

1 **Coupling of wave and circulation models in coastal-ocean**
2 **predicting systems: A case study for the German Bight**

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10 **Abstract**

11 This study addresses the impact of coupling between wave and circulation models on the
12 quality of coastal ocean predicting systems. This is exemplified for the German Bight and its
13 coastal area known as the Wadden Sea. The latter is the area between the barrier islands and
14 the coast. This topic reflects the increased interest in operational oceanography to reduce
15 prediction errors of state estimates at coastal scales, which in many cases are due to
16 unresolved nonlinear feedback between strong currents and wind-waves. In this study we
17 present analysis of wave and hydrographic observations, as well as results of numerical
18 simulations. A nested-grid modelling system is used to produce reliable nowcasts and short-
19 term forecasts of ocean state variables, including waves and hydrodynamics. The data base
20 includes ADCP observations (taken from the BSH) and continuous measurements from data
21 stations. The individual and combine effects of wind, waves and tidal forcing are quantified.
22 The performance of the forecast system is illustrated for the cases of several extreme events.
23 The combined role of wave effects on coastal circulation and sea level are investigated by
24 considering the wave-dependent stress and wave breaking parameterization. Also the
25 response, which the circulation exerts on the waves are tested for the coastal areas. The
26 improved skill of the coupled forecasts compared to the non-coupled ones, in particular
27 during extreme events, justifies the further enhancements of coastal operational systems by
28 including wave effects into circulation models.

1 **1 Introduction**

2 In the last decade the north European coasts were affected by severe storms which caused
3 serious damages in the North Sea coastal zones. Additionally, different human activities, e.g.
4 offshore wind power industry, oil industry and coastal recreation necessitate information
5 about the sea state in the coastal ocean with high resolution in space and time. There seems to
6 be a consensus that high-quality predictions of extreme events like storm surges and flooding
7 caused by storms could substantially contribute to avoid or minimize human and material
8 damages and losses. Therefore reliable wave forecasts and long term statistics of extreme
9 wave conditions are of utmost importance for the coastal areas. In many coastal areas the need
10 for reliable risk assessments increases the demand of precise coastal predictions. This cannot
11 be achieved by further neglecting the wave-current interaction in coastal ocean operational
12 forecasting.

13 Waves-current interaction is recently an important issue in the field of coastal ocean
14 forecasting (Roland and Arduin, 2014, Bolaños et al., 2014). Understanding this process is
15 of utmost importance on the road of fully integrating the atmospheric, wave and ocean models
16 and their further coupling with biological, morphological, and hydrographical forecasting
17 systems. The uncertainties in most of the presently used models results from the nonlinear
18 feedback between the currents, water level variations and wind-waves, which can no longer
19 be ignored, in particular in the coastal zone. The joint impact of surges, currents and waves is
20 strongly inter-related (Wolf et al., 2011, Brown et al., 2011) and those cannot be considered
21 separately for coastal ocean predictions.

22 The ocean waves affect not only the sea level but also the currents and mixing, the latter being
23 of utmost importance for the sediment dynamics (Lettmann et al, 2009). Prandle et al. (2000)
24 demonstrated the need of accounting for surface waves with a significant wave height larger
25 than one meter in the sediment modelling. This is of big importance for sediment dynamic
26 and other ecosystem processes (Wolf and Prandle, 1999). These authors showed also that the
27 effects of waves add to the ones due to surges and tides; on the other side the waves'
28 characteristics are affected by the changes of sea level height due to tides and wind.

29 The main effects of waves that are commonly considered in the coupled modelling are due to
30 radiation stress and Stoke drift. Babanin et al. (2010) showed that interaction of turbulence
31 and bottom stress is also very important.

1 Wave-current interaction has been a topic of many studies recently (Ardhuin et al., 2008,
2 Mellor, 2003; 2008; 2011; Kumar et al., 2012; Michaud et al. 2012, Zodiatis et al. 2015).
3 Mellor (2003, 2005, 2008) extended the radiation stress formulation based on the linear wave
4 theory of Longuet-Higgins and Stewart (1964). Bennis and Ardhuin (2011) questioned the
5 method of Mellor and suggested the use of Lagrangian mean framework leading to the so
6 called vortex force. Vortex force method has been implemented in ROMS-SWAN (Kumar et
7 al., 2012; Lane et al., 2007; McWilliams et al., 2004; Uchiyama et al., 2010). Moghimi et al.
8 (2013) compared critically the two approaches claiming that the radiation stress formulation
9 showed unrealistic offshore directed transport in the wave shoaling regions; on the other hand
10 the results of longshore circulations performed similarly for both methods. Aiki and
11 Greatbatch (2013, 2014) proved that the radiation stress formulation of Mellor is applicable
12 for small bottom slopes. Bolaños et al. (2011, 2014) demonstrated the importance of wave-
13 current interactions in a tidally dominated estuary and showed that the inclusion of wave
14 effects through 3D radiation stress improves the velocity in the study area. They also
15 compared the different radiation stress methods and concluded that for the tidally dominated
16 area the 3D version of radiation stress produces better results than the 2D version. Polton et
17 al. (2005) found that accounting for the Stokes-Coriolis forcing results in encouraging
18 agreement between model and measurements of the mixed layer. Janssen (2012) showed
19 positive impact of wave breaking to the daily cycle of sea surface temperature. Later Breivik
20 et al. (2015) demonstrated reduced bias between modelled and measured water temperature
21 by incorporating the Stoke-Coriolis forcing, turbulence induced by breaking waves and ocean
22 side stress in the NEMO model at global ocean scale. Weber et al. (2006) estimated that the
23 wave induced stress is about 50% of the total atmospheric stress for moderate to strong wind.
24 Wolff et al. (2011) studied the effects of waves on hydrodynamics; Brown et al. (2013)
25 considered the wave effects on the storm surges; Roland et al. (2009) studies wave effects on
26 water level for the Adriatic Sea. The importance of ocean depth and velocity variations for the
27 simulated waves in the estuaries is analysed by Pleskachevsky et al. (2011) and Lin and Pierre
28 (2003). However, within the framework of practical coastal ocean forecasting, the interactions
29 between waves and currents are still not yet enough considered.

30 In this study we will address the coupling between wave and circulation models for coastal
31 ocean prediction systems on the example of the German Bight. We do not plan to analyse the
32 role of different parameterizations used. Rather we will demonstrate the areas of

1 improvements of coastal ocean predictions due to coupling between wave and hydrodynamic
2 models.

3 The structure of the paper is as follows. The wave and hydrodynamic models and the
4 processes of their interaction are described in Section 2. Section 3 addresses the effects of
5 hydrodynamics on wave model performance, while in Section 4 we discuss the effects of
6 waves on hydrodynamics and improvement of short-term forecast; followed finally by
7 concluding remarks.

8

9 **2 Model Description**

10 **2.1 Hydrodynamical Model**

11 The General Estuarine Transport Model (GETM, Burchard and Bolding, 2002) was used in
12 this study to simulate the circulation. This model solves the primitive equations for
13 momentum, temperature, salinity, and water level. The model set up described here uses the
14 k- ϵ turbulence closure to solve for the turbulent kinetic energy k and its dissipation rate ϵ .
15 Horizontal discretization was done on a spherical grid. The coarse resolution North Sea–
16 Baltic Sea (3 nautical miles and 21 σ -layers) outer model was described in more detail by
17 Staneva et al. (2009); see also Fig. 1 of for the maps of model domains. The sea surface
18 elevation at the open boundary was generated using 13 tidal constituents obtained from the
19 satellite altimetry via the OSU Tidal Inversion Software (Egbert and Erofeeva, 2002). The
20 model was forced by atmospheric fluxes computed from bulk aerodynamic formulas. These
21 formulas used model-simulated sea surface temperature, 2-m air temperature, and relative
22 humidity together with 10-m winds from atmospheric analysis data. This information was
23 derived from the regional model COSMO-EU operated by the German Weather Service
24 (DWD; Deutscher Wetter Dienst) with a horizontal resolution of 7 km. River runoff data were
25 provided by the German Federal Maritime and Hydrographic Agency (BSH; Bundesamt für
26 Seeschifffahrt und Hydrographie). A set up for the German Bight based on the same model
27 with about 1-km horizontal resolution was nested in the coarser domain model as explained
28 by Staneva et al. (2009). Further downscaling to the scales of the Wadden Sea coastal areas
29 was implemented in nested area in the German Bight resolved with 200 m horizontal
30 resolution. All model configurations account for flooding and drying, which is a fundamental
31 dynamic process in the Wadden Sea.

2.2 Wave Model

WAM is a third generation wave model which solves the wave transport equation explicitly without any presumptions on the shape of the wave spectrum. The basic physics and numerics of the WAM Cycle 4 wave model, which is described in Komen et al. (1994) and Guenther et al. (1992) are kept in the new release WAM 4.5.3. In the coupled model system, the source function integration scheme of Hersbach and Janssen (1999) and the reformulated wave model dissipation source function (Bidlot et al., 2005), later reviewed by Bidlot et al. (2007) and Janssen (2008) are incorporated. Additionally, depth induced wave breaking (Battjes and Janssen, 1978) has been included as source function. Depth and/or current fields can be non-stationary. It is crucial for strongly tidally forced shallow areas, like the German Bight one, that model grid points can fall dry and refraction due to spatially varying current and depth is accounted for. These modifications are of utmost importance for the improvement of wave modelling results in the coastal areas such as the Wadden Sea.. The wave model code is freely available under <http://mywave.github.io/WAM/>.

Similar to the circulation model, the open boundary conditions for the German Bight WAM are taken from the regional WAM set-up for the North Sea area (with a spatial resolution: of ca. 5 km). The German Bight wave model has the same horizontal resolution and uses the same topography as the circulation model GETM. The driving wind fields are the same as the ones used in the hydrodynamical model. The required boundary information at the open boundaries of the North Sea model is derived from the regional wave model EWAM for Europe that is running twice a day in the operational wave forecast routine of the DWD. Within the framework of Coastal Observing System for Northern and Arctic Seas (COSYNA), a pre-operational wave and hydrodynamic forecast system has successfully been implemented and is running continuously since December 2009 providing hindcast and forecast data freely available on COSYNA web site under <http://www.coastlab.org>.

2.3 Coupled model implementation and periods of analyses

The original version of GETM was modified to account for the depth dependent radiation stress and Stokes drift. The terms were calculated from the integrated wave parameters according to Mellor (2008, 2011) and Kumar et al. (2011). The gradients of the radiation stresses serve as additional explicit wave forcing in the momentum equations for the horizontal velocity components. Here the Stokes drift components have been subtracted from

1 the wave processes in order to transfer it to the Eulerian framework. Moghimi et al. (2013)
2 studied the effects of the two different approaches utilising the radiation stress (Mellor 2011)
3 and vortex force (Ardhuin et al. 2008) using GETM –SWAM coupled models and showed
4 that the results for the longshore-directed transport are similar for both formulations. Recently
5 Aiki and Greatbatch (2013) showed that the radiation stress parameterization is applicable for
6 small bottom slopes and Grashorn et al. (2015) showed that radiation stress formalism is
7 applicable for shallow area like the German coastal ones. They also demonstrated that the
8 criterion proposed by Mellor (2013) to test the applicability of the radiation stress method
9 gives reasonable results in this region. This gives us a confidence that despite the know
10 limitations of the radiation stress formalism it is well applicable for our study area.
11 Additionally, the bottom friction modifications as dependent upon bottom roughness and
12 wave properties (Styles and Glenn, 2000) have been implemented. Turbulent kinetic energy
13 due to wave friction (wave breaking/white capping and bottom dissipation) that is wave
14 enhanced turbulence has also been taken into consideration (Pleskachevsky et al., 2011).

15 In order to demonstrate the impact of wave-current interaction on coastal model simulations
16 we performed two different experiments. In the first one the wave model WAM and the
17 circulation model GETM have been run separately (we will further refer to it as non-coupled
18 run). The results have been compared with the GETM-WAM coupled model system, in which
19 all wave-hydrodynamic processes described above are considered. We will further refer to it
20 as the coupled model run). Three case studies have been analysed here, which we consider
21 interesting in terms of both atmospheric conditions/extreme events and observational data
22 availability.

23 The first analyses period is in July 2011, which was a calm weather period. Two different
24 wind regimes were dominating the atmospheric state in July, 2011, which will be addressed
25 separately.

26 The next two analyses periods are chosen such as to address the effects of two of the most
27 severe storm surges affecting our study region in the last hundred years. The first storm surge
28 is the Britta storm of 31 October–1 November 2006 causing serious damages for the off-shore
29 infrastructures and shipping in the North Sea region. Britta storm was characterized by a deep
30 low-pressure centre that moved on a trajectory from north of Scotland to western Norway and
31 then eastwards through the Baltic Sea. Severe storm surge damages occurred in the East
32 Frisian Wadden Sea. Extreme sea level during this storm-surge is considered as a 100-year

1 event (Madsen et al., 2007). In addition to the storm surge, unusually high waves have been
2 measured in the southern North Sea developing on northern North Sea and propagated
3 southward under the influence of strong north winds with a long fetch. The Britta storm has
4 been given particular attention in our analyses for the types of changes that may occur during
5 single event (Bartholomä et al., 2009; Lettmann et al., 2009; Stanev et al., 2009; Grashorn et
6 al., 2015).

7 The second extreme event that we consider here is the winter storm Xavier on the 5th and 6th
8 of December, 2013 causing severe flooding and devastation along the German North Sea
9 coast. Besides of extreme high water levels along the coasts extreme sea state conditions have
10 been observed causing serious erosion of dunes and sand-displacements on the barrier islands.

11

12 **3 Impact of hydrodynamics on waves**

13 **3.1 Model validation**

14 At the buoy ‘Elbe’, which is located in the open sea (water depth about 21 m, see the middle
15 panel of Fig. 1), two different wind regimes occurred between 1. July and 10. July, 2011 (Fig.
16 2). From July 1st to 5th the dominating north-western wind did not change its direction (see
17 the red line in Fig. 2b). However wind speed increased from 7.7 m/s on 1st July to a
18 maximum of 15 m/s on 3. July (Fig.2c). The decrease of wind speed to moderate values after
19 5 of July, 2011 (less than 5 m/s) was accompanied by changing wind direction. The variations
20 of water depth and currents are tidally dominated (Fig. 2a) and not much influenced by the
21 wind during the whole period. The observed significant wave height (Fig. 2d) and the wave
22 direction (Fig. 2f) are generally in a good agreement with the measurements for both the wave
23 model only and the coupled wave-circulation one. A clear tidal signal can be seen in the wave
24 periods in the coupled model simulations, which accounted for the varying currents. It is
25 noteworthy that in addition to current refraction, the tidal water level variations and depth
26 refraction play a strong role in tidal-dominated seas like the North Sea. This well replicates
27 the available measurements (blue dots on Fig. 2e). Consequently difference of the SD of tm1
28 period decreases from 0.439s in the non-coupled run to 0.397s in the coupled one and the bias
29 (model-measurement) decreases from 0.245s to 0.174s, respectively (see Table 1). The bias
30 and SD of the significant wave height (H_s) are small in both runs demonstrating that the wave
31 models fit well with the observations.

1 The frequency wave spectra from the Elbe buoy and the two runs are shown in Fig. 3 for the
2 first 5 days in July during the strong wind event. Similarly to Fig. 2, the patterns of wave
3 spectra from the measurements and those of the coupled model run are in a very good
4 agreement (compare the top and bottom panel of Fig. 3). This is not the case for the non-
5 coupled wave model (the middle panel in Fig. 3). The tidal currents are mainly affecting the
6 tail of the spectra, whereas the energy around the peak is not much different in all three
7 panels. The statistical analysis of the observations and simulations (see Table 1) clearly
8 demonstrates the improvement of the quality of coupled wave-circulation model forecasts for
9 the German Bight in comparison to the non-coupled one.

10 **3.2 Spatial patterns**

11 To quantify the impact of currents, including water depth hydrodynamics on the results of
12 wave model, the standard deviation (SD) of H_s and the mean period (tmI), of the coupled run
13 normalized by the mean values of the non-coupled wave model are shown in Fig. 4. The
14 horizontal patterns are given as one month average for July, 2011. In the open North Sea area
15 there are no significant differences between the coupled and non-coupled wave modes for
16 both H_s and tmI . However, along the coastal areas, where currents and water level change
17 rapidly under the influence of tides, the impact of coupling seems to be significant. Within the
18 German Bight coastal areas the SD of H_s goes up to 30%, mainly due to the changes in water
19 depth. The SD of tmI is about 10-15% in the coastal area. In particular, in the South-East of
20 the German Bight, where the rivers Elbe and Weser are entering, the impact of coupling on
21 tmI period spreads much further off-shore.

22 Interesting to notice are several relatively small areas, mainly located on the tidal inlets where
23 the SD of tmI reaches values of up to 30%. These areas are characterized by strong currents,
24 up to 1.5 m/s (see Staneva et al., 2009), often parallel to the waves inducing a large Doppler
25 shift. The large SD in the entrance of the Jade Bay (located in the east Frisian Wadden Sea
26 which is the southern German Bight area with coordinates 8.25°E , 53.5°N and water depth 6
27 $\text{m} \pm 1 \text{ m}$) reveals that the wave variables H_s and tmI increase substantially during northerly
28 wind periods (inducing local wave growth, longer effective fetch) and opposing currents
29 (responsible for wave blocking and Doppler shift).

1 **4 Impact of waves on hydrodynamics**

2 **4.1. Analyses for the periods of extreme events**

3 In this section we demonstrate the role of coupling by analysing the impact of waves on
4 hydrodynamics during several extreme events. Sea level variability in four locations (ST1-
5 ST4, see Fig.1 for their geographical locations) are analysed along the German coast for the
6 period including the extreme event Xavier on 06.12.2013 (see description in Section 2). The
7 observations and simulations are shown in Fig. 5 for the tide gauge observations (black line),
8 coupled wave-circulation model simulations (coupled run- red line) and the non-coupled run
9 (circulation model only, blue line). During normal meteorological conditions, the coupled and
10 non-coupled models fit well with the tide gauge data. However, during the storm Xavier, the
11 sea level predicted by the hydrodynamic model only is underestimated with more than 40 cm.
12 It appears that the sea level predictions of the coupled model are closer to the measurements
13 (compare the red and black lines). This demonstrates the importance of wave-current
14 interactions also for the hydrodynamics. The Root Mean Square Errors (RMSE) between
15 observations and coupled model have been significantly reduced compared with the RMSE
16 differences between the observations and circulation only model for all coastal locations
17 (Table 2). Predictions of storm events with coupled models could be of utmost importance for
18 many coastal applications dealing with risk analyses (e.g. off-shore wind industry, oil
19 platform operations, etc.) where higher accuracy is needed. This justifies the consideration of
20 waves in operational forecasting.

21 **4.2 Spatial patterns**

22 In order to give an idea of the spatial distribution of the effects resulting from coupling we
23 show in Figure 6 the differences of sea surface elevation between the coupled and circulation
24 only model for 3.12.2013 at 01:00 UTC (normal meteorological situation, left panel) and
25 06.12.2013- 01:00 UTC (extreme event, right panel). The wave-induced parameterization
26 increases the average water level, which is more pronounced in the coastal area. In the open
27 North Sea the effects of coupling are almost negligible. During normal conditions the
28 difference of the sea level due to the coupling of circulation and wave models reaches a
29 maximum of 10-15 cm in the area of Elbe Estuary. However, during the storm Xavier, the
30 differences of simulated sea level when considering waves are more than 30 cm along the
31 whole German coast. In some of the Wadden Sea areas the increase of water level in the
32 simulations taking into consideration the wave-current interactions was above half meter. The

1 results shown here are indicative that the uncertainties in most of the presently used non-
2 coupled operational models result from the missing nonlinear feedback between strong tidal
3 currents and wind-waves. This can no longer be ignored in the operational oceanography, in
4 particular in the coastal zone where the wave-circulation interplay seems to be dominant. The
5 statistical analyses of simulated sea level elevation (SLE) versus tide gauge data over the
6 German Bight (Table 2) show that the coupling improves significantly the ocean predictions
7 for the whole German coastal area. The RMSEs during the calm conditions are small in both
8 coupled and circulation model only. However during the extreme events the RMSE of sea
9 surface elevation are significantly reduced when considering ocean-waves interactions.

10 In the following we will demonstrate the effect of coupling on the storm Britta on 1st of
11 November, 2011. During this storm event (see Fig. 7a), significant wave height over 10 m has
12 been simulated in the open North Sea (close to the north-western boundary). The East Frisian
13 Wadden Sea area was exposed to waves with a magnitude of about 6-7 m. Only 2 days later
14 significant wave height dropped to 4 m within the German Bight (Fig. 7b). As an example of
15 the impact of wave effects we show the dissipation of surface turbulent kinetic energy in the
16 German Bight area at the peak of the storm at 03:00 UTC on 1st of November (Fig. 7c) and
17 under calm meteorological conditions (Fig. 7d). Along the coast dissipation rates exceed 0.06
18 m^2/s^2 , which is about 100 times larger than under normal meteorological conditions.

19 Predictions of both zonal and meridional velocity have been also improved due to the
20 coupling between the waves and circulation during Storm Britta (see Fig. 8). The zonal
21 velocity has been under-estimated in the circulation only model (green line) and got closer to
22 the ADCP data for the coupled wave-circulation model (red line). There is also a very good
23 correlation between the differences of the predicted velocity and significant wave height (Fig.
24 8, bottom patterns). During the Britta storm when the significant wave height reached almost
25 8m in the coastal station the difference of the zonal velocity between the coupled run and the
26 hydrodynamic model was more than 40 cm/s. The transport along the coastal area has been
27 also increased in the coupled runs (the differences of the zonal velocity between both runs
28 being above 35 cm/s). These results are indicative that coupled hydrodynamics and wave
29 models could be of significant importance for further Lagrangian drift applications e.g. for
30 search and rescue operations as well as oil-spill analyses. The effect of wave-current
31 interactions on Lagrangian particle transport has been investigated in Röhrs et al. (2012,
32 2014).

1 Vertical section of the intensification of the longshore currents during the Britta storm is
2 shown on Fig. 9 (the location of the section is plotted in Fig. 1). Not only does the longshore
3 velocity increases but also its vertical structure has been changed through the effects of
4 coupling. Similar behaviour has been also observed by Grashorn et al. (2015).

5

6 **5 Conclusions**

7 Wave and hydrodynamic hindcast and forecast for the North Sea and German Bight are of
8 great importance for the management of coastal zones, ship navigation, off-shore wind
9 energy, naval operations etc. Storms and waves which they generate have direct impact on the
10 coastal and marine environment. The population living in the coastal areas is recently
11 concerned with the impacts of erosion and flooding, and actions aiming at better predictions,
12 impact assessments of minimization of damages are of greatest importance. Some driving
13 forces that cause serious damages on coastal environment are due to the wave conditions.
14 Their absolute and relative impact can be estimated by using coastal models. In this paper we
15 demonstrated the improvements of coastal ocean predictions due to consideration of waves-
16 current interaction for the North Sea and German Bight regions.

17 The state-of the art wave (WAM) and hydrodynamic (GETM) models coupled interactively
18 demonstrate here one step on the road to improving the ocean state estimates and predictions
19 in the coastal areas. Improved forecast statistics once considering coupling is being
20 demonstrated for both wave and circulation models.

21 The coupled system presented here enables to provide reliable predictions as well as analyse
22 long term changes of wave and circulation conditions, including extreme events. The
23 performance of the forecasting system was illustrated for the cases of several extreme events
24 along with the effects of ocean waves on coastal circulation. For our study area it can be
25 coincided that the use of radiation stress parameterization produced physically reasonable
26 results. However, the different wave-induced formalisms lead to different limitations and no
27 general recommendation should be performed. The improved skill resulting from the recent
28 coupled model developments, in particular during storms, justifies further enhancements of
29 the both forecast applications at operational services and long-term hindcasts and climate
30 analyses for the North Sea and the German Bight.

31

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Table 1: Statistics of the validation. Additionally to mean and standard deviation the coefficients of a linear regression are given.

	'Elbe'				'Hoernum Tief'			
	hs [m]		tm1 [s]		hs [m]		tm1 [s]	
mean meas.	1.10		4.36		0.33		2.43	
	WAM	WAM-GETMI	WAM	WAM-GETM	WAM	WAM-GETM	WAM	WAM-GETM
bias	0.004	-0.025	0.245	0.174	-0.073	-0.120	0.326	0.150
SD	0.164	0.171	0.439	0.397	0.117	0.136	0.350	0.293
slope	1.051	1.085	0.982	1.026	0.779	0.835	0.322	0.574
intercept	-0.061	-0.068	-0.169	-0.285	0.146	0.174	1.323	0.886

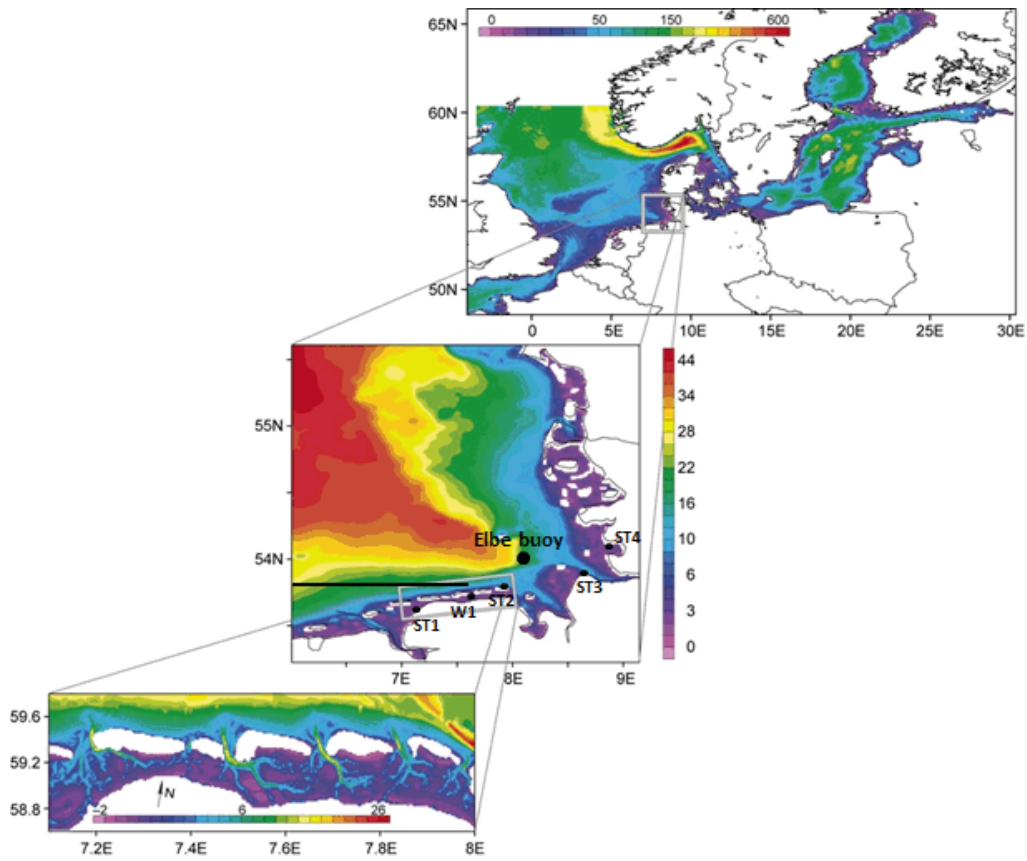
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3 Table 2: Elevation amplitude (cm) Root-Mean Square Errors (RMSE) and mean errors
4 (model-observations) for the coupled wave-circulation model and GETM model only for the
5 tide gauge data from British Oceanographic Data Centre (BODC) over the German Bight area

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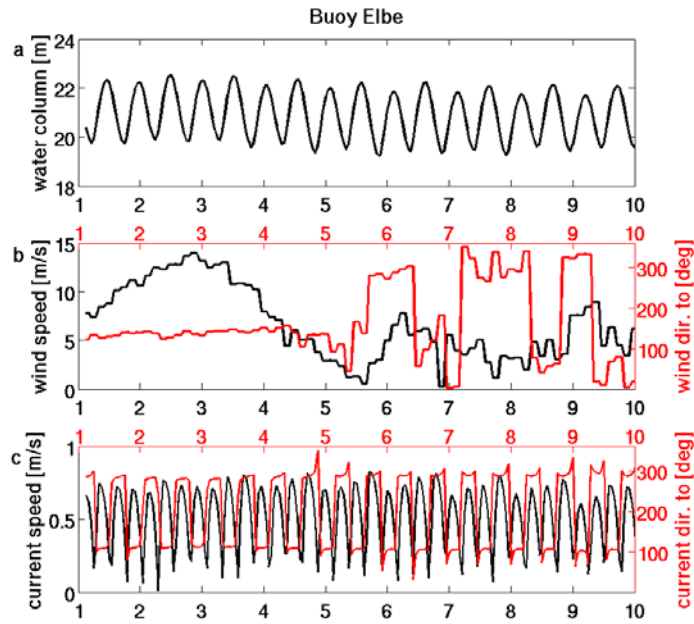
	RMSE		MEAN Error	
	WAM-GETM	GETM	WAM-GETM	GETM
Period1 (01.12.2013-12.12.2013)	12.4	19.4	-7.6	-11.5
Period2 (01.12.2013-05.12.2013)	11.8	15.2	-6.6	-10.4
Period3 (06.12.2013-07.12.2013)	13.6	22.7	-8.5	-18.5

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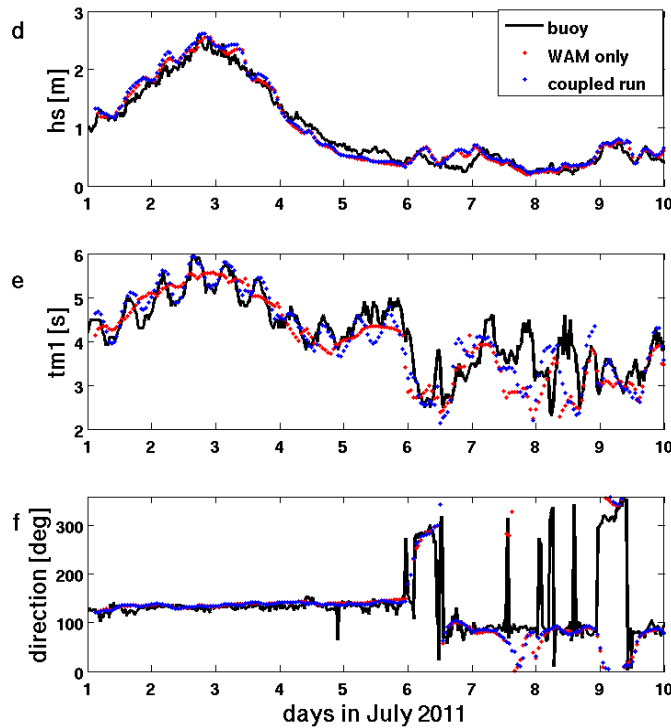


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Figure 1. Nested grid model domains for the North Sea (top pattern), German Bight (middle pattern) and East-Frisian Wadden Sea (bottom pattern). The spatial resolution is: 3 nm, 1 km and 200 m, respectively. The geographical location of stations and sections analysed later are shown as well.



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3 Figure 2: Time series at the buoy Elbe station (see Fig. 1 for its location) from 01.07.2011 to
 4 10.07.2011) of: (a): water column [m], (b) wind speed [m/s] (black line-left axis) and wind
 5 direction [deg.] (red line, left axis; (c) surface current magnitude (black line-left axis) and
 6 current direction (red line, left axis) (d) significant wave height [m]; (e) mean period-tm1 [s];
 7 and (f) wave direction [%]. For the patterns (d-f) black line corresponds to the buoy
 8 measurements, red dots– coupled model simulations, blue – wave model only.

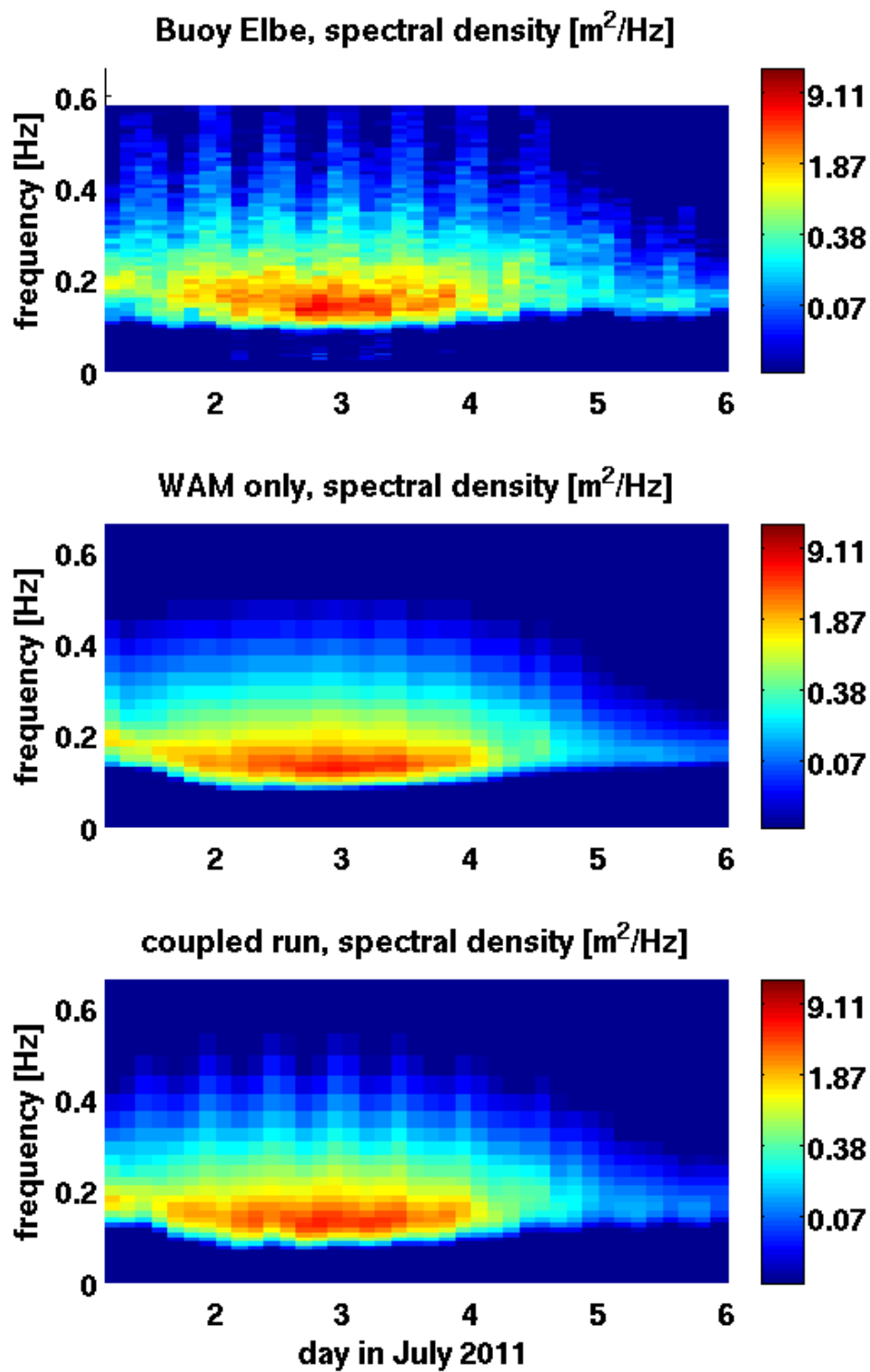
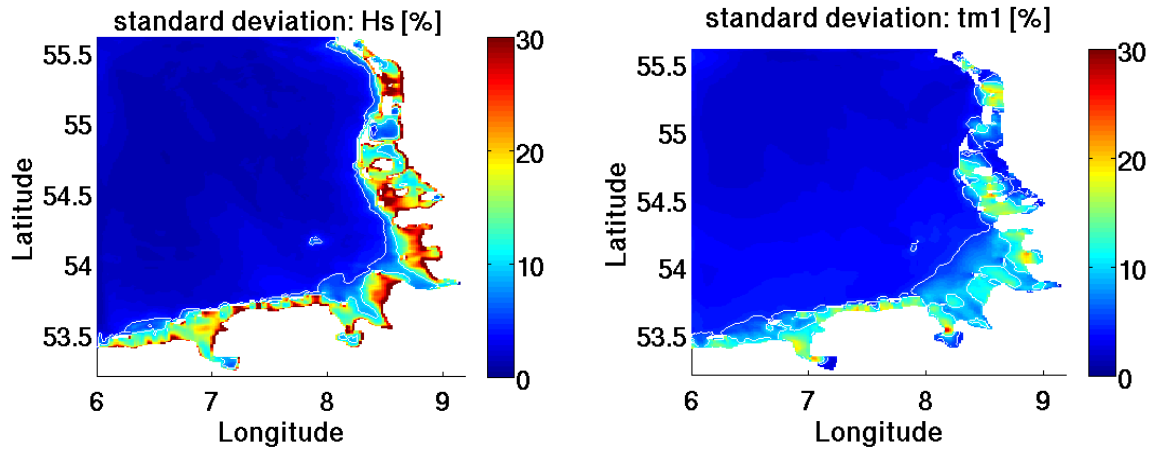


Figure 3. Comparison of measured (top) and computed values of the spectral energy density at the buoy ‘Elbe’ (see Fig. 1 for its location).

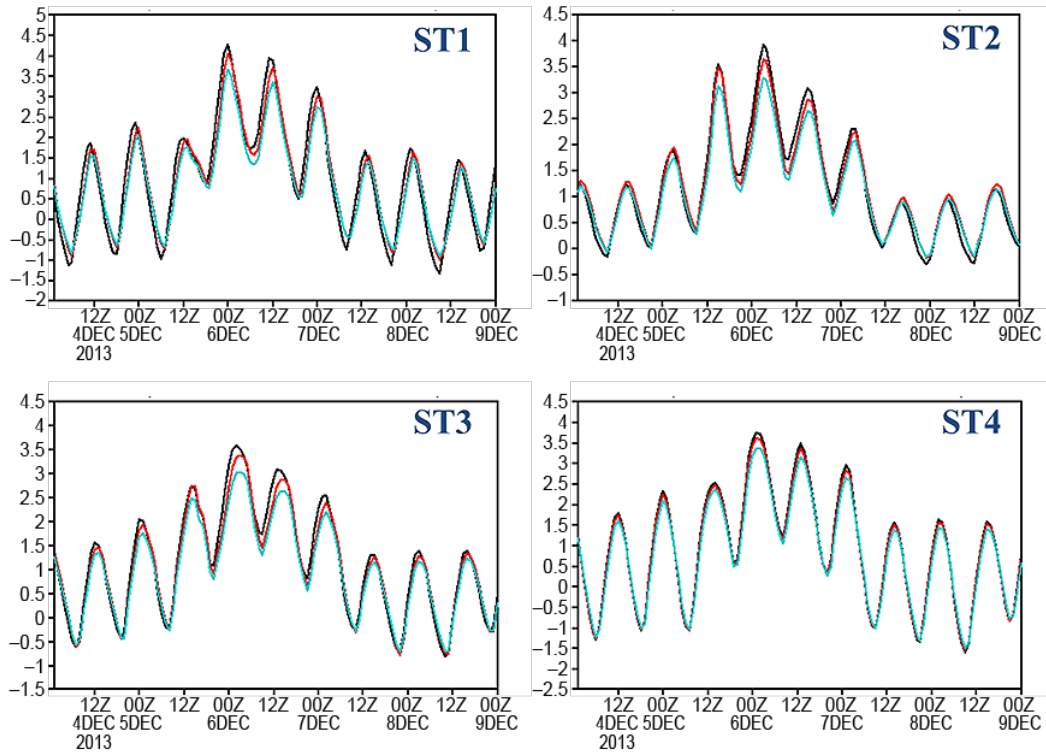


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Figure. 4. Impact of hydrodynamics on waves: Normalized standard deviation (estimated as the difference between the control run and the coupled run relative to the control run values) of significant wave height (H_s , left) and mean period ($tm1$, right) between coupled wave-circulation model and wave model only. Averaging is for one month (July 2011). The 5% and 10% isolines are plotted with white lines.

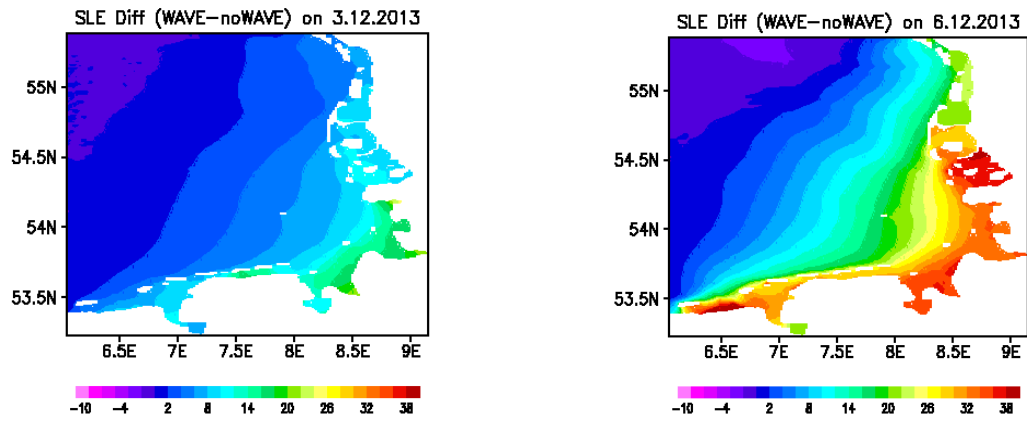
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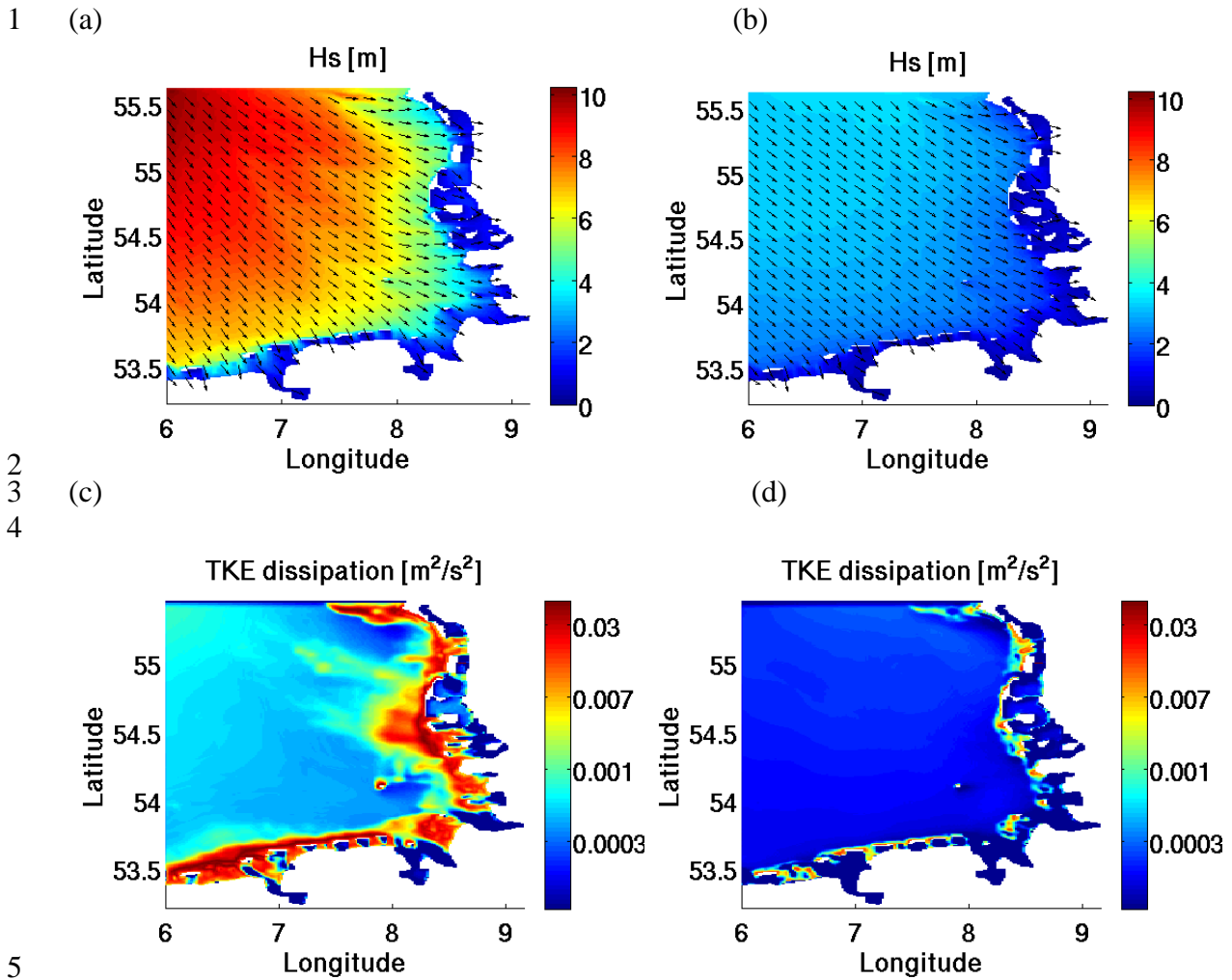
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Figure 5: Time series of Sea Level Elevation (SLE) in [m] at four coastal stations of the German Bight (ST1-ST4, see Fig. 1 for the locations). Black line: tide gauge observations, red line: coupled wave-circulation model (WAM-GETM) and green line only circulation model (GETM).



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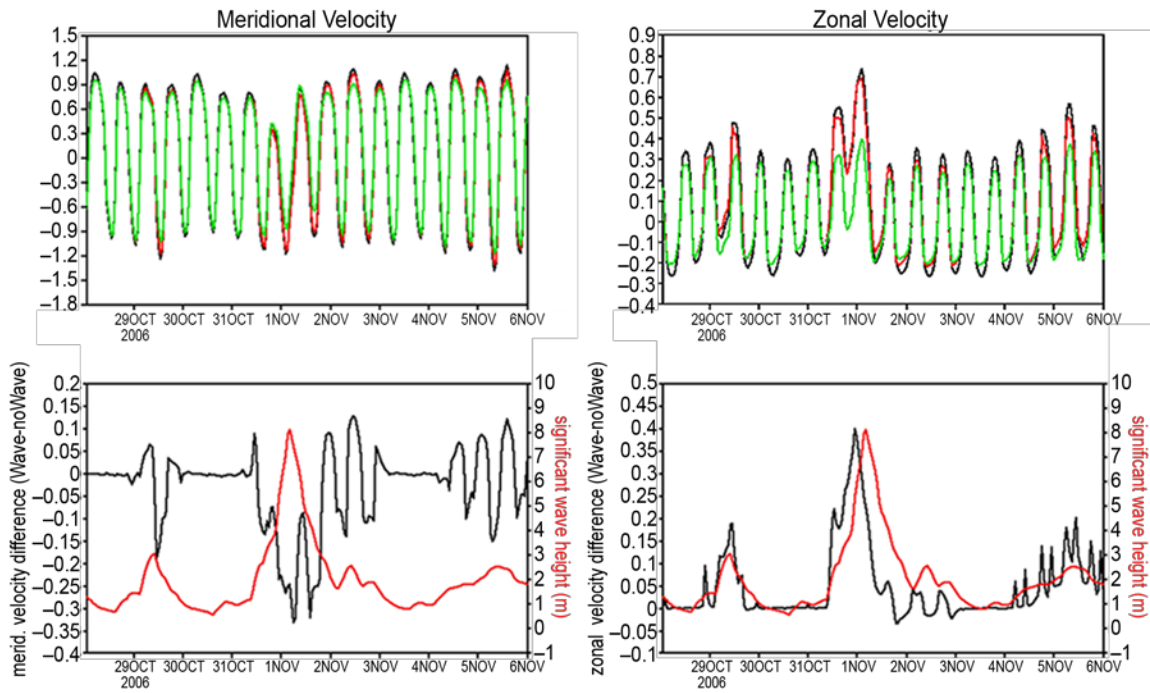
Figure 6: Sea level elevation (SLE) difference [cm] between coupled wave-circulation model (WAM-GETM) and circulation only model (GETM) for the German Bight on 03.12.2013 01:00 UTC (left) and during the storm Xavier on 06.12.2013, 01:00 UTC.



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7 Figure7. (a) Significant wave height [m] in the German Bight during the peak
 8 of storm Britta on 01.11.2006 03:00 UTC (b) (a) Significant wave height [m] in
 9 the German Bight during normal meteorological conditions on 03.11.2006
 10 03:00 UTC (c) TKE distribution in the German Bight during storm Britta on
 11 01.11.2006 03:00 UTC (d) TKE distribution in the German Bight during
 12 normal meteorological conditions on 03.11.2006 03:00 UTC.

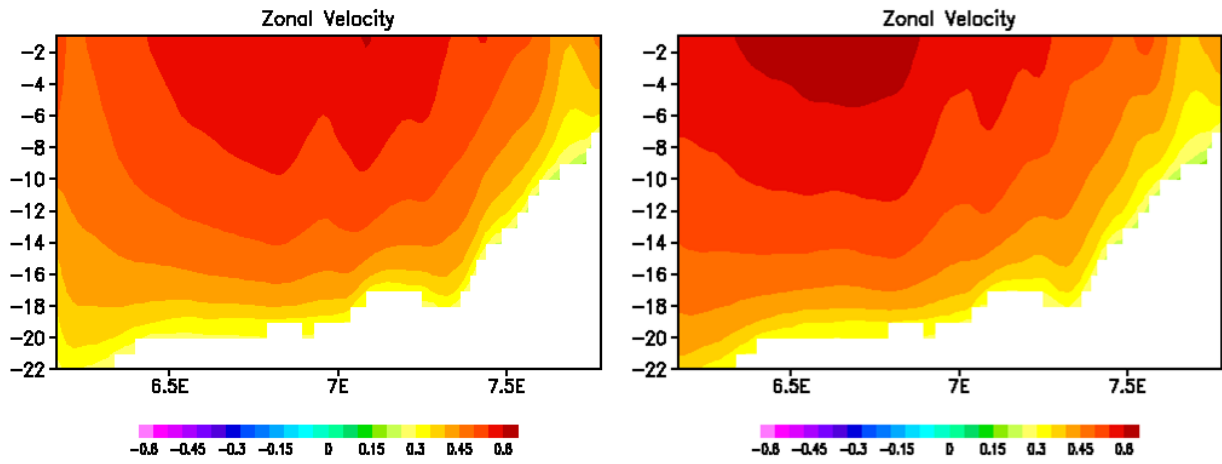
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Figure 8: Top: Meridional (left) and zonal (right) velocity time series [m/s] on station W1 (see Fig. 1 for its location) from measurements (black line), coupled wave-circulation model (red line) and hydrodynamic only model (green line) during storm Britta. Bottom: Differences between the coupled and non-coupled model simulations of meridional (left) and zonal (right) velocity [m/s]-black line and significant wave height [m]-red line.

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Figure 9: Zonal velocity vertical section [m/s] during Britta on 01.11.2006 03:00 UTC (the location of the section is shown on Fig. 1) from the hydrodynamic only model (left) and coupled model (right).