1	FINAL MANUSCRIPT
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3	MS No.: os-2019-11
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8	USING AN ADVANCED NEURAL CLASSIFIER
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19 ABSTRACT

We processed daily ocean-color satellite observations to construct a monthly climatology of 20 phytoplankton pigment concentrations in the Senegalo-Mauritanian region. Our proposed new method 21 22 primarily consists of associating, in well-identified clusters, similar pixels in terms of ocean-color parameters and in situ pigment concentrations taken from a global ocean database. The association is 23 24 carried out using a new Self Organized Map (2S-SOM). Its major advantage is to allow taking into 25 account the specificity of the optical properties of the water by adding specific weights to the different 26 ocean color parameters and the in situ measurements. In the retrieval phase, the pigment concentration 27 of a pixel is estimated by taking the pigment concentration values associated with the 2S-SOM cluster presenting the ocean-color satellite spectral measurements, which are the closest to those of the pixel 28 29 under study according to some distance. The method was validated by using a cross-validation 30 procedure. We focused our study on the fucoxanthin concentration, which is related to the abundance 31 of diatoms. We showed that the fucoxanthin starts to develop in December, presents its maximum 32 intensity in March when the upwelling intensity is maximum, extends up to the coast of Guinea in 33 April and begins to decrease in May. The results are in agreement with previous observations and recent in situ measurements. The method is very general and can be applied in every oceanic region. 34

36 1 - INTRODUCTION

37

38 Phytoplankton are the basis of the ocean food web and consequently drive the ocean productivity. 39 They also play a fundamental role in climate regulation by trapping atmospheric carbon dioxide (CO₂) through gas exchanges at the sea surface, and consequently lowering the rate of anthropogenic increase 40 41 in the atmosphere of CO₂ concentration by about 25% (Le Quéré et al, 2018). With the growing interest 42 in climate change, one may ask how the different phytoplankton populations will respond to changes 43 in ocean characteristics (temperature, salinity, acidity) and nutrient supply, which presents an 44 important societal impact with respect to both climate and fisheries, with a possible effect on fish 45 grazing phytoplankton via the marine food chain.

Methods for identifying phytoplankton have greatly progressed during the last two decades. 46 47 Phytoplankton were first described by microscopy. Microscopy is time consuming and is unable to identify picoplankton. Imaging flow cytometry (IFC) has renewed microscopic methods, thanks to the 48 49 speed at which they are able to characterize phytoplankton in a water sample (IOCCG report n°15, 2014). An alternative method is the analysis of seawater samples by high-performance liquid 50 51 chromatography (HPLC) which is widely used to categorize broad phytoplankton groups such as PFT 52 or PSC (Jeffreys et al, 1997, Brewin et al, 2010, Hirata et al, 2011). HPLC enables the identification of 25 to 50 pigments within a single analysis, which is much easier and faster to conduct than 53 microscopic observations (Sosik, H.M et al, 2014). Each phytoplankton group is associated with 54 55 specific diagnostic pigments, and a conversion formula, the so-called "Diagnostic Pigment Analysis" 56 can be derived to estimate the percentage of each group from the pigment measurements (Vidussi et 57 al, 2001; Uitz et al, 2010). HPLC measurements are now recognized as the standard for calibrating and validating satellite-derived chlorophyll-a concentration and for mapping groups of phytoplankton 58 59 (IOCCG report n°15, 2014).

The use of satellite ocean color sensor measurements has permitted to map the ocean surface at a daily frequency. Satellite sensors measure the sunlight, at several wavelengths, backscattered by the ocean. The downwelling sunlight interacts with the seawater through backscattering and absorption in such a manner that the upwelling radiation transmitted to the satellite ('water-leaving' reflectance) contains information related to the composition of the seawater. The light transmitted to the satellite depends on the phytoplankton cell shape (backscattering), its pigments (absorption), the dissolved matter (e.g. CDOM).

This upwelling radiation, the so-called remotely sensed reflectance $\rho_w(\lambda)$, is determined by the spectral absorption *a* and backscattering (b_b (m⁻¹)) coefficients of the ocean (pure water and various particulate

$$\rho_w(\lambda) = G b_b(\lambda) / (a(\lambda) + b_b(\lambda))$$
(1)

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where $(a \text{ (m}^{-1}))$ is the sum of the individual absorption coefficients of water, phytoplankton pigments, colored dissolved organic matter, and detrital particles, $(b_b \text{ (m}^{-1}))$ depends on the shape of the phytoplankton species. *G* is a parameter mainly related to the geometry of the situation (sensor and solar angles) but also to environmental parameters (wind, aerosols).

In the open ocean far from the coast (in case-1 waters), the light seen by the satellite sensor mainly
contains information on phytoplankton abundance and diversity. Ocean-color measurements have
been first used intensively to estimate chlorophyll-*a* concentration (*chl-a* in the following) in the
surface waters of the ocean, marginal seas and lakes. (*Longhurst et al.*, 1995; *Antoine et al.*, 1996; *Behrenfeld and Falkowski*, 1997; *Behrenfeld et al.*, 2005; *Westberry et al.*, 2008).

82 It has been shown that it is also possible to extract additional information such as phytoplankton sizeclasses (PSC) by using some relationship between chlorophyll concentration and PSC (Uitz et al., 2006; 83 84 *Ciotti and Bricaud*, 2006; *Hirata et al.*, 2008; *Mow and Yoder*, 2010). These algorithms try to establish a relationship between the *chl-a* concentration and the *chl-a* concentration fractions associated with 85 each of the three PSC. Some of them (Uitz et al, 2006; Aiken et al., 2009) break-down the chl-a 86 abundance into several ranges for each of which a specific relationship is computed. Others (Brewin 87 et al, 2010; Hirata et al, 2011) are based on a continuum of chl-a abundance. Studies have also been 88 done to estimate the phytoplankton groups (PFT) by taking into account spectral information 89 90 (Sathyendranath et al., 2004, Alvain et al., 2005, 2012; Hirata et al., 2011; Ben Mustapha et al, 2013; 91 Farikou et al, 2015). This is of fundamental interest to the understanding of the phytoplankton behavior 92 and to modeling its evolution.

Due to highly non-linear relationship linking the multispectral ocean color measurements with the pigment concentrations, we proposed a neural network clustering algorithm (2S-SOM) able to deal with multi variables linked by complex relationships. The 2S-SOM algorithm is well adapted to this complex task by weighting the different inputs. The clustering algorithm was calibrated on a restricted database composed of remote sensed observations co-located with measurements taken in the global ocean.

99 In the present paper, we propose the retrieval of the major pigment concentrations from satellite ocean 100 color multi-spectral sensors in the Senegalo-Mauritanian upwelling, which is an oceanic region off the 101 coast of West Africa where а strong seasonal upwelling occurs (Figure 1). 102



Figure 1: Mauritania and Senegal coastal topography. The land is in brown and the ocean depth is
represented in meters by the color scale on the right side of the figure. The UPSEN stations are shown
at the bottom left cartoon of the figure.

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The Senegalo-Mauritanian upwelling is one of the most productive eastern boundary upwelling systems (EBUS) with strong economic impacts on fisheries in Senegal and Mauritania. Since the region has been poorly surveyed in situ, we have chosen to extract pertinent biological information from ocean-color satellite measurements. The region has been intensively studied by analysis of SeaWiFS ocean-color data and AVHRR sea-surface temperature as reported in *Demarcq* and *Faure* (2002), *Sawadogo et al.* (2009), *Farikou et al.* (2013, 2015), *Ndoye et al.* (2014) and more recently by *Capet et al.* (2017) with in situ observations.

The paper is articulated as follows: in section 2, we present the data we used (in situ and remote sensing observations). The mathematical aspect of the clustering method (2S-SOM) is detailed in section 3. In section 4 we present the methodological results. The spatio-temporal variability of the fucoxanthin and chl-a concentration in the Senegalo-Mauritanian upwelling region are presented in section 5, as well as the results of the oceanic UPSEN campaigns. In section 6 we discuss the results and the method. A conclusion is presented in section 7.

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126 **2- MATERIALS**

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In this study we used three distinct datasets: the first was used to calibrate the method, the second to conduct a climatological analysis of the Senegalo-Mauritanian upwelling region and the third was obtained during the oceanographic UPSEN campaign. These datasets are composed of satellite remote sensing observations and in-situ measurements.

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133 2.1 The calibration data base (DPIG)

The calibration database (DPIG) comprises in situ pigment measurements co-located with satellite
 ocean-color observations done by the SeaWiFS (Sea-viewing, Wide-Field-of-view Sensor).

136 This DPIG is composed of 515 matched satellite observations and in situ measurements made in the 137 global ocean (mainly in the North Atlantic and the equatorial ocean; Ben Mustapha et al., 2014). The 138 match-up criteria were quite severe: we used satellite pixel situated at a distance less than 20km from the in situ measurement in a time window of +/-12h. The geographic distribution of the 515 coincident 139 140 in situ and satellite measurements is shown in Fig. 2. Matchup procedure between in situ and satellite observation is a crucial question to estimate remote sensing algorithms. If the parameters of the 141 142 procedure are too severe, the number of collocated data is 143



144 145

Figure 2: Geographic positions of the 515 in situ and satellite collocated measurements of the
DPIG database.

dramatically decreasing. If the parameters are too large, it is the accuracy of the matching, which is
decreasing. We accordingly chose some compromise. Usually people use a matchup window of 3X3 pixels

- (*Alvain et al*, 2005) which corresponds to a distance less than 20km between the satellite pixel and in
 situ measurement, since we deal with level 3 satellite observations whose pixel is of the order of 9X9km.
 This criterium refers to the typical length of ocean variability (*Levy et al*, 2012; *Levy*, 2003)
- 154
- 155 In Figure 3 we present the R^2 coefficient between the in situ *chl-a* a and the SeaWiFS *chl-a* a computed
- by using the OC4V4 algorithm (*O'Reilly et al*, 2001) for the DPIG collocated observations. We remark
- 157 that the two measurements are in good agreement at global scale. Each data of DPIG is a vector
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Figure 3: Dispersion diagram of DPIG chl-a computed from the SeaWiFS observations using the
OC4V4 algorithm versus in situ chl-a. The coefficient of vraisemblance R² and the RMSE (Root Mean
Square Error) were computed in mg m⁻³

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having 17 components (five ocean reflectance ($\rho_w(\lambda)$ and $Ra(\lambda)$ at five wavelengths (412, 443, 490,

- 167 510 and 555nm), SeaWiFS *chl-a*, five in situ pigment ratios and in situ *chl-a* concentration). The in
- 168 situ *chl-a* a concentration ranges between 0.007 and 3 mg m^{-3} (see Table 1).
- 169 The five $Ra(\lambda)$ are defined following Alvain et al, (2012 :
- 170 $Ra(\lambda) = \rho_W(\lambda) / \rho_{Wref}(\lambda, chl-a)$ (2)
- where the parameter $\rho_{wref}(\lambda, chl_a)$ is an average reflectance depending on the *chl-a* concentration only which was computed according to the procedure reported in *Farikou et al*, 2015. *Ra*(λ) is a non-

- dimensional parameter which depends on the *chl-a* abundance at second order and is mainly sensitive
 to the secondary pigments (*Alvain et al*, 2012).
- 175

176 The DPIG database thus provides information on the existing links between the pigment composition

- and the SeaWiFS measurements. The pigment composition are defined by the pigment ratios whichare non-dimensional variables of the form in the present study:
- 179 Pigment Ratio=DP/T*chl-a*
- 180 which is defined as the ratio of the diagnostic pigment (DP) versus the total *chl-a* 181 (T*chl-a* = *chl-a* +divinyl *chl-a*, according to *Alvain et al.*, 2005).
- 182

The pigments of the DPIG and their statistical characteristics are given in Table 1. The statistical tests presented in Figure 3 (R^2 and RMSE) and in Table 1 (MEAN, STD, MIN, MAX) were computed in mg m⁻³.

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- 187

	RDIVINY A	RPERID	RFUCO	R19HF	RZEAX	CHLORO IN SITU
MEAN	0.1414	0.0272	0.1248	0.1859	0.1696	0.5292
STD	0.1584	0.0196	0.0971	0.0996	0.2063	0.5720
MIN	0.0037	0.0035	0.0053	0.0066	0.0027	0.007
MAX	0.8889	0.2027	0.8514	0.7654	1.5574	2.9980

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193 2.2 The Senegalo-Mauritanian upwelling satellite data (DSAT)

The satellite dataset we processed to retrieve the pigment concentration consist of five $\rho_w(\lambda)$ and five *Ra(\lambda)* at five wavelengths (412, 443, 490, 510 and 555nm), and the SeaWiFS *chl-a* concentration observed in the Senegalo-Mauritanian upwelling region (8°N-24°N, 14°W-20°W; Figure 3) during 11 years (1998-2009) by SeaWiFS. This data set is here below denoted *DSAT*.

- 198 The satellite observations ($\rho_w(\lambda)$ and *chl-a* concentration) were provided by NASA with a resolution
- 199 of nine kilometers. Due to the presence of Saharan dusts in this region, very few estimations of satellite
- 200 $\rho_w(\lambda)$ and in situ *chl-a* were available, and some satellite estimations of *chl-a* could present strong over-
- estimations (*Gregg et al*, 2004). For this reason, we reprocessed the $\rho_w(\lambda)$ and *chl-a* data with an

atmospheric correction algorithm developed specifically for Saharan dust (*Diouf et al*, 2013,
 http://poac.locean-ipsl.upmc.fr) in order to improve the satellite observations.

(3)

<sup>Table 1: Pigments of the DPIG and their statistical characteristics: STD (Standard Deviation), MIN
(minimum value), MAX (maximum value).</sup>

205 2.3 The UPSEN database

Recently, some HPLC measurements were made in the Senegalo-Mauritanian region during two 206 207 oceanographic cruises (UPSEN campaigns) of the oceanographic ship "Le Suroit" from 7 to 17 March 208 2012 and from 5 to 26 February 2013 as reported in Ndoye et al, (2014); Capet et al, (2017). The goal 209 was to study the dynamics and the biological variability of the Senegalo-Mauritanian upwelling. During these campaigns, in-situ HPLC measurements were carried out. We expected to be able to co-210 211 locate them with the ocean-color VIIRS (Visible Infra-red Imaging Radiometer Suite) sensor 212 observations whose wavelengths are close to those of the SeaWiFS. Unfortunately, we were only able 213 to process satellite observations made on 21 February 2013 due to the presence of clouds and Saharan 214 aerosols the other days. We processed the satellite observations provided by the VIIRS sensor at four 215 wavelengths (443, 490, 510, 555 nm) for pixels in the vicinity of the ship stations (within a distance 216 of 20km) and observed in a time window of +/-12h, and for which the satellite *chl-a* was less than 3 mg m⁻³, which is the limit of validity of our method imposed by the range of *chl-a* observed in DGIP 217 218 (mean of 0.52 mg m⁻³). Only five stations off Cabo Verde peninsula fitted these requirements (see Figure 1 for their positions). 219

220 **3 - THE PROPOSED METHOD (2S-SOM)**

Classification methods were applied for retrieving geophysical parameters from large databases in several studies including weather forecasting (*Lorenz*, 1969; *Kruizinga and Murphy*, 1983), short-term climate prediction (*Van den Dool*, 1994), downscaling (*Zorita and von Storch*, 1999), reconstruction of oceanic pCO₂ (*Friedrichs and Oschlies.*, 2009), and of *chl-a* concentration under clouds (*Jouini et al*, 2013). In the present study, we used a new neural network classifier, which is an extension of the SOM algorithms.

227 3-1 The SOM clustering

The SOM algorithms (*Kohonen*, 2001) constitute powerful nonlinear unsupervised classification methods. They are unsupervised neural classifiers, which have been commonly used to solve environmental problems (*Cavazos*, 1999; *Hewitson et al*, 2002; *Richardson et al*, 2003; *Liu et al*, 2005, 2006; *Niang et al*, 2003, 2006; *Reusch et al*, 2007). The SOM aims at clustering vectors $\mathbf{z}_i \in \mathbb{R}^N$ of a multidimensional database \mathbf{D} . Clusters are represented by a fixed network of neurons (the SOM map), each neuron *c* being associated with the so-called referent vector $\mathbf{w}_c \in \mathbb{R}^N$ representing a cluster. The self-organizing maps are defined as an undirected graph, usually a rectangular grid of size $p \times q$. This 235 graph structure is used to define a discrete distance (denoted by δ) between two neurons of the p x q rectangular grid which presents the shortest path between two neurons. Each vector z_i of **D** is assigned 236 to the neuron whose referent w_c is the closest, in the sense of the Euclidean distance: w_c is called the 237 projection of the vector z_i on the map. A fundamental property of a SOM is the topological ordering 238 provided at the end of the clustering phase: close neurons on the map represent data that are close in 239 240 the data space. The estimation of the referent vectors w_c of a SOM and the topological order is achieved through a minimization process in which the referent vectors w are estimated from a learning data set 241 (The DPIG data base in the present case). The cost function is shown in Annex: 242

- The SOMs have frequently been used in the context of completing missing data (*Jouini et al*, 2013), so the projected vectors z_i may have missing components. Under these conditions, the distance between a vector $z_i \in D$ and the referent vectors w_c of the map is the Euclidean distance that considers only the existing components (the Truncated Distance or *TD* hereinafter).
- 247

248 *3-2 The 2S-SOM Classifier*

249 In the present case, we used the 2S-SOM algorithm, which is a modified version of the SOM, very powerful in the case of a large number of variables. It automatically structures the variables having 250 251 some common characters into conceptually meaningful and homogeneous blocks. The 2S-SOM takes advantage of this structuration of **D** and the variables into different blocks, which permits an automatic 252 253 weighting of the influence of each block and consequently of each variable. The block weighting 254 facilitates the clustering procedure by considering the most pertinent variables. The vectors of DPIG 255 defined in section 2 can be decomposed in four blocks. The essence of this decomposition in blocks is that each of the 17 components of the DPIG vectors gathered information with a different physical 256 influence in the classification phase. The composition of each block is done as follows: 257

- *First Block* (B1) comprises the five pigment in-situ concentration ratios (divinyl chlorophyll-a,
 peridinin, fucoxanthin, 19'hexanoyloxyfucoxanthin, zeaxanthin concentration ratios). The pigment
 ratios are defined in Eq. 3.
- 261 **Second Block** (B2) comprises the water-leaving reflectance $\rho_w(\lambda)$ at the five SeaWiFS wavelengths
- 262 **Third Block** (B3) comprises the five $Ra(\lambda)$,
- 263 *Fourth Block* (B4) comprises two variables: The in situ and the SeaWiFS *chl-a* concentrations.

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The 2S-SOM is able to deal with a large quantity of variables, choosing those that are the most significant for the classification and neutralizing those which are the least significant. This is done by

- estimating weights on the blocks and the variables. We fully describe the 2S-SOM algorithm in Annex.
 In the following we use a simplified version of 2S-SOM in which only the blocks are weighted.
- 269

270 3.3 The calibration phase

- Similarly to the standard SOM, the 2S-SOM is determined through a learning phase by using a more complex cost function (see Annex) that estimate for each neuron, in addition to the referent vector, a weight (α) for each block. For a neuron *c*, we define the weights α_{cb} of each block *b* (*b* = 1....4).
- At the end of the calibration phase, each element z_i of the dataset DPIG is associated with a referent *w_c* whose components are partitioned into four blocks. In the present study, the 2S-SOM map is represented by a two-dimensional (9x18=162) grid that represents the partition of the DPIG dataset
- 277 into different classes. Each class provided by the 2S-SOM is associated with a so-called referent vector
- 278 w_c with $c \in \{1,...,162\}$. The size of the map has been determined by using the procedure provided by
- the SOM software available at : http://www.cis.hut.fi/projects/somtoolbox/download/.
- 280

281 3.4 The Pigment retrieval

- 282 In the second phase, which is an operating phase, we estimated the pigment concentration ratios of a pixel PX_m from its satellite ocean-color sensor observations only. The 11 ocean color satellite 283 284 observations (5 $\rho_w(\lambda)$, 5 $Ra(\lambda)$, and *chl-a*) of pixel PX_m were projected onto the 2S-SOM using the 285 Truncated Euclidian Distance (section 3.1). We select the neuron c associated with a referent vector whose the 11 ocean-color parameters are the closest to those observed by the satellite sensor. The 286 287 pigment ratios of PX_m are those associated with the neuron c. At the end of the assignment phase, each pixel PX_m of a satellite image is associated with a referent vector w_c , which has 6 pigment 288 289 concentration ratios among its 17 components. The flowcharts of the method (2S-SOM learning and 290 pigment retrieval) are presented in Figure 4.
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Figure 4: Flowchart of the method: top panel - Learning phase; bottom panel – operational phase which consists in pigment retrieval and the determination of the α_{cb} block parameters.

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298 4 - METHODOLOGICAL RESULTS

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300 4-1 Statistical validation of the method

The validation of the method was focused on the retrieval of the fucoxanthin ratio, which is a 301 302 characteristic of diatoms, but the same procedure could be applied to any pigment. The hyper-303 parameter μ (see Annex) was optimized in order to retrieve that ratio, while η was set constant since 304 only the block were weighted in the present study. Due to the small amount of data in the DPIG, we estimated the accuracy of the fucoxanthin retrieval by a cross-validation procedure, which is a 305 powerful procedure in statistics. The principle is the following: we learned 30 2S-SOM using 30 306 different learning datasets L_i constituted of 90% of DPIG taken at random, and then computed 307 statistical estimator on the retrieved quantities using 30 test datasets (10% of DPIG). The algorithm 308 was as follows: 309

- 310 *i*=1 30
- 311 1. determination at random of a learning dataset L_i (90% of DPIG) and a test dataset TL_i (10% of 312 DPIG)
- 313 2. training of a 2S-SOM map M_i using L_i (see section 3.2 and 3.3).
- 314 3. Validation using TL_i according to the procedure described in section 3.4
- 315 4 Estimation of the $RMSE_i$ and R^2_i on TL_i between the estimated and observed fucoxanthin ratios
- 316 *end*

317 Computation of the mean RMSE and
$$R^2$$
 (R^2 , RMSE = $\frac{1}{30} \sum_{i=1}^{I=30} R^2 i$, RMSEi)

- 318
- The flowchart of the cross-validation procedure is presented in Figure 5.
- 320



323 Figure 5: Flowchart of the cross-validation procedure for 30 partitions of the DPIG database.

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Statistical parameters (R^2 coefficients, RMSE and P-values) of the cross validation between the DPIG in situ pigments and the pigments given by the 2S-SOM averaged for the 30 2S-SOM realizations, which are presented in table 2, show the good performance of the method.

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	R ²	RMSE (MG M ⁻³)	PVAL
CHLA SOM	0.84	0.22	0.001
DVCHLA	0.60	0.02	0.001
FUCO	0.87	0.02	0.001
PERID	0.81	0.01	0.001

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331

Table 2: Statistical parameters (R^2 coefficients, RMSE and P-values) of the cross validation between the DPIG in situ pigments and the pigments given by the 2S-SOM averaged for the 30 2S-SOM realizations.

337 4-2 Analysis of the topology of the 2S-SOM

As explained in sections 3-2 and 3-3, the referent vector components ($w_c \in R^{17}$), which are estimated 338 during the learning phase, are partitioned in four blocks B1, B2, B3 and B4. The hyper parameters μ 339 340 was tuned in order to favor the accuracy of the retrieval of the fucoxanthin ratio. We recall that all the pigment ratios are estimated during the calibration phase, but in the present paper attention was focused 341 342 fucoxanthin ratio selecting the on the when parameter μ. In Figure 6, we

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344 345

Figure 6: 2S-SOM Map. From left to right and top to bottom, values of the referent vectors for $\rho_w(490)$, Ra(490), SeaWiFS chl-a, and fucoxanthin, peridinin, divinyl Ratios. The number in each neuron indicates the amount of DPIG data captured at the end of the learning phase, the values indicated by the color bars are centered-reduced and non-dimensional values.

350

present six of the referent vector components of the 2S-SOM map. These components are $\rho_w(490)$,

352 Ra(490), SeaWiFS chl-a, and the ratios of fucoxanthin, which is a specific diatom pigment, and of

peridinin and divinyl. They exhibit a coherent topological order, the components having close values 353 being close together on the topological map. The remaining eleven components (not shown) exhibit 354 the same coherent topological order. One can observe a very good topological order for the fucoxanthin 355 356 ratio that was favored by the determination of the hyperparameter μ . Moreover, the bottom right region in the 2S-SOM map (Figure 6) may correspond to the diatoms with a good confidence since high 357 fucoxanthin is associated with high chlorophyll concentration and low peridinin. This is endorsed in 358 section 5 by looking at the geographical location of the different pigment concentrations (figures 8, 10, 359 11). Another important remark is that the value of each component presents a large range of variation 360 361



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Figure 7: 2S-SOM map. Weights (α_{cb}) of the four block parameters determined at the end of the learning phase; from left to right and top to bottom: ρ_w , Ra, Pigment, SeaWifs chl-a. The color bars show the % of the weight estimated by 2S-SOM, a value of 1 or 0 indicating that the data in the neuron are assembled with respect to that block only.

of the same order as the range of variation found in the DPIG variables. It means that the 2S-SOM
map has captured most of the variability of the dataset.

- Figure 6 shows a strong link between the values of the referent vectors for fucoxanthin and *chl-a* (high
- fucoxanthin and *chl-a* values, at the bottom right of the 2S-SOM) while fucoxanthin is high and *chl-a*
- 373 low for the referent vectors at the bottom left of the 2S-SOM. Additional information will be provided
- by the Ra(490) values when the fucoxanthin is less closely linked to the chlorophyll.
- Besides, for each neuron, the 2S-SOM provides a weight for each block (α_{cb}) and each variable (β_{cbj}). For a given neuron *c* the weights (α_{cb}) of the blocks are normalized, their sum being 1. A value of 1
- 377 for one block (and therefore a value of 0 for the other blocks) indicates that the data in the neuron are
- gathered with respect to that block only because there is too much noise in the variables in the other
 blocks. By examining the weights on the map, one can see which block most influences the link
 between the satellite measurements and the pigment ratios.
- In Figure 7, we present the α_{cb} values estimated during the learning phase of the 4 blocks (B1, B2, B3, 381 B4). For some neurons, only the blocks related to the reflectance and the reflectance ratio are used for 382 383 the definition of the neuron, while the weights for the two other blocks (pigments and *chl-a*) are null, indicating that for these neurons, in situ observations and SeaWiFS chl-a are more noisy than the 384 385 reflectance. These neurons correspond to very small *chl-a* concentrations, which are estimated with large error. Besides, we remark that high α values for *chl-a* corresponds to high *chl-a* concentration 386 values (bottom right of the *chl-a* panel in figure 7 and figure 6 respectively). For these cases, the 387 388 clustering assembled data that mainly depend on *chl-a* concentration.
- 389 390

391 5 - GEOPHYSICAL RESULT

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In the present study, we apply the 2S-SOM (section 3), which explicitly makes a weighted use of the data according to their specificity (ocean-color signals or in situ observations) to retrieve the fucoxanthin concentration from remote sensed data in the Senegalo-Mauritanian upwelling region where in situ measurements are lacking. According to the good results of the cross-validation method as shown in section 4.1, we expect that the 2S-SOM will provide pertinent results in a region which has been poorly surveyed.

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401 5-1 The pigment estimation from SeaWiFS observations in the Sénégalo-Mauritanian upwelling 402 region

We decoded the DSAT database (section 2-3) using the 2S-SOM for 11 years (1998-2009) of SeaWiFS 403 404 data observed in the Senegalo-Mauritanian upwelling region (8°N-24°N, 14°W-20°W). This study was done according to the retrieval phase described in section 3.4. For each day, we projected the 11 405 SeaWiFS observations (5 $\rho_w(\lambda)$, 5 $Ra(\lambda)$ and *chl-a*) of each pixel PX_m on the 2S-SOM. At the end of 406 407 the assignment phase, each pixel of a satellite image was associated with 6 pigment concentration 408 ratios. The underlying assumption is that the link between the remote sensing information and the pigment ratios of a pixel is this provided by the selected referent w_c . Thanks to the topological order 409 provided by the 2S-SOM, we expect that the best neurons chosen during the retrieval would give 410 accurate concentration ratios. In Figures 8, 10 and 11 we present the fucoxanthin concentration ratio 411 restitution for three different days and the associated SeaWiFS Chlorophyll images (1 and 6 January, 412 and 28 February 2003). Due to the limited size of the DPIG, the range of the ratio learned for the 413



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Figure 8: A) chl-a concentration, (B) fucoxanthin ratio, (C) aerosol optical thickness, (D) peridinin
for 1 January 2003. Panels (B) and (D) show that a second-order information was retrieved, which is
correlated with the chl-a concentration (A) but not equivalent. The aerosol optical thickness (C) does
not seem to contaminate the estimated parameters (fucoxanthin and peridinin ratios).

the fucoxanthin is between 0.3% and 20% with a mean of 10% and the *chl-a* content is between 0.5
mg m⁻³ and 3 mg m⁻³. The statistical estimator we used cannot extrapolate what has not been learned,



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426 Figure 9: SST for 2 January 2003. Note the well-marked upwelling (cold temperature) north of 13°N.





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Figure 10: (A) chl-a concentration, (B) fucoxanthin ratio, (C) aerosol optical thickness, (D) peridinin for 6
January 2003. Panels (B) and (D) show that a second-order information was retrieved, which is correlated

432 with the chl-a concentration (A) but is not equivalent. It is found that the aerosol optical thickness (C) does

433 not contaminate the estimated parameters (fucoxanthin and peridinin ratios).

- and for that raison we flagged the pixels in the SeaWiFS images that have a *chl-a* concentration greater
- 435 than 3. mg m⁻³.
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Figure 11: (A) chl-a concentration, (B) fucoxanthin ratio, (C) aerosol optical thickness, (D) Peridinin for 28 February 2003. Panels (B) and (D) show that a second order information was retrieved, which is correlated with the chl-a concentration (A) but is not equivalent. It is found that the aerosol optical thickness (C) does not contaminate the estimated parameters (fucoxanthin and peridinin ratios). The position of the NSB and OFB boxes are figured out by black square boxes.

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Regarding the images obtained for 1 January 2003 in the Senegalo-Mauritanian region (Fig 8A, B, C, D), we observe that the *chl-a* (Fig 8A) is very high at the coast and decreases offshore in accordance with the upwelling intensity as shown in the SST image (Fig 9). Moreover, we observed a persistent well-marked *chl-a* pattern south of the Cap Vert peninsula in form of a "W", which is the signature of a baroclinic Rossby wave (*Sirven et al*, 2019).

- 450 Except in the southern part of the region, the AOT (Aerosol Optical Thickness) is low, which means
- that the atmospheric correction of the reflectance is quite small, which gives confidence in the ocean-
- 452 color data products. The fucoxanthin concentration is maximum at the coast and decreases offshore as
- does the *chl-a* concentration, in agreement with the works of *Uitz et al.*, (2006, 2010). Fucoxanthin
- 454 presents coherent spatial patterns. Peridinin concentration is somewhat complementary to that of

fucoxanthin, with the low fucoxanthin concentration area corresponding to high peridinin
concentration area (northern part of Figs 8B, D). This behavior is also observed in Figure 10 (6 January
2003) and in Figure 11 (28 February, 2003) endorsing the analysis shown in Figure 8.

For 28 February, we selected two square box regions (Fig. 11), one near the coast (NSB, long [-20°, -18°], lat [12°,14°]) and the other about 800 km offshore (OFB, long [-28°, -26°], lat [12°,14°]). NSB waters correspond to upwelling waters while OFB waters correspond to oligotrophic waters. We projected the eleven ocean color parameters of the NSB and OFB pixels on the 2S-SOM map.





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Figure 12: *Reflectance spectra (in blue) captured the 28 February by six neurons whose referent vector spectra are in yellow: top line, for pixels in the NSB region (long. [-20°, -18°], lat. [12°, 14°]); bottom line, for pixels in the OFB region (long. [-28°, -26°], lat. [12°, 14°]).*

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Figure 12 presents the reflectance spectra (in blue) captured by three neurons of the 2S-SOM corresponding to pixels located in the NSB region (*top line*) and those captured by three neurons corresponding to pixels located in the OFB region (*bottom line*). The reflectance spectra of the associated referent vectors w are in yellow. The satellite reflectance spectra match the referent vector spectra; moreover the fucoxanthin ratio varies inversely with the mean value of the spectrum: the higher the fucoxanthin ratio, the smaller the mean value of the spectrum. The pigment concentration is greater near the coast.

We note a strong difference between the shape and the intensity of the near-shore (NSB) and offshore (OFB) spectra. The OFB spectra present mean values higher than those of the NSB spectra. This is due to the fact that NSB spectra were observed in a region where diatoms are abundant, as shown by

the high value of fucoxanthin concentration in this region (Figs 8, 10, and 11), which is a proxy for 482 diatoms along with higher *chl-a* concentration. In Figure 12, we note the lower values of the coastal 483 spectra at 443 nm, which can be interpreted as a predominant effect of spectral absorption by 484 phytoplankton pigments and CDOM. The different spectra are close together in the OFB region and 485 more disperse in the NSB region. This can be explained by the fact that the OFB region corresponds 486 to Case-1 waters while the NSB region waters are close to Case-2 waters and are influenced by the 487 488 variability of near shore process like turbidity or presence of dissolved matters, and dynamical 489 instabilities.





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Figure 13: Box plot of the weights of the selected neurons during the decoding of the 28 February
data. From left to right, weights of blocks B1, B2, B3, B4. Top panel, in the NSB region (long. [-20°,
-18°], lat. [12°, 14°]); bottom panel, in the OFB region (long. [-28°, -26°], lat. [12°, 14°]).

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We analyzed the weights of the blocks for the neurons selected in the analysis of the costal (NSB) and offshore (OFB) boxes. Figure 13 presents the box plot of the weight α_{cb} corresponding to the neurons belonging to the four blocks (B1, B2, B3, B4), with the constrain that the sum of the weights of a neuron is 1; a weight α larger than 0.25 indicates the predominance of a block in the learning for the classification (see section 3.5). It is clear that the weights for pixels near the coast (Fig 13, top panel) are different from those for offshore pixels (Fig. 13, bottom panel). As already mentioned in section 4.3 and also shown in Figure 7, the weights of the 2S-SOM play a significant role in the 2S-SOM

topology and consequently in the pigment retrieval. The weights of blocks B1 and B4 that take into 505 506 account the influence of the pigment ratios and the chlorophyll content in the retrieval are very low for 507 the offshore (OFB) oligotrophic region and more important for the coastal (NSB) region. The weights 508 of the blocks B2 and B3, which take into account the influence of the reflectance $(\rho_w(\lambda), Ra(\lambda))$, 509 dominate for the offshore regions. In coastal waters, the weights of all the blocks are used, with a smaller influence of B3, which is associated with R_a . This gives information on the role played by the 510 511 different variables on the classification in waters having different phytoplankton concentration and 512 composition. Besides it shows the automatic adaptation of the 2S-SOM to the environment in order to 513 optimize the clustering efficiency with respect to a classical SOM.

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Figure 14: Monthly fucoxanthin concentration averaged for an 11- years (1998-2009) for December
(A), March (B) and May (C).

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519 In order to study the seasonal variability of the fucoxanthin concentration with some statistical 520 confidence in the Senegalo-Mauritanian upwelling region, we constructed a monthly climatology for 521 an 11-year period (1998–2009) of the SeaWiFS observations by summing the daily pixels of the month 522 under study. The resulting climatology is presented in Figure 14 for December (Fig. 14a), March (Fig. 523 14b), and May (Fig 14c), which correspond to the most productive period (Fig. 14c). The fucoxanthin concentration, and consequently the associated diatoms, presents a well-marked seasonality. 524 525 Fucoxanthin starts to develop in December North of 19°N, presents its maximum intensity in March 526 when the upwelling intensity is maximum, extends up to the coast of Guinea (12°N) in April and 527 begins to decrease in May where it is observed north of Cabo Verde peninsula (15°N) in agreement 528 with the observations reported by Farikou et al, (2015) and Demarcq and Faure, (2000).

Figure 15 shows the fucoxanthin (in green) and the *chl-a* (in blue) concentrations computed from satellite observations for an 11-year period of SeaWiFS observations in the NSB region. There is a good correlation in phase between these two variables but not in amplitude (a good coincidence of peak occurrence but weak correlation in peak amplitude) showing that the relationship between





Figure 15: *chl-a (in blue) and fucoxanthin (in green) concentrations for near-shore pixels (in the NSB region).*

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fucoxanthin and *chl-a* is complex as mentioned by *Uitz et al*, (2006). In particular, there is a weak peak
in fucoxanthin in October 2001, which is not correlated with a *chl-a* peak.

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542 5-2 Analysis of the UPSEN campaigns

Figure 16 shows, for every UPSEN stations 1, 2, 3, 5a and 5b (see figure 1 for their geographical 543 544 position), the averaged in-situ UPSEN spectrum (in blue), the referent spectrum (in red) of the 2S-SOM neuron captured by the collocated satellite VIIRS sensor observations. The referent spectrum 545 546 is the mean of the different spectra captured by that neuron during the learning phase. Among these different spectra, there is one (black curve in figure 16) which is the closest to the UPSEN 547 548 spectrum. Obviously, the black curve is closer to the blue curve than the red one which is flatten 549 due to the averaging process. These three spectra are close together showing the good functioning of the 2S-SOM. 550





Figure 16: For ship stations 1, 2, 3, 5a and 5b, we show the averaged spectrum of the in situ spectra of the UPSEN station in blue; the spectrum of the referent vector (in red) of the 2S-SOM neuron, which has captured the closest satellite observations to the UPSEN station; among the different spectra constituting the referent spectrum, the spectrum of the learning database (DGIP) that is the closest to the averaged satellite spectra is shown in black. In the rectangular cartoons, we show the position of the UPSEN station, the number of the neuron of the 2S-SOM which has captured the satellite observation, the Rfuco of the referent vector, the Rfuco_{DGIP} of the closest DGIP and the in situ Rfucoupsen.

Their shapes are close to these observed in the NSB region (Figure 12) but their intensity is lower 565 meaning that their waters are more absorbing than the NSB waters due to a higher pigment 566 concentration. In fact, the UPSEN stations were located close to the coast (figure 1) in the Hann bight 567 568 south off the Cap Verde peninsula, which is very rich in phytoplankton pigments. In table 3, we present the fucoxanthin ratios associated with the referent vectors (Rfuco2S-SOM), the closest DPIG fucoxanthin-569 570 ratios captured by the neuron of the referents and the fucoxanthin-ratios measured during the UPSEN campaign. We note that the fucoxanthin ratios of the in-situ measurements are in the range of the DPIG 571 (see table 1), which allows a good functioning of the 2S-SOM estimator. The pigment ratios obtained 572 573 from ocean-color observations through the 2S-SOM are close to pigment concentrations measured at 574 the ship stations, which confirms the validity of the method we have developed. We remark that the 575 best 2S-SOM estimate of fucoxanthin ratio with respect to the UPSEN in-situ measurement is given 576 at station 5b which is the farthest off the coast. These results endorse the climatological study of the 577 Senegalo-Mauritanian upwelling region we have done with the 2S-SOM (section 5.1).

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UPSEN STATION	REFERENT N°	RFUCO	RFUCO	RFUCO
		2S-SOM	DPIG	UPSEN
STAT 1 17.3E 14.5 N	126	0.213	0.236	0.378
STAT 2 17.2E 14.4 N	126	0.213	0.236	0.391
STAT 2 17.2E 14.5 N	126	0.213	0.236	0.436
STAT 5A 17.5E 14.5 N	126	0.213	0.171	0.299
STAT 5B 17.5E 14.5 N	143	0.242	0.258	0.295

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Table 3: For ship stations 1, 2, 3, 5a and 5b of the UPSEN campaigns, we show the referent captured by the VIIRS observations, the fucoxanthin-ratio associated with this referent (Rfuco-2S-SOM), the fucoxanthin-ratio of the closest DPIG fucoxanthin-ratio captured by the neuron of the referent and the fucoxanthin-ratio measured in situ during the UPSEN campaign

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The 2S-SOM method gives pigment concentrations that are close to those obtained by in situ observations. The method could be applied to a large variety of other parameters in the context of studying and managing the planet Earth. The major constraint to obtaining accurate results is to deal with a learning data set that statistically reflects all the situations encountered in the observations processed. Due to its construction, the method cannot be used to find values beyond the range of the learning data set.

597 **6 - DISCUSSION**

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599 Machine learning methods are powerful methods to invert satellite signals as soon as we have adequate 600 database to support the calibration. Several technics have been used for retrieving biological 601 information from ocean color satellite observations. First, studies employed multilayer perceptrons 602 (MLP), which are a class of neural networks suitable to model transfer function (*Thiria* et al, 1993). 603 Gross et al, (2000, 2004) retrieved chl-a concentration from SeaWiFS, Bricaud et al, (2006) modeled 604 the absorption spectrum with MLP, Raitsos et al, 2008 and Palacz et al, 2013 introduced additional 605 environmental variables in their MLPs such as SST in the retrieval of PSC/PFT from SeaWiFS, which 606 improved the skill of the inversion. Another suitable procedure was to embed NN in a variational inversion, which is a very efficient way when a direct model exists (Jamet et al, 2005; Brajard et al, 607 608 2006a,b; Badran et al, 2008). Statistical analysis of absorption spectra of phytoplankton and of pigment 609 concentrations were conducted by *Chazottes* et al, (2006, 2007), by using a SOM.

In the present study, due to the fact that the learning dataset was quite small (515 elements), we used an unsupervised neural network classification method, which is an extension of the SOM method well adapted to dealing with a small database whose elements are very inhomogeneous. We clustered available satellite ocean-color reflectance at five wavelengths and their derived products, such as chlorophyll concentration, and the associated in situ pigment ratios.

615 The major points of this study are as follows:

- The clustering was carried out by developing a new neural classifier, the so-called 2S-SOM, which 616 presents several advantages with respect to the classical SOM. As in the SOM, we defined clusters 617 618 that assemble vectors, which are close together in terms of a specified distance. This classifier was 619 learned from a worldwide database (DPIG) whose vectors are ocean-color parameters observed by 620 satellite multi-spectral sensors and associated pigment concentrations measured in situ. In the 621 operational phase, SeaWiFS images are decoded, allowing the estimation of the pigment concentration ratios. The major advantage of 2S-SOM with respect to the classical SOM is to cluster 622 623 variables having similar physical significance in blocks having specific weights. The weights attributed to the four blocks are computed during the learning phase and vary with the quality of the 624 625 variables and with respect to their location on the ocean (near the coast or offshore). This permits to 626 modulate the variable influence in the cost function, which makes the clustering more informative 627 than that provided by the SOM. The block decomposition provides useful scientific information. For 628 offshore, the weight analysis allowed us to show that more influence is given to the reflectance ratios 629 $Ra(\lambda)$ and less to the *chl-a* and pigment concentrations; on the contrary near the coast the weights

indicate a more active use of the pigment composition and the *chl-a* concentration. Therefore, the
 resulting 2S-SOM clustering therefore at best takes into account the information that belongs to the
 specific water content.

633 - The 2S-SOM decomposes the DPIG into a large number of significant ocean-color classes allowing 634 reproduction of the different possible situations encountered in the dataset we analyze. Besides, we assume that the relationship between the pigment concentration and the remote sensed ocean-color 635 636 observations is independent on the location, which is justifiable since the relationship depends on the 637 optical properties of ocean waters through well-defined physical laws which are region-independent. 638 This also endorses the fact that we used a global database to retrieve pigments in a definite region. 639 On the contrary, the different phytoplankton species vary from one region to another making the 640 relationship between pigment ratio and phytoplankton species strongly depending on the region. This 641 justifies the fact we focused our study on the pigment retrieval rather than on the PSC or PFT, as 642 mentioned above. Moreover, most of the recent phytoplankton in situ identifications have been made 643 using pigment measurements with the HPLC method (*Hirata et al*, 2011). It is therefore more natural 644 to retrieve the pigment concentrations, which is the quantity we measured, than the associated PSC 645 or PFT, which are estimated from the pigment observations through complex non-linear and region-646 dependent algorithms (Uitz et al, 2006). Due to the characteristics of the DPIG, the method can 647 retrieve pigment concentration patterns over a large range $(0.02 - 2 \text{ mg m}^{-3})$.

- We were able to analyze the pigment concentration in the Senegalo-Mauritanian region by processing 648 satellite ocean color observations with the 2S-SOM. We found an important seasonal signal of 649 650 fucoxanthin concentration with a maximum occurring in March. We evidenced a large offshore 651 gradient of fucoxanthin concentrations, the near shore waters being richer than the offshore ones. We 652 showed that the offshore region waters correspond to Case-1 waters, while the near shore waters are 653 close to Case-2 waters and are influenced by the variability of near shore process like turbidity, or 654 the presence of dissolved matters. The UPSEN measurements show that the pigment ratios of the 655 Senegalo-Mauritanian region are in the range of the DPIG database used to calibrate the method, which justifies the use of the 2S-SOM algorithm to investigate this region. 656

We used daily satellite observations to construct a monthly climatology of pigment concentrations
of the Senegalo-Mauritanian upwelling region, which has been poorly surveyed by oceanic cruises.
Due to the highly non-linear character of the algorithms for determining the pigment concentrations
from satellite measurements, it is mathematically more rigorous to apply these algorithms to daily
satellite data and to average this daily estimate for the climatology period under study, than to
estimate them from the satellite data climatology, as many authors have done (*Uitz et al., 2010*; *Hirata et al., 2011*). We found that Fucoxanthin starts developing in December North of 19°N,

664 presents its maximum intensity in March when the upwelling intensity is maximum, extends up to 665 the coast of Guinea (12°N) in April and begins to decrease in May

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Another important aspect of our study concerns the validity of our results. The 2S-SOM method has
 been validated by focusing the retrieval accuracy on the fucoxanthin ratio, by using a cross-validation
 procedure. These results were qualitatively confirmed by two other independent studies.

- We first applied a cross validation procedure (see section 4.1), which is powerful technique for
 validating models (*Kohavi*, 1995; *Varma* and *Simon*, 2006). We learned 30 different 2S-SOM using
 30 different learning dataset determined at random from the DPIG dataset (each learning dataset
 representing 90% of DPIG) and 30 test datasets (10% of DPIG). By averaging the results, we found
 that the 2S-SOM method retrieves the fucoxanthin concentration with a good score (see the
 statistical parameters in table 2) which confirms the pertinence of the method.
- 676 - We then found that our fucoxanthin climatology is in agreement with in situ observations of phytoplankton reported in Blasco et al. (1980) in March to May 1974 off the coast of Senegal during 677 678 the JOINT I experiment. These authors analyzed 740 water samples collected with Niskin bottles at 136 stations extending along a line at 21°40'N (in the northern part of the studied region) from 0 679 680 to 100 km offshore. The samples were taken at several depths (mostly at 100, 50, 30, 15, 5 m). 681 Phytoplankton cells were counted and identified by the Utermohl inverted microscope technique (Blasco, 1977). These authors found that diatoms reach their maximum concentration in April–May 682 683 and are the most abundant group in that period, whereas the other cells predominate in March. Similar microscope observations have been reported in the ocean area south of Dakar by A. Dia 684 (1985) during several ship surveys in February–March 1982–1983. 685
- Our method is also in agreement with the monthly eleven years climatology presented in *Farikou et al*, (2015) who used a modified PHYSAT method to retrieve the *PFT* in the Senegalo-Mauritanian region.
- The pigment concentrations provided by the 2S-SOM from the VIIRS sensor observations are in
 qualitative agreement with the in-situ measurements done at five stations during the two UPSEN
 campaigns in 2012 and 2013, showing that the method is able to function in waters where the
 pigment concentrations are quite high (fucoxanthin ratios of the order 0.4).
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698 **7 - CONCLUSION**

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700 We developed a new neural network clustering method, the so-called 2S-SOM algorithm to retrieve 701 phytoplankton pigment concentration from satellite ocean color multi spectral sensors. The 2S-SOM 702 algorithm is a SOM specifically designed to deal with a large number of heterogeneous components 703 such as optical and chemical measurements. The major advantage of 2S-SOM with respect to the 704 classical SOM is to cluster variables having similar significance in blocks having specific weights. 705 The weights attributed to the blocks during the learning phase vary with the quality of the variables in 706 the classification. This permits to modulate the variable influence in the cost function, which makes 707 the clustering more informative than that provided by the SOM. Besides, the block weighting provides 708 useful information on the functioning of the classification by permitting to identify the variables which 709 control it. It also allows us to better understand the dynamics of the phytoplankton communities.

710 The 2S-SOM method is efficient and rapid as soon as the calibration is done, since it uses elementary 711 algebraic operations only. The 2S-SOM method is like a piecewise regression that takes advantage of 712 the unsupervised classification of the SOM. We decomposed the DPIG database into quite a large number of partitions (9x8=162) when comparing our study to other studies (*Uitz et al*, 2006, 2012). 713 714 The validity of the method has been controlled through a cross validation procedure and confirmed by 715 three qualitative studies. Statistical parameters (R^2 coefficients, RMSE and P-values) of the cross-716 validation between the DPIG in situ pigments and the pigments given by the 2S-SOM averaged for the 717 30 2S-SOM realizations presented in table 2, show the good performance of the method. It must be noticed that the performance mainly depends on the size of the learning set used to calibrate the 2S-718 719 SOM. This set must include all the situations encountered in the pigment retrieval. The larger the 720 learning set, the better the method performs. Due to its generic character and its flexibility, the method 721 could be used to determine a large variety of measures done with satellite remote sensing 722 observations.

723 In this work, the method was applied to study the seasonal variability of the fucoxanthin concentration in Senegalo-Mauritanian upwelling region. We showed a large offshore gradient of fucoxanthin, the 724 725 higher concentration being situated near the shore. We were able to construct a monthly climatology for an 11-year period (1998–2009) of the SeaWiFS observations by summing the daily pixels of the 726 727 month under study in a region which was poorly surveyed by oceanic cruises. The fucoxanthin 728 concentration, and consequently the associated diatoms, present a well-marked seasonality (Figure 10). 729 Fucoxanthin starts developing in December North of 19°N, presents its maximum intensity in March 730 when the upwelling intensity is maximum, extends up to the coast of Guinea (12°N) in April and 731 begins to decrease in May where it is observed north of Cabo Verde peninsula (15°N), in agreement

- with the observations reported by *Farikou et al*, (2015) and *Demarcq and Faure*, (2000). The UPSEN
 campaign results endorse the validity of the study of the Senegalo-Mauritanian upwelling region done
 with the 2S-SOM.
- 735

736 Acknowledgments

The study was supported by the projects CNES-TOSCA 2013-2014 and 2014-2015. The water-leaving 737 reflectances were obtained from the SeaWiFS daily reflectances, $\rho obsTOAw(\lambda)$, provided by the 738 NASA/GSFC/DAAC observed at the top of the atmosphere (TOA) and processed with the SOM-NV 739 algorithm (Diouf et al., 2013) from 1998 to 2010. They are available at the web site: 740 http://poacc.locean-ipsl.upmc.fr/. The DPIG data base was kindly provided by Dr. S. Alvain. We thank 741 Dr. Alban Lazar and Dr. E. Machu for providing in situ data measured during the UPSEN experiments 742 as well as stimulating discussions for their interpretation. We also thank Ray Griffiths for editing the 743 744 manuscript.

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747 **References**

- Alvain S, Moulin C., Dandonneau Y. and Breon F. M. : Remote sensing of phytoplankton groups in
 case-1 waters from global SeaWiFS imagery. Deep-Sea Res. Part1, V 52 (11), pp 1989-2004,
 2005.
- Alvain, S. Loisel H. and Dessailly D. : Theoretical analysis of ocean color radiances anomalies and
 implications for phytoplankton group detection. Optics Express, V 20 (2), 2012.
- Antoine D., André J. M., Morel A. : Oceanic primary production : Estimation at global scale from
 satellite (Coastal Zone Color Scanner) chlorophyll. Global Biogeochem Cy. V 10, pp 57-69, 1996.
- Badran F., Berrada M., Brajard J., Crepon M., Sorror C., Thiria S., Hermand J.P., Meyer M.,
- Perichon L., Asch M. : Inversion of satellite ocean colour imagery and geoacoustic
- characterization of seabed properties : Variational data inversion using a semi-automatic adjoint

approach J. Marine Systems, V 69, pp 126-136, 2008

- Behrenfeld M. J., Boss E., Siegel D.A., Shea D.M. : Carbon-based ocean productivity and
 phytoplankton physiology from space. Global Biogeochem. Cy. V 19, GB1006,
 doi:10.1029/2004GB002299, 2005
- Behrenfeld M. J., and Falkowski P.G. : Photosynthetic rates derived from satellite base chlorophyll
 concentration. Limnol. Oceanogr, V 42, pp 1-20, 1997
- Ben Mustapha Z. S., Alvain S., Jamet C., Loisel H. and Desailly D.: Automatic water leaving radiance
 anomalies from global SeaWiFS imagery: application to the detection of phytoplankton groups in
 open waters. Remote Sens. Environ., vol 146, pp 97-112, 2014.
- Blasco D. : Red tide in the upwelling region of Baja California. Limnol. Oceanogr. vol 22, pp 255263, 1977
- Blasco D., Estrada M. and Jones B. : Relationship between the phytoplankton distribution and
 composition and the hydrography in the northwest African upwelling region, near Cabo Corbeiro.
 Deep-Sea Res. , vol 27A, pp 799-821, 1980.
- 772 Bracher A., Bouman HA, Brewin RJW, Bricaud A, Brotas V, Ciotti AM, Clementson L, Devred E,
- Di Cicco A, Dutkiewicz S, Hardman-Mountford NJ, Hickman AE, Hieronymi M, Hirata T, Losa
- SN, Mouw CB, Organelli E, Raitsos DE, Uitz J, Vogt M and Wolanin A : Obtaining
- Phytoplankton Diversity from Ocean Color: A Scientific Roadmap for Future Development.
- Front. Mar. Sci. 4:55. doi: 10.3389/fmars.2017.00055, 2017
- 777 Brajard J., Jamet C., Moulin C. and Thiria S. : Atmospheric correction and oceanic constituents
- retrieval with a neuro-variational method. Neural Networks, Vol 19(2), p178-185, 2006

- Brajard J., Jamet C., Moulin C. and Thiria S : Neurovariational inversion of ocean color images. J.
 Atmos. Space Res. Vol 38, n 2, pp 2169-2175, 2006
- 781 Brewin R. J. W., Hardman-Mountford N. J., Lavender S. J., Raitsos D. E., Hirata T., Uitz J., et al. :
- An inter-comparison of bio-optical techniques for detecting dominant phytoplankton size class
- from satellite remote sensing. Remote Sens. Environ. 115, 325–339. doi:
- 784 10.1016/j.rse.2010.09.004, 2011
- Brewin R. J. W., Sathyendranath S., Hirata, T., Lavender, S.J., Barciela, R., Hardman-Montford, N.J.
 A three-component model of phytoplankton size class for the Atlantic Ocean. Ecol. Model. vol 22, pp 1472-1483, 2010.
- Bricaud A., Mejia C., Blondeau Patissier D., Claustre H., Crepon M. and Thiria S. : Retrieval of
 pigment concentrations and size structure of algal populations from absorption spectra using
 multilayered perceptrons. Applied Optics Mars 2007 vol 46 n°8., 2006
- Capet X., Estrade, P., Machu, E., Ndoye, S. et al. : On the Dynamics of the Southern Senegal
 Upwelling Center: Observed Variability from Synoptic to Superinertial Scales : J. Phys. Oceanogr.
- 793 vol **47** (1), pp 155-180, 2017
- Cavazos T. : Using Self-Organizing Maps to Investigate Extreme Climate Events: An Application to
 Wintertime Precipitation in the Balkans. J. Climate, vol 13, 1718–1732, 2000.
- Chazotte A., Crepon M., Bricaud A., Ras J. and Thiria S. : Statistical analysis of absorption spectra
 of phytoplankton and of pigment concentrations observed during three POMME cruises using a
 neural network clustering method. Applied Optics, 46 (18), 3790-3799, 2007
- Chazottes A., Bricaud A., Crepon M. and Thiria S. : Statistical analysis of a data base of absorption
 spectra of phytoplankton and pigment concentrations using self-organizing maps. Appl. Opt. 45,
 801 8102-8115, 2006
- 802 Ciotti A. and Bricaud A. : Retrievals of a size parameter for phytoplankton and spectral light absorption
- by colored detrital matter from water-leaving radiances at SeaWiFS channels in a continental shelf
 region off Brazil. Limnol. Oceangr. Methods, vol 4, pp 237-253, 2006.
- Demarcq H. and Faure V. : Coastal upwelling and associated retention indices from satellite SST.
 Application to Octopus vulgaris recruitment. Oceanografica Acta, vol 23, pp 391-407, 2000.
- 807 Dia A. Biomasse et biologie du phytoplancton le long de la petite côte sénégalaise et relations avec
- l'hydrologie. Rapport interne N°44 du CRODT, Réf: 0C000798, 1981-1982. On line on the web
 site:http://www.sist.sn/gsdl/collect/publi/index/assoc/HASH2127.dir/doc.pdf
- 810 Diouf D., Niang A., Brajard J., Crepon M. and Thiria S. : Retrieving aerosol characteristics and sea-
- surface chlorophyll from satellite ocean color multi-spectral sensors using a neural-variational
- 812 method. Remote Sens. Environ. vol 130, pp 74-86, 2013.

- 813 Farikou O., Sawadogo S., Niang A., Brajard J., Mejia C., Crépon M. and Thiria S. : Multivariate
- analysis of the Sénégalo-Mauritanian area by merging satellite remote sensing ocean color and SST
 observations. J. Environ. Earth Sci. vol 5 (12), pp 756-768, 2013
- 816 Farikou O., Sawadogo S., Niang A., Diouf D., Brajard J., Mejia C., Dandonneau Y., Gasc G., Crepon
- 817 M., and Thiria S. : Inferring the seasonal evolution of phytoplankton groups in the Sénégalo-
- 818 Mauritanian upwelling region from satellite ocean-color spectral measurements, J. Geophys. Res.
- 819 Oceans, vol **120**, pp 6581-6601, 2015.
- Friedrich T. and Oschlies A. : Basin-scale pCO2 maps estimated from ARGO float data : A model
 study, J. Geophys. Res., vol 114, C10012, doi: 10. 1029/2009JC005322, 2009.
- B22 Gordon H. R. : Atmospheric correction of ocean color imagery in the Earth Observing System era. J.
- 823 Geophys. Re. Atmospheres, vol **102**(D14), pp 17081-17106, 1997.
- Hewitson B.C. and Crane R. G. : Sef organizing maps : application to synoptic climatology. Climate research, vol 22, pp 13-26, 2002
- 826 Gross L., Frouin R., Dupouy C., Andre J. M. and Thiria S. : Reducing biological variability in the
- retrieval of chlorophyl_a concentration from spectral marine reflectance. Applied Optics, Vol. 43
 Issue 20 pp. 4041, 2004
- 829 Gross L., Thiria S., Frouin R., Mitchell B.G : Artificial neural networks for modeling transfer
- function between marine reflectance and phytoplankton pigment concentration J. Geophys. Res.
 Vol 105,no.C2, pp3483-3949, february 15, 2000.
- Hirata T., Aiken J., Hardman-Mountford N., Smyth T. J. and Barlow R.G. : An absorption model to
- determine phytoplankton size classes from satellite ocean color, Remote Sens. Environ. vol 112, pp
 3153-3159, 2008.
- 835 Hirata T., Hardman-Mountford N.J., Brewin R.J.W., Aiken J., Barlow R., Suzuki K., Isada T., Howell
- 836 E., Hashioka T., Noguchi-Aita M. and Yamanaka Y. : Synoptic relationships between surface
- 837 chlorophyll-*a* and diagnostic pigments specific to phytoplankton functional types. Biogeosciences,
- 838 vol **8** (2): pp 311-327, 2011.
- Jamet C., Thiria S., Moullin C., Crepon M. : Use of a neural inversion for retrieving Oceanic and
 Atmospheric constituents for Ocean Color imagery : a feasability study.
- doi:10.1175/JTECH1688.1, J. Atmos. Ocean. Techno. :/ Vol. 22, No. 4, pp. 460–475, 2005
- 842 Jeffreys S.W. and Vesk M. : Phytoplankton Pigment in Oceanography : Guidelines to Modern
- 843 Methods, UNESCO, Paris, ed S. W. Jeffery, R.F.C. Mantoura and S. W. Wright, Introduction to
- marine phytoplankton and their pigment signatures, pp 33-84, 1997.
- Jouini M., Lévy M., Crépon M. and Thiria S. : Reconstruction of ocean color images under clouds using a neuronal classification method. Remote Sens. Environ. vol **131**, pp 232-246, 2013

- 847 Kohavi R. : A study of cross-validation and bootstrap for accuracy estimation and model selection.
- 848 Proceedings of the Fourteenth International Joint Conference on Artificial Intelligence. San Mateo,
- 849 CA: Morgan Kaufmann ed.. **2** (12): pp 1137–1143, 1995.
- 850 Kohonen T : Self-organizing maps (3rd ed.). Springer, Berlin Heidelberg New York. 2001
- Kruizinga S. and Murphy A : Use of an analogue procedure to formulate objective probabilistic
 temperature forecasts in the Netherlands. Mon. Wea. Rev., vol 111, pp 2244–2254, 1983.
- Le Quéré et al, (2018) : Global Carbon Budget 2018, Earth Syst. Sci. Data, 10, 2141–2194, 2018 ;
 https://doi.org/10.5194/essd-10-2141-2018
- Lévy M., D. Iovino, L. Resplandy, P. Klein, G. Madec, A.-M. Tréguier, S. Masson, K. Takahashi, Large-scale
 impacts of submesoscale dynamics on phytoplankton: Local and remote effects, Ocean Modelling, 77–93,
 2012
- Levy, M., Mesoscale variability of phytoplankton and of new production: Impact of the large-scale nutrient
 distribution, J. Geophys. Res., 108(C11), 3358, doi:10.1029/2002JC001577, 2003.
- Liu Y. and Weisberg R. H. : Patterns of ocean current variability on the West Florida Shelf using the self-organizing map, J. Geophys. Res., **110**, C06003, doi:10.1029/2004JC002786, 2005
- Liu Y., Weisberg R. H., and He R. : Sea surface temperature patterns on the West Florida Shelf using growing hierarchical self-organizing maps, J. Atmos. Oceanic Technol., vol **23**(2), pp 325–338, 2006
- Longhurst A. R., Sathyendranath S., Platt T., Caverhill C. : An estimation of global primary production
- in the ocean from satellite radiometer data. J. Plank. Res. vol 17, pp 1245-1271, 1995
- Lorenz E. N : Atmospheric predictability as revealed by naturally occurring analogs. J. Atmos. Sci.,
 vol 26, pp 639–646, 1969
- Morel A. and Gentili G. : Diffuse reflectance of oceanic waters. III. Implication of bidirectionality for
 the remote-sensing problem. Appl. Opt. vol 35, pp 4850-4862, 1996.
- 870 Mouw C. B. and Yoder J. A. : Optical determination of phytoplankton size composition from global
- 871 SeaWiFS imagery. J. Geophys. Res. vol 115, C12018, doi:10.1029/2010JC006337, 2010.
- Ndoye S., Capet X., Estrade P., Sow B., Dagorne D., Lazar A., Gaye A. and Brehmer P. : SST patterns
- and dynamics of the southern Senegal-Gambia upwelling center. J. Geophys. Res. Oceans, vol 119,
 pp 8315–8335. 2014
- Niang, A., Gross, L., Thiria, S., Badran, F., & Moulin, C. Automatic neural classification of ocean
- 876 colour reflectance spectra at the top of atmosphere with introduction of expert knowledge.
- 877 Remote Sens. Environ, vol 86, pp 257–271, 2003.
- Niang A., Badran F., Moulin C., Crépon M. and Thiria S. : Retrieval of aerosol type and optical
- thickness over the Mediterranean from SeaWiFS images using an automatic neural classification
- method. Remote Sens. Environ. vol 100, pp 82-94, 2006.

- 881 O'Reilly, J.E., Maritorena, S., Siegel, D. A., O'Brien, M. C., Toole, D., Mitchell, B. G., Kahru, M.,
- 882 Chavez, F. P., Strutton, P., Cota, G. F., Hooker, S. B., McClain, C. R., Carder, K. L., Muller-
- 883 Karger, F., Harding, L., Magnuson , A., Phinney, D., Moore, G.F., Aiken, J., Arrigo, K. R.,
- Letelier, R., and Culver, M. Ocean color chlorophyll a algorithms for SeaWiFS, OC2 and
- 885 OC4: Version 4. In S. B. Hooker, and E. R. Firestone (Eds), SeaWiFS postlaunch calibration and
- validation analyses: Part 3. NASA Tech. Memo. 2000-206892, vol. 11(pp.9-23). Greenbelt, MD:
- 887 NASA Goddard Space Flight Center. 2001.
- 888 Palacz A. P., St. John, M. A., Brewin, R. J.W., Hirata, T., and Gregg, W.W. : Distribution of
- phytoplankton functional types in high-nitrate low-chlorophyll waters in a new diagnostic
- ecological indicator model. Biogeosciences 10, 7553–7574. doi: 10.5194/bg-10-7553, 2013.
- Raitsos D. E., Lavender, S. J., Maravelias, C. D., Haralambous, J., Richardson, A. J., and Reid, P.
- C. : Identifying phytoplankton functional groups from space: an ecological approach. Limnol.
 Oceanogr. 53, 605–613. doi: 10.4319/lo.2008.53.2.0605, 2008
- Reusch D. B., Alley, R. B., and Hewitson, B. C : North Atlantic climate variability from a self-
- organizing map perspective, J. Geophys. Res., vol **112**, D02104, doi:10.1029/2006JD007460, 2007.
- Sathyendranath S., Watts S., L., Devred E., Platt T., Caverhill C. M., and Maass H. : Discrimination
 of diatom from other phytoplankton using ocean-colour data, Mar. Ecol. Prog. Ser., vol 272, pp 59–
 68, 2004.
- Sirven J., Mignot J., Crépon M. : Generation of Rossby waves off the Cap Verde Peninsula: the role
 of the coastline . Ocean Sci., 15, 1–24, 2019
- Sosik, H.M.; Sathyendranath, S.; Uitz, J.; Bouman, H.; Nair, A. In situ methods of measuring
 phytoplankton functional types. In Phytoplankton Functional Types from Space. IOCCG report, No.
 15; Sathyendranath, S., Ed.; IOCCG: Dartmouth, NS, Canada, pp. 21–38, 2014.
- 904 Uitz J., Claustre H., Morel A. and. Hooker S.B : Vertical distribution of phytoplankton communities
 905 in open ocean: an assessment based on surface chlorophyll. J. Geophys. Res. 111, C08005,
 906 doi:10:1029/2005JC003207. 2006
- 907 Uitz J., Claustre H., Gentili B. and Stramski D. : Phytoplankton class-specific primary production in
 908 the world's ocean: seasonal and interannual variability from satellite observations. Global
 909 Biogeochem. Cycles, vol 24, GB 3016, doi:10:1029/2009GB003680, 2010
- Van den Dool H. : Searching for analogs, how long must we wait? Tellus, vol **46A**, pp 314–324, 1994.
- 911 Varma, S., Simon, R. : Bias in error estimation when using cross-validation for model selection; BMC
- 912 Bioinformatics. vol 7. PMC 1397873 . PMID 16504092. doi:10.1186/1471-2105-7-91, 2006

- 913 Vidussi F., Claustre H., Manca B. B., Luchetta A. and Marty J. C. : Phytoplankton pigment distribution
- 914 in relation to upper thermocline circulation in the eastern Mediterranean sea during winter. J.
- 915 Geophys. Res., vol 106, pp 19,939-19,956, 2001.
- 916 Westberry T., Behrenfeld M.J., Siegel D. A. and Boss E.: Carbon-based productivity modeling with
- 917 vertically resolved photoacclimatation. Global Biogeochem. Cycles, vol 22, GB2024,
- 918 DOI:10.1029/2007GB003078, 2008
- 919 Zorita E. and Von Storch H. : The Analog Method as a Simple Statistical Downscaling Technique:
- 920 Comparison with More Complicated Methods. Journal of Climate, vol 12, pp 2474-2489, 1999.

922 ANNEX 1 923 A1 Cost function of the SOM 924 Let us recall the following notation: 925 $\boldsymbol{D} = \{\boldsymbol{z}_1, \cdots, \boldsymbol{z}_i, \cdots, \boldsymbol{z}_K\}$ the dataset composed of K vectors $\boldsymbol{z}_i \in \mathbb{R}^N$ 926 $W = \{w_1, \dots, w_c, \dots, w_c\}$ the set of weights $w_c \in \mathbb{R}^N$ where $C = p \times q$ is the size of the SOM. 927 The w_c of the SOM are estimated by minimizing a cost function of the form 928 929 $J_{SOM}^{T}(\chi, W) = \sum_{i=1}^{K} \sum_{c=1}^{p \times q} K^{T} \left(\delta(c, \chi(z_{i})) \right) ||z_{i} - w_{c}||^{2},$ 930 (A.1)where c indices the neurons of the SOM map, χ is the allocation function that assigns each element z_i 931 of **D** to its referent vector w_c which is of the form $\chi(\mathbf{z}_i) = \arg \min_c ||\mathbf{z}_i - \mathbf{w}_c||^2$, 932 $\delta(c, \chi(\mathbf{z}_i))$ is the discrete distance on the SOM between a neuron if index *c* and the neuron allocated 933 to observation \mathbf{z}_i , and K^T a kernel function parameterized by T that weights the discrete distance on 934 the map and decreases during the minimization process. T acts as a regularization term (Kohonen, 2001, 935 Niang et al, 2003). In the present case K^T is of the form : 936 $K^{T}(\delta) = (1/T)K(\delta/T)$, where K is the gaussian function of mean 0 and standard deviation 1. 937 The cost function (A.1) takes into account the proper inertia of the partition of the data set D and 938 939 ensures that its topology is preserved. 940 A2 Definition of the Algorithm 2S-SOM 941 The 2S-SOM algorithm is an extension of the Self-Organizing maps (SOM, Kohonen, 2001) based on 942 the K-mean method (Ouattara et al., 2014, https://www.theses.fr/179489704). It automatically 943 944 structures the variables having some common characters into conceptually meaningful and 945 homogeneous blocks during the learning phase. The 2S-SOM takes advantage of this structuration of 946 **D** and the variables into B different blocks, which permits an automatic weighting of the influence of each block and consequently of each variable in the classification phase. The 2S-SOM is based on a 947 modification of the cost function of the SOM algorithm. For a neuron of index c, we define the weights 948 α_{cb} of each block b (b = 1, ..., B) and the weights β_{cbj} of the variables j ($j = 1, ..., P_b$) in this block, 949

950 where P_b is the number of variable in the block indexed by b. The vectors of weighs are denoted

951 $\boldsymbol{\alpha} = \{\alpha_{cb}\}_{1 \le c \le C, 1 \le b \le B}$ and $\boldsymbol{\beta} = \{\beta_{cbj}\}_{1 \le c \le C, 1 \le b \le B, 1 \le j \le P_b}$

952 The new cost function is:

953
$$J_{2S-SOM}^{T}(\boldsymbol{\chi}, \boldsymbol{W}, \boldsymbol{\alpha}, \boldsymbol{\beta}) = \sum_{c} \left(\sum_{b=1}^{B} \left(\sum_{zi \in D} \alpha_{cb} K^{T} \left(\delta(c, \boldsymbol{\chi}(z_{i})) \right) d_{\beta_{cb}}(i) + J_{cb} \right) + I_{c} \right),$$
(A.2)

954 with

955
$$d_{\beta_{cb}}(i) = \sum_{j=1}^{P_b} \beta_{cbj} \, (z_{ib}^j - w_{ib}^j)^2, \tag{A.3}$$

where *c* indices the neurons of the 2S-SOM map.

957 under the two constraints:

958
$$\sum_{b=1}^{B} \alpha_{cb} = 1; \alpha_{cb} \in [0,1] \ \forall c, 1 \le c \le C$$
 (A.4)

959 and

960
$$\sum_{j=1}^{P_b} \beta_{cbj} = 1; \beta_{cbj} \in [0,1], \forall c, 1 \le c \le C; \forall b, 1 \le b \le B.$$

961 I_c and J_{cb} are used to regularize the weights α and β . They are defined as negative entropies weighted 962 by μ for the blocks and η for the variables of each block

963

964
$$I_c = \mu \sum_{b=1}^{P_b} \alpha_{cb} log(\alpha_{cb})$$
 (A.6)

965 and

966
$$J_{cb} = \eta \sum_{j=1}^{B} \beta_{cbj} log(\beta_{cbj})$$
(A.7)

The topological conservation properties of 2S-SOM are influenced by the weights α_{cb} and β_{cbj} in the classification through the hyper-parameters μ , η and the neighborhood parameter T.

The weights α_{cb} and β_{cbj} respectively indicate the relative importance of blocks and variables in the neurons. Thus, the greater the weight of a block *b* or a variable *j*, the more the block or the variable contributes to the definition of the class (or neuron) in the sense that it makes it possible to reduce the variability of the observations in the cell and in its close neighborhood. For a high value of η and a fixed one for μ , the β_{cbj} in a block are equal to $1/P_b$. In this case, only the blocks are modified according to their capacity to define the neurons. In this context, the 2S-SOM then makes possible to weight the different blocks for each neuron

- 976 For high values of μ , I_c is large. The minimization of J_{cb} forces all its coefficients to become 977 equal. For a fixed value of η , the α_{cb} associated with the blocks are all equal to 1/B. In this case, 978 only the β_{cbi} of the variables inside the blocks weight the neurons
- When μ and η tend to very large values, the blocks are equiprobable as well as the variables.
 Thus, the 2S-SOM algorithm is comparable to the SOM.
- 981

982 A3 How the 2S-SOM algorithm works:

- 983 For fixed μ and η , the learning of the 2S-SOM algorithm is as follows:
- 984 <u>Step 0:</u> Initialization with iteration of the algorithm SOM, by setting α and β to homogeneous 985 values.
- The optimization of J_{2S-SOM}^{T} is carried out through an iterative process composed of three steps (1, 2, and 3) presented below.
- 988 <u>Step 1:</u> The w_c referents, the weights α and β are known and fixed, the observations are assigned 989 to the neurons by respecting the assignment function:

990
$$c(zi) = \chi(z_i) = \arg\min_{c \in C} \left(\sum_{r \in C} K^T(\delta(r,c)) \left(\sum_{b=1}^B \alpha_{cb} d_{\beta_{cb}}(i) \right) \right)$$
(A.8)

- 991
- 992 <u>Step 2:</u> Updating the neuron centers (the *w_c* referents) according to the formula of the SOM
 993 algorithm.
- 994
- 995 <u>Step 3:</u> the assignment function and the referents w_c being fixed, α and β are determined 996 according to the equations (A.9, A.10, A.11, A.12), by minimizing the cost function 997 J_{2S-SOM}^T with respect to α and β under the constraints (A.4) and (A.5).

998
$$\alpha_{cb} = \frac{\exp\left(\frac{-\psi_{cb}}{\mu}\right)}{\sum_{b=1}^{B} \exp\left(\frac{-\psi_{cb}}{\mu}\right)}$$
(A.9)

999 with

1000
$$\psi_{cb} = \sum_{zi \in D} K^{T} (\delta(\chi(z_{i}), c)) d_{\beta_{cb}}(i)$$
(A.10)

1001

and

1002
$$\beta_{cbj} = \frac{\exp\left(\frac{-\Phi_{cbj}}{\eta}\right)}{\sum_{b=1}^{p_b} \exp\left(\frac{-\Phi_{cbj}}{\eta}\right)}$$
(A.11)

1003 with

1004
$$\Phi_{cbj} = \sum_{zi\in D} \alpha_{cb} K^T(\chi(z_i), c) (z_{ib}^j - w_{cb}^j)^2$$
(A.12)

1005

1006 This algorithm is repeated by sampling the hyper-parameters μ and η until convergence.

Finally, at the convergence, the 2S-SOM provides on the one hand a topological map allowing to visualize the data, and on the other hand a weight system for the neurons of the map allowing us to interpret the role of the different variables and to choose those that are the most significant for the classification and to neutralize those which are the least significant.

1012 1013 1014 1015	FIGURE CAPTION
1016 1017 1018 1019	Figure 1 : Mauritania and Senegal coastal topography. The land is in brown and the ocean depth is represented in meters by the color scale on the right side of the figure. The UPSEN stations are shown at the bottom left cartoon of the figure.
1020 1021 1022	Figure 2 : Geographic positions of the 515 in situ and satellite collocated measurements of the DPIG database.
1023 1024 1025 1026	Figure 3: Dispersion diagram of DPIG chl-a computed from the SeaWiFS observations using the OC4V4 algorithm versus in situ chl-a. The coefficient of vraisemblance R^2 and the RMSE (Root Mean Square Error) were computed in mg m ⁻³
1027 1028 1029	Figure 4: Flowchart of the method: top panel - Learning phase; bottom panel – operational phase which consists in pigment retrieval and the determination of the α_{cb} block parameters.
1030 1031	Figure 5 : Flowchart of the cross-validation procedure for 30 partitions of the DPIG database.
1032 1033 1034 1035 1036	Figure 6 : 2S-SOM Map. From left to right and top to bottom, values of the referent vectors for $\rho_w(490)$, $Ra(490)$, SeaWiFS chl-a, and fucoxanthin, peridinin, divinyl Ratios. The number in each neuron indicates the amount of DPIG data captured at the end of the learning phase, the values indicated by the color bars are centered-reduced and non-dimensional values.
1037 1038 1039 1040 1041	Figure 7: 2S-SOM map. Weights (α_{cb}) of the four block parameters determined at the end of the learning phase; from left to right and top to bottom: ρ_w , Ra, Pigment, SeaWifs chl-a. The color bars show the % of the weight estimated by 2S-SOM, a value of 1 or 0 indicating that the data in the neuron are assembled with respect to that block only.
1042 1043 1044 1045 1046	Figure 8 : A) chl-a concentration, (B) fucoxanthin ratio, (C) aerosol optical thickness, (D) peridinin for 1 January 2003. Panels (B) and (D) show that a second-order information was retrieved, which is correlated with the chl-a concentration (A) but not equivalent. The aerosol optical thickness (C) does not seem to contaminate the estimated parameters (fucoxanthin and peridinin ratios).
1047 1048	Figure 9 : SST for 2 January 2003. Note the well-marked upwelling (cold temperature) north of 13°N.
1049 1050 1051 1052	Figure 10 : (A) chl-a concentration, (B) fucoxanthin ratio, (C) aerosol optical thickness, (D) peridinin for 6 January 2003. Panels (B) and (D) show that a second-order information was retrieved, which is correlated with the chl-a concentration (A) but is not equivalent. It is found that the aerosol optical thickness (C) does not contaminate the estimated parameters (fucoxanthin and peridinin ratios).

1054

1055

retrieved, which is correlated with the chl-a concentration (A) but is not equivalent. It is found that 1056 1057 the aerosol optical thickness (C) does not contaminate the estimated parameters (fucoxanthin and peridinin ratios). The position of the NSB and OFB boxes are figured out by black square boxes. 1058 1059 1060 Figure 12 : Reflectance spectra (in blue) captured the 28 February by six neurons whose referent vector spectra are in yellow: top line, for pixels in the NSB region (long. [-20°, -18°], lat. [12°, 1061 14°]; bottom line, for pixels in the OFB region (long. [-28°, -26°], lat. [12°, 14°]). 1062 1063 Box plot of the weights of the selected neurons during the decoding of the 28 February 1064 Figure 13 1065 data. From left to right, weights of blocks B1, B2, B3, B4. Top panel, in the NSB region (long. [-20°, -18°], lat. [12°, 14°]); bottom panel, in the OFB region (long. [-28°, -26°], lat. [12°, 14°]). 1066 1067 Figure 14 : Monthly fucoxanthin concentration averaged for an 11- years (1998-2009) for December 1068 (A), March (B) and May (C). 1069 1070 1071 Figure 15 : . chl-a (in blue) and fucoxanthin (in green) concentrations for near-shore pixels (in the NSB region). 1072 1073 Figure 16: For ship stations 1, 2, 3, 5a and 5b, we show the averaged spectrum of the in situ 1074 spectra of the UPSEN stations in blue; the spectrum of the referent vector (in red) of the 2S-SOM 1075 1076 neuron, which has captured the closest satellite observations to the UPSEN station; among the 1077 different spectra constituting the referent spectrum, the spectrum of the learning database (DGIP) 1078 that is the closest to the averaged satellite spectra is shown in black. In the rectangular cartoons, we 1079 show the position of the UPSEN station, the number of the neuron of the 2S-SOM which has 1080 captured the satellite observation, the Rfuco of the referent vector, the Rfuco_{DGIP} of the closest DGIP 1081 and the in situ Rfucoupsen.

Figure 11 : (A) chl-a concentration, (B) fucoxanthin ratio, (C) aerosol optical thickness,

(D) Peridinin for 28 February 2003. Panels (B) and (D) show that a second order information was

1082

1088

10831084 Table Caption

1085
1086 Table 1 : Pigments of the DPIG and their statistical characteristics: STD (Standard Deviation), MIN
1087 (minimum value), MAX (maximum value).

- 1089Table 2 : Statistical parameters (R^2 coefficients, RMSE and P-values) of the cross validation between1090the DPIG in situ pigments and the pigments given by the 2S-SOM averaged for the 30 2S-SOM1091realizations
- 10921093Table 3 : For ship stations 1, 2, 3, 5a and 5b of the UPSEN campaign, we show the referent captured1094by the VIIRS observations, the fucoxanthin-ratio associated with this referent (Rfuco-2S-SOM), the1095fucoxanthin-ratio of the closest DPIG fucoxanthin-ratio captured by the neuron of the referent and the1096fucoxanthin-ratio measured in situ during the UPSEN campaign.

1098	Author Contribution
1099	Dr N'Dye Niang and Maurice Ouattara provided the 2S-SOM code, Khalil Yala processed the data
1100	and did the computations with the 2S-SOM, Sylvie Thiria, Michel Crepon and Julien Brajard
1101	analyzed the results, Carlos Mejia and Roy El Hourany did the statistical tests presented in tables and
1102	figure 13. Prof. Sylvie Thiria conceived and supervised the study.
1103	
1104	
1105	Code/data availability
1106	The satellite data (ocean color and SST) are available at the web site:
1107	http://poacc.locean-ipsl.upmc.fr/.
1108	The DPIG data base was kindly provided by Dr. S. Alvain (Severine.alvain@univ-littoral.fr)
1109	The UPSEN data are available at : alban.lazar@locean-ipsl.upmc.fr
1110	The 2S-SOM code is available on request at: <u>carlos.mejia@locean-ipsl.upmc.fr</u>
1111	
1110	
1112	
1113	Short summary
1114	The paper is a contribution to the study of the phytoplankton pigment climatology from satellite
1115	ocean colour observations in the Sénégalo-Mauritanian upwelling, which is a very productive region
1116	where in situ observations are lacking. We processed the satellite data with an efficient new neural
1117	network classifier. We were able to provide the climatological cycle of diatoms. This study may have
1118	an economic impact on fisheries thanks to a better knowledge of phytoplankton dynamics.

1119 1120