Ms. Ref. No.:	os-2016-70
Title:	Synoptic fluctuation of the Taiwan Warm Current in winter on the East China Sea shelf
Journal:	Ocean Science

Dear Referees,

Thank you very much for your comments on our manuscript. All the comments (RF#1 comments, RF#2 comments, RF#3 first comments and RF#3 further comments) have been responded one by one with detailed explanations, which were enclosed after this letter.

Best regards,

Jiliang Xuan, Daji Huang, Thomas Pohlmann, Jian Su, Bernhard Mayer, Ruibin Ding, Feng Zhou
December 15, 2016

Response to the RF#1 comments

Comments: The minor comments are, 1. There are too many keywords, I suggest that "Taiwan Warm Current, Taiwan Strait Current, Kuroshio, Zhe-Min Coastal Current, East China Sea" are enough.

Author's response: We agree that seven keywords are too many. Therefore, we keep the first five keywords which are more important and more frequently used in the manuscript than the last two keywords.

Author's changes in manuscript: Line 41-43: The last two keywords "Zhe-Min Coastal Current, wind" were deleted.

Comments: 2. The model domain and open boundaries are not clear, they may be drawn or indicated in figure 1 or 2.

Author's response: We now have drawn the model domain and the open boundaries in Fig. 2 (Left panel).

Author's changes in manuscript: Line 896: We added a left panel in Fig.2 to show model domain, open boundaries and grids. The Fig. 2 will be attached in this reply.

Comments: 3. In figures 10&11, can you calculate the confidence level of the correlation?

Author's response: We used the MATLAB function "corr" to compute the declining indicator, i.e., the p value which can directly be related to the confidence level. E.g., if the confidence level is larger than 95%, the p value is less than 0.05.

Author's changes in manuscript: Line 954: We changed the statement "whereby R has statistical significance when the p value is less than 0.05." to "whereby R has statistical significance and the confidence level is larger than 95% when the p value is less than 0.05.".

Comments: 4. A new paper by Liu et al. (2016, Numerical simulation of the Kuroshio intrusion into the South China Sea by a passive tracer. Acta Oceanologica Sinica, 2016, 35(9), 1-12.) used tracer simulation to study the Kuroshio intrusion, their model domain covered some of yours. I suggest in discussion of your Fig.14b, comparison between both model results is beneficial.

Author's response: Thanks for the information of the new paper by Liu et al. (2016). We have carefully compared the MITgcm results from Liu et al. (2016) and the FVCOM results from this manuscript. Their results nice complement our study, since they provide the information on the origin of the TSC, which is missing in our paper. Liu et al. (2016) showed that the winter TSC originated in a small branch of Kuroshio intrusion from the Luzon Strait. Their Fig. 14 shows that the winter TSC will flow into both the ECS shelf and the Kuroshio region showing significant synoptic fluctuations.

Author's changes in manuscript: Line 527: We referred to "Liu et al. (2016)" in the text and discussed about the route of TSC from the two models. A discussion was added after "… the TSC also had an effect on the distribution of Kuroshio water north of Taiwan.": "Liu et al. (2016) showed that the winter TSC originated from a small branch of Kuroshio intrusion into the Luzon Strait. Our results complement this picture, since they show that most TSC particles flow into the TWC offshore branch under the influence of cross-shore fluctuation."

Response to the RF#2 comments

Comments: 1. The title '... in winter on the East China Sea Shelf'. In fact only results from the winter (Feb) 2009 were analyzed and presented in the paper and no attempt was made to extend to other winters. I, therefore, suggest that the title to be changed to '…in the winter of 2009…', or similar.

Author's response: We have now analyzed the TWC structures in other winters, i.e., of the years 2010 to 2013, and added a "Figure 14" to the manuscript. Results show that the general TWC structures in the other winters were similar to that in winter 2009. Therefore, we have kept the original title "... in winter on the East China Sea Shelf".

Author's changes in manuscript:

Line 974: A "Figure 14" was added to show the mean currents and synoptic fluctuations in the winters of the years 2010 to 2013.

Line 450-458: The following discussion was added: "The simulated results in the winters of the years 2010 to 2013 (Fig. 14) show that general structures of the TWC in the other winters were similar to that in winter 2009 (Fig. 5 and Fig. 9), which indicates that the results from the winter 2009 can be regarded as representative for the winter situation. The two TWC branches and the two areas of strong fluctuations were present in all the winters from of 2009 to 2013, although their strength showed a certain inter- annual variability in accordance with the changing surface forcing and boundary fluxes.".

Comments: 2. Line 133-135, on the model, 'the river discharge of the Changjiang and Huanghe…'. In the immediate region of the study, there are other important rivers, e.g., Qiantang, etc.; did the authors include them; if not, why not.

Author's response: Other rivers were not included in this study, because the discharges for the other rivers are very small compared to the Changjiang and Huanghe discharge.

Author's changes in manuscript: Line 150-152: We added the statement explaining why other rivers were not included: "Other rivers were not included because of their small discharge rates, e.g., the Qiantang River, with the largest runoff from the Zhejiang coast, has a climatological mean discharge rate in winter of about 230 m^3/s, which is nearly negligible compared to the Changjiang winter discharge of about 11500 m^3/s.".

Comments: 3. Validation of the model. The model results were validated in the coastal regional of the East China Sea shelf and good agreements between measurements and model were obtained. However this is limited to the shallow coastal region. It is well known that the FVCOM uses a sigma co-ordinate and prone to errors in the region of steep topography. Therefore authors should at least caution readers that in the slope region, such as the Kuroshio intrusion (line 465-470), model results are less reliable unless a good validation is provided.

Author's response: We agree that our results in the region of steep topography may not have the same accuracy because of the sigma co-ordinates used in FVCOM, in particular at in the shelf break area where Kuroshio intrusion occurs. However, the synoptic fluctuations, which we focus on, are mainly located on the shelf, where results could be nicely validated. Nevertheless the validation for the offshore area could not be performed properly. Therefore, we added a statement regarding the simulation results in the slope region.

Author's changes in manuscript: Line 534-535: we added the following statement: "However, because FVCOM uses sigma co-ordinates in the vertical which are prone to errors in regions of steep topography, our results may underestimate the fluctuations at the shelf break, in particular to the northeast of Taiwan where Kuroshio intrusion occurs.".

Comments: 4. Line 264. '… in wintertime, both branches flowed on the isobaths, which is fully in accordance with the conservation of potential vorticity'. It implies 'cross isobaths flow is not following the pv conservation law', which is not correct. btw, 'on the isobaths' - > 'along the isobaths'.

Author's response: We agree that our statement "… is fully in accordance with the conservation of potential vorticity" is wrong.

Author's changes in manuscript: Line 296: We omitted this statement. At this place, the main argument also holds without referring to the PV-conservation.

Response to the RF#3 first comments

Comments: 1. Regrettably, their analysis, interpretation and discussion are found to be incoherent and devoid of strong/convincing physical reasonings probably due to lack of comprehensive understanding of winter monsoonal flows.
Author's response: We have revised the manuscript according to the comments in the following two aspects: 1) unified the topic, definitions and terms throughout the manuscript, e.g., specified our topic which is about the mean currents and synoptic fluctuations in the study time (January and February 2009). 2) Explained the physical reasoning in detail, e.g., explained that other factors (wind stress and bottom stress) also play important roles on the dynamics in the Taiwan Strait.
Author's changes in manuscript: A number of revisions have been made according to the following major comments and minor comments.

Comments: 2. Some key findings of the following articles may be helpful to enhance this study:- (i) Hong, Huasheng, et al. "An overview of physical and biogeochemical processes and ecosystem dynamics in the Taiwan Strait." Continental Shelf Research 31.6 (2011): S3-S12. (ii) Hu, Jianyu, et al. "Review on current and seawater volume transport through the Taiwan Strait." Journal of Oceanography 66.5 (2010): 591-610.
Author's response: Thanks for the information. We have carefully studied the references and quoted them to show the dynamics of the Taiwan Strait Current.
Author's changes in manuscript: in **line 107-109**, we added a statement: "Hong et al. (2011) and Hu et al. (2010) summarized that the temporal and spatial variation of TSC is modulated by strong monsoon forcing, complex topography and circulation in the northern South China Sea as well as coastal water input and the Kuroshhio intrusion".

Comments: 3. In addition, definitions and terms such as north of Taiwan, inshore area, inshore branch, offshore branch, alongshore, cross shore and cross shelf are noted to cause confusion when some of these terms are used interchangeably at times.
Author's response: We have now unified the terms in two aspects: 1) Since the study area is on the ECS shelf, we changed "cross-shelf" to "cross-shore" throughout the manuscript; 2) explained that the inshore branch is associated with the inshore area.
Author's changes in manuscript: We changed "cross-shelf" to "cross-shore" throughout the manuscript. Moreover, in **line 82-83**, we defined that "the inshore area is the area between the 30 and 100 m isobaths where the TWC inshore branch dominates".

Comments: 4. Quoting literature review without further elaboration to strengthen a point is insufficient. Some figures are hard to see, not properly captioned and without the unit specified for the parameter. There exists a number of structural and grammatical errors in the language used.
Author's response: We have revised the manuscript according to the comments.

Author's changes in manuscript: A number of revisions have been made according to the following major comments and minor comments.

Comments: 5. Finally, this is merely a case study (for January and February 2009) and the conclusions drawn are only applicable for this specific late-winter case. As such, a major revision of this manuscript, inclusive of its title, is needed before it can be considered for publication.
Author's response: We have now analyzed the TWC structures in the following four winters as well, i.e., December-March of the years 2010 to 2013, and added a "Figure 14" to the manuscript. Results show that the general TWC structures in the other winters were similar to that in January-February of 2009. Therefore, the results from the January-February 2009 can be regarded as representative for the winter situation.
Author's changes in manuscript: **Line 975:** A "Figure 14" was added to show the mean currents and synoptic fluctuations in the winters of the years 2010 to 2013.
Line 451-458: The following discussion was added: "The simulated results in the winters of the years 2010 to 2013 (Fig. 14) show that general structures of the TWC in the other winters were similar to that in winter 2009 (Fig. 5 and Fig. 9), which indicates that the results from the winter 2009 can be regarded as representative for the winter situation. The two TWC branches and the two areas of strong fluctuations were present in all the winters from of 2009 to 2013, although their strength showed a certain inter- annual variability in accordance with the changing surface forcing and boundary fluxes".

Major comments
Comments: 1-1. Are November to March the winter months?
Author's response: We have now added a definition of winter months from December to March, according to the critical value of local air temperature (< 10 ℃ in winter) in the East China Sea (Chen et al. 1992). Detail information could be seen in the reference: Shangji Chen, Weihuan He, Tiyu Yao, et al., 1992, Discrimination of ocean hydroclimate seasons in the China Seas, Acta Oceanologica Sinica, 14(6), 1-11 (In Chinese).
Author's changes in manuscript: **Line 50:** We defined the wintertime as December to March.

Comments: 1-2. How could you explain the weak mean velocity of the winter TWC on the ECS shelf (lines 51-52)?
Author's response: We agree that the statement of "weak mean velocity" is not clear. We have changed "weak mean velocity" to "weak mean surface velocity" according to Qiu and Imasato (1990)'s study. Qiu and Imasato (1990)'s study is also quoted to show the climatological structure of the surface current in the East China Sea, which is mapped by averaging GEK current data available over 1/5° × 1/5° resolution boxes from 1953 to 1984.
Author's changes in manuscript: **Line 51-52:** We changed the statement "… its

weak mean velocity superimposed by pronounced small-scale spatial and synoptic temporal variations" to "… its weak mean surface velocity, according to a climatological structure of the surface current in the ECS mapped by Qiu and Imasato (1990)".

Comments: 1-3. What are the dominant physical factors that cause the fluctuations of the TWC to have periods between 3 and 15 days (lines 55-58)?
Author's response: The physical mechanism of the TWC fluctuations is a question to be solved in our manuscript. In the former studies, the wind was recognized as a main physical factor over most continental shelves. In the coastal area of the ECS, Huang et al. (2016) have shown that the wind was a main physical factor which caused the temporal variation of the wintertime currents at the synoptic scale. However, as the Hong et al. (2011)'s study (mentioned by the referee) said that "Due to its complex bottom topography, alternating monsoon forcing and conjunction of several current systems [such as the Zhejiang–Fujian (Zhe–Min) Coastal Current, the Kuroshio intrusion and the extension of the South China Sea Warm Current], the physical and biogeochemical processes and ecosystem dynamics in the Taiwan Strait vary significantly both in space and in time", the dominant physical factors which cause the TWC fluctuations on the whole shelf of the ECS may be very complicated. Therefore, we added an argument to raise the question that the dominant physical factors of the TWC fluctuations are still lack of study.
Author's changes in manuscript: Line 58-63: The following argument was added: "Huang et al. (2016) have shown that the wind was a main physical factor which caused the temporal variation of the wintertime currents at the synoptic scale in the coastal area of the ECS. However, the dominant physical factors of the TWC fluctuations are still lack of study, regarding that the fluctuations on the whole shelf of the ECS may be more complicated due to the complex bottom topography, alternating monsoon forcing and conjunction of several current systems such as the Kuroshio Current, the Taiwan Strait Current (TSC) and the Zhe-Min Coastal Current (ZMCC)."

Comments: 1-4. What do you mean by "the intermittency of the TWC in winter" (line 60)?
Author's response: We agree that the statement of "the intermittency of the TWC in winter" is not clear. Zhu et al. (2004)'s study demonstrated that the TWC has an episodic wintertime feature. This means the TWC intermittently occurs in winter. Therefore, we referred Zhu et al. (2004)'s result to illustrate the episodic wintertime feature of the TWC.
Author's changes in manuscript: Line 67-69: We changed "some recent observations have shown that the intermittency of the TWC in winter reaches maximum velocity variations larger than 0.2 m/s (Zhu et al., 2004; Zeng et al., 2012)" to "some recent observations have shown that the TWC has an episodic wintertime feature (Zhu et al., 2004) and the intermittency of the TWC in winter has an amplitude as large as 0.2 m/s (Zeng et al., 2012)".

Comments: 1-5. Under what synoptic condition can the TSC be considered as an upstream flow of the TWC (line 96)?

Author's response: Obviously it is the northeastward TSC that could be an upstream flow of the TWC. In addition, we have referred Hong et al. (2011)'s overview to explain the physical factors of the TSC variations.

Author's changes in manuscript: **Line 106:** Changed "The TSC" to "The northeastward TSC"; **Line 107-109:** we added a result of Hong et al. (2011)'s overviews "Hong et al. (2011) and Hu et al. (2010) summarized that the temporal and spatial variation of TSC is modulated by strong monsoon forcing, complex topography and circulation in the northern South China Sea as well as coastal water input and the Kuroshhio intrusion.".

Comments: 1-6. What is the physical significance of inserting "Takahashi … the annual (?) variation of the TWC … the propagation of vorticity anomalies …" (lines 98-100)?

Author's response: We have now added a statement to explain the significance of Takahhashi and Morimoto (2013)'s result, which further demonstrated that the fluctuation of TWC was associated with its upstream currents such as the TSC.

Author's changes in manuscript: **Line 111-114:** we revised the statement as following: "Takahashi and Morimoto (2013) pointed out that the temporal variation of the TWC is characterized by the propagation of vorticity anomalies originating from northeast of the Taiwan Strait, which further demonstrated that the fluctuation of TWC was associated with its upstream currents such as the TSC".

Comments: 2-1. It is obvious that your case study is for January and February 2009. Hence, your climatological (Years of climatological period are not mentioned in your manuscript) and observational deductions must refer only to these late-winter months.

Author's response: We have now analyzed the TWC structures in the following four winters as well, i.e., December-March of the years 2010 to 2013, and added a "Figure 14" to the manuscript. Results show that the general TWC structures in the other winters were similar to that in January-February of 2009. Therefore, the results from the January-February 2009 can be regarded as representative for the winter situation.

Author's changes in manuscript: **Line 975:** A "Figure 14" was added to show the mean currents and synoptic fluctuations in the winters of the years 2010 to 2013.

Line 452-458: The following discussion was added: "The simulated results in the winters of the years 2010 to 2013 (Fig. 14) show that general structures of the TWC in the other winters were similar to that in winter 2009 (Fig. 5 and Fig. 9), which indicates that the results from the winter 2009 can be regarded as representative for the winter situation. The two TWC branches and the two areas of strong fluctuations were presented in the all winters from of 2009 to 2013, although their strength showed a certain inter- annual variability in accordance with the changing surface forcing and boundary fluxes".

Comments: 2-2. Apart from defining near-coast, inshore and offshore areas based on isobaths, can you offer an explanation why the TWC inshore and offshore branches only dominate in those specific isobaths (lines 75-78)?

Author's response: We have now added an explanation that the TWC inshore and offshore branches mainly occur in those specific isobaths is due to the conservation of potential vorticity, according to the hydrographic data analysis and numerical interpretation by Su and Pan (1987).

Author's changes in manuscript: **Line 87-88:** We added the explanation that "According to the hydrographic data analysis and numerical interpretation by Su and Pan (1987), the TWC inshore and offshore branches mainly occur in those specific isobaths is due to the conservation of potential vorticity".

Comments: 3-1. Lines 103-113 under Introduction should be moved to Data and Methods (suggest to change from Methods and validation) section.

Author's response: We still kept this paragraph under the section of Introduction for the following reasons: 1) statements in the paragraph are an introduction of three popular methods in studying the synoptic fluctuations of the wintertime TWC, while the first two methods were not used in this manuscript because the observation data are limited in terms of temporal and spatial coverage; 2) statements in the paragraph indicated that the numerical simulation provide a promising approach for studying the synoptic fluctuations, which gave the background for the next section of Methods and validation. In addition, the section name "Methods and validation" was also kept because the model validation is a very important part in the manuscript since we need reliable numerical simulations to investigate the TWC fluctuations.

Author's changes in manuscript: Nothing has been changed, due to the above arguments.

Comments: 4-1. Please provide sufficient details on model setup, configuration, data used, forces and boundary conditions. As the model is run fully in three dimension, time steps for baroclinic and barotropic runs should be defined separately.

Author's response: We have carefully examined the model setup, configuration and driving forces and added some essential information.

Author's changes in manuscript: **Line 136:** We changed "an unstructured-grid FVCOM" to "a 3-D unstructured-grid (Fig. 2, left panel) FVCOM". **Line 148:** We changed "The river discharge" to "The daily-mean river discharge". **Line 150-153:** We added a statement "Other rivers were not included because of their small discharges, e.g., the Qiantang River, with the largest runoff from the Zhejiang coast, has a climatological mean discharge in winter of about 230 m^3/s, which is nearly negligible compared to the Changjiang winter discharge of about 11500 m^3/s". **Line 166-167:** We changed "The model time step was 90 seconds" to "The model time step was 15 seconds for the 2-D barotropic mode and 90 seconds for the 3-D baroclinic mode".

Comments: 5-1. Lines 143-147: Not clear. How could you obtain the hindcast outputs for late winter 2009 when you simulated the model using 2009 to 2013 data with three years of spin-up?

Author's response: We have now added the detailed information for the spin-up years (2006-2008) and the initial conditions (the initial temperature and salinity were taken from the Hybrid Coordinate Ocean Model and initial velocity was set to zero).

Author's changes in manuscript: **Line 162-164:** We added the detailed information for the spin-up years and the initial conditions as following: "… simulation from 2009 to 2013 are used, following three spin-up years (2006-2008) initiated with the temperature and salinity taken from the Hybrid Coordinate Ocean Model and velocity setted to zero".

Comments: 6-1. Lines 156-171: Write-up on validations is vague.

Author's response: We have rewritten the validations of circulation structure and boundary fluxes.

Author's changes in manuscript: **Line 179:** We emphasized that "The FVCOM has reproduced almost all of the known circulation structure in the ECS in winter". **Line 191-193:** We highlighted the validation of the volume transport through the Taiwan Strait that "The simulated transports were accurate enough to reproduce volume transport (1.22 Sv) through the Taiwan Strait which is closer to the observation value (1.20 Sv) from Isobe (2008) than former model results".

Comments: 7-1. Lines 186-188: "The cross-shore component (Figs 3c and 3d) is much … spatial pattern. It flows offshore in the upper layer and onshore in the lower layer at Station 1." What is seen in the figure is different from what is expressed here.

Author's response: It is because the color scales in Fig. 3 are not same between the alongshore components and cross-shore components.

Author's changes in manuscript: **Line 908-909:** We added a notation that "Note, an enlarged color scale is used for the cross-shore component to have a clear view of its weak structure.".

Comments: 8-1. Errors are found in labelling Figure 4 and the explanation given on simulated and observed results is not clear.

Author's response: We have now revised the label and explanation of Figure 4.

Author's changes in manuscript: **Line 912:** We changed "Validations of the wintertime TWC fluctuations" to "Variations of the inshore branch of TWC during January and February 2009". **Line 224:** We changed "Synoptic fluctuations of the wintertime TWC were also validated against the mooring results (Fig. 4)" to "Synoptic fluctuations of the TWC inshore branch during January and February 2009 were also validated against the mooring results (Fig. 4)".

Comments: 8-2. If the alongshore component is nearly one order of magnitude larger than cross-shore component, how could their fluctuation magnitudes be comparable (lines 206-207).

Author's response: Apparently "in the mean condition" was missing in the first sentence "the alongshore component is nearly one order of magnitude larger than cross-shore component".

Author's changes in manuscript: Line 234: We revised the statement as following: "…the alongshore component is nearly one order of magnitude larger than the cross-shore component in the mean condition …"

Comments: 9-1. Line 235: "Second, according to the hydrostatic …" This is ambiguous, please rewrite it.

Author's response: We have rewritten the sentence.

Author's changes in manuscript: Line 261: We changed "Second, according to the hydrostatic assumption used in FVCOM [as shown in Eq. (2)], the pressure is integrated from depth z to the sea surface" to "Second, according to the hydrostatic approximation used in FVCOM [as shown in Eq. (2)], the pressure gradient is given as the product of density times the gravitational acceleration".

Comments: 10-1. Line 245: please define the mathematical form of the wind stress (at the sea surface (ta) and sea bottom (tb)) used in the model.

Author's response: We have now defined the mathematical form of the wind stress and bottom stress.

Author's changes in manuscript: Line 274: Eq. (5) has been revised to show the mathematical form of the wind stress $\underbrace{\rho_a C_D \left|\vec{U}\right| \vec{U}}_{\tau_a}$ and bottom stress $\underbrace{-k_b \left|\vec{U_b}\right| \vec{U_b}}_{\tau_b}$.

Comments: 11-1. Lines 263-264: Please elaborate this statement - "…which is fully in accordance with the conservation of potential vorticity".

Author's response: We agree that our statement "… is fully in accordance with the conservation of potential vorticity" is not correct due to the influence of frictions.

Author's changes in manuscript: Line 296: We omitted this statement.

Comments: 12-1. Lines 271: Explain why different cooling exists in both areas.

Author's response: The differential cooling between coastal and offshore area is due to the different heat capacity of water columns with different water depths.

Author's changes in manuscript: Line 305-306: We explained that "Assuming a relatively spatially homogeneous heat loss, a different cooling occurs, due to the smaller heat capacity of the shallow coastal water compared to the deep offshore water".

Comments: 13-1. Lines 275-278: "The fact that the depth of the subsurface VMV … the effects of baroclinicity and wind friction …" Explain in detail this key finding.

Author's response: We have explained this point (the effects of baroclinicity and wind friction on the inshore branch are stronger than for the offshore branch) in **line 303:** 1) The northerly wind in winter weakens the northward TWC, particularly in the upper layer, which leads to the formation of the subsurface VMV; 2) Differential cooling, due to different heat capacity of the water columns, as explained in the previous response is responsible for a stronger cooling of the coastal shallow waters ccompared to the offshore deep waters. In addition, through analyzing the vertical structure of the inshore branch (P2, Fig. 7) in the following section (3.2 Synoptic fluctuations, Line 351-362), we also reached the conclusion that the fluctuations induced by wind and surface cooling are stronger in the inshore branch than that in the offshore branch.

Author's changes in manuscript: We hope that the clarification with respect to the previous comment (12-1) helps to make our point more clear. Moreover, we have modified the text at lines 310-313 and tried put a stronger focus on our main findings.

Comments: 14-1. Lines 311-312: "The currents fluctuated … occurred episodically". What episodic events you are referring to?

Author's response: From the result of the standard deviations of the currents (Fig. 7), we can infer that the TWC inshore branch occurs episodically, although specific episodic events could not be identified. However, our manuscript is intended to focus on the synoptic variations (with about 3-15 days periods) and their general impact on the water transports, which is different from the traditional concerns only on mean structures. Hence, specific events have not been highlighted in the manuscript except, for two extreme events in the discussion.

Author's changes in manuscript: **Line 347:** We changed "The currents fluctuated in the alongshore direction …" to "As deduced from the standard deviation, the currents fluctuated significantly in the alongshore direction …"

Comments: 15-1. Line 320: Suggest you calculate the mixed layer depth based on: "Lorbacher K, Dommenget D, Niiler PP, Kohl A. Ocean mixed layer depth: A subsurface proxy of ocean atmosphere variability. J Geophys Res-Ocean. 2006; 111(C7): 1978–2012. doi: 10.1029/2003JC002157."

Author's response: Thank you for your information about how to calculate the upper mixed layer depth from observed temperature, salinity or potential density. In this manuscript, we kept the original method, which used the critical Richardson number, because it is based on the dynamics of instability which include both effect of the potential density and vertical current shear and is also implemented in the FVCOM. This method is also widely used, see e.g. Mellor and Durbin (1975), Grachev et al., 2013, and Richardson et al., (2013).

Mellor, G. L. and Durbin, P A.: The structure and dynamics of the ocean surface mixed layer. J. Phys. Oceanogr., 5(4), 718–728, 1975.

Grachev, A. A., Andreas, E. L., Fairall, C. W., Guest, P. S., and Persson, P. O. G.: The critical Richardson number and limits of applicability of local similarity theory in the stable boundary layer, Boundary-layer meteorology, 147(1), 51–82, 2013.

Richardson, H., Basu, S., and Holtslag, A. A. M.: Improving stable boundary-layer height estimation using a stability-dependent critical bulk Richardson number, Boundary-layer meteorology, 148(1), 93–109, 2013.

Author's changes in manuscript: **Line 356-359:** We have quoted the references as the following: "The depths of the upper mixed layer were determined by a Richardson number criterion (Mellor and Durbin, 1975; Grachev et al., 2013; Richardson et al., 2013), where the critical Richardson number equals 0.25 in this paper [as in Xuan et al. (2012)]"

Comments: 16-1. Lines 341-346: Hard to follow your explanation in the figures. Please plot them in different depths.
Author's response: We have now added the 30, 50, 70, 100 and 200 m isobaths in Fig. 9.
Author's changes in manuscript: Line 947: We redrawn the Fig. 9 by adding the 30, 50, 70, 100 and 200 m isobaths.

Comments: 17-1. Under "3.3 Dynamic diagnostics", you argued (based on Figure 12) that the Coriolis force is mainly balanced by the total pressure in both branches, … in the wintertime TWC. This is not convincing for the Taiwan Strait.
Author's response: We agree that other factors also play important roles on the dynamics in the Taiwan Strait, especially the wind stress and bottom stress. Both the studies of Guo et al. (2003, Fig. 13) and this manuscript (Fig. 12) showed the important effects of wind stress and bottom stress in the Taiwan Strait.
Author's changes in manuscript: Line 414: We have added the effects of wind stress and bottom stress in the Taiwan Strait as follows: "the wind-induced surface friction plays an important role in the TWC, especially in the inshore area and the Taiwan Strait (Fig. 12c). The bottom friction has an impact north of Taiwan and in the shallow Taiwan Strait, in particular when a significant Kuroshio intrusion enhances the bottom flow (Fig. 12d)."

Comments: 18-1. Lines 395-400: I am not convinced. You argue that "… the TSC mainly caused variations in the barotropic pressure gradients, which further …" As I know barotropic pressure gradients is generated by a sloping sea surface and the pressure gradient is depth independent. Please clarify.
Author's response: We agree that barotropic pressure gradients are generated by a sloping sea surface and the pressure gradient is depth independent. In **line 436**, we have explained, that "The mechanism can be interpreted as follows. When a larger TSC intrusion occurred, the isobaric slope tilted downward from south to north, generating a cross-shore current from the coastal area to the offshore area. On the contrary, when the TSC intrusion was weak, the Kuroshio intrusion from offshore to inshore dominated north of Taiwan".
Author's changes in manuscript: Line 435: We changed "the TSC mainly caused variations in the barotropic pressure gradients" to "in the shallow coastal area the TSC mainly caused variations in the depth-independent barotropic pressure gradients".

Comments: 19-1. Lines 404-408: Confusing. Why is the negative Coriolis force associated with a northerly wind?
Author's response: We agree that the statement of negative Coriolis force is not clear. Obviously it is the northwestward Coriolis force that indicates the southwestward

current.
Author's changes in manuscript: Line 445: We changed "negative" to "northwestward direction".

Comments: 20-1. Lines 430-435: What is numerical tracer simulation and how is it connected with tracer assimilation? Please elaborate with their physical applications.
Author's response: We have now added a detailed statement for the numerical tracer simulation and its associated physical applications.
Author's changes in manuscript: Line 486-489: We modified the text as follows: "A numerical tracer simulation was used to analyze the role of the cross-shore fluctuation in the transport of the TSC water and the Kuroshio water north of Taiwan. In order to demonstrate the characteristics of the flow patterns more clearly, artificial tracers are released in the model domain and transported by the velocity field provided by the FVCOM simulation. The tracer running was part of the FVCOM simulation; therefore, all the above mentioned dynamics were involved, e.g., tide, wind, and boundary forces."

Comments: 21-2. Lines 491-497: Did the episodic occurrence relate to the surge and lull periods of the late winter?
Author's response: We have updated the response to this comment in the further comments from the referee:
We agree that both the wind surges and wind-induced synoptic fluctuations are the short-term events caused by strong wind. Wind surges are often recognized as a wind-induced sea level fluctuation, especially the sea level rise along coastal line or at the coast area, which are much concerned in the coastal engineering for disaster prevention or mitigation. In contrast, synoptic fluctuation is frequently used on the continent shelf as a phenomenon with a period of 3-15 days. In addition to the sea level, other physical parameters (e.g., currents, temperature, and salinity) are also concerned in the synoptic fluctuation studies. The synoptic fluctuation is also mainly caused by wind and can be well explained by coastal trapped wave (Huthnance et al. 1986; Brink, 1991). Regarding the different concerns of the wind surges and synoptic fluctuations associated with region and parameters, we put our topic to synoptic fluctuations instead of wind surges.
Actually when the extremely short-term events of wind surges occur in the coastal area, there must be synoptic fluctuation on the continent shelf. The synoptic events, with time scales of 3-15 days, play a dominant role on the horizontal advective transports. According to Ledwell et al. (1998) synoptic variations are much more effective on the horizontal transport than variations on shorter time scales. Therefore, we decided to focus on the synoptic events to investigate the mass transports in the season with the strongest synoptic wind, i.e., winter situation form December to March in the ECS.
Author's changes in manuscript: Line 59: we added two references of Brink. (1991) and Huthnance et al. (1986) to highlight the well-known synoptic fluctuations over the continental shelves, which can be explained theoretically by coastal trapped wave.

Line 468-471: we added a discussion to show that the synoptic events play the dominated role for the horizontal transports.

Comments: 22-3. Lines 506-519: Description and explanation are vague.
Author's response: We have now added some information on the description and explanation in the discussion of the offshore transports along the latitudes 26.5 °N and 28 °N.
Author's changes in manuscript: Line 573: We have added a quote of Fig. 9f to specify the locations of the offshore transports. **Line 579-580:** We changed the statement "… played an important role in cross-shelf material exchange. Several mechanisms have been used to explain this process" to "… played an important role in cross-shelf material exchange, but the mechanisms of the offshore-penetrating fronts are still under debate".

Comments: 23-1. Lines 552-558: Confusing. Suggest to use the late winter monsoonal flow patterns during the surge and lull periods as a basis to recast your findings.
Author's response: we kept the statement of synoptic fluctuations because it is different from the surge and lull periods. Here we concluded the roles of synoptic fluctuations on the alongshore and cross-shore transports in two areas, north of Taiwan and in the inshore area, respectively.
Author's changes in manuscript: Line 621: See our response to comment 21-2. Moreover, we changed "Due to these fluctuations" to "Due to the fluctuation north of Taiwan".

Minor Comments:
Comments: 1: Line 10 : Is it 10 meter wind above sea surface ?
Author's response: Yes, it is.
Author's changes in manuscript: Line 154: We have shown the source of wind data as following: "the 3-hourly wind stress and 10 m wind speed data was from the ERA-40 re-analysis (Uppala et al., 2005)".

Comments: 2: Line 30: should be "When a strong TSC intrudes towards the north of Taiwan, …"
Author's response: Thanks.
Author's changes in manuscript: Line 30: We changed "When a larger TSC intrudes north of Taiwan" to "When a strong TSC intrudes to the north of Taiwan".

Comments: 3: Line 55: "continental shelf" instead of "continental shelves"
Author's response: Since the "continental shelves" is related to the "some features common", we kept the plural "shelves".
Author's changes in manuscript: Nothing has been changed.

Comments: 4: Line 120, "To investigate the currents … ECS shelf, an unstructured-grid FVCOM was developed for …" should change to "… a 3d unstructured- grid FVCOM is developed for …"
Author's response: Agree.
Author's changes in manuscript: Line 136: We changed "an unstructured-grid … was developed" to "a 3-D unstructured-grid … is developed".

Comments: 5: Lines 124-125: "A regional refinement of the resolution (approximately 3 km) was specified …" should be replaced by "A regional … 3km is specified …".
Author's response: Agree.
Author's changes in manuscript: Line 138: We changed "was" to "is".

Comments: 6: Line 125: Add (GEBCO) after "The General Bathymetric Chart of the Ocean".
Author's response: Agree.
Author's changes in manuscript: Line 140: We added "(GEBCO)".

Comments: 7: Line 125: Please specify number of grid points (m*n*l) used in model configuration.
Author's response: We have now added the cell number (n), which is the most important index in estimating the FVCOM quality.
Author's changes in manuscript: Line 141: We changed "Twenty vertical layers" to "Twenty vertical layers with 76954 triangle cells".

Comments: 8: Line 130: Usually tide is used as boundary conditions and is not a driving force.
Author's response: We kept the definition of tides as driving forces, because according to our understanding external forces can been imposed not only at the air-sea boundary (wind stress, heat fluxes), but also at lateral boundaries.
Author's changes in manuscript: Nothing has been changed.

Comments: 9: Section 2.2 Validation of the mean currents and synoptic fluctuations: "The mean current was …". What you mean by "the mean current" ? Amend "was" to " is".
Author's response: We have now specified the mean currents as the Kuroshio Current, the TWC, and the ZMCC.
Author's changes in manuscript: Line175: Revised as following: "The mean currents, e.g., the Kuroshio Current, the TWC, and the ZMCC".

Comments: 10: Line 180: "…for the observational period … " Please specify the period.
Author's response: We have specified the period: "… for the observational period, which was from January 1 to February 28, 2009".

Author's changes in manuscript: We changed "…for the observational period … " to "…for the observational period , which was from January 1 to February 28, 2009"

Comments: 11: Line 180: "We defined the alongshore direction from southwest (218o) to northeast (38o), which is …" It is very confusing please amend it.
Author's response: We have now revised the statement.
Author's changes in manuscript: Line 207-208: We changed the statement as follows: "Using the same method as in Huang et al. (2016), we defined the positive alongshore current directing from southwest (218°) to northeast (38°)".

Comments: 12: Line 210: Section 2.3: "The Empirical …" should be written as "The Empirical …, as a statistical method, has been used to understand the synoptic fluctuations of the …".
Author's response: Agree.
Author's changes in manuscript: Line 240: We changed "The Empirical Orthogonal Function (EOF) method (Emery and Thomson, 2001)" to "The Empirical Orthogonal Function (EOF) method (Emery and Thomson, 2001), as a statistical method".

Comments: 13: Page 11, Section "2.4 Momentum analysis" should be worded as "2.4 The governing equations".
Author's response: We kept the statement of "2.4 Momentum analysis". As a method name, the "2.4 Momentum analysis" is more appropriate than the "2.4 The governing equations". In addition, it is consistent with the name of upper section "2.3 EOF analysis of synoptic fluctuations".
Author's changes in manuscript: Nothing has been changed.

Comments: 14: Page 18, Section 3.3 Dynamic diagnostics: "The seasonal mean of the water …". What is exact period of this mean.
Author's response: Thanks.
Author's changes in manuscript: Line 410: We changed "The seasonal mean" to "The wintertime (January and February 2009) mean".

Comments: 15: Page 19, line 405: "It indicates that strong winter monsoon can weaken …". Replace "weaken" with an appropriate word.
Author's response: We have changed the statement of this sentence.
Author's changes in manuscript: Line 447: We changed "strong winter monsoon can weaken or even stop the TWC" to "strong northerly monsoon in winter can reduce or even stop the northeastward TWC".

Comments: 16: Page 20, line 430: "Two section, one in the Taiwan Strait and another in …" show transection.
Author's response: We have added the quote of Fig. 15 which showed the two sections.

Author's changes in manuscript: Line 492: We changed the statement as follows: "Two sections, one in the Taiwan Strait (Fig. 15a, black dots) and another in the East Taiwan Channel (Fig. 15b, black dots)".

Comments: 17: Page 21, line 450: "Figure 14b shows the tracers …". Is it tracers or traces?
Author's response: Thanks. It is traces.
Author's changes in manuscript: Line 510: We changed "tracers" to "traces".

Response to the RF#3 further comments

Comments: (1) Cold air outbursts from the Siberian High during winter are responsible for the episodic wind surges over the East China Sea. The outbursts are particularly strong in the late winter months of January and February. Under such episodic event, wind-induced current along the East China Sea prevents the intrusion of the TSC and causes the eastward shift of the strong warm Kuroshio current. This affects inshore and offshore flows.

Author's response:

We fully agree with your comment on the short-term events (wind surges and synoptic fluctuation), which originate from the Siberian cold air and have significant effects on the alongshore and cross-shore transports in the ECS.

Relating to our use of term "synoptic fluctuation" instead of the term "wind surges" suggested by you in the previous round comment, we agree that both the wind surges and wind-induced synoptic fluctuations are the short-term events caused by strong wind. Wind surges are often recognized as a wind-induced sea level fluctuation, especially the sea level rise along coastal line or at the coast area, which are much concerned in the coastal engineering for disaster prevention or mitigation. In contrast, synoptic fluctuation is frequently used on the continent shelf as a phenomenon with a period of 3-15 days. In addition to the sea level, other physical parameters (e.g., currents, temperature, and salinity) are also concerned in the synoptic fluctuation studies. The synoptic fluctuation is also mainly caused by wind and can be well explained by coastal trapped wave (Huthnance et al. 1986; Brink, 1991). Regarding the different concerns of the wind surges and synoptic fluctuations associated with region and parameters, we put our topic to synoptic fluctuations instead of wind surges.

Actually when the extremely short-term events of wind surges occur in the coastal area, there must be synoptic fluctuation on the continent shelf. The synoptic events, with time scales of 3-15 days, play a dominant role on the horizontal advective transports. According to Ledwell et al. (1998) synoptic variations are much more effective on the horizontal transport than variations on shorter time scales. Therefore, we decided to focus on the synoptic events to investigate the mass transports in the season with the strongest synoptic wind, i.e., winter situation form December to March in the ECS.

Author's changes in manuscript: Line 59: we added two references of Brink. (1991) and Huthnance et al. (1986) to highlight the well-known synoptic fluctuations over the continental shelves, which can be explained theoretically by coastal trapped wave. Line 469-472: we added a discussion to show that the synoptic events play the dominated role for the horizontal transports.

Comments: (2) The study area is a continental shelf. Topography and bathymetry as well as strong winds give rise to drag in the shallow shelf. Hence, the use of conservation of potential vorticity and geostrophic balance are invalid in the area concerned.

Author's response: We agree that the PV is not a strictly conservative quantity on the continent shelf. However, our conclusions are not all based on the PV conservation argument. Actually, there are only two places in the manuscript where it is mentioned, i.e., when we show the wintertime averaged distribution of the TWC under frictionless conditions and for flows with a minor meridional extension. At both places, the main argument also holds without referring to the PV-conservation.
Author's changes in manuscript: We have omitted the references to the "conservation of potential vorticity" in lines 87 and 296.

Comments: (3) For the baroclinic model run to respond to the lateral boundary conditions such as variations of temperature, salinity and currents, tens of seconds of spin-up time is totally wrong as this can cause the model to blow up.

Author's response: We agree with your worry. Obviously, we did not make it clear enough that the initial conditions are ramped-up over a period of 30 days and that at the lateral boundaries a sponge layer was used. Therefore our simulation did not blow up as our results did not show any unstable signals.

Author's changes in manuscript: In line 164-166, we have added these two points to the model description and also refer to the paper of Chen et al. (2008), where all the details are given:
"The initial conditions are ramped-up over a period of 30 days and at the lateral boundaries a sponge layer was used with the same method as Chen et al. (2008)."

References:

Brink, K. H.: Costal trapped waves and wind-induced currents over the continental shelf, Annual Review of Fluid Mechanics, 23, 389-412, 1991.

Chen, C., Xue, P., Ding, P., Beardsley, R. C., Xu, Q., Mao, X., Gao, G., Qi, J., Li, C., Lin, H., Cowles, G., and Shi, M.: Physical mechanisms for the offshore detachment of Changjiang Diluted Water in the East China Sea, J. Geophys. Res., 113, C02002, doi:10.1029/2006JC003994, 2008.

Huthnance, J. M., Mysak, L. A., and Wang, D. P.: Coastal trapped waves, in: Mooers C.N.K. (ed.) Baroclinic Processes on Continental Shelves. Coastal and Estuarine Sciences, 3, pp. 1-18. Washington DC, American Geophysical Union, 1986.

Ledwell, J. R., Watson, A. J., and Law, C. S.: Mixing of a tracer in the pycnocline, J. Geophys. Res., 103(C10), 21499–21529, doi:10.1029/98JC01738, 1998.

Synoptic fluctuation of the Taiwan Warm Current in winter on the East China Sea shelf

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Highlights

- Synoptic fluctuations of the wintertime Taiwan Warm Current appear mainly in two areas: north of Taiwan and the inshore area
- Synoptic fluctuation is mainly driven by the Taiwan Strait Current north of Taiwan and by wind in the inshore area
- Large Taiwan Strait Current intrusion generates a cross-shore transport from the coastal area to the offshore area
- Winter monsoon affects the alongshore transport of Taiwan Warm Current water between the 30 and 100 m isobaths
- Winter monsoon affects the cross-shore transport of Taiwan Warm Current water at the latitudes 26.5 °N and 28 °N

Abstract. The seasonal mean and synoptic fluctuation of the wintertime Taiwan Warm Current (TWC) were investigated using a well validated finite volume community ocean model. The spatial distribution and dynamics of the synoptic fluctuation were highlighted. The seasonal mean of the wintertime TWC has two branches: an inshore branch between the 30 and 100 m isobaths and an offshore branch between the 100 and 200 m isobaths. The Coriolis term is much larger than the inertia term and is almost balanced by the pressure gradient term in both branches, indicating the geostrophic balance of the mean current. Two areas with significant fluctuations of the TWC were identified during wintertime. One of the areas is located to the north of Taiwan with velocities varying in the cross-shore direction. These significant cross-shore fluctuations are driven by barotropic pressure gradients associated with the intrusion of the Taiwan Strait Current (TSC). When a ~~larger~~ strong TSC intrudes~~–~~ to north of Taiwan, the isobaric slope tilts downward from south to north, leading to a cross-shore current from the coastal area to the offshore area. When the TSC intrusion is weak, the cross-shore current to the north of Taiwan is directed from offshore to inshore. The other area of significant fluctuation is located in the inshore area, extending in the region between the 30 and 100 m isobaths. The fluctuations are generally strong in the alongshore direction, in particular at the latitudes 26.5 °N and 28 °N where they are important for the local cross-shore transports. Wind affects the synoptic fluctuation through episodic events. When the northeasterly monsoon prevails, the southward Zhe-Min Coastal Current dominates the inshore area associated with a deepening of the mixed layer. When the winter monsoon is weakened or the southerly wind prevails, the northward TWC dominates in the inshore area.

Keywords:

Synoptic fluctuation, East China Sea, Taiwan Warm Current, Taiwan Strait Current, Kuroshio

1 Introduction

On the East China Sea (ECS) shelf, the mean path of the Taiwan Warm Current (TWC) has two branches: the inshore branch along the 50 m isobath and the offshore branch along the 100 m isobath (Su and Pan, 1987). The summer TWC has been well studied because the current is stationary and strong, with an average speed of 0.3 m/s (Guan, 1978; Fang et al., 1991; Isobe, 2008; Yang et al., 2011, 2012). The spatial structure and temporal variation of the wintertime (December to March) TWC are less known due to its surface weak mean velocity, according to a climatological structure of the surface current in the ECS mapped by Qiu and Imasato (1990) ~~superimposed by pronounced small-scale spatial and synoptic temporal variations~~.

The wintertime TWC on the ECS shelf shows synoptic fluctuations (Cui et al., 2004; Zhu et al., 2004; Zeng et al., 2012; Huang et al., 2016). These synoptic fluctuations show some features common with those over other continental shelves, i.e., they have periods between 3 and 15 days and are associated with coastal sea level changes, which can be explained by local winds or by coastal trapped waves (Huyer, 1990; Brink, 1991; Huthnance et al., 1986). Huang et al. (2016) have shown that the wind was a main physical factor which caused the temporal variation of the wintertime currents at the synoptic scale in the coastal area of the ECS. However, the dominant physical factors of the TWC fluctuations are still lack of study, regarding that the fluctuations on the whole shelf of the ECS may be more complicated due to the complex bottom topography, alternating monsoon forcing and conjunction of several current systems such as the Kuroshio Current, the Taiwan Strait Current (TSC) and the Zhe-Min Coastal Current

(ZMCC). These synoptic fluctuations are also known to influence the regional material transport, especially when the amplitude of the fluctuations is comparable to, or even larger than, the mean current. On the ECS shelf, some recent observations have shown that the TWC has an episodic wintertime feature (Zhu et al., 2004) and the intermittency of the TWC in winter has an amplitude ~~reaches maximum velocity variations~~as large~~r~~ as~~than~~ 0.2 m/s (~~Zhu et al., 2004;~~ Zeng et al., 2012). Moreover, it has been observed that the intermittency of the TWC in winter causes a cross-~~shelf~~ shore current which is closely linked to the alongshore component (Huang et al., 2016). Therefore, we focus on studying the spatial patterns of synoptic fluctuations to better understand the role of the wintertime TWC on the cross-~~shelf~~ shore water exchange.

A comparison between the wintertime climatological density (Fig. 1a) and synoptic density distributions observed during two surveys (Figs. 1b and 1c) suggests that two distinct areas with significant synoptic fluctuations exist. The climatological density is taken from the Generalized Digital Environment Model (GDEM, Carnes, 2009) data, and the two surveys were carried out in February 2007 by two research vessels. Because the isopycnal lines are closely related to geostrophic currents, we can infer the strength of the TWC from the horizontal gradient of the isopycnals between 24-σ_t and 25-σ_t contours (Fig. 1a). This accounts for the fact that in winter the water mass of TWC is located in this density range [according to the hydrography analysis of Su et al. (1994)]. The two-branch structure of the TWC can be inferred from the wintertime climatological density. In this paper, we defined that the near-coast area is the area between the coast and 30 m isobath where the ~~Zhe-Min Coastal Current (~~ZMCC~~)~~ occurs; the inshore area is the area between the 30 and 100 m isobaths where the TWC inshore branch dominates; and the offshore area is the region between the 100 and 200 m isobaths where the TWC offshore branch prevails.

According to the hydrographic data analysis and numerical interpretation by Su and Pan (1987), the TWC inshore and offshore branches mainly occur in those specific isobaths. However, these two branches were missing during the two synoptic surveys (Figs. 1b and 1c), indicating strong synoptic fluctuations of the TWC on the ECS shelf. Furthermore, the density anomalies between the two surveys and the GDEM data (Figs. 1d and 1e) indicate that the most significant fluctuations are located north of Taiwan and in the inshore area. Both surveys show negative density anomalies north of Taiwan, indicating that the TWC was weak and that less low-density coastal water was transported to the ECS shelf during the observational periods. The density anomalies in the inshore area show different patterns for the two synoptic surveys, with a positive anomaly in the first survey (Fig. 1d) and a negative anomaly in the second (Fig. 1e), indicating a much stronger synoptic fluctuation in the inshore area.

Figure 1

Candidate factors for driving these synoptic fluctuations are local wind, surface cooling, and the upstream currents of the Kuroshio Current and the ~~Taiwan Strait Current (~~TSC~~)~~. As discussed by Huyer (1990), wind is often considered as the major driving mechanism of synoptic fluctuations of the wintertime TWC. The northeasterly monsoon wind in winter blows against the northeastward TWC and produces a southwestward ZMCC (Chuang and Liang, 1994; Oey et al., 2010). Zhu et al. (2004) suggested that the occurrence and duration of the TWC are associated with the meandering of the Kuroshio Current north of Taiwan. The northeastward TSC, as an upstream flow of the TWC, also influences the synoptic fluctuation of the wintertime TWC. Hong et al. (2011) and Hu et al. (2010) summarized that the temporal and spatial variation of TSC is modulated by strong monsoon forcing, complex topography and circulation in the northern South China Sea as well as coastal water input and the Kuroshhio intrusion. Guan and Fang (2006) showed evidence that the TSC and the TWC merge in the area between the Taiwan Strait and the Zhe-Min coastal region. Takahashi and Morimoto (2013) pointed out that the ~~annual~~ temporal variation of the TWC is characterized by the propagation of vorticity anomalies originating from northeast of the Taiwan Strait, which further demonstrated that the fluctuations of TWC was associated with its upstream currents such as the TSC.

To explore the spatial distribution of synoptic fluctuations of the wintertime TWC on the ECS shelf, current data with high resolution in both space and time are required. Previous studies on the wintertime TWC were based on cruise surveys (Su and Pan, 1987; Chen et al., 1994; Chen and Wang, 1999), anchored mooring observations (Zhu et al., 2004; Zeng et al., 2012; Huang et al., 2016) and numerical simulations (Guo et al., 2003, 2006; Yang et al., 2011, 2012; Xuan et al., 2012, 2016). The observation data are limited in terms of temporal and spatial coverage; hence, they cannot fully reveal the synoptic fluctuations of the TWC and their regional differences. Numerical simulations provide a promising approach for studying the overall structure and driving mechanisms of synoptic fluctuations of the TWC in more detail.

In this study, the Finite Volume Coastal Ocean Model (FVCOM; Chen et al., 2003) is used to investigate synoptic fluctuations and their mechanisms of the wintertime TWC. The rest of this paper is organized as follows. In Sect. 2, we provide a description of methods and validation. The mean distribution, synoptic fluctuations, and dynamic diagnostics of the wintertime TWC are given in Sect. 3. The impact of synoptic fluctuation on water exchange is further discussed in Sect. 4, followed by conclusions in Sect.

5.

2 Methods and validation

2.1 Model configuration

To investigate the currents (TWC, Kuroshio Current, ZMCC, etc.) and their synoptic fluctuations on the ECS shelf, a 3-D~~n~~ unstructured-grid (Fig. 2, left panel) FVCOM i~~wa~~s developed for the entire Bohai, Yellow, and East China Seas (part of the Japan/East Sea, and part of the Pacific Ocean). A regional refinement of the resolution (approximately 3 km) ~~was~~ is specified around the ECS shelf break at the 200 m isobaths, where a strong excursion of the Kuroshio Current also occurs. The General Bathymetric Chart of the Oceans (GEBCO)~~(Smith and Sandwell, 1997)~~ provides high-resolution (approximately 1 km) bathymetric data (Smith and Sandwell, 1997). Twenty vertical layers with 76954 triangle cells were specified in the water column in a sigma-stretched coordinate system.~~ A detailed description of model configuration can be found in Xuan et al. (2016).~~

The driving forces of the numerical simulation include tides, river discharge, surface heat fluxes, wind, and open boundary conditions. Harmonic constants of 11 major tidal constituents (M_2, S_2, N_2, K_2, K_1, O_1, P_1, Q_1, M_4, MS_4, and MN_4) were used; these are based on the Oregon State University global inverse tidal model TPXO.7.0 (Egbert et al., 1994; Egbert and Erofeeva, 2002). The daily-mean river discharge of the Changjiang and Huanghe were taken from publicly available observation data at the Datong hydrometric station (http://yu-zhu.vicp.net/). Other rivers were not included because of their small discharge~~ rate~~s, e.g., the Qiantang River, with the largest runoff from the Zhejiang coast, has a climatological mean discharge ~~rate~~ in winter of about 230 m^3/s, which is nearly negligible compared to the Changjiang winter discharge of about 11500 m^3/s. The daily-mean heat fluxes were from the objectively analyzed air–sea fluxes (Yu and Weller, 2007), and the 3-hourly wind stress and 10 m wind speed data was from the ERA-40 re-analysis (Uppala et al., 2005). The open boundary conditions, including daily temperature, salinity, and velocities at the Taiwan Strait, the western Pacific Ocean, and the Japan/East Sea, were obtained from the Hybrid Coordinate Ocean Model (Bleck, 2002) and interpolated onto the FVCOM model grid points. The temporal resolution of all the driving force fields is better than or equal to one day, which is essential to resolve synoptic fluctuations.

The hindcast outputs of sea surface height, temperature, salinity, and velocities for the five years of simulation from 2009 to 2013 are used, following three spin-up years (2006-2008) initiated with the temperature and salinity taken from the Hybrid Coordinate Ocean Model and velocity setted to zero~~following three years of spin up from the initial climatological conditions~~. The initial conditions are ramped-up over a period of 30 days and that at the lateral boundaries sponge layer was used with the same method as Chen et al. (2008). The model time step was ~~90~~ 15 seconds for the 2-D barotropic mode of and 90 seconds for the 3-D baroclinic mode. All of the output fields were processed with a tidal filter (Godin, 1972) to remove tidal oscillations (considering that the major time scale of synoptic fluctuations in this study area is 3–15 days).

Since the currents in 2009 could partly be validated by means of available observational data (see Sect. 2.2), the currents from January 1 to February 28, 2009 were selected for analysis of the wintertime TWC.

2.2 Validation of the mean currents and synoptic fluctuations

The mean currents, e.g., the Kuroshio Current, the TWC, and the ZMCC, ~~was~~ were calculated by averaging the outputs of January and February 2009. We validated the mean currents~~flow~~ in terms of ~~overall~~ circulation structure, boundary fluxes, and coastal currents.

The FVCOM has reproduced almost all of the known circulation~~overall~~ structure in the ECS in winter. T~~of t~~he surface mean currents~~flow~~ (Fig. 2) shows three major currents~~ in the ECS in winter~~: the Kuroshio Current, the TWC, and the ZMCC. The Kuroshio Current, with a speed of about 1 m/s, enters the ECS just northeast of Taiwan and flows along the shelf break up to the northern area and ultimately leaves the ECS through the Tokara Strait. Both the route and strength of the Kuroshio are comparable with those reported in the literature (Guan, 1978; Qiu and Imasato, 1990). The TWC has two northward branches, one inshore (between the 30 and 100 m isobaths) and another offshore (between the 100 and 200 m isobaths), which is consistent with Su and Pan (1987). The southward directed ZMCC in the nearshore area from the Changjiang Estuary to the Taiwan Strait agrees well with that reported in previous studies (Guan and Mao, 1982; Zeng et al., 2012).

The simulated volume transports across the Taiwan Strait, the East Taiwan Channel, the Tsushima Strait, the Tokara Strait, and the shelf break of the 200 m isobath were validated using results from the literature (Table 1). The simulated transports were accurate enough to reproduce volume transport (1.22 Sv) through the Taiwan Strait which is closer to the observation value (1.20 Sv) from Isobe (2008) than former model results. The volume transports across the Taiwan Strait and the Tokara Strait, and the cross-~~shelf~~ shore exchange, affected the path and magnitude of the TWC. The annual mean transport across the 200 m isobath toward the shelf is 1.66 Sv, which is balanced by the inflow from the Taiwan Strait (1.22 Sv) and the outflow through the Tsushima Strait (2.85 Sv).

Figure 2

Table 1

Figure 3 shows a comparison between simulation and observation results for the alongshore currents and the cross-shore currents on the ECS shelf. The observational data were obtained from four mooring surveys (Fig. 2, red stations) off the Zhe-Min coast (Zeng et al., 2012). The observed and simulated currents were both averaged for the observational period, which was from January 1 to February 28, 2009. Using the same method as in Huang et al. (2016), ~~W~~we defined the positive alongshore current directing~~ion~~ from southwest (218°) to northeast (38°), which is the mean tangential direction of the isobaths on the southwestern shelf of the ECS. The positive cross-~~shelf~~ shore direction is the mean normal direction of the isobaths from northwest (308°) to southeast (128°). The alongshore components (Figs. 3a and 3b) show that the ZMCC flows southwestward parallel to the coast in winter, with a maximum speed of 0.15 m/s along the 30 m isobath. The TWC flows northeastward with a speed of 0.05 m/s, and the core is located in the lower layer at about 50 m at Station 4. The cross-shore component (Figs. 3c and 3d) is much weaker than the alongshore components, and it shows a complex spatial pattern. It flows offshore in the upper layer and onshore in the lower layer at Station 1. Moreover, it mainly flows onshore at Station 2, and it flows offshore in the entire water column at Stations 3 and 4. Altogether, the simulated pattern and magnitude both of the alongshore and cross-shore components are in good agreement with the observations. However, there are some differences between the observed and simulated results; for example, the simulated ZMCC occupies a broader space than that in the observations. This may have been caused by the relatively low number of observational stations.

Figure 3

Synoptic fluctuations of the ~~wintertime~~ TWC inshore branch during January and February 2009 were also validated against the mooring results (Fig. 4). Since the TWC shows a strong signature at Station 4, the time series of the alongshore currents and cross-shore currents in the whole water column of Station 4 were used for the validation. To eliminate the influence of local effects, the simulated currents were averaged in a 10 × 10 km^2 area around Station 4. Both the observed and simulated results show that the TWC fluctuates with a period of 3–15 days. The simulated TWC (Fig. 4a, warm color) appeared stronger (> 0.1 m/s) on Jan. 7, Jan. 12, Jan. 18, Jan. 21, Jan. 26, Jan. 29, Feb. 10, Feb. 14, Feb. 19, Feb. 22, and Feb. 25, which agrees well with data from the observations (Fig. 4b). The time series of the simulated cross-shore component (Fig. 4c) are virtually in phase with the observations (Fig. 4d). In contrast to the anisotropic feature for the mean currents (Fig. 3), i.e., that the alongshore component is nearly one order of magnitude larger than the cross-shore component in the mean condition, the magnitude of the cross-shore fluctuations is comparable to the alongshore fluctuations.

Figure 4

2.3 EOF analysis of synoptic fluctuations

The Empirical Orthogonal Function (EOF) method (Emery and Thomson, 2001), as a statistical method, has been~~was~~ used to understand~~analyze~~ synoptic fluctuations of the wintertime TWC. The simulated currents from Jan. 1 to Feb. 28, 2009 were selected and their anomalies were calculated. Then, using the Matlab EOF-function, the current vectors were separated into several orthogonal modes to show the spatial and temporal variations. Because the first two leading modes explain 91 % of the total variance, only these two modes were used for the analysis.

The spatial distributions of the two leading EOF modes were used to analyze the regional difference of the synoptic fluctuations. To investigate the driving force of the two EOF modes, the temporal variation was compared to the potential influence factors, such as wind, upstream currents, and net surface heat flux.

2.4 Momentum analysis

The driving mechanisms of the synoptic fluctuations were further analyzed using the momentum equation. First, the momentum balance as implemented in FVCOM (Chen et al., 2003) is shown in Eq. (1). The three terms on the left hand side represent local acceleration, Coriolis acceleration, and advection, respectively, and the three terms on the right hand side represent pressure gradient, friction, and diffusion, respectively.

$$\frac{\partial \vec{V}}{\partial t} - 2\vec{\Omega} \times \vec{V} + (V \bullet \nabla \vec{V}) = -\frac{1}{\rho_0}\nabla P + \frac{\partial}{\partial z}\left(K_m \frac{\partial \vec{V}}{\partial z}\right) + \vec{F}\,, \qquad (1)$$

where $\vec{V}$ ~~V~~ is velocity, $\vec{\Omega}$ ~~Ω~~ is the Earth's rotation angular velocity, ρ_0 ~~ρ_0~~ is the average density, P is pressure, K_m is the vertical eddy viscosity coefficient, and $\vec{F}$ ~~F~~ is horizontal diffusion.

Second, according to the hydrostatic approximation~~assumption~~ used in FVCOM [as shown in Eq. (2)], the pressure gradient is given as the product of density times the gravitational acceleration~~integrated from depth z to the sea surface~~. This results in Eq. (3), which indicates that pressure gradient can be decomposed into the effects of the barotropic and baroclinic components, as shown in Eq. (4).

$$\frac{\partial P}{\partial z} = \rho g \,, \tag{2}$$

$$P_z = \int_z^{\eta} \rho g dz = \int_z^{\eta} (\rho_0 + \rho') g dz = \rho_0 g (z + \eta) + \int_z^{\eta} \rho' g dz \,, \tag{3}$$

$$\nabla \vec{P} = \rho_0 g \nabla \eta + \nabla \left(\int_z^{\eta} \rho' g dz \right), \tag{4}$$

where ρ ~~ρ is~~ is density, ρ' ~~ρ'~~ is density anomaly, g is the gravitational acceleration, and η ~~η~~ is sea surface height.

Finally, the momentum equation is vertically integrated to estimate momentum balance for the water column. Since the horizontal diffusion is a comparably small term, it is neglected for simplicity.

$$\underbrace{\int_{-H}^{0} \frac{\partial \vec{V}}{\partial t}}_{Acceleration} + \underbrace{\int_{-H}^{0} -2\vec{\Omega} \times \vec{V}}_{Coriolis} + \underbrace{\int_{-H}^{0} (V \bullet \nabla \vec{V})}_{Advection} = \underbrace{\underbrace{-gH\nabla\eta}_{Barotropic} - \underbrace{\int_{-H}^{0} \nabla \left(\int_z^{\eta} \rho' g dz \right)}_{Baroclinic}}_{Total\ Pressure} + \underbrace{\rho_a C_D \left|\vec{U}\right| \vec{U}}_{\tau_a} \underbrace{- k_b \left|\vec{U_b}\right| \vec{U_b}}_{\tau_b}$$

$$\underbrace{\int_{-H}^{0} \frac{\partial \vec{V}}{\partial t}}_{Acceleration} + \underbrace{\int_{-H}^{0} -2\vec{\Omega} \times \vec{V}}_{Coriolis} + \underbrace{\int_{-H}^{0} (V \bullet \nabla \vec{V})}_{Advection} = \underbrace{\underbrace{-gH\nabla\eta}_{Barotropic} - \underbrace{\int_{-H}^{0} \nabla \left(\int_z^{\eta} \rho' g dz \right)}_{Baroclinic}}_{Total\ Pressure} + \underset{Wind}{\tau_a} - \underset{Bot}{\tau_b} \,, \tag{5}$$

where τ_a ~~τ_a~~ is wind stress and τ_b ~~τ_b~~ is bottom stress, ρ_a is the density of air, $\vec{U}$ is the wind speed at 10 m above sea surface, C_D is a drag coefficient at the sea surface (which varies with wind speed $\vec{U}$), k_b is a bottom friction coefficient (k_b =0.005), and $\vec{U_b}$ is the simulated velocity at the bottom.

3 Results

3.1 Mean distribution of TWC in winter

Since the observational results (Su and Pan, 1987; Zeng et al., 2012) show that both branches of the wintertime TWC are flowing in the subsurface, we use the vertical maximum velocity (VMV) and its corresponding depth as two indices to quantify the strength of the subsurface currents (Fig. 5).

As stated above, the distribution of the VMV shows two branches of the TWC (Fig. 5a). The inshore branch (Fig. 5a, blue arrow of IB), which was located between the 30 and 100 m isobaths, followed a straight route from the northwest of Taiwan to the northern ECS shelf. The offshore branch (Fig. 5a, blue arrow of OB) existed near the 100 m isobath and had a horizontal structure with two meanders. The two meanders turn to the cross-shore direction along latitudes 26.5 °N and 28 °N. These two branches are further illustrated in the distributions of current speed along the six cross-TWC sections (S1-S6), which were located at critical points in the two meanders (Fig. 6). From the VMV structure, it can be inferred that the intrusions of the TSC and the Kuroshio Current both affected the origin of the offshore branch (Fig. 6, S1–S3). ~~In wintertime, both branches flowed along the isobaths, which is in accordance with the conservation of potential vorticity under frictionless conditions and for flows with a minor meridional extension.~~

We further examined the subsurface current core using the depth of the VMV (Fig. 5b). We found that the VMV of the TWC was located 40–60 m below the surface at the inshore branch and 20–40 m below the surface at the offshore branch. Figure 6 shows the VMV positions in the subsurface layer; it also illustrates that the depth of the subsurface VMV in the inshore branch was deeper than that in the offshore branch. The northerly wind in winter weakens the northward TWC, particularly in the upper layer, which leads to the formation of the subsurface VMV. Assuming a relatively spatially homogeneous heat loss, a d~~D~~ifferent cooling occurs, due to the smaller heat capacity of the shallow ~~makes the~~ coastal ~~shallow~~ water compared to~~cooler than~~ the deeper offshore~~deep~~ waters; hence generating a horizontal density gradient leading to a southeastward vertical current shear according to the thermal wind relationship, resulting in an increasing southwestward flow component from surface to bottom, which in turn weakens the northeastward flow of the TWC inshore branch. Therefore, ~~T~~the fact that the depth of the subsurface ~~VMV~~ current core in the inshore branch is greater than that in the offshore branch indicates that the effects of baroclinicity and wind friction on the ~~TWC in the coastal~~ inshore ~~area~~ branch are stronger than ~~those in~~ the offshore branch~~area~~.

The magnitude of the wintertime TWC was obtained by flux analysis. Two dividing lines (Fig. 5a, red lines) were defined as the boundaries for the ZMCC, the TWC inshore branch, and the TWC offshore branch, which had the weakest flows. The flux of each branch (Fig. 5c) was calculated using the horizontal integration between the boundaries and the vertical integration in the water column. The inshore branch intensifies along its way and becomes significant north of 26.5 °N, showing particularly strong flow velocities between 27.5 and 28.0 °N. In this area, the subsurface current was much stronger from S4 to S5 than in the other areas (Fig. 6). The flux in the entire offshore branch was large, particularly north of Taiwan.

Figure 5

Figure 6

3.2 Synoptic fluctuations

The observations (Fig. 4) have demonstrated that the synoptic fluctuation in the TWC inshore branch (near 121.5 °E, 27.0 °N) is significant. We further investigated the regional difference of fluctuations in the two TWC branches in winter 2009 using the following three steps: (i) two regions with significant fluctuations are identified by the current standard deviations of the VMV (Fig. 7) and the corresponding temporal variation of vertical structures at their extremes (Fig. 8); (ii) each of the two significant fluctuations is decomposed into EOF components (Fig. 9), and (iii) the influence factors, such as wind, upstream currents, and net surface heat flux, are investigated by examining their correlations with the first two leading EOF components (Figs. 10 and 11).

The current standard deviations (Fig. 7) shows that prominent fluctuations occurred in two regions: north of Taiwan and the inshore area. The standard deviations of VMV at the two regions were larger than 0.1 m/s (comparable to the mean currents). In the area north of Taiwan, the fluctuation was located in the origin area of the TWC offshore branch. The fluctuation in this region was in phase with the fluctuation in the Taiwan Strait, indicating that the TSC played an important role in generating the fluctuation north of Taiwan (to a greater extent than did the Kuroshio intrusion). The TWC fluctuation had a strong magnitude in the cross-shore direction, which means the fluctuation transported the water north of

Taiwan to both the inshore and offshore branches. In the inshore area, the fluctuation was located in a wide region between the 30 and 100 m isobaths, where the southwestward flowing ZMCC and the northeastward directed TWC meet. As deduced from the standard deviation, ~~T~~the currents~~-~~ fluctuated significantly in the alongshore direction, indicating that the TWC inshore branch occurred episodically.

Figure 7

The vertical structures of the fluctuations north of Taiwan and in the inshore area at two representative points and their relation with upper mixed layer depth are further analyzed (Fig. 8). The major component (the alongshore current) of the TWC in each of the two regions (P1 and P2, Fig. 7) is used to show the vertical structure of the fluctuation. The depths of the upper mixed layer were determined by a Richardson number criterion (Mellor and Durbin, 1975; Grachev et al., 2013; Richardson et al., 2013) ~~[as in Xuan et al. (2012)]~~, where the critical Richardson number equals 0.25 in this paper [as in Xuan et al. (2012)]. The mean depth of the upper mixed layer north of Taiwan (20 m) was much shallower than the mean depth in the inshore area (42 m). However, the TWC (Fig. 8, warm color) fluctuated with significant variations of the upper mixed layer depth (Fig. 8, gray lines) in both areas. When the upper mixed layer deepened, the northeastward TWC (Fig. 8, warm color) was weakened or even replaced by the southwestward ZMCC, and vice versa. Hence, wind and surface cooling, which both drive the mixed layer depth, can affect the TWC fluctuation.

Figure 8

The TWC fluctuations were further decomposed into EOF modes. The first two leading EOF modes account 54% and 37% of the total variances (Fig. 9), which were associated with the two prominent fluctuations north of Taiwan and in the inshore area (Fig. 7). Both EOF modes had a maximum fluctuation larger than 0.2 m/s (comparable to the mean currents). The spatial pattern of the first EOF mode (EOF1, Fig. 9a) shows that the fluctuation continued from the Taiwan Strait to the area north of Taiwan, indicating that the fluctuation north of Taiwan was related to the TSC and not to the Kuroshio Current. The alongshore component also showed a strong fluctuation in the Taiwan Strait, which means that the TSC episodically intruded the shelf. The cross-~~shelf~~ shore component revealed a great fluctuation north of Taiwan that was larger than 0.1 m/s. This cross-shore fluctuation impacted on the trajectory of the TWS water, synoptically flowing into the TWC inshore branch, offshore branch, or Kuroshio Current.

The spatial pattern of the second EOF mode (EOF2, Fig. 9b) shows a synoptic fluctuation in the inshore area. The fluctuation mainly varied in the alongshore direction, which indicates the episodic occurrence of the TWC inshore branch. The area with alongshore fluctuation (Fig. 9d) larger than 0.1 m/s was located between the 30 and 100 m isobaths, which demonstrates that the TWC could also episodically affect this area. In addition, there were cross-shore fluctuations in the inshore area (Fig. 9f), mostly along the latitudes 26.5°N and 28°N. The latitudes of great cross-shore fluctuations agreed well with the latitudes where the TWC offshore branch of the mean currents (Fig. 5a) turned to the cross-shore direction. This indicated that the cross-shore transports were most significant at the latitudes 26.5°N and 28°N, according to both the mean currents and the synoptic fluctuations.

Figure 10 shows the temporal variation of EOF1 and its relation with north-south component of wind speed, net surface heat flux, the TSC, and the Kuroshio Current. We found a close correlation between EOF1 and TSC (R = 0.86), demonstrating that the TSC played the most important role in generating the TWC fluctuation north of Taiwan. The EOF1 and TSC were positively correlated, meaning that a larger TSC intrusion north of Taiwan leads to a cross-shore current from the coastal area to the offshore area and that a weak TSC intrusion causes a cross-shore current from offshore to inshore north of Taiwan.

Figure 11 shows the temporal variation of EOF2 and its relation with the north-south component of wind speed, net surface heat flux, the TSC, and the Kuroshio Current. It can be seen that EOF2 and wind are well correlated (R = 0.89), indicating the important role of wind in generating the TWC fluctuation in the inshore area. The northerly monsoon would greatly enhance the southwestward ZMCC, which replaces the northeastward TWC in the inshore area. Together with the effect of net surface heat flux, the stronger northerly monsoon during Jan. 5-13, Jan. 19-25 and Feb. 16-18 causes the deepening of the mixed layer (P2, Fig. 8).

Figure 9

Figure 10

Figure 11

3.3 Dynamic diagnostics

The ~~seasonal~~ wintertime (January and February 2009) mean of the water column momentum balance (Fig. 12) is used to show the overall distribution of the fundamental forces over the ECS shelf. The Coriolis force (Fig. 12a) is mainly balanced by the total pressure (Fig. 12b) in both branches, indicating the dominant role of geostrophic balance in the wintertime TWC. ~~–~~However, the wind-induced surface friction plays an important role in the TWC, especially in the inshore area and the Taiwan Strait (Fig. 12c). The bottom friction has an ~~an~~ impact north of Taiwan and in the shallow Taiwan Strait, in particular when ~~under the condition of~~ significant Kuroshio intrusion enhances the bottom flow (Fig. 12d). The effects of advection and acceleration are predominantly local indicated by mostly incoherent small scale distributions (Figs. 12e, 12f), so they can be ignored when studying the large-scale current of the wintertime TWC.

Figure 12

The variation of the driving forces at two representative points P1 and P2 were used to analyze the dynamics of synoptic fluctuations north of Taiwan and in the inshore area. Regarding the results from the EOF analysis, the three force terms, namely Coriolis, total pressure, and wind (Fig. 13), were selected to investigate the effect of the TSC on the fluctuation north of Taiwan (Fig. 9a) and the effect of wind on the fluctuation in the inshore area (Fig. 9b).

In the area north of Taiwan, the cross-shore fluctuations were induced by the TSC intrusion. The variation of alongshore Coriolis force (Fig. 13a, black line) was much greater than the cross-shore Coriolis force

(Fig. 13b, black line), which means that the fluctuation north of Taiwan was mainly in the cross-shore direction. The Coriolis force (Fig. 13a, black line) was mainly balanced by the total pressure (Fig. 13a, blue line), which means the currents fluctuations north of Taiwan are dominated by geostrophic balance. As mentioned in Sect. 3.2, the TWC fluctuation north of Taiwan was associated with the TSC rather than with the Kuroshio Current. Therefore, in the shallow coastal area the TSC mainly caused variations in the depth-independent barotropic pressure gradients, which further generated the cross-shore fluctuation. The mechanism can be interpreted as follows. When a larger TSC intrusion occurred, the isobaric slope tilted downward from south to north, generating a cross-shore current from the coastal area to the offshore area. On the contrary, when the TSC intrusion was weak, the Kuroshio intrusion from offshore to inshore dominated north of Taiwan.

Wind friction (Figs. 13c and 13d) was a fundamental factor in generating the fluctuations in the inshore area. Although the geostrophic balance dominated in the inshore branch for most of the time, the episodically strong winter monsoon had an important role in generating the TWC fluctuations. The ~~negative~~northwestward direction Coriolis force (Fig. 13c, black line) shows that the southwestward ZMCC occurred on Jan. 12, Jan. 22, and Feb. 14, 2009 and was associated with a northerly wind (Fig. 13c, red line). It indicates that strong ~~winter~~ northerly monsoon in winter can ~~weaken~~ reduce or even stop the northeastward TWC in the inshore area, causing the intermittency of the TWC inshore branch.

Figure 13

4 Discussion

The simulated results in the winters of the years 2010 to 2013 (Fig. 14) show that general structures of the TWC in the other winters were similar to that in winter 2009 (Fig. 5 and Fig. 9), which indicates that the results from the winter 2009 can be regarded as representative for the winter situation. The two TWC branches and the two areas of strong fluctuations were presented in the all ~~the~~ winters from of 2009 to 2013, although their strength showed a certain inter- annual variability in accordance with the changing surface forcing and boundary fluxes.

Figure 14

The wintertime TWC, which is manifested by with two subsurface branches and significant synoptic fluctuations, has a very different structure when compared with the stationary and surface summertime TWC reported in previous studies (Guan, 1978; Fang et al., 1991; Isobe, 2008). The synoptic events, with time scales of 3-15 days, play a dominant role on the horizontal advective transports. According to Ledwell et al. (1998) synoptic variations are much more effective on the horizontal transport than variations on shorter time scales. The synoptic fluctuations modulate the spatial structure of the wintertime TWC, especially when their magnitudes are comparable with that of the mean currents, such as the two prominent fluctuations north of Taiwan and in the inshore area (Fig. 7). Therefore, the two prominent fluctuations will be discussed next in terms of their contributions to the alongshore and cross-shore transports.

4.1 Cross-shore transport north of Taiwan induced by the TSC

In the area north of Taiwan, the TSC intrusion generated strong fluctuations of the TWC in the cross-shore direction (Fig. 9a). When a larger TSC intrusion occurred, the isobaric slope tilted downward from south to north, generating a cross-shore current from the coastal area to the offshore area. Compared to the reported summer route that transports Taiwan Strait water to the inshore area between the 30 and 100 m isobaths (Guan, 1978; Fang et al., 1991; Isobe, 2008; Yang et al., 2011, 2012), our results showed that most Taiwan Strait water was transported to the TWC offshore branch and to the Kuroshio area as a result of the cross-shore fluctuations induced by the synoptic TSC intrusion.

A numerical tracer simulation was used to analyze the role of the cross-shore fluctuation in the transport of the TSC water and the Kuroshio water north of Taiwan. In order to demonstrate the characteristics of the flow patterns more clearly, artificial tracers are released in the model domain and transported by the velocity field provided by the FVCOM simulation. The tracer running was part of the FVCOM simulation; therefore, all the above mentioned dynamics were involved, e.g., tide, wind, and boundary forces. The release location and start date of the particles were configured as follows. Two sections, one in the Taiwan Strait (Fig. 15a, black dots) and another in the East Taiwan Channel (Fig. 15b, black dots), were selected as the source locations for the water masses of the TSC and the Kuroshio, respectively. The particles were released on January 1, 2009 and tracked until March 31, 2009 (a total of 90 days). ~~The tracer assimilation was part of the FVCOM simulation; therefore, all the above mentioned dynamics were~~

~~involved, e.g., tide, wind, and boundary forces.~~

Figure 15a shows the trace~~r~~s originating from the TSC area. Unlike the traditional route, where the TSC water flows from the Taiwan Strait to the inshore area between the 30 and 100 m isobaths, most particles (Fig. 15a, gray lines) were concentrated in the offshore branch under the effect of cross-shore fluctuation. Two particles were selected to show the inshore route (Fig. 15a, red line) and offshore route (Fig. 15a, blue line), with both passing the area north of Taiwan. When the two particles arrived at the area north of Taiwan, the behavior of the tracers, according to specific velocity conditions (Fig. 15c), was very different: a northwestward transport occurred on Jan. 25 for the inshore particles (Fig. 15c) and a northeastward transport occurred on Feb. 12 for the offshore particles (Fig. 15c). The velocity conditions in the area north of Taiwan corresponded to the variation of the Taiwan Strait flux (Fig. 10), which shows that the Taiwan Strait flux on Feb. 12 was much greater than on Jan. 25. Therefore, it can be concluded that the TSC intrusion induced an offshore transport north of Taiwan.

Figure 15b shows the trace~~r~~s originating from the Kuroshio area. In the same way as the TSC water, the Kuroshio water was also transported to the northern shelf via both the inshore branch and the offshore branch. The separation of the two branches north of Taiwan was caused by cross-shore fluctuations of the currents. When the two particles arrived at the area north of Taiwan, a northwestward transport occurred on Feb. 2 for the inshore particles (Fig. 15c) and a northeastward transport occurred on Feb. 12 for the offshore particles (Fig. 15c). This means that the offshore transport induced by the TSC also had an effect on the distribution of Kuroshio water north of Taiwan. Liu et al. (2016) showed that the winter TSC originated from a small branch of Kuroshio intrusion into the Luzon Strait. Our results complement this picture, since they show that most TSC particles flow into the TWC offshore branch under the influence of cross-shore fluctuation.

Figure 15

Our results may underestimate the impact of Kuroshio intrusion on the fluctuation of the TWC northeast of Taiwan, especially at the seasonal and interannual time scales. Wei et al. (2013) demonstrated that the annual and interannual variations of the Kuroshio volume transport are large. In addition, Zhou et al. (2015) pointed out that the annual and interannual variations of the Kuroshio intrusion northeast of Taiwan are prominent. Liu et al. (2014b) presented supportive evidence that the Kuroshio intrusion, from east of Taiwan to the onshore area north of Taiwan, is closely related to the Kuroshio volume transport. This relation between the Kuroshio intrusion and the Kuroshio volume transport had been interpreted by Su and Pan (1987) as the β-effect because of the sudden change in topography northeast of Taiwan. Our results show that the intra-seasonal variation of the Kuroshio intrusion and the Kuroshio volume transport was negligible compared with the TSC variation at the same time scale, indicating that the synoptic fluctuation of TWC north of Taiwan is mainly induced by the TSC. However, because FVCOM uses sigma co-ordinates in the vertical which are prone to errors in regions of steep topography, our results may underestimate the fluctuations at the shelf break, in particular to the northeast of Taiwan where Kuroshio intrusion occurs.

4.2 Water exchange in the inshore area induced by wind

In the inshore area, the synoptic fluctuations of the TWC (Fig. 9b) caused by wind were generally strong in the alongshore direction and regionally important (along the latitudes 26.5 °N and 28 °N) in the cross-shore direction. The alongshore fluctuations showed that the TWC inshore branch occurred episodically. This episodic occurrence of the TWC agrees with the results from a previous study based on four mooring surveys off the Zhe-Min coast (Zeng et al., 2012). The mechanism of the episodic occurrence of the TWC was mainly associated with the winter monsoon, which agrees with the analysis of observational data by Huang et al. (2016). However, the overall magnitude of the TWC fluctuation, and its role on the cross-shore flux, are still not fully understood due to the short-term nature of the observational data.

We investigated the magnitude of TWC fluctuation, and its role on the water exchange, in the inshore area. Previous studies (Su and Pan, 1987; Zeng et al., 2012) show that the TWC flows between the 50 and 100 m isobaths, whereas the ZMCC water dominates the coastal area west of the 50 m isobath in the surface layer. As mentioned when discussing Figure 9d, the strongest TWC could reach the coastal area as close as the 30 m isobath, being stronger than those reported in the literature. Moreover, the area with large fluctuations spanned the area between the 30 and 100 m isobaths (Fig. 9b), indicating that water exchange between the ZMCC water and the TWC water exists in the area between the 30 and 100 m isobaths.

The episodic occurrence of the TWC inshore branch is directly related to the relative importance of the southwestward ZMCC (Fig. 16, blue arrows) and the northeastward TWC (Fig. 16, red arrows). In this paper, only wind-induced synoptic fluctuations are considered in the relations to the episodic events and no short-term extreme storm events. When the winter monsoon (the northerly wind) prevails, the ZMCC occupies most of the inshore area and the TWC inshore branch weakens (Fig. 16a). On the contrary, the

TWC inshore branch can intrude into the near-coast area under southerly wind conditions (Fig. 16b). The boundary between the coastal current and the TWC may shift from the 100 m isobaths to the 30 m isobath in the cross-shore direction, covering the entire area of the TWC inshore branch.

Our results further reveal that strong wind-induced cross-shore fluctuations occur in the inshore area (Fig. 9f). This cross-shore fluctuation has a significant ecological impact because of the connected nutrient transport (Zhao and Guo, 2011). Ren et al. (2015) observed a cross-shore flux in the inshore area, which was triggered by the transition of northeasterly to southwesterly monsoonal winds. Their observed features can be further interpreted with our result that wind-induced fluctuations can affect the cross-shore water transport in the inshore area.

Largest cross-shore fluctuations were located at the latitudes 26.5 °N and 28 °N (Fig. 9f), which agreed well with the latitudes where the TWC offshore meanders occurred in the mean currents (Fig. 5a). This also indicates that the offshore transports were most significant along the latitudes 26.5 °N and 28 °N according to both the mean currents and the synoptic fluctuations. The offshore transport may be associated with the offshore-penetrating fronts of coastal water in the ECS. Many remote-sensing images (He et al. 2010; Bai et al. 2013) have exhibited offshore-penetrating fronts that crossed the 70 m isobath and played an important role in cross-~~shelf~~ shore material exchange, but the~~. Several~~ mechanisms of the offshore-penetrating fronts are still under debate~~ have been used to explain this process~~. Yuan and Qiao (2005) pointed out that both downwelling- and upwelling-favorable winds are associated with the occurrence of the offshore-penetrating front. Ren et al. (2015) suggested that the penetrating front is generated by the transition of northeasterly to southwesterly monsoonal winds. Wu (2015) suggested that the offshore-penetrating front is the response of buoyant coastal water to an along-isobath undulation of the ambient pycnocline, which is controlled by a temperature stratification of the water column. Our study offers a new interpretation, i.e., that the penetrating front is generated through the wind-induced fluctuations and the TWC offshore meanders.

Figure 16

5 Conclusions

The FVCOM model was able to reproduce the wintertime TWC in 2009 reasonably well, as shown by a validation in terms of the overall structure of the surface mean currents~~flow~~, the ECS boundary fluxes, and data from four mooring stations. The validation showed that the simulated TWC was comparable to the observed results, not only in terms of the mean currents but also in terms of the synoptic fluctuations.

The wintertime TWC showed two branches: one inshore and another offshore. The inshore branch covered an area between the 30 and 100 m isobaths and flowed northward via a straight route. The offshore branch was located between the 100 and 200 m isobaths and showed two prominent meanders. It was shown that the Coriolis force was balanced by the pressure gradient in both branches, indicating the dominant role of the geostrophic balance for the mean current in both branches.

Two regions with significant synoptic fluctuations, north of Taiwan and the inshore area, were investigated using the EOF method. The first two leading modes explained 91% of the total variance.

EOF1 showed that fluctuations occurred in the cross-shore direction south of 26°N. These fluctuations were mainly associated with variation of the TSC flux. EOF2 showed significant fluctuation between the 30 and 100 m isobaths. These fluctuations caused the episodic existence of the TWC inshore branch in the alongshore direction and cross-shore fluctuations mainly at latitudes 26.5°N and 28°N, which were mainly associated with the variation of wind speed.

We also studied the different dynamic reasons for the fluctuations in the two regions. In the area north of Taiwan, the TSC and Kuroshio converged to initiate the TWC. A barotropic pressure anomaly was generated by TSC intrusion from the Taiwan Strait causing a barotropic pressure gradient in the alongshore direction; this explains why the synoptic fluctuations in this area occurred in the cross-shore direction. Additionally, the wind had a strong effect on the synoptic fluctuations in the inshore area. The northeasterly monsoon enhanced the southwestward ZMCC and replaced the TWC in the inshore area. This situation is reversed during the southwest monsoon.

The synoptic fluctuations north of Taiwan and in the inshore area are important for both the alongshore and cross-shore transports. Due to the~~se~~ fluctuation north of Taiwan~~s~~, the mixed water of the TSC and the Kuroshio was transported to both the inshore area and the offshore area, whereas most Taiwan Strait water was transported to the offshore area in winter. The inshore fluctuation not only caused an episodic occurrence of the TWC in the alongshore direction, which affected the alongshore transport of ZMCC water and TWC water between the 30 and 100 m isobaths, but also impacted the cross-shore transports along latitudes 26.5°N and 28°N.

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Table Captions

Table 1: Annual-mean volume transports (Sv = 10^6 m^3/s) through various sections. The sections are shown in Figure 2 using blue dashed lines.

Figure Captions

Figure 1: Density (σ_t, kg/m^3) distributions at 50 m depth derived the GDEM climatological data in February (a), an ocean survey from Feb. 1–27, 2007 (b), and an ocean survey from Feb. 3–16, 2007 (c), with the density anomalies between the GDEM data and the two surveys (d and e). The two blue arrows indicate the two TWC branches in winter. The 30, 50, 70, 100 and 200 m isobaths are indicated with grey lines in panel a.

Figure 2: The FVCOM model grid (Left) and the surface mean flow in the ECS in winter (Right). The colors in the left panel show the grid length (km). The letters a, b, and c indicate the three open boundaries at the Taiwan Strait, the northwest Pacific Ocean, and the Japan/East Sea, respectively. The blue dashed lines show some important straits around shelf boundary, including the Taiwan Strait (TWS), the East Taiwan Channel (ET), the Tsushima Strait (TUS), the Tokara Strait (TOS), and shelf break at the 200 m isobath. The red rectangle shows the study area of the wintertime TWC. The four red numbers off the Zhe-Min coast shows the four mooring sites observed from Jan. 5 to Feb. 28, 2009.

Figure 3: Validations of the wintertime TWC (warm color) along the section off the Zhe-Min coast (the short line with four red numbers in Figure 2): (a) observed alongshore currents; (b) simulated alongshore currents; (c) observed cross-shore currents; (b) simulated cross-shore currents.

Figure 4: Validations of the wintertime TWC fluctuations: (a) observed alongshore currents; (b) simulated alongshore currents; (c) observed cross-shore currents; (d) simulated cross-shore currents. The

observation data comes from Station 4 in Figure 1 and the simulated data has the same position and period as the observation data.

Figure 5: a) Distribution of flow axes in the ECS in winter. The black arrows show the maximum velocity (m/s) in the vertical profile (VMV) and the color shows the speed of the VMV. The two blue arrows with label IB and OB represent the flow axes of the inshore branch and offshore branch, respectively. The red line DL1 represents the dividing line between the coastal current and inshore branch, and the red line DL2 separates the two TWC branches. b) Depth (m) of flow axes in the ECS. Sections S1–S6 were selected to study the wintertime TWC. c) Flux of inshore branch (blue) and offshore branch (red) at different latitudes. Dashed lines show the positions of Sections. S1–S6. Note, the scale is not linear.

Figure 6: Distributions of current speed along the six sections S1–S6 in winter. The blue arrow on the left indicates the inshore branch according to the velocity cores from section S3 to S6. The blue arrow on the right indicates the offshore branch according to the velocity cores from section S2 to S6. TSC is the Taiwan Strait Warm Current.

Figure 7: Current standard deviation in the layer of the VMV. The color shading shows the magnitude of the current standard deviation. The two blue arrows indicate the two TWC branches. The red curve indicate the area where the current standard deviation is larger than 0.1 m/s and their representative points (P1 and P2) are selected for later analysis.

Figure 8: Variation of alongshore currents (m/s) for the entire water column north of Taiwan (P1) and in

the inshore area (P2) and their relation with upper mixed layer depth. The positive velocity (warm color) indicates the occurrence of the TWC. The gray solid lines show the depth of the upper mixed layer.

Figure 9: The spatial pattern of the first (EOF1; left) and second (EOF2; right) leading modes of the VMV in the ECS: (a) EOF1 currents, (b) EOF2 currents, (c) EOF1 alongshore component, (d) EOF2 alongshore component, (e) EOF1 cross-shore component, and (f) EOF2 cross-shore component. The 30, 50, 70, 100 and 200 m isobaths are indicated with grey lines.

Figure 10: Temporal variation of EOF1, north-south component of wind speed, surface net heat flux, and TSC flux along the TWS section, and Kuroshio flux along the ET section. Their linear correlation coefficients R and time-lags are also indicated in each panel. The p value is a declining indicator which indicates the impact significance of the linear correlation coefficients R whereby R has statistical significance and the confidence level is larger than 95% when the p value is less than 0.05.

Figure 11: Temporal variation of EOF2, north-south component of wind speed, surface net heat flux, and TSC flux along the TWS section, and Kuroshio flux along the ET section. Their linear correlation coefficients and time-lags are also indicated in each panel.

Figure 12: The effects of Coriolis force (a), total pressure (b), surface friction (c), bottom friction (d), advection (e), and local acceleration (f) for water column in winter according to Eq. (5) (units: 10-4 m2/s2). The two blue arrows indicate the two TWC branches. The two triangles indicate the two regions with significant fluctuation north of Taiwan (P1) and in the inshore area (P2).

Figure 13: Variations in Coriolis force, total pressure, and wind in the alongshore direction at P1 (a), the cross-shore direction at P1 (b), the alongshore direction at P2 (c), and the cross-shore direction at P2 (d) according to Eq. (5).

Figure 14: Mean currents (upper panels) and synoptic fluctuations (EOF1 in middle panels and EOF2 in bottom panels) in winters of 2010-2013. The black arrows show the velocity (m/s) in the layer of VMV and the color shows the current speed. The two blue arrows with label IB and OB represent the flow axes of the inshore branch and offshore branch, respectively.

Figure 15: Traces of TSC water (a) and Kuroshio water (b) in winter, with the variation of surface currents at the original location of P1 (c). The black dots represent the release locations of tracers. The gray lines show the entire trajectories of the tracers. The red lines and blue lines are selected trajectories, which are close to the inshore branch and offshore branch, respectively. The dates show the times that selected tracers reached the origin location P1; note that the location of P1 is not fixed but varies with time. The numbers are the depths of the tracers, which are labeled at an interval of six days. The two black arrows represent the two TWC branches.

Figure 16: The VMV under the northerly wind (a) and southerly wind (b). Panel (c) shows the variation of wind in winter. Blue vectors and red vectors show the southwestward coastal current and the northeastward TWC, respectively. Gray contours indicate the 30, 50, 70, and 100 m isobaths. The two black arrows represent the two TWC branches. The green ellipse indicates the inshore area with

significant fluctuation.

Table 1: Annual-mean volume transports (Sv = 10^6 m^3/s) through various sections. The sections are shown in Figure 2 using blue dashed lines.

Section	Present model	Previous estimates
Taiwan Strait	1.22	1.2 (Isobe, 2008)
		1.8 (Wang et al., 2003)
		1.09 (Wu and hsin, 2005)
		1.03 (Yang et al., 2011)
		1.72 (Guo et al., 2006)
		0.5 (Hung et al., 2003)
		1.10 (Liu et al., 2014b)
Tsushima Strait	2.85	2.65 (Isobe, 2008)
		3.03 (Guo et al., 2006)
		2.70 (Yang et al., 2011)
		2.52 (Liu et al., 2014b)
200m isobath	1.66	1.46 (Guo et al., 2006)
		0.87 (Liu et al., 2014a)
		3.0 (Teague et al., 2003)
		2.74 (Lee and Matsuno, 2007)
East Taiwan Channel	22.71	21.50 (Johns et al., 2001)
		23.00 (Teague et al., 2003)
		23.83 (Guo et al., 2006)
		28.4 (Hsin et al., 2013)
		21.37 (Yang et al., 2011)
		20.74 (Liu et al., 2014b)
Tokara Strait	23.20	23.4 (Feng et al., 2000)
		20.00 (Teague et al., 2003)
		20.66 (Yang et al., 2011)
		24.42 (Liu et al., 2014b)

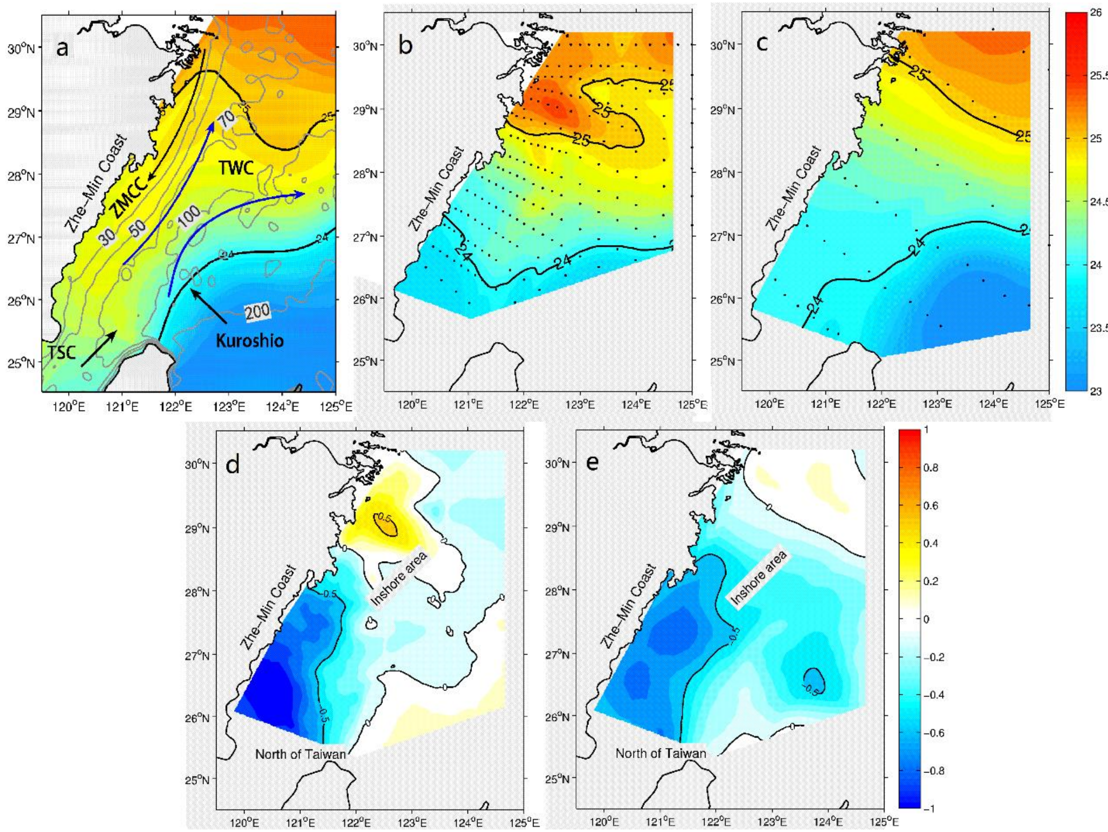
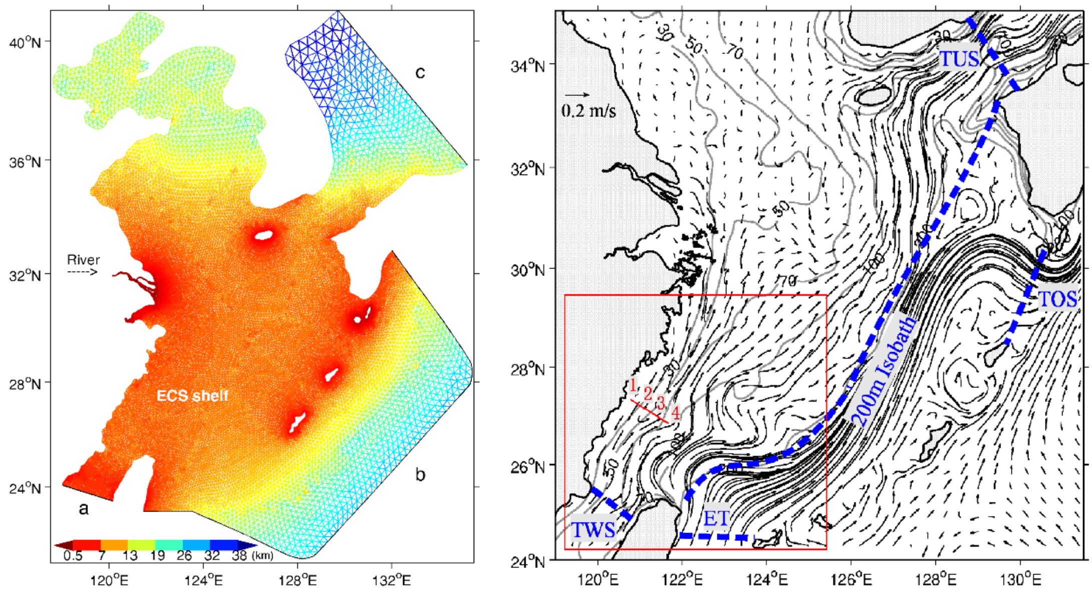


Figure 1: Density (σ_t, kg/m^3) distributions at 50 m depth derived the GDEM climatological data in February (a), an ocean survey from Feb. 1–27, 2007 (b), and an ocean survey from Feb. 3–16, 2007 (c), with the density anomalies between the GDEM data and the two surveys (d and e). The two blue arrows indicate the two TWC branches in winter. The 30, 50, 70, 100 and 200 m isobaths are indicated with grey lines in panel a.



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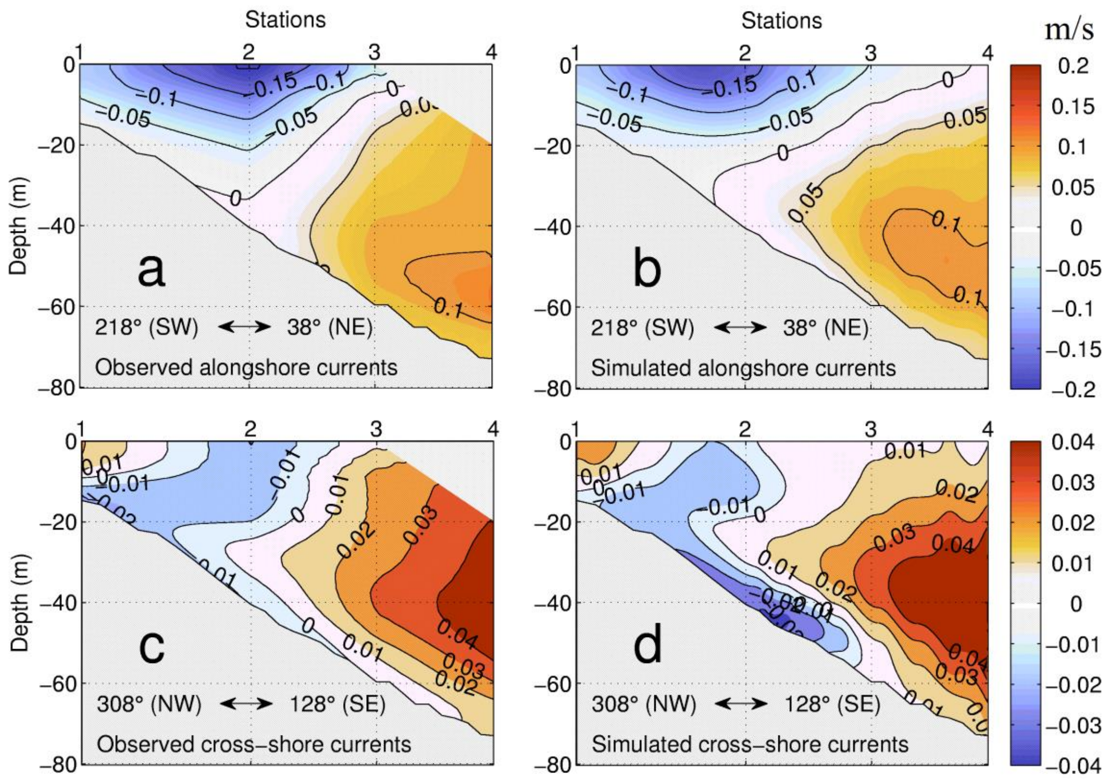


Figure 3: Validations of the wintertime TWC (warm color) along the section off the Zhe-Min coast (the short line with four red numbers in Figure 2): (a) observed alongshore currents; (b) simulated alongshore currents; (c) observed cross-shore currents; (b) simulated cross-shore currents. Note, an enlarged color scale is used for the cross-shore component to have a clear view of its weak structure.

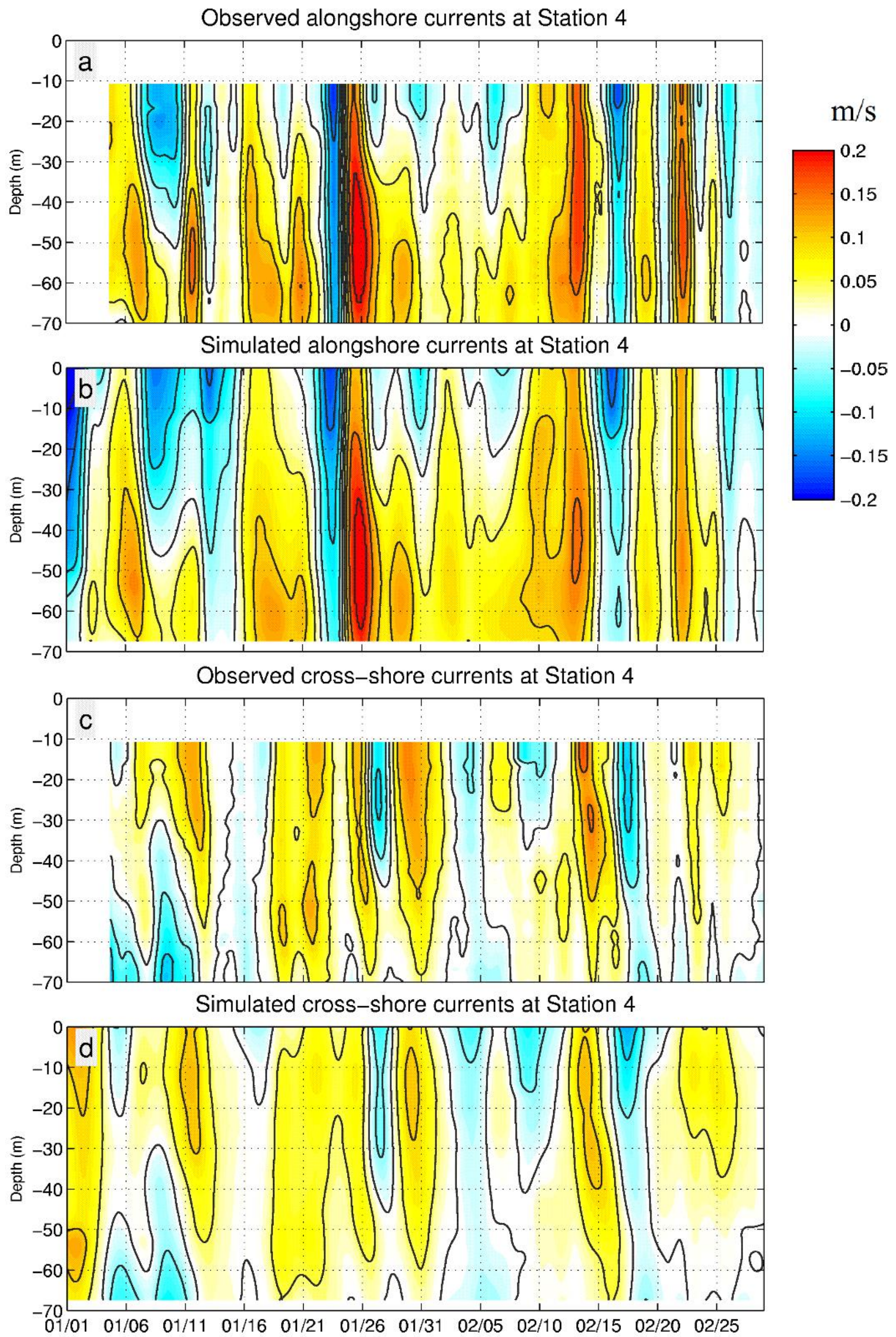
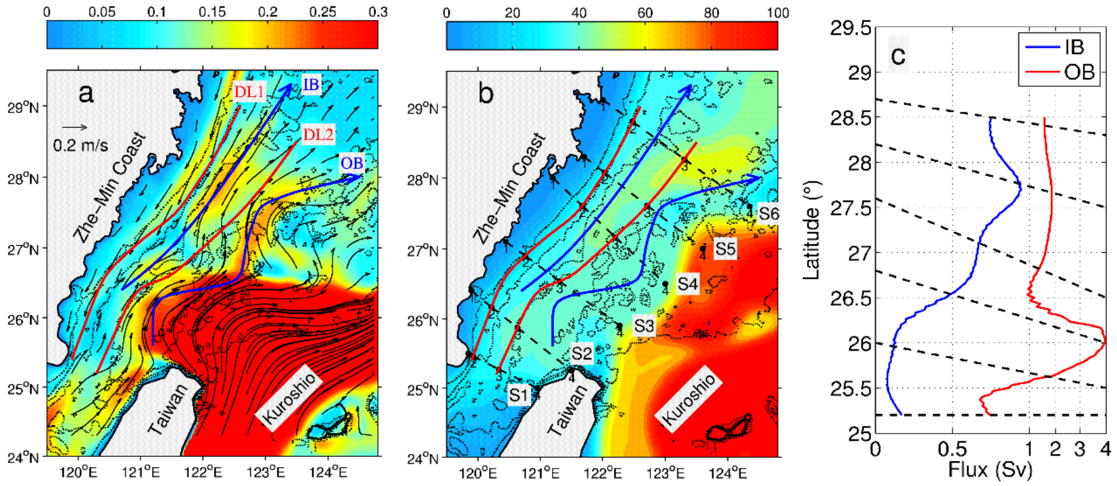


Figure 4: ~~Validations~~Variations of the ~~wintertime~~ inshore branch of TWC during January and February 2009~~fluctuations~~: (a) observed alongshore currents; (b) simulated alongshore currents; (c) observed cross-shore currents; (d) simulated cross-shore currents. The observation data comes from Station 4 in Figure 1 and the simulated data has the same position and period as the observation data.



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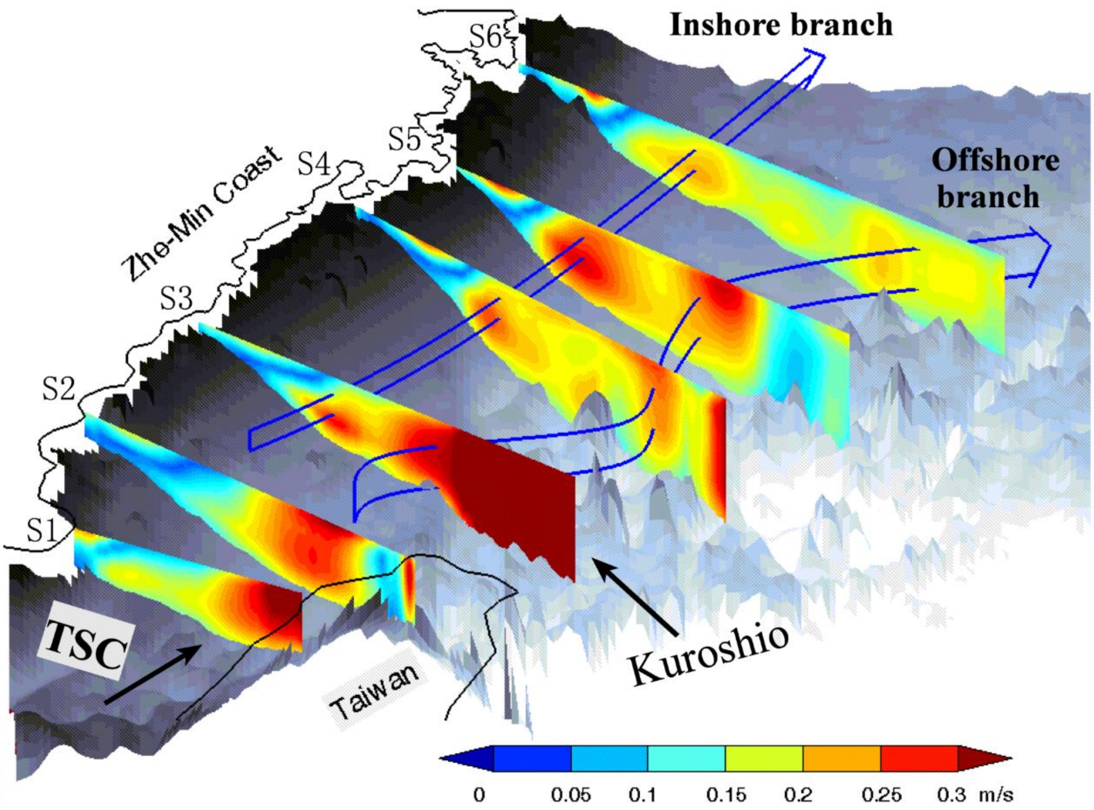

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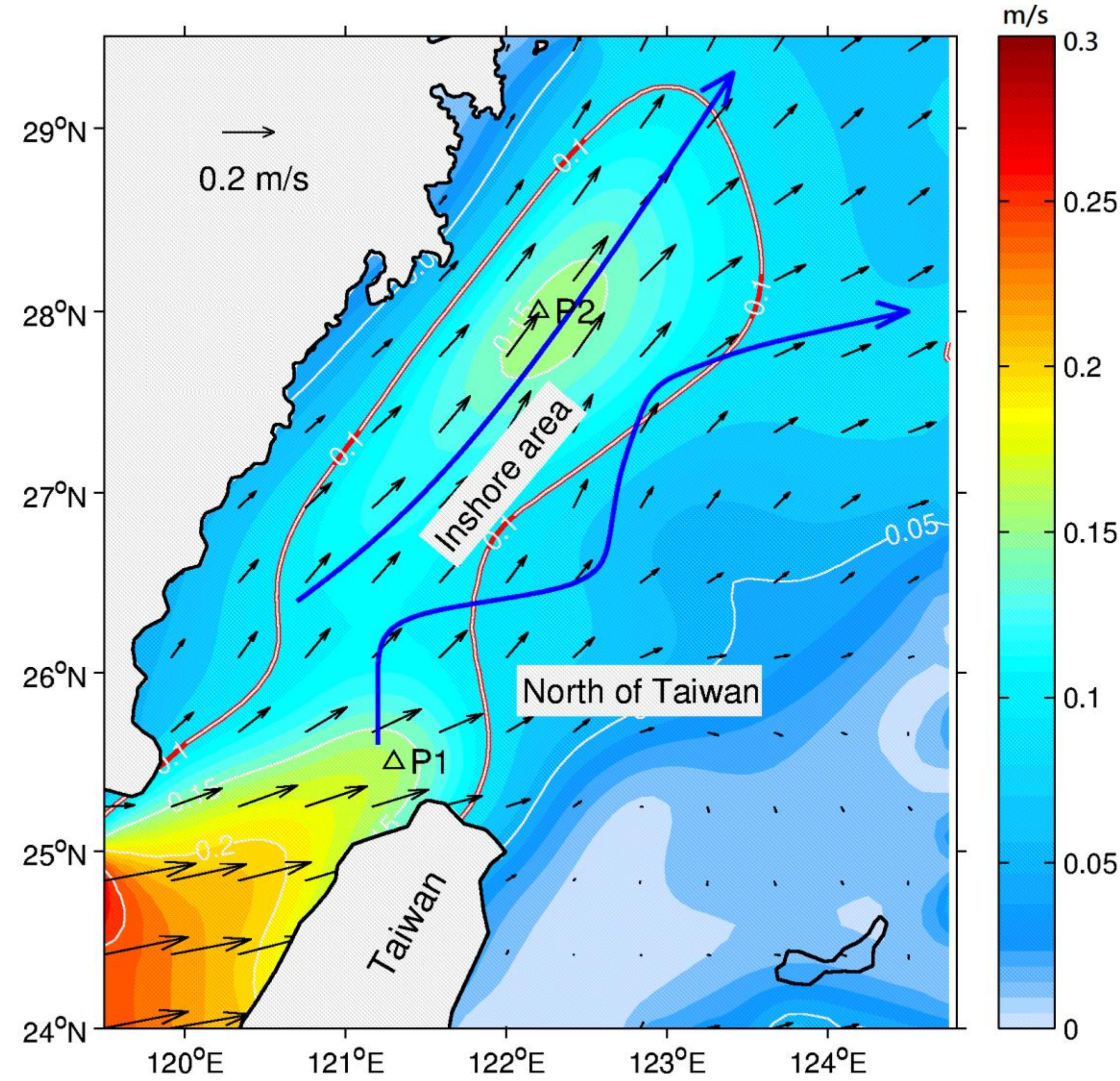

Figure 7: Current standard deviation in the layer of the VMV. The color shading shows the magnitude of the current standard deviation. The two blue arrows indicate the two TWC branches. The red curve
935 indicate the area where the current standard deviation is larger than 0.1 m/s and their representative points (P1 and P2) are selected for later analysis.

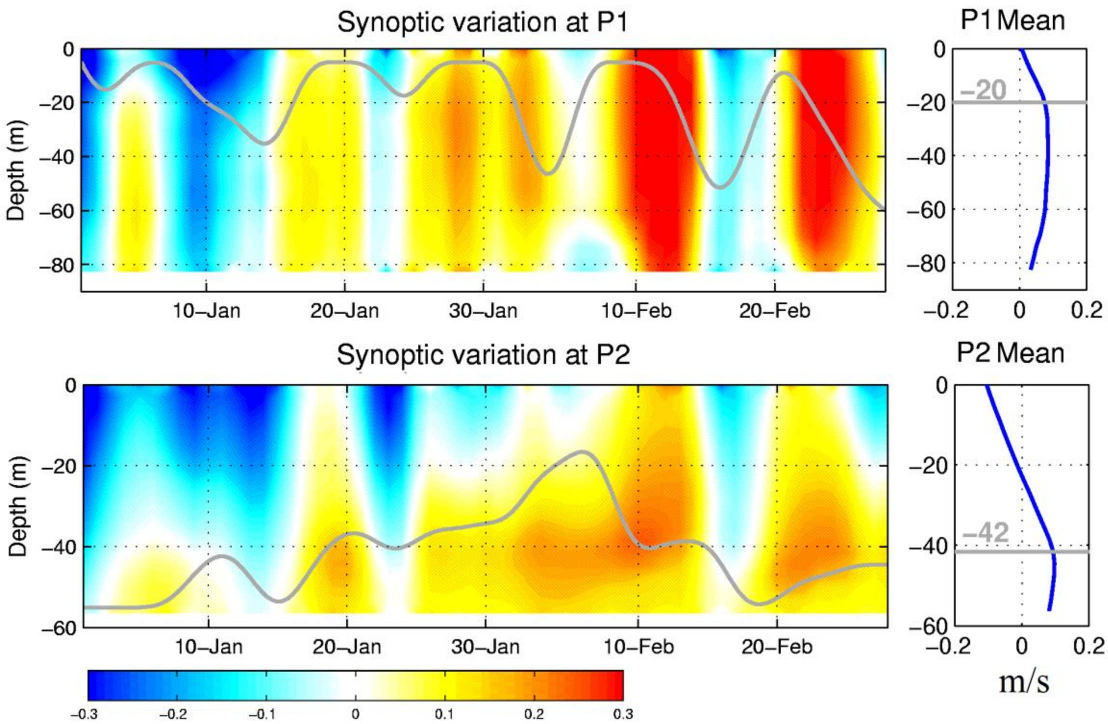

Figure 8: Variation of alongshore currents (m/s) for the entire water column north of Taiwan (P1) and in the inshore area (P2) and their relation with upper mixed layer depth. The positive velocity (warm color) indicates the occurrence of the TWC. The gray solid lines show the depth of the upper mixed layer.

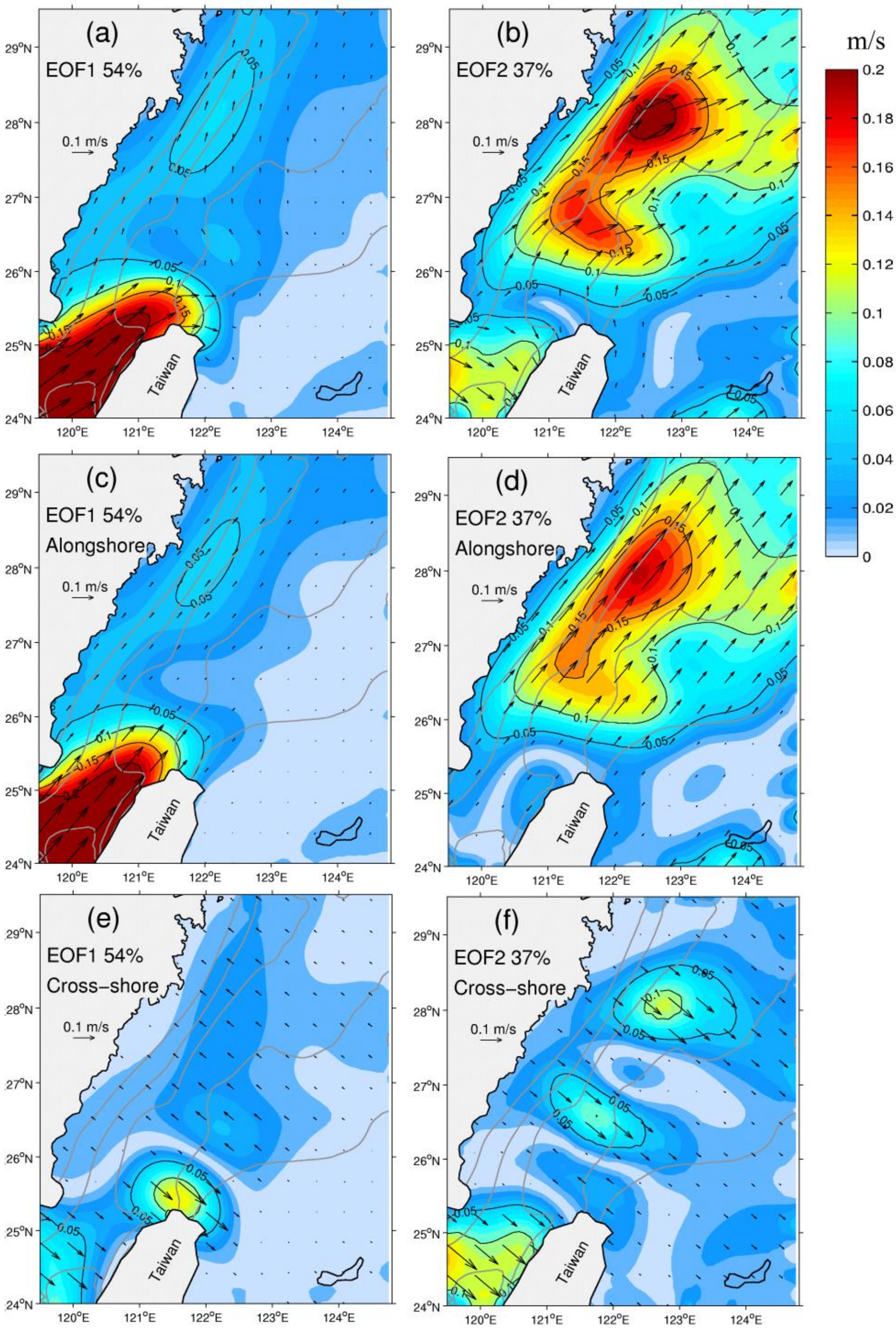

Figure 9: The spatial pattern of the first (EOF1; left) and second (EOF2; right) leading modes of the VMV in the ECS: (a) EOF1 currents, (b) EOF2 currents, (c) EOF1 alongshore component, (d) EOF2 alongshore component, (e) EOF1 cross-shore component, and (f) EOF2 cross-shore component. The 30, 50, 70, 100 and 200 m isobaths are indicated with grey lines.

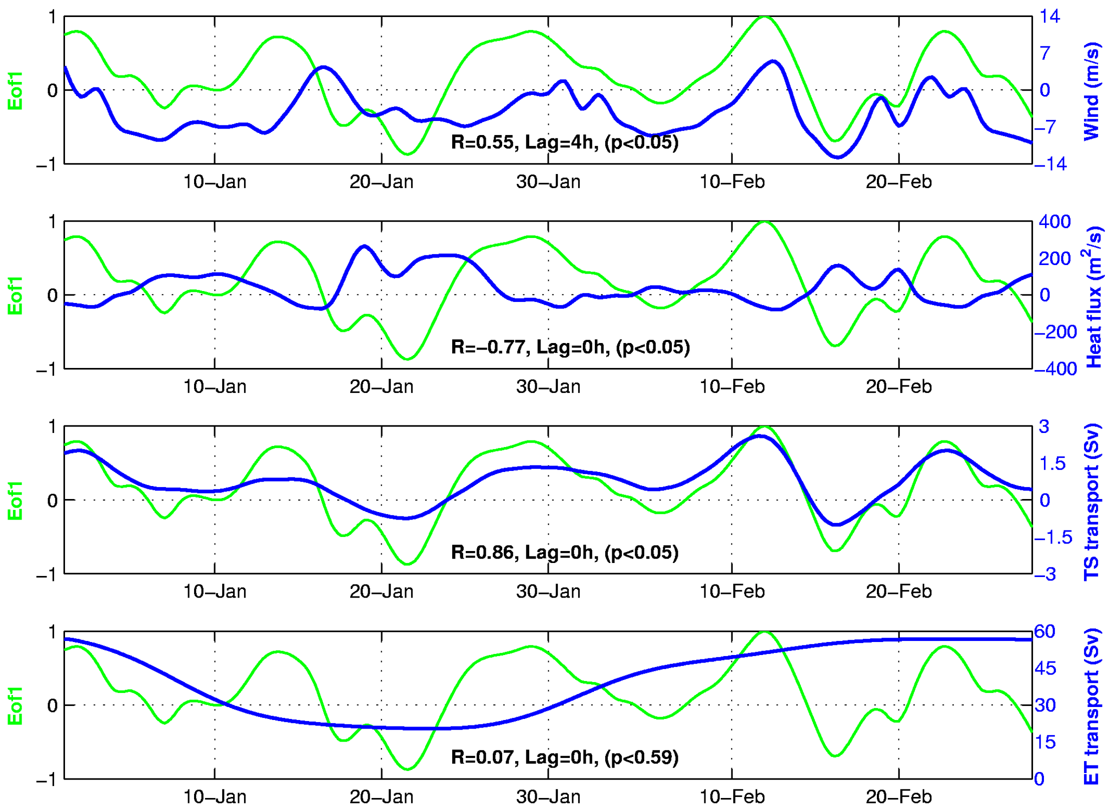

Figure 10: Temporal variation of EOF1, north-south component of wind speed, surface net heat flux, and TSC flux along the TWS section, and Kuroshio flux along the ET section. Their linear correlation coefficients R and time-lags are also indicated in each panel. The p value is a declining indicator which indicates the impact significance of the linear correlation coefficients R whereby R has statistical significance and the confidence level is larger than 95% when the p value is less than 0.05.

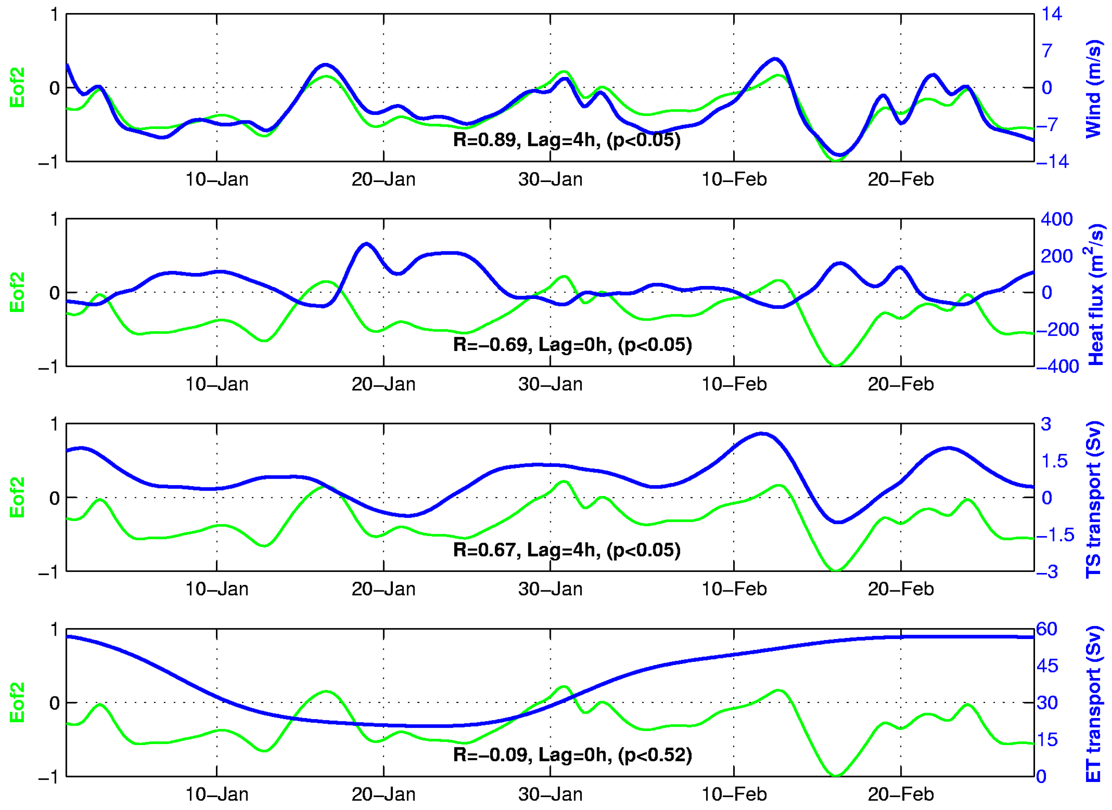

Figure 11: Temporal variation of EOF2, north-south component of wind speed, surface net heat flux, and TSC flux along the TWS section, and Kuroshio flux along the ET section. Their linear correlation coefficients and time-lags are also indicated in each panel.

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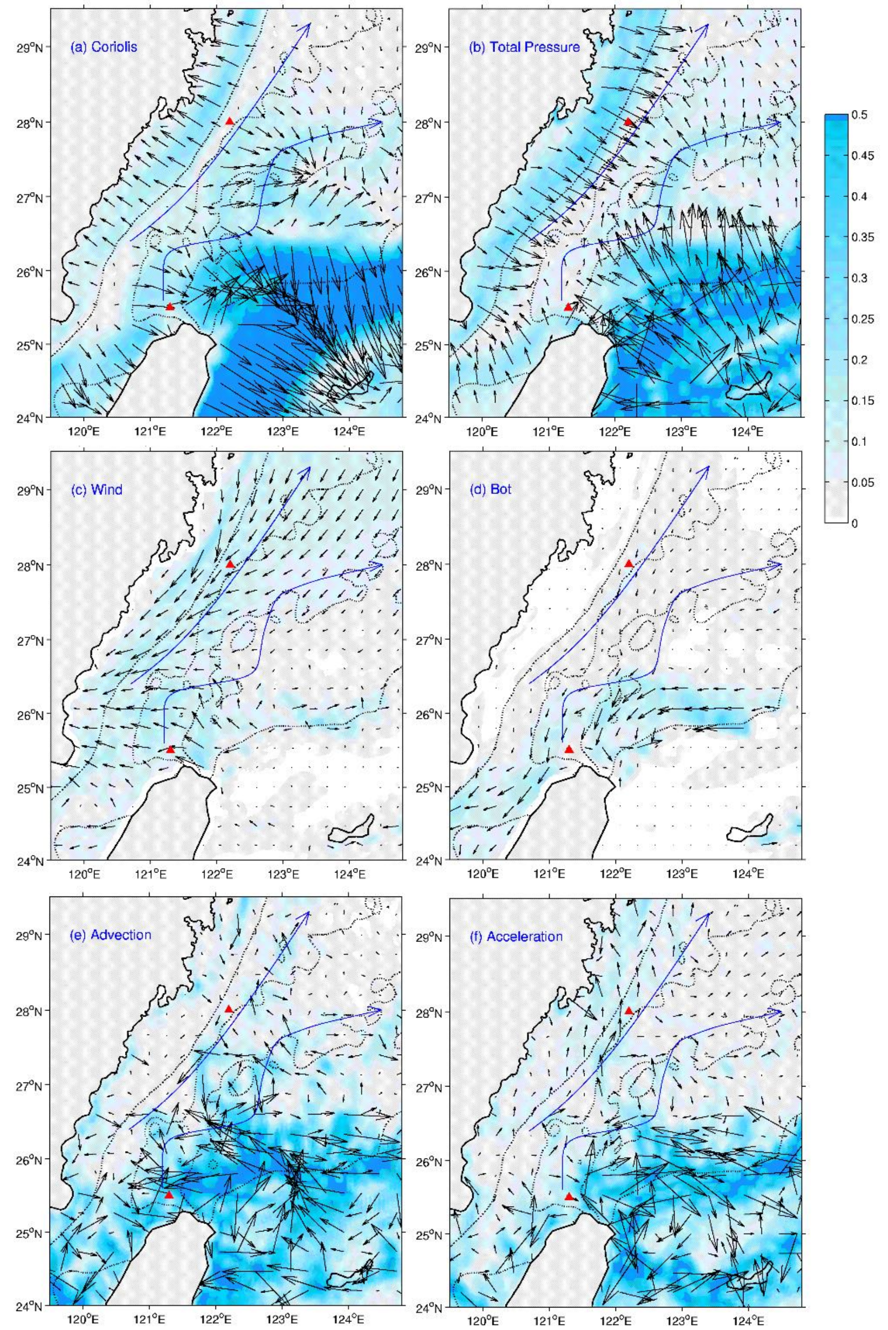

Figure 12: The effects of Coriolis force (a), total pressure (b), surface friction (c), bottom friction (d), advection (e), and local acceleration (f) for water column in winter according to Eq. (5) (units: 10^{-4} m^2/s^2). The two blue arrows indicate the two TWC branches. The two triangles indicate the two regions with significant fluctuation north of Taiwan (P1) and in the inshore area (P2).

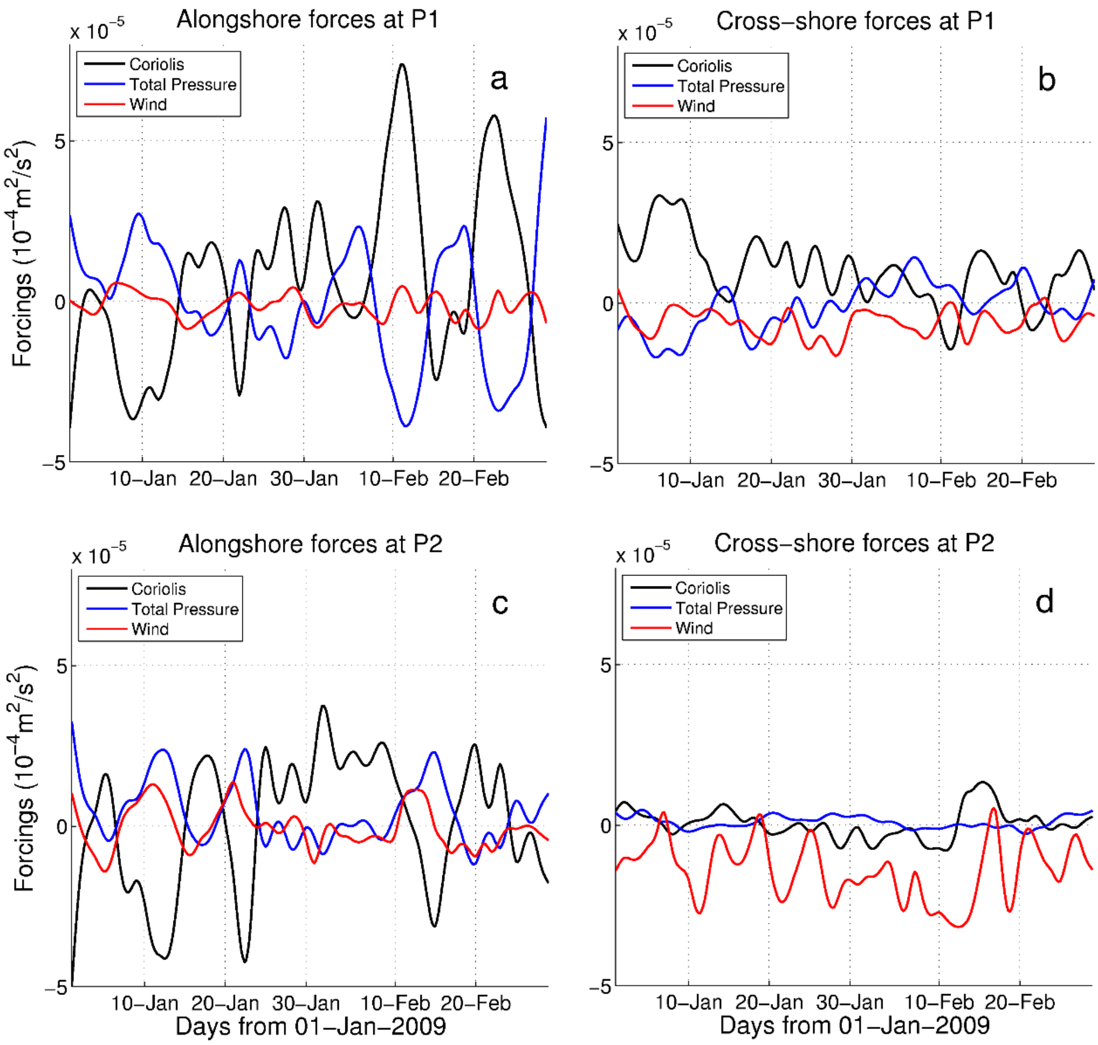

Figure 13: Variations in Coriolis force, total pressure, and wind in the alongshore direction at P1 (a), the cross-shore direction at P1 (b), the alongshore direction at P2 (c), and the cross-shore direction at P2 (d) according to Eq. (5).

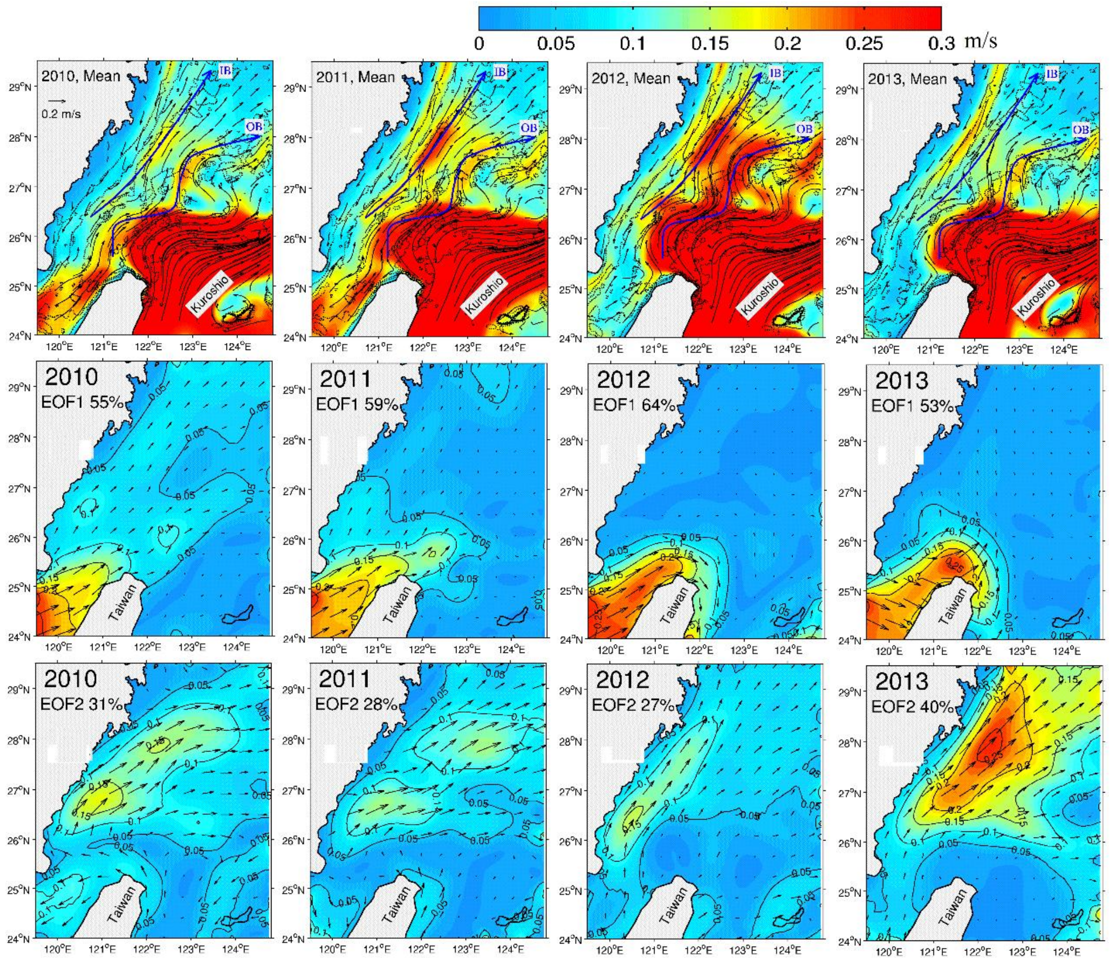

Figure 14: Mean currents (upper panels) and synoptic fluctuations (EOF1 in middle panels and EOF2 in bottom panels) in winters of 2010-2013. The black arrows show the velocity (m/s) in the layer of VMV and the color shows the current speed. The two blue arrows with label IB and OB represent the flow axes of the inshore branch and offshore branch, respectively.

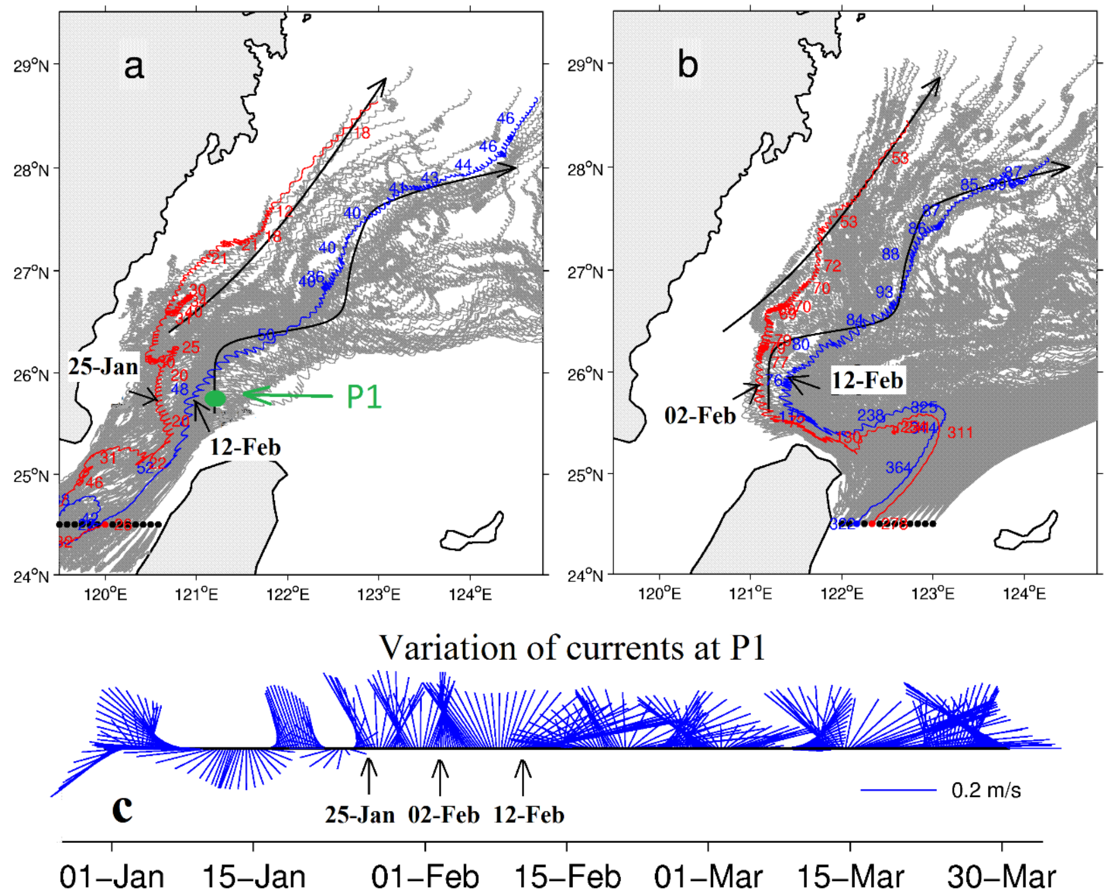

Figure 15: Traces of TSC water (a) and Kuroshio water (b) in winter, with the variation of surface currents at the original location of P1 (c). The black dots represent the release locations of tracers. The gray lines show the entire trajectories of the tracers. The red lines and blue lines are selected trajectories, which are close to the inshore branch and offshore branch, respectively. The dates show the times that selected tracers reached the origin location P1; note that the location of P1 is not fixed but varies with time. The numbers are the depths of the tracers, which are labeled at an interval of six days. The two black arrows represent the two TWC branches.

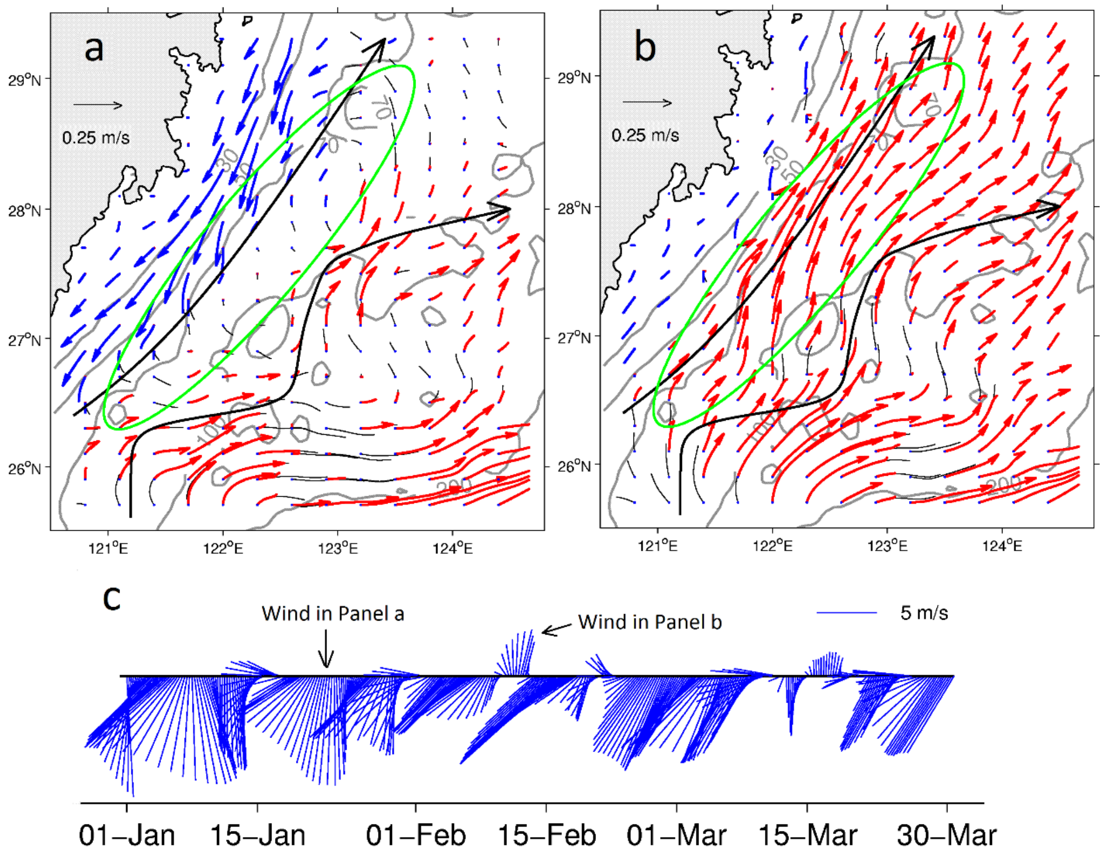

Figure 16: The VMV under the northerly wind (a) and southerly wind (b). Panel (c) shows the variation of wind in winter. Blue vectors and red vectors show the southwestward coastal current and the northeastward TWC, respectively. Gray contours indicate the 30, 50, 70, and 100 m isobaths. The two black arrows represent the two TWC branches. The green ellipse indicates the inshore area with significant fluctuation.