., Ezcurra, A., Elias, A., Diaz Argandona,
 636 J., Errasti, I.: Downscaling of surface moisture flux and precipitation in the
 637 Ebro Valley (Spain) using analogues and analogues followed by random

638 forests and multiple linear regression. *Hydrol. Earth Syst. Sci.* 15 (6), 1895–
639 1907, 2011.

640 Jianping H., Yuhong Y., Shaowu W and Jifen C.: An analogue-dynamical long-
641 range numerical weather prediction system incorporating historical evolution.
642 *Q. J. R. Meteorol. Soc.*, 119, pp.547-565, 1993.

643 Kaplan, D. M. and Lekien, F.: Spatial interpolation and filtering of surface current
644 data based on open-boundary modal analysis, *Journal of Geophysical*
645 *Research: Oceans*, 112, <https://doi.org/10.1029/2006JC003984>, c12007,
646 2007.

647 Kohut, J.T., Glenn, S.M.: Improving HF radar surface current measurements with
648 measured antenna beam patterns. *J. Atmos. Oceanic Technol.* 20 (9), 1303–
649 1316, 2003.

650 Lana, A., Marmain, J., Fernández, V., Tintoré, J., Orfila A.: Wind influence on
651 surface current variability in the Ibiza Channel from HF Radar. *Ocean*
652 *Dynamics*, 66: 483. <https://doi.org/10.1007/s10236-016-0929-z>, 2016

653 Langodan S., Cavaleri L., Vishwanadhapalli Y., Pomaro A., Bertotti L., Hoteit I.:
654 Climatology of the Red Sea - Part 1: the wind. *Int. J. Climatol.* 37: 4509–4517.
655 DOI: 10.1002/joc.5103, 2017.

656 Liu Y., Weisberg, R.H., and Merz, C.R.: Assessment of CODAR SeaSonde and
657 WERA HF Radars in Mapping Surface Currents on the West Florida Shelf.
658 *Journal of atmospheric and oceanic technology*, Vol 31., pp 1363:1382, 2014.

659 Lorenz, E. N.: Atmospheric Predictability as Revealed by Naturally Occurring
660 Analogues. *Journal of Atmospheric Sciences*, Volume 29, pp 636-646, 1969.

661 Martin, M.L., Valero, F., Pascual, A., Sanz, J., and Frias, L.: Analysis of wind power
662 productions by means of an analog model. *Atmos. Res.* 143, 238–249, 2014.

663 Oke, P. R., Allen, J. S., Miller, R. N., Egbert, G. D., and Kosro P. M.: Assimilation
 664 of surface velocity data into a primitive equation coastal ocean model, J.
 665 Geophys. Res., 107, 3122, doi:10.1029/2000JC000511, 2002.

666 Orfila A., Molcard, A., Sayol, J.M., Marmain, J., Bellomo, L., Quentin, C., and
 667 Barbin, Y.: Empirical Forecasting of HF-Radar Velocity Using Genetic
 668 Algorithms IEEE Transactions on Geoscience and Remote Sensing, Vol. 53,
 669 No. 5, 2015.

670 Ohlmann, C., White, P., Washburn, L., Emery, B., Terrill, E., Otero, M.:
 671 Interpretation of coastal HF radar-derived surface currents with high-
 672 resolution drifter data. J. Atmos. Oceanic Technol. 24 (4), 666–680, 2007.

673 Paduan, J.D., and Rosenfeld, L.K.: Remotely sensed surface currents in Monterey
 674 Bay from shore-based HF radar (coastal ocean dynamics application radar. J.
 675 Geophys. Res. 101 (C9), 20669–20686, 1996.

676 Paduan, J.D., and Shulman, I.: HF radar data assimilation in the Monterey Bay area.
 677 J. Geophys Res. 109:C07S09, 2004.

678 Paduan, J.D., Kim, K.C., Cook, M. S., and Chavez, F.P.: Calibration and Validation
 679 of Direction-Finding High-Frequency Radar Ocean Surface Current
 680 Observations. IEEE Journal of oceanic engineering, Vol. 31, No. 4, 2006.

681 Paduan, J.D., and Washburn, L.: High-Frequency Radar Observations of Ocean
 682 Surface Currents. Annual Rev. Marine. Sci. 2013.5:115-136, 2013.

683 Prince, K., X. and Goswami, B., N.: An Analog Method for Real-Time Forecasting
 684 of Summer Monsoon Subseasonal Variability. Monthly weather review, Vol
 685 135, pp: 4149-4160. <https://doi.org/10.1175/2007MWR1854.1>, 2007.

686 Ren L., Miaro, J., Li Y., Luo, X., Li J. and Hartnett, M.: Estimation of Coastal
 687 Currents Using a Soft Computing Method: A Case Study in Galway Bay,
 688 Ireland. Mar. Sci. Eng., 7(5), 157; <https://doi.org/10.3390/jmse7050157>, 2019

689 Roarty, H., Cook, T., Hazard, L., George, D., Harlan, J., Cosoli, S., Wyatt, L.,
 690 Alvarez Fanjul, E., Terrill, E., Otero, M., Largier, J., Glenn, S., Ebuchi, N.,
 691 Whitehouse, B., Bartlett, K., Mader, J., Rubio, A., Corgnati, L., Mantovani,
 692 C., Griffo, A., Reyes, E., Lorente, P., Flores-Vidal, X., Saavedra-Matta, K.J.,
 693 Rogowski, P., Prukpitikul, S., Lee, S.H., Lai, J.W., Guerin, C.A., Sanchez, J.,
 694 Hansen, B. and Grilli, S.: The Global High Frequency Radar Network. *Front.*
 695 *Mar. Sci.* 6:164. doi: 10.3389/fmars.2019.00164, 2019 (in press).
 696 Schmidt, R.: Multiple emitter location and signal parameter estimation. *IEEE Trans.*
 697 *Antennas Propag.*, 34, 276–280, doi:10.1109/TAP.1986.1143830, 1986.
 698 Schott F., Frisch, A.S., Leaman, K., Samuels, G., Popa Fotino, I.: High-Frequency
 699 Doppler Radar Measurements of the Florida Current in Summer 1983. *J. Geo.*
 700 *Research*, Vol 90, No C5, pp 9006:9016, 1985.
 701 Seubert, S., Fernández-Montes, S., Philipp, A., Hertig, E., Jacobeit, J., Vogt, G.,
 702 Paxian, A., Paeth, H.: Mediterranean climate extremes in synoptic
 703 downscaling assessments. *Theor. Appl. Climatol.* 117 (1–2), 257–275, 2014.
 704 Shao, Q. and Li, M.: An improved statistical analogue downscaling procedure for
 705 seasonal precipitation forecast . *Stoch Environ Res Risk Assess* 27, pp.: 819-
 706 830. <https://doi.org/10.1007/s00477-012-0610-0>, 2013.
 707 Sofianos, S. S., and Johns, W. E.: An oceanic general circulation model (OGCM)
 708 investigation of the Red Sea circulation: 2. Three- dimensional circulation in
 709 the Red Sea. *Journal of Geophysical Research*, 108(C3), 3066.
 710 <https://doi.org/10.1029/2001jc001185>, 2003.
 711 Solabarrieta, L., Rubio, A., Castanedo, S., Medina, R., Charria, G., Hernández, C.:
 712 Surface water circulation patterns in the southeastern Bay of Biscay: new
 713 evidences from HF radar data. *Cont Shelf Res* 74:60–76
 714 doi:10.1016/j.csr.2013.11.022, 2014.

715 Solabarrieta, L., Rubio, A., Cárdenas, M., Castanedo, S., Esnaola, G., Méndez, F.J.,
 716 Medina, R., and Ferrer, L.: Probabilistic relationships between wind and
 717 surface water circulation patterns in the SE Bay of Biscay. *Ocean Dyn.*, 65,
 718 1289–1303, doi:10.1007/s10236-015-0871-5, 2015.

719 Solabarrieta, L., Frolov, S., Cook, M., Paduan, J., Rubio, A., González, M., Mader,
 720 J., and Charria, G.: Skill Assessment of HF Radar-Derived Products for
 721 Lagrangian Simulations in the Bay of Biscay. *J. Atmos. Oceanic Technol.*, 33,
 722 2585–2597, doi: 10.1175/JTECH-D-16-0045.1, 2016.

723 Stanev, E.V., Schulz-Stellenfleth, J., Staneva, J., Grayek, S., Seemann, J. and
 724 Petersen, W.: Coastal observing and forecasting system for the German Bight
 725 – estimates of hydrophysical states. *Ocean Sci.*, 7, 569–583, 2011
 726 doi:10.5194/os-7-569-2011, 2011.

727 Ullman, D.S., O'Donnell, J., Kohut, J., Fake, T., Allen, A.: Trajectory prediction
 728 using HF radar surface currents: Monte Carlo simulations of prediction
 729 uncertainties. *J. Geophys. Res.* 111 (C12005), 1–14, 2006.

730 Vilibić, I., Šepić, J., Mihanović, H., Kalinić, H., Cosoli, S., Janeković, I., Žagar, N.,
 731 Jesenko, B., Tudor, M., Dadić, V. and Ivanković, D.: Self-organizing maps-
 732 based ocean currents forecasting system. *Scientific Reports* 6, 22924, 2016.

733 Yao, F., Hoteit, I., Pratt, L. J., Bower, A. S., Zhai, P., Kohl, A., and Gopalakrishnan,
 734 G.: Seasonal overturning circulation in the Red Sea: 1. Model validation and
 735 summer circulation, *J. Geophys. Res. Oceans*, 119, doi:10.1002/
 736 2013JC009004, 2014.

737 Yao, F., Hoteit, I., Pratt, L. J., Bower, A. S., Kohl, A., Gopalakrishnan, G., and
 738 Rivas, D.: Seasonal overturning circulation in the Red Sea: 2. Winter
 739 circulation, *J. Geophys. Res. Oceans*, 119, 2263–2289, doi:10.1002/
 740 2013JC009331, 2014.

741 Zarokanellos, N. D., Kürten, B., Churchill, J. H., Roder, C., Voolstra, C. R.,
 742 Abualnaja, Y., and Jones, B. H.: Physical mechanisms routing nutrients in the
 743 central Red Sea. *Journal of Geophysical Research: Oceans*, 122.
 744 <https://doi.org/10.1002/2017JC013017>, 2017a.

745 Zarokanellos, N. D., Papadopoulos, V. P., Sofianos, S. S., and Jones, B. H.: Physical
 746 and biological characteristics of the winter-summer transition in the Central
 747 Red Sea. *Journal of Geophysical Research: Oceans*, 122, 6355–6370.
 748 <https://doi.org/10.1002/2017JC012882>, 2017b.

749 Zhan, P., Subramanian, A. C., Yao, F., and Hoteit, I.: Eddies in the Red Sea: A
 750 statistical and dynamical study, *J. Geophys. Res. Oceans*, 119, 3909–3925,
 751 doi:10.1002/2013JC009563, 2014.

752 Zelenke B. C.: An Empirical Statistical Model Relating Winds and Ocean Surface
 753 Currents. Master of Science in Oceanography - Thesis, Oregon State
 754 University, 2005.

755 Zorita E, and von Storch H.: The analog method as a simple statistical downscaling
 756 technique: comparison with more complicated methods. *J Climate* 12:2474–
 757 2489, 1999.

758 **TABLES**759 *Table 1: Characteristics of the previously developed STP works based on HFR data.*

Authors	Approach	Needs continuous training period	Complementary data required?	Region of application	Reliable forecast period
Zelenke 2005	EOF + bilinear regression model	Yes	Wind	Oregon coast	48 hours
Frolov et al. 2012	EOF + linear auto regression model	Yes	Wind and tides (optional)	Monterey Bay, California	48 hours
Barrick et al., 2012	Constant linear trend model applied to OMA modes	Yes	Wind	Finnmark, Norway	12 hours
Orfila et al. 2015	EOF+Genetic Algorithm	Yes	No	Toulon, France	48 hours
Solabarrieta et al. 2016	Frolov et al., 2012	Yes	No	Bay of Biscay	48 hours
Vilibić et al., 2016	SOM+neural network +winds	Yes	Wind	Northern Adriatic Sea	72 h
Ren et al., 2019	Random Forest (RF) classification algorithm	No	Tide and Wind	Galway Bay, Ireland	59 h
This paper: L-STP	Analogue finding	No	No	Bay of Biscay and the Central Red Sea	48 h

760

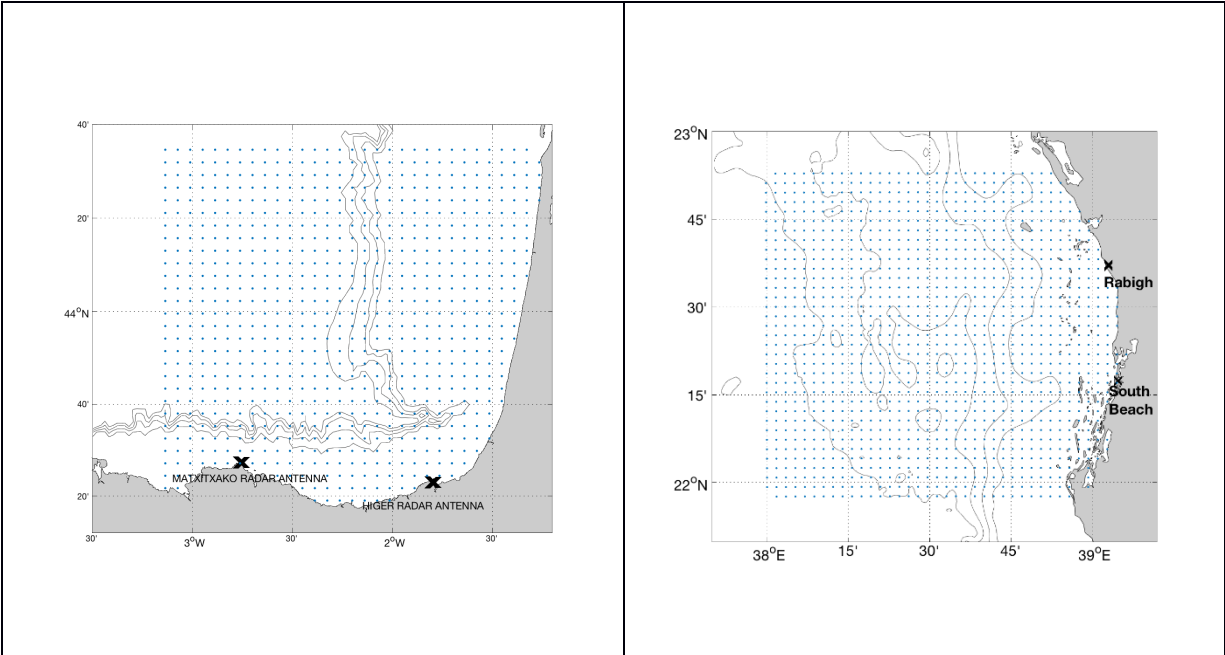
761 Table 2: Correlation coefficient values between winner ε_{ANL} and δ_{STP} and between ε_{ANL} and
762 δ_{PRS} , after 6, 12, 24, 36 and 48 hours of simulation.

	6 hours	12 hours	24 hours	36 hours	48 hours
$R^2 \varepsilon_{ANL} - \delta_{STP}$	0.19	0.37	0.55	0.56	0.54
$R^2 \varepsilon_{ANL} - \delta_{PRS}$	0.07	0.11	0.03	0.01	0.04
ε_{ANL} [km2], for the inflection point between δ_{STP} and δ_{PRS}	-	714	774	857	1027
% of ε_{ANL} (accumulative) for the previous line	-	81	84	87	95

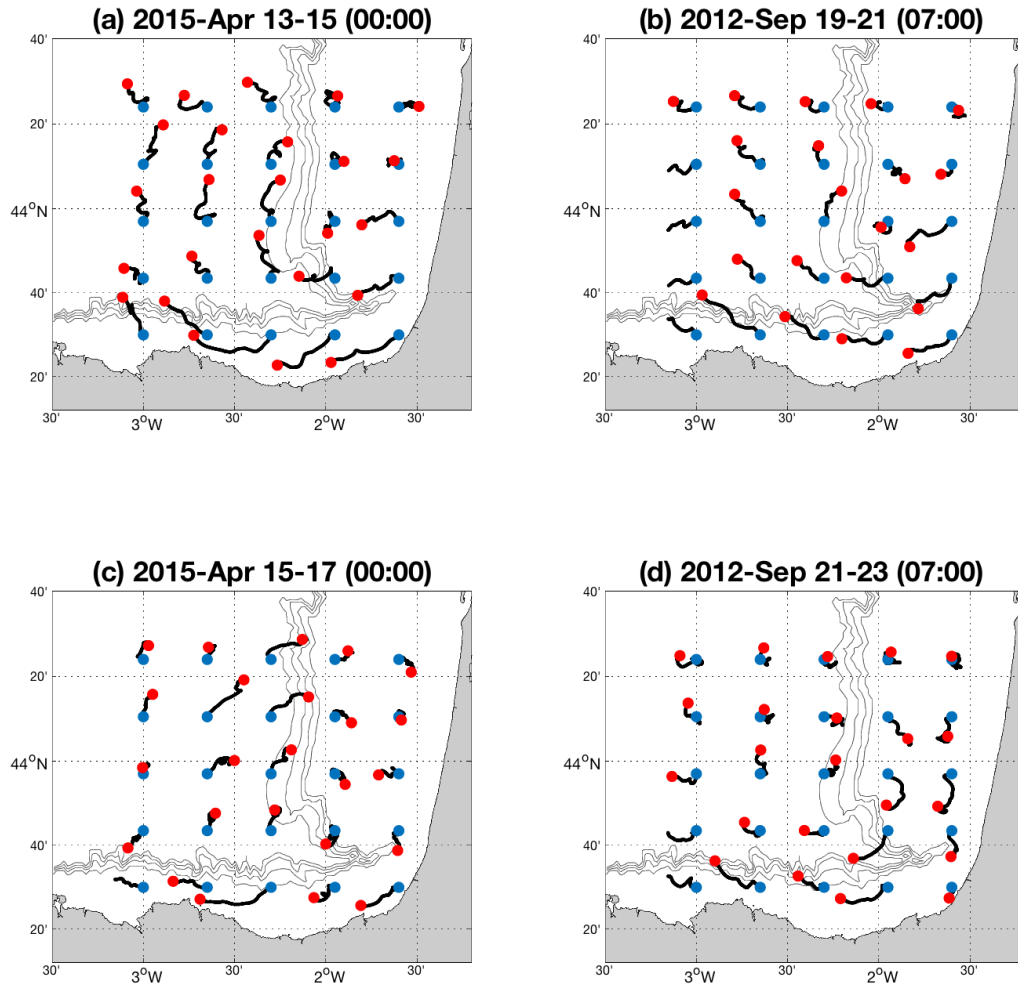
763

FIGURES

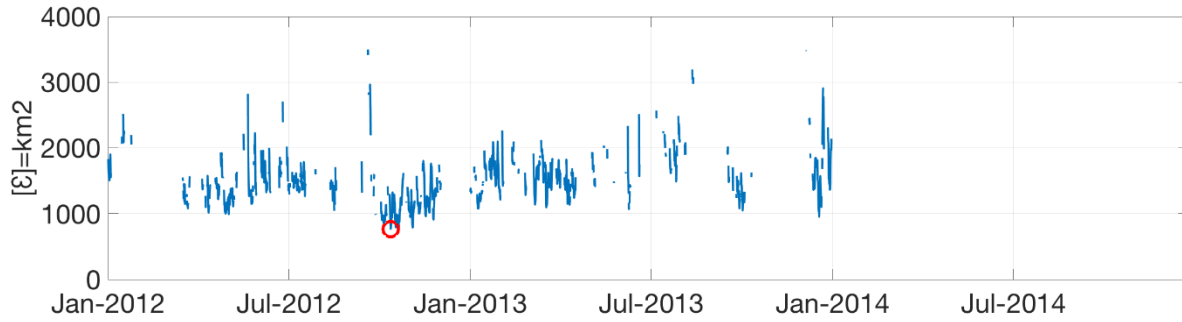
Figure 1: (Left) HFR system of the BoB. (Right) HFR system of the central Red Sea. Blue dots represent the data points and the black cross are the HFR antenna positions



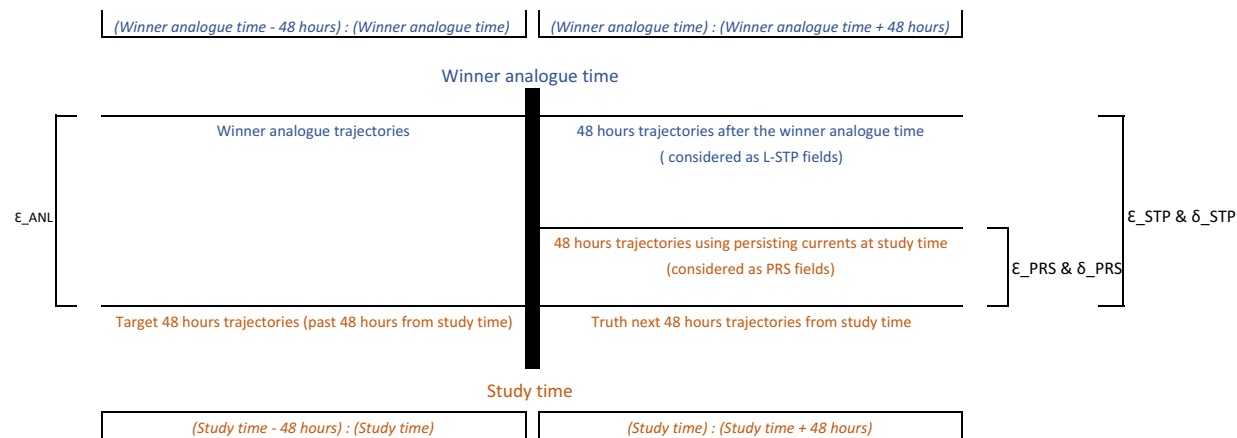
772 *Figure 2: (1) 15-Apr-2015 00:00 example of the developed methodology applied to*
773 *the BoB HFR system. (a) The past 48 hours of target field of test period (b) The*
774 *analogue having the lowest error, (c) The truth trajectories for the forecast period*
775 *(d) the STP trajectories. The initial positions of the particle trajectories are*
776 *indicated by the blue dots, and the red dots indicate the position after 48 hours.*



778 *Figure 3: Example for the test period: 15-Apr-2015 00:00; errors for the whole*
 779 *Lagrangian catalogue fields of the BoB HFR System, restricted to the $\delta_{cg} = 10$ km*
 780 *condition. The red dot indicates the occurrence date and the error of the winner*
 781 *analogue (19-Sep-2012 07:00).*



783 *Figure 4: Scheme of the analogue selection and L-STP forecast assessment process.*

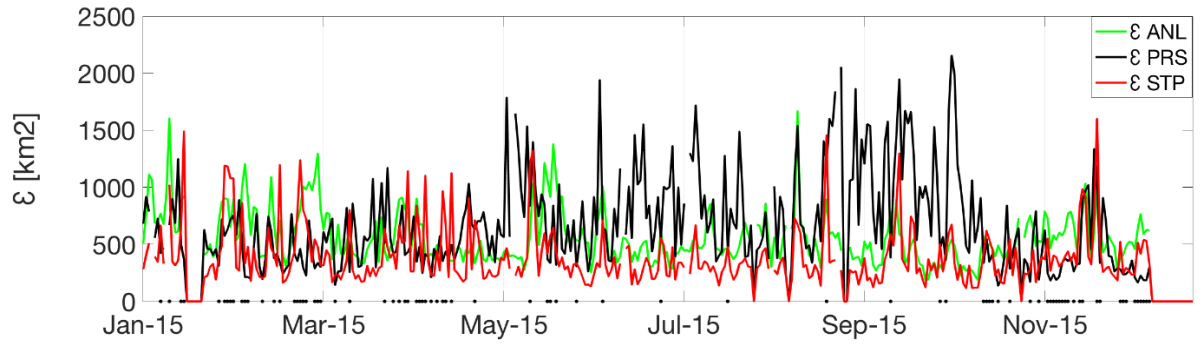


ϵ_{ANL} is used to select the winner/best analogue (min ϵ)

ϵ_{STP} , ϵ_{PRS} , δ_{STP} and δ_{PRS} are used to validate the methodology and estimate the final error or separation distances between truth and L-STP/PRS trajectories

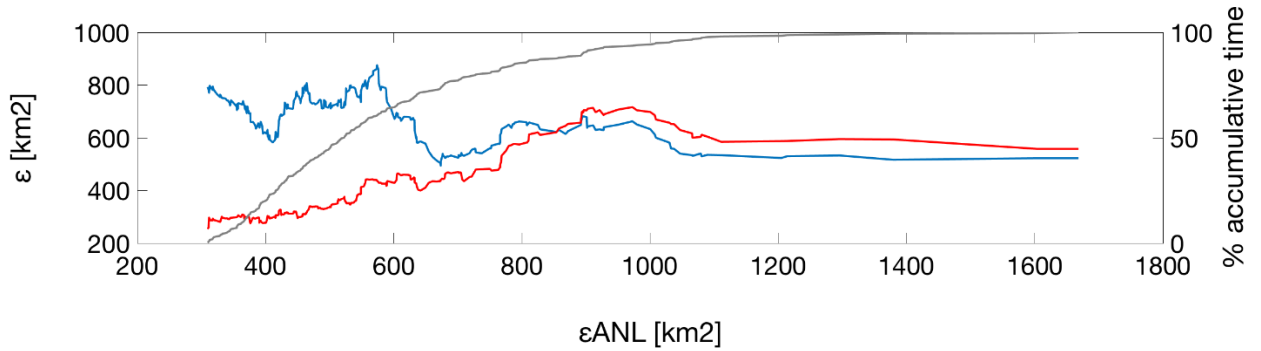
785

786 *Figure 5: errors of the hourly winner analogue for 2015 (ϵ_{ANL}), together with the*
787 *ϵ_{STP} and ϵ_{PRS} . The black dots over the timeline shows the times when ϵ_{STP} is higher*
788 *than ϵ_{PRS}*

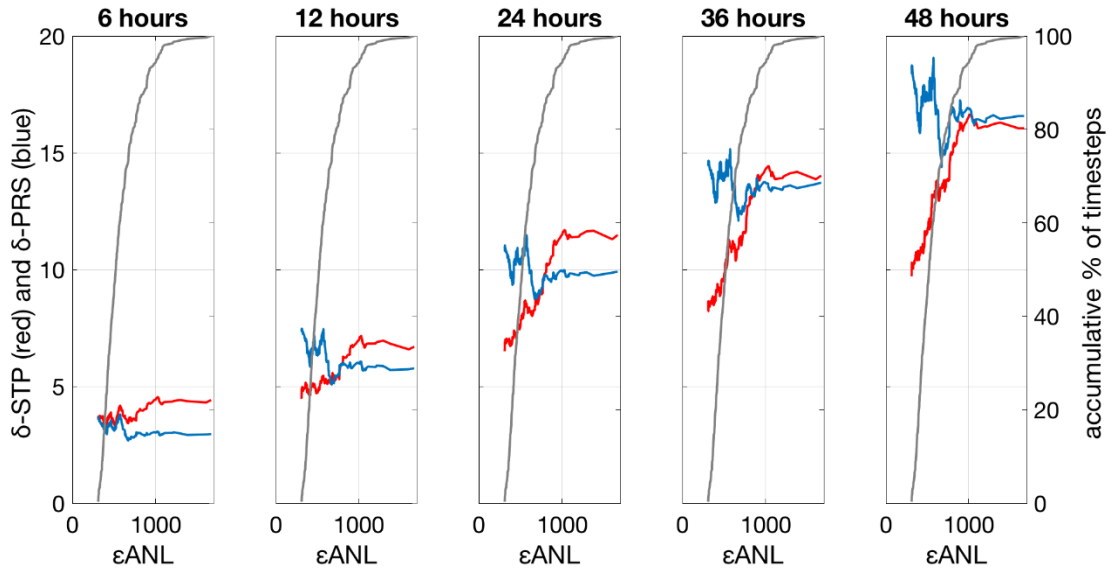


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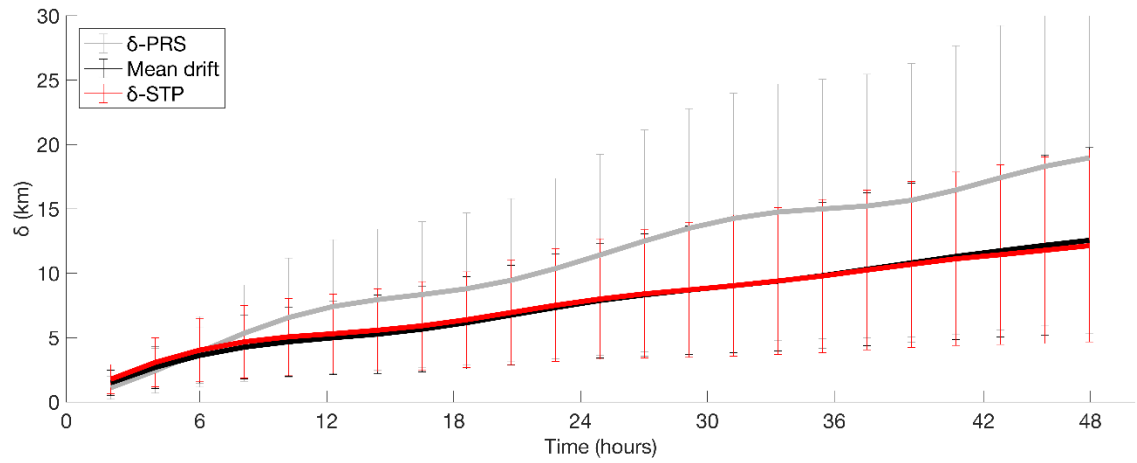
790 Figure 6: X axis shows the ε_{ANL} , ordered from minimum to maximum, for the winner
 791 analogue for the test year 2015. Left Y axis indicates ε_{STP} (red) and ε_{PRS} (blue) for
 792 the corresponding ε_{ANL} . Right Y axis indicates the % of the accumulative comparison
 793 times as shown by the black solid line. Dashed vertical line indicates the crossing
 794 point between ε_{STP} and ε_{PRS} ($\varepsilon_{ANL}=853\text{Km}^2$).



796 Figure 7: Left Y axis indicates δ_{STP} (red) and δ_{PRS} (blue) for the corresponding
 797 ϵ_{ANL} , after 6, 12, 24, 36 and 48 hours. Right Y axis is the cumulative % of timesteps
 798 in the computation of the mean errors, as indicated by the black line in the plots. X
 799 axis is the ϵ_{ANL} , ordered from minimum to maximum, for the winner analogue for the
 800 test year 2015.

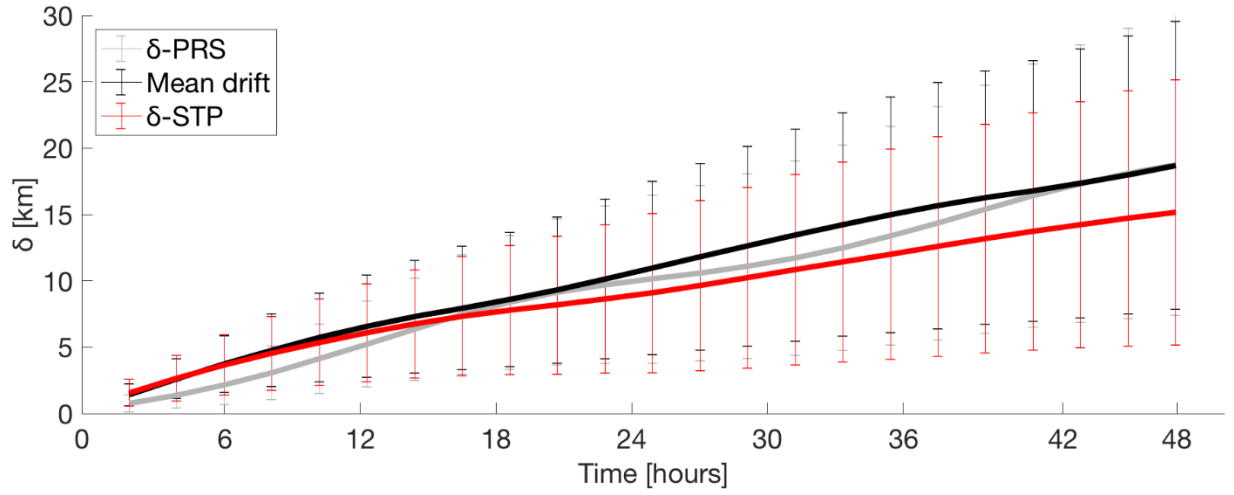


802 *Figure 8: Time evolution of the mean separation δ_{STP} and δ_{PRS} [km] between truth*
803 *and forecast trajectories using truth and STP/PRS currents and the mean drift, with*
804 *BoB system data, for 2015.*



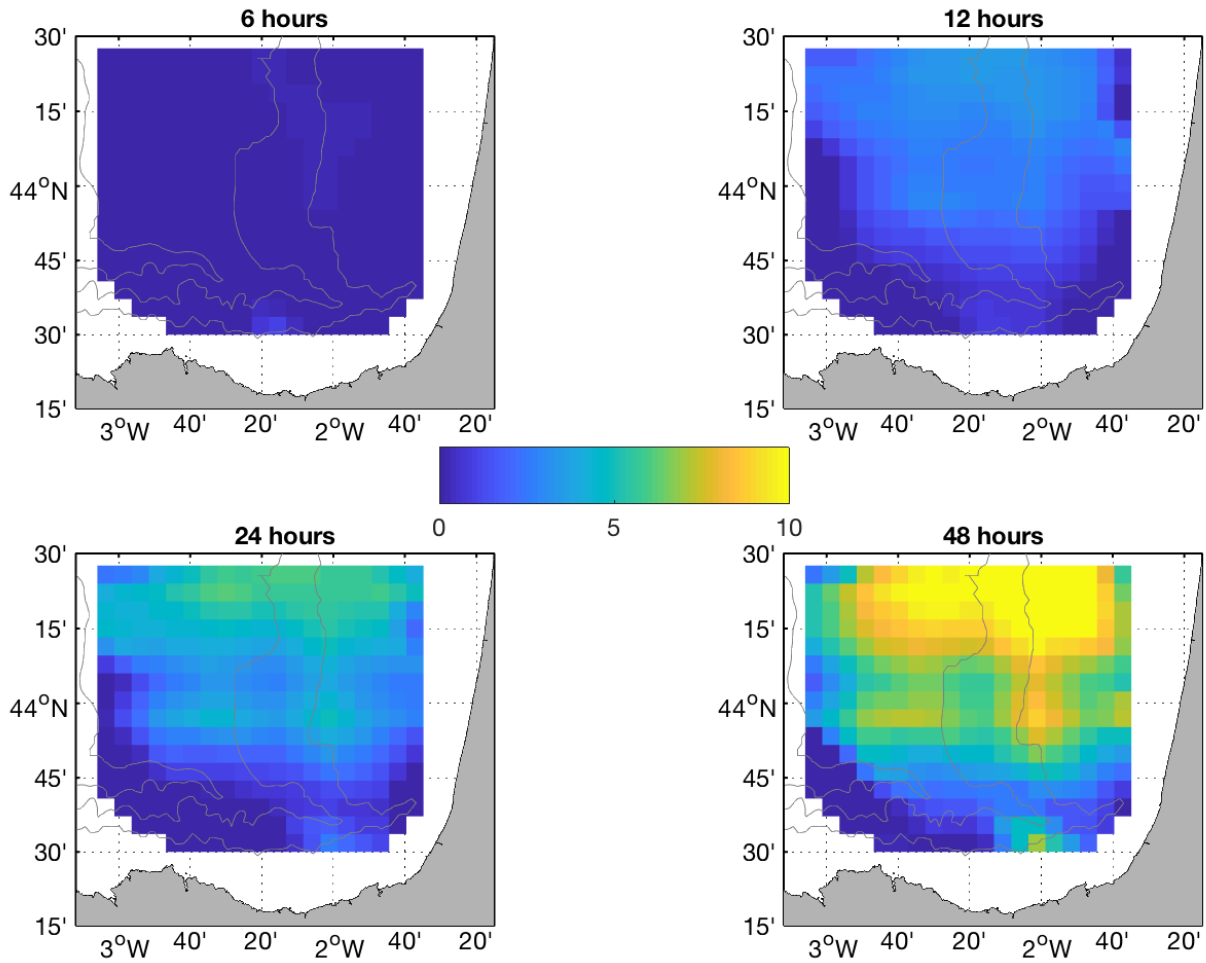
805

806 *Figure 9: Time evolution of the mean separation distances δ_{STP} and δ_{PRS} [km]*
 807 *between real and forecast trajectories using truth and STP/PRS currents and the*
 808 *mean drift, with the Red Sea HFR system data, for July 2017 to October 2018.*



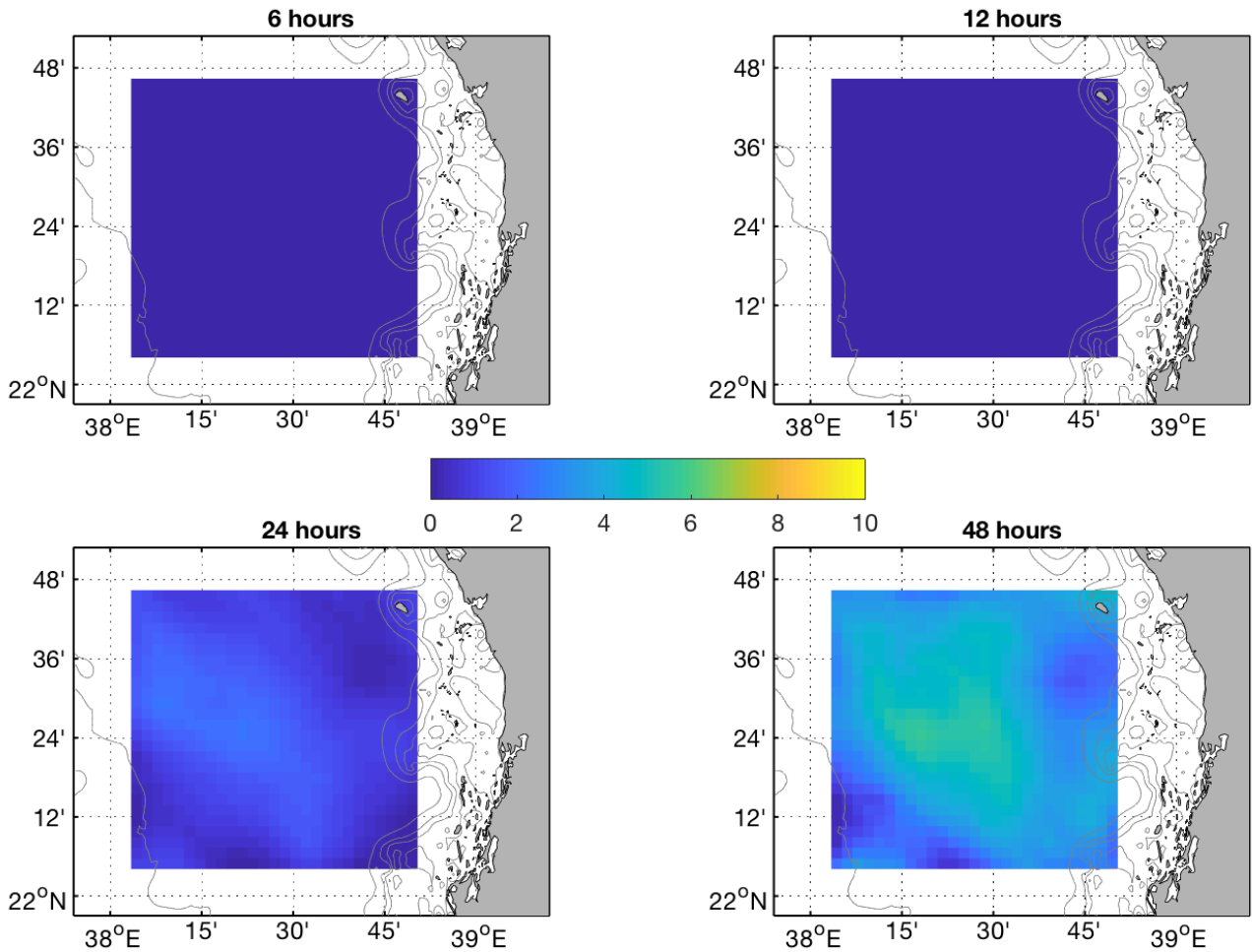
809

810 *Figure 10: Spatial distribution of separation distances [km] between trajectories*
811 *using L-STP and persistent currents at 6, 12, 24 and 48 hours, for the BoB HFR*
812 *System.*



813

814 *Figure 11: Spatial distribution of separation distances [km] between trajectories*
815 *using L-STP and persistent currents at 6, 12, 24 and 48 hours, for the Red Sea HFR*
816 *system.*



817