

## Response to Reviewer #1

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We do thank the Reviewer for carefully considering our work, and for her/his interesting remarks that contributed to improve the manuscript. After consideration of the comments contained in the report, we have largely amended the paper and better presented our results.

For the sake of clarity, new parts are written in color red in the revised version.

Here below, we reply to the Reviewer's comments point by point, better discussing those issues that may not look very clear. We also list what we changed in the paper.

Reviewers criticisms are here marked in blue, bold face. Our answers in black. The changes made in the manuscript in red.

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**Rev: The authors use the FSLE metrics computed from surface drifters as benchmark. More details on the drifter data set should be given. In particular, are they CODE sampling the upper 1 meter, or are they drogued at 15 m?**

The drifters are drogued at 15 m. In the revised version, we added details about the drifter data set.

In the revised version, we added the following sentence:

**“Drifter data belong the Mediterranean Sea-In-Situ Near Real Time Observations (database INSITU\_MED\_NRT\_OBSERVATIONS\_013\_035 available on <http://marine.copernicus.eu/>). These are surface buoys, drogued at a nominal depth of 15 m.”**

**Rev: The authors compare the drifter results with results from MFS trajectories at the surface and they find a significant difference, since the model results show a plateau after approximately 30-40 km while drifter results keep increasing. This is actually a known result, as discussed in details in a number of papers (e.g. Haza et al., 2012), and it is due to the lack of resolution of the model. The authors mention this as a possible reason, but I think they should expand and introduce the references.**

We agree with the Reviewer. In the revised version, we acknowledge that the reason for the observed discrepancy is associated to both the lack of spatial and temporal resolutions of the numerical model.

In the revised version, we added the following sentence:

**“As it has been previously observed, this discrepancy is due to both the coarse spatial resolution and the time averaging of any mesoscale model (see e.g. Haza et al. (2012) and references therein).”**

**Rev: More interesting is the discussion on the effects of vertical shear, even though there is a conceptual problem in the way it is presented in the paper. The authors use the same FSLE metric from surface drifters as benchmark, but surface drifters (especially if CODE!!!) do not feel the effects of vertical shear on scales of approx. 40 m as considered in the model and even if they would, it would be a vertically averaged effect, very different from the buoyant particle case of the model. As a consequence, in my opinion this comparison is seriously flawed.**

This is actually a delicate point of our work that we will try to make clearer. We do not claim that surface drifters feel the effects of vertical shear, since obviously they cannot. The reason why drifter dispersion is important in this context is the following. When simulating **horizontal** tracer dispersion within an ocean model, one has to decide if the tracer is fully 3D (e.g., as chlorophyll scalar field in the mixed layer), or if it is constrained to a 2D surface (e.g. as a surface drifter).

Numerical simulations – in the absence of any sub-grid scale model- show a significant difference between the two cases, even at quite small vertical separations: for the MFS surface drifter case, a flat FSLE ( $\lambda(r) \sim r^0$ ) is observed for separations up to the mesoscale range, indicating a constant exponential rate of trajectory separation; for the MFS Serie I case (a 3D MFS), a FSLE increasing at any given small separation ( $\lambda(r) \sim r^a$ ).

We claim that this behavior is an artifact of the finite resolution of the ocean model, i.e. due to an anomalous temporal persistence of the velocity vertical gradients, and we support this hypothesis by comparison with observational velocity profile data. We also show that such “anomalous shear effect” events are triggered as soon as an elementary sub-grid-scale vertical mixing model is added to the ocean model in order to simulate 3D motions typical of the mixed layer.

Now, the question arises on the proper small-scale ocean model velocity field able to simulate horizontal dispersion of a tracer **having 3D structure** down to, say, the bottom of the mixing layer. The solution to this problem, we believe, is very difficult to find, at least at a kinematic parameterization level.

Finally, another important aspect has yet to be considered. The MFS ocean model underestimates mesoscale turbulent dispersion. The primary correction, in our opinion, is to restore mesoscale horizontal dispersion by comparison with drifter data.

Clearly, mesoscale eddies are not pure 2D structures, but have a certain vertical development in the mixed layer. This implies that mesoscale turbulent dispersion, e.g. the Richardson’s regime, is not a property of the surface layer only, but belongs to a whole vertical range of ocean layers. If one applies such “primary correction” for mesoscale eddies, she/he realizes that the effect is to have an efficient dispersion that **covers** the one due to mean vertical shear.

So for the ocean situation studied in our work, restoring horizontal mesoscale dispersion, in agreement with experimental drifter data, automatically eliminates the problem of the anomalous shear effects. We stress that this is what occurs in the case of the Mediterranean Sea dispersion analysed in the present work: but it might not be true in general, and the problem of realistically modelling Lagrangian transport could require different solutions.

Since this is an important point in our work, in the revised version we added the following paragraphs to make this clear from the very beginning.

In particular:

At the beginning of the new Section 3.1

“We discuss different sets of numerical simulations based on the velocity configurations of the MFS model, also supplemented by the use of the kinematic model to described poorly resolved motions. Kinematic models can be adapted to the different dispersion regimes, namely exponential separation, turbulent dispersion, and standard diffusion. Their implementation hence depends on the specific dynamics and specific range of scales that one wants to describe. Here, we compute transport properties by introducing statistical Lagrangian motions for the mixed layer motions (3D KLM), and separately for the poorly mesoscale motions (2D KLM). By doing so, we demonstrate that i) small-scale motions enabling tracer pairs to explore the whole mixed layer do not modify MFS horizontal dispersion properties, in reason of the anomalous persistence of vertical velocity

gradients in the MFS model; ii) differently, the horizontal relative separation resulting from the introduction of the 2D KLM is fast enough to encompass the anomalous shear effect produced by the MFS solution.”

#### At the end of the new Section 3.2

In Series I, the effect of the vertical shear onto the horizontal dispersion comes from the MFS model only. The associated FSLE curves clearly indicates that vertical shear is able to promote horizontal dispersion. Neutrally buoyant tracers moving at different depths experience velocity differences: as a result they start to separate already at very small scales. In Series II, the 3D KLM terms are switched on, and particles vertically explore the whole mixed layer. The obtained FSLE curve is very similar to that of Series I, and in particular it results to be slightly below this last. This finding is somehow surprising since, thanks to the introduction of small-scale turbulent-like motions, tracer pairs can explore the whole mixed layer. However the fluctuations of the 3D KLM do not substantially modify the horizontal pair dispersion, and actually they make it slightly slower in the present case. This implies that the dominant effect is the one associated to the vertical shear.

In Series III, both the 3D and the 2D KLM are switched on. The resulting FSLE is larger than that of series II at any scale. This means that the most important dynamical correction to the MFS model is that associated to the 2D KLM. Indeed, the dispersion effect induced by the mesoscale eddies inserted in the 2D KLM covers any other horizontal dispersion effects, including the one associated to the anomalous persistence of vertical gradients in the MFS model.

We can summarise the results of the numerical simulations as follows. By comparing mesoscale dispersion of the bare MFS model with Mediterranean drifter data, one sees that real drifter pair dispersion follows a turbulent-like type of behavior, whereas model trajectories separate more slowly and at a nearly constant rate. Adding vertical mixing to the ocean model, e.g. in the form of the 3D kinematic model, may trigger a type of shear dispersion which is affected by an anomalous persistence of the vertical velocity gradients, as discussed in Sec. 2. Since we estimate that the anomalous persistence is an artifact of the poor temporal resolution of MFS, the adoption of the 3D KLM only does not seem an appropriate choice. On the other hand, adding a two dimensional kinematic model, one finds that the anomalous shear dispersion effects become practically negligible, being hidden by the more energetic dispersion processes occurring at the mesoscales. Clearly, mesoscale eddies are not pure 2D structures, but they have a certain vertical development in the mixed layer. This implies that mesoscale turbulent dispersion is not a property of the surface layer only, but belongs to a whole vertical range of ocean layers. By adding the 2D KLM for mesoscale eddies, one realises that the effect is to have an efficient dispersion that covers the one due to mean vertical shear. Finally, it is worth recalling that, as discussed in Lacorata et al. (2014), the FSLE measured for surface drifters follows the Richardson diffusion behaviour  $\lambda(r) \propto r^{-2/3}$  for  $r \in [10,100]$  km: this is consistent with the observed dispersion rates in the GLAD experiment, which spans however a much wider range of scales (Poje, 2014)”

Also in the conclusion, we added the following paragraph:

“Now, the question arises on the proper small-scale ocean model velocity field able to simulate horizontal dispersion of a tracer having 3D structure in the mixed layer and below. The solution to this problem is very difficult, mainly because experimental data of 3D tracer dispersion are not easily available. Different modelling solutions can be adopted to account for different problems, depending whether mesoscales, sub-mesoscales or small-

scales are the relevant range of scales in the dispersion problem. Even the poor man's procedure of fitting numerical simulations to observations would require an ad hoc mixed layer modelling which is not straightforward to implement."

**Rev: What is interesting is the comparison between the model results, i.e. the MFS surface, the MFS at 2 levels and the case with KLM, that show the effects of shear for a model tracer distributed in the first 50 m. Results indicate that adding a Lagrangian Kinematic Model that parametrizes the action of 3d turbulence does not increase the effects of shear (it actually tends to reduce them in some cases...). The author rational for this, if I understand correctly, is that model vertical shear is unrealistically high because correlation time scales are too long compared with data. This point should be expanded and better explained.**

The Referee is right: adding a 3D Kinematic model does not modify the measured effect of shear in comparison with the bare MFS model. This happens because in MFS vertical velocity gradients are anomalously persistent, and the chaotic mixing introduced by the small-scale 3D KLM is sub-dominant with respect of the shear effect of the MFS model. However, the anomalous persistence is no longer a problem when a 2D KLM for mesoscale motions is added. Now it is the latter which gives the dominant contribution, as mentioned above.

In the revised version, we try to make this very important point clearer (read above paragraphs).

**Rev: Overall, the paper needs in my opinion some restructuring. My suggestion is the following:- a) The first step should be showing ADCP data to characterize vertical shear in 2 Mediterranean Sea sites. b) ADCP results are compared to MFS results, indicating that correlation times are too long in the model c) The 3d KLM is introduced to parametrize the effects of turbulence and introduce short time scales d) the effects of shear are studied comparing model results: first surface MFS, then cseries I and II. e) I would remove the part on drifter comparison, because I am not sure it adds anything. It can be mentioned as previous results.**

We have reshaped the presentation of results following the Reviewer suggestion. Now the ADCP data and their comparison with MFS estimates are presented in Section 2. Numerical results of Lagrangian dispersion are presented in Section 3, where the 3D and 2D KLM model formulations are now discussed in detail and the comparison of different series results is expanded. We kept however the drifter data results, since in our opinion and as explained above, these are relevant for the aims of the paper.

**Rev: More in details: where did they get the KLM 3d results?**

In the revised version we added all details that are crucial for the understanding of the present work, trying to avoid repetitions with previous works. In particular, the 3D KLM model equation (eq. 7) and parameters (eq. 8) are now explained. The 2D KLM model equation (eq. 9) and parameters (eq. 10) are also explained.

**Rev: The discussion in the Introduction should be more focused and the specific scales of interest should be mentioned.**

In the revised manuscript, in the Introduction, we added the following paragraphs:

"In this paper, we focus on the role of vertical shear as important mechanism promoting

the horizontal diffusion in the ocean. By vertical shear, we mean the vertical variation of the velocity horizontal components. The approach here considered consists in combining observative and model data to assess the effect of vertical shear for the tracer horizontal relative dispersion. Observative data come from Acoustic Doppler Current Profilers (ADCP), deployed in the South Mediterranean. Numerical data come from the Mediterranean sea Forecasting System model, and are supplemented with the use of deterministic kinematic models (Palatella et al., 2014; Lacorata et al., 2014), to parameterise poorly resolved mesoscale motions, or unresolved processes in GCMs. The Kinematic Lagrangian Model (KLM) here adopted can be two dimensional, to better account for the horizontal dispersion due to mesoscale eddies, or three dimensional, to simulate vertical turbulent-like motions in the ocean mixed layer. Both dynamics are often underestimated in General Circulation Models (GCM). Even if our primary interest is in the former situation, we will discuss both.

The paper is organized as follows. In Sec. 2, we compare in-situ observations of vertical velocity gradients with measures obtained from MFS. The comparison highlights that velocity gradients correlation times derived from MFS are considerably larger than the observed ones: such anomalous temporal persistence of the vertical shear is responsible of an enhanced relative dispersion, which is possibly an artifact of the low temporal resolution of the model. In Sec. 3, we discuss the relative dispersion properties of neutrally buoyant tracers by means of numerical simulations. We show that, by a suitable implementation of the kinematic model, the anomalous shear effect can be overcome. Section 4 contains the final remarks and perspectives.”

***Rev: The choice of the parameters for the 3D KLM should be discussed and motivated.***

In the revised paper, it is explicitly described the 3D KLM and its motivation, namely to reproduce the mixed-layer turbulent like dynamics. In particular, one now reads:  
“The small spatial scales  $l_i$  and their associated fast time-scales  $t_i \sim l_i / A$  are chosen to reproduce, on average, the dynamical properties within the mixed layer”

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A short list of the main changes is the following:

- Title has been changed.
- Introduction: some general paragraphs have been removed, while more specific comments and references have been added to make it more focused;
- Section 2 : it contains the analysis of the ADCP data, and the comparison with MFS estimates;
- Section 3: it contains the numerical simulation results. The 3D KLM model equation (eq. 7) and parameters (eq. 8) are explained. The 2D KLM model equation (eq. 9) and parameters (eq. 10) are explained. The FSLE discussion in sub-section 3.2 has been expanded;
- Conclusions: some summarizing paragraphs have been removed to avoid repetition. A paragraph containing some perspective for this work has been added.
- References: reference list has changed to make it more focused on the topic of the paper.