1	Improved Sea Level record over the satellite altimetry era
2	(1993-2010) from the Climate Change Initiative project
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29 Abstract

30 Sea level is one of the 50 Essential Climate Variables (ECVs) listed by the Global Climate Observing 31 System (GCOS) in climate change monitoring. In the last two decades, sea level has been routinely 32 measured from space using satellite altimetry techniques. In order to address a number of important 33 scientific questions such as: 'Is sea level rise accelerating?', 'Can we close the sea level budget?', 34 'What are the causes of the regional and interannual variability?', 'Can we already detect the 35 anthropogenic forcing signature and separate it from the internal/natural climate variability?', and 36 'What are the coastal impacts of sea level rise?', the accuracy of altimetry-based sea level records at 37 global and regional scales needs to be significantly improved. For example, the global mean and 38 regional sea level trend uncertainty should become better than 0.3 and 0.5 mm /year, respectively 39 (currently of 0.6 and 1-2 mm/year). Similarly, interannual global mean sea level variations (currently uncertain to 2-3 mm) need to be monitored with better accuracy. In this paper, we present various 40 41 respective data improvements achieved within the European Space Agency (ESA) Climate Change 42 Initiative (ESA CCI) project on 'Sea Level' during its first phase (2010-2013), using multi-mission 43 satellite altimetry data over the 1993-2010 time span. In a first step, using a new processing system 44 with dedicated algorithms and adapted data processing strategies, an improved set of sea level 45 products has been produced. The main improvements include: reduction of orbit errors and wet/dry 46 atmospheric correction errors, reduction of instrumental drifts and bias, inter-calibration biases, 47 intercalibration between missions and combination of the different sea level data sets, and an improvement of the reference mean sea surface. We also present preliminary independent 48 49 validations of the SL_cci products, based on tide gauges comparison and sea level budget closure 50 approach, as well as comparisons with ocean re-analyses and climate model outputs.

51 **1. Introduction**

52 Global warming in response to the anthropogenic green-house gases emissions has already shown 53 several visible consequences, among them the increase of the Earth's mean air temperature and ocean heat content, melting of glaciers, and loss of ice masses from glaciers and the Greenland and 54 55 Antarctica ice sheets. Ocean warming and land ice melting in turn are causing sea level to rise, with 56 potentially negative impacts in many low-lying regions of the world. The precise measurement of sea 57 level changes as well as its different components, at global and regional scales, is an important issue for a number of reasons. It provides information on how the climate system and its different 58 components respond to global warming and on the relative contributions of anthropogenic forcing 59 60 and natural/internal climate variability. This also allows validating the climate models developed for 61 projecting future changes as the models are supposed to correctly reproduce present-day and

recent-past changes. The Global Climate Observing System (GCOS) has recently defined a set of 50 62 climate variables (called Essential Climate Variables -ECVs-) that need to be precisely monitored on 63 64 the long-term in order to improve our understanding of the climate system, its functioning and its 65 response to anthropogenic forcing, as well as to provide constraints for climate modelling (GCOS, 66 2011). In 2010, the European Space Agency (ESA) developed a new program, the Climate Change 67 Initiative (CCI), dedicated to reprocessing a set of 13 ECVs currently observed from space; among them, the satellite altimetry-based sea level ECV. The objective of the CCI sea level project (called 68 69 SL cci below) was to produce a consistent and precise sea level record covering the last two 70 decades, based on the reprocessing of all satellite altimetry data available from all missions 71 (including the ERS-1&2 and Envisat missions, in addition to the TOPEX/Poseidon, Jason-1&2 and 72 Geosat Follow-on (GFO) missions). During the 1st phase of the project, that lasted 3 years from 2011 73 to 2013, satellite altimetry data from 7 altimeter satellites have been reprocessed by the SL_cci 74 consortium. Improved satellite orbits have been computed for all satellites except TOPEX/Poseidon 75 and GFO using up-to-date force models and an improved reference frame realization. Updated 76 geophysical corrections adapted to each satellite mission have been implemented after being 77 evaluated and selected. Other improvements concern the reduction of instrumental drifts and biases 78 (in particular for the Envisat mission), a new calculation of the mean sea surface used as reference, 79 the method used for geographical averaging of sea surface height data and the reduction of 80 systematic bias between missions. The main SL_cci products computed during the phase 1 consist of: 81 (1) a Global Mean Sea Level (GMSL) time series at monthly interval between January 1993 and 82 December 2010, and (2) a global gridded sea level time series (resolution 0.25°x0.25°) at the same time interval. 83

This paper thus intends to provide a global overview of the main results obtained in the frame of the SL_cci project. We firstly describe the validation protocol (section 2) that has been applied to evaluate and select the algorithms and corrections used (section 3) to generate the SL_cci products (described in section 4). Then, section 5 and 6 are focused on the assessment and the error characterization.

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2. Definition of a formal validation protocol

The altimetry data processing system used to compute sea level (or the Sea Surface Height/SSH) integrates a number of components: the altimeter range measurement (Range), the satellite orbit height (Orbit) and the instrumental and geophysical corrections. The estimation of these components needs additional information coming from different domains as orbitography (a force model) for the precise orbit determination, geodesy (geoid, mean sea surface, global isostatic adjustment (GIA), etc.), atmosphere (pressure, wind, dry and wet troposphere, etc.), and ocean (ocean tides, sea state, etc.). This information may be eventually linked together either directly or indirectly. Because of these complex interactions, sea level estimates (i. e., $SSH = Orbit - Range - \sum_{i=0}^{N} Correction_i$) are provided with different standards. In practice, an optimized sea level calculation requires a large number of algorithms and corrections that need to be rigorously validated and regularly updated.

100 In the framework of the SL_cci project, we developed a new formal validation protocol which 101 allowed us to evaluate the impact of new altimeter corrections or standards on a sea level record of 102 climate quality, i.e., precise enough for climate studies. It consists in comparing the new altimeter 103 corrections with corrections designed as a reference through their impact on the sea level 104 calculation. This was done using a common set of validation diagnoses defined in such a way that 105 they fulfil the sea level accuracy and precision requirements. The validation diagnoses are distributed 106 into 3 distinct families allowing the assessment of altimetry data with complementary objectives:

- 107 (1) the "global internal analyses" with the aim of checking the internal consistency of a specific
 108 mission related-altimetry system by analyzing the computed sea level, its instrumental
 109 parameters (from altimeter and radiometer) and associated geophysical corrections,
- (2) the "global multi-mission comparisons" allowing evaluation of the coherence between two
 different altimetry systems through comparison of SSH data,
- (3) the "altimetry-in-situ data comparison" dedicated to the computation of the sea level differences between altimeter data and in-situ sea level measurements; e.g., from tide gauges or Argo-based steric sea level data (Valladeau et al., 2012); this 3rd approach allows for the detection of potential drifts or jumps in the long-term sea level time series.

For each family, several validation diagnoses have been defined using elementary statistical approaches (e.g., mean, standard deviation, linear regression) and data representation (e.g., global mean time series, maps, histograms, periodograms, etc.). Other tests based on altimeter correction differences, sea surface height differences at satellite track crossovers, sea level anomalies, etc. were also performed. The list of all the diagnoses and their specification is described in detail in the Product Validation Plan (PVP) report of the SL_cci project (see appendix for all referenced SL_cci reports available on the SL_cci website).

The analyses of these diagnoses were performed for different spatial (global mean and regional sea level, mesoscale) and temporal scales (Figure 1, left panel): long term >10 years, interannual 2-5 years, and periodic signals -annual, semi-annual scales. These spatio-temporal scales were chosen according to the sea level user requirements document (SL_cci User Requirements Document, 2010) presented in the last section. This formal validation protocol allows us to determine, for each spatial 128 and temporal scale, the level of impact (i.e. low or strong) of the new altimetry corrections on the 129 sea level calculation (Figure 1, right panel). For instance, if a new altimetry correction causes a GMSL 130 trend > 0.15 mm/yr (over a period > 10 years), we consider that the impact is strong, whereas if the 131 trend effect is in the range 0.05-0.15 mm/yr, it is assumed low, and negligible below 0.05 mm/yr.

132 Our goal is also to check whether the new altimeter corrections improved or degraded the sea level estimates for each time scale. Most of the time, it was possible to clearly detect either improvement 133 or degradation (illustrated Figure 1, left panel, with the symbols "+" or "-" meaning improvement or 134 135 degradation). For example, increased consistency between GMSL trends derived from two different 136 altimetry missions or from in-situ measurements demonstrates that the accuracy/precision of sea 137 level data has been improved. In only a few cases, the diagnoses were inconclusive. This occurred 138 when errors of altimetry missions are of the same order of magnitude or correlated (e.g. same error 139 for the regional mean sea level trends). In these rare cases, thorough investigations could be 140 conducted through a 'case by case' approach. When no obvious conclusion could be reached, the sea 141 level differences due to the new correction were then allocated to the altimetry error budget (see section 6). 142

143 Thanks to this formal validation protocol, the impact of all altimeter corrections could be described 144 through a homogeneous approach and is therefore comparable between each other. The table 145 presented in Figure 1 (left panel) allows us to provide easily and quickly relevant information about 146 the impact of each correction on the sea level products.

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3. Development, validation and selection of new altimeter corrections and algorithms

149 In this section, we present applications of the formal validation protocol described in section 2. An important output of the SL_cci project was the development of new altimetry corrections 150 151 (mentioned in section 2) and algorithms (e.g. for merging data from different altimetry missions). A total of 42 new corrections/algorithms were evaluated within the project using the validation 152 153 protocol described above. The reference standards were those used for AVISO products (Dibarboure 154 et al., 2011) at the beginning of the SL cci project. In order to select the best corrections, a "selection 155 meeting" was held in Toulouse in May 2012 gathering a team of international experts in satellite 156 altimetry, not involved in the SL_cci project. The new corrections were then selected on the 157 condition that they led to improvements in sea level calculation. In the rare cases where the new 158 processing did not improve the results or, even worse, led to deterioration, a conservative approach 159 was applied and finally, the former corrections were unchanged.

160 Table 1 presents the new selected corrections for each component and altimetry missions (for detailed information, see "SL_cci Validation Report, Executive Summary", 2013). One of the most 161 162 dramatic improvements comes from the use of ERA-interim reanalyses (from the European Centre 163 for Medium-Range Weather Forecasts -ECMWF-; Dee et al., 2011) instead of operational ECMWF 164 fields to calculate the dry tropospheric and other dynamical atmospheric corrections. Applying our 165 validation protocol, we noted strong improvements at mesoscale and regional spatial scales, over the 166 first altimetry decade (1993-2003) (Carrere et al., 2014; "SL cci Validation reports, Atmospheric 167 corrections"). The GMSL error reduction (Figure 2, top) obtained from crossover analyses is of the 168 order of 2.5 cm on the early years of altimetry era (1993-1995). Then, the error decreases linearly 169 until 2004, and remains stable close to 0, during recent years. The improvement observed in the first 170 decade (1993-2003) is stronger at high latitudes (6 cm) where the atmospheric pressure and wind 171 fields have strong high frequency variability. Looking at regional sea level trends (Figure 2), significant 172 trend differences are observed (> 1 mm/yr) mainly in the South Pacific Ocean below 50°S latitude.

Similarly, the model-based wet tropospheric correction was also strongly improved (until 1 cm error reduction on the GMSL) before 2002 using ERA-interim instead of ECMWF operational fields (Legeais et al., 2014). While not as good as the wet troposphere corrections derived from the on-board microwave radiometers (MWR), the ERA-Interim wet tropospheric correction allows us to better characterise the uncertainty of wet troposphere content over the long term (Thao et al., 2014 ; Legeais et al., 2014). However, this was not used in the sea level calculation where the radiometerbased corrections were preferred.

180 In parallel, the radiometer-based corrections have been improved using combined estimates from 181 valid on-board MWR values, Global Navigation Satellite Systems (GNSS) measurements and ECMWF 182 model (ERA Interim fields) in areas where the MWR measurements are degraded due to, e.g., land or 183 ice contamination or instrument malfunction (Fernandes et al., 2010, 2014). This new correction, 184 called GNSS-derived Path Delay (GPD), computed for all ESA and reference missions, brings 185 improvements mainly in coastal areas and in the polar regions. In Figure 3, the sea level error 186 reduction is plotted versus the distance to the coast using the new GPD corrections instead of the 187 reference radiometer-based corrections. For almost all missions, except Jason-2 which already 188 benefits from an improved coastal radiometer correction (Brown et al., 2009), there is a significant 189 SSH error reduction, close to 1 cm between 20 and 40-50 km from the coast. Improvements have 190 also been noticed in the open ocean, especially for TOPEX data (Fernandes et al., 2014) where 191 radiometer data gaps degrade the interpolation process. Finally, the GPD corrections have been 192 selected for all altimeter missions because of the noted improvement in the sea level calculation at 193 short and long time scales, mainly in coastal and polar regions.

194 Orbit error is the main source of the error for the long-term sea level evolution at oceanic basin 195 scales (Couhert et al., 2014). Strong efforts have been made within the SL_cci project to develop new 196 orbit solutions (Rudenko et al., 2014) and to compare them with external solutions provided by other 197 projects. The International Terrestrial Reference Frame (ITRF) realisation (Altamimi, 2011) and the 198 Earth gravity field model used in the orbit computation are crucial as far as the quality of orbit 199 solutions is concerned. After analyzing all orbit solutions for all the missions, the REAPER combined 200 orbit solutions (Rudenko et al., 2012) have been selected for ERS-1 and ERS-2, with the new CNES 201 GDR-D orbit solutions (Couhert et al., 2014) being selected for the Jason-1, Jason-2 and Envisat 202 missions. Strong effects were observed on the regional sea level trend, in the range of 1-2 mm/yr, 203 with large patterns at hemispheric scale when using static and time variable Earth gravity field 204 models for orbit computation (Figure 4). Thanks to cross-comparisons between altimetry missions 205 (Ollivier et al., 2012) and with in-situ measurements (Valladeau et al., 2012), we have demonstrated 206 that these new orbit solutions dramatically improved the regional sea level trends. Furthermore, this 207 inter-comparison, using different orbit solutions, provided interesting information on the orbit 208 sensitivity to the choice of the Earth gravity field model (Rudenko et al., 2014).

209 In addition to these major improvements, other corrections were also selected, although their 210 impact on the sea level estimate was lower. These concern the ionospheric correction with the use of 211 the NIC09 (New Ionosphere Climatology) model for ERS-1 (Scharroo et al., 2010), the GOT4.8 212 (Geocentric Ocean Tide) ocean tide solution (Ray et al., 2013) and the DTU10 (Danish Technical University) mean sea surface (Andersen et al., 2010) for all missions. In addition, we also benefited 213 from the reprocessing of Envisat and Jason-2 level-2 products "GDR V2.1" (Ollivier et al., 2012) and 214 215 "GDR-D" (Philipps et al., 2013). This allowed us to increase the data coverage (mainly for Envisat) and 216 to improve the sea-state bias corrections along with instrumental bias and drift corrections. For the 217 latter, the impact is strong for Envisat since a global instrumental drift of about 2 mm/yr was identified and corrected in the altimeter range (Thibaut et al., 2010; Roca et al., 2009; Garcia and 218 219 Roca, 2010). It is worth mentioning that the SL_cci project contributed to correct this anomaly, while 220 Envisat was not designed for climate studies but rather mesoscale variability.

The last new algorithm developed and selected aims at better combining the different sea level time series from TOPEX, Jason-1 and Jason-2 at regional scale. Thanks to the verification phase between these missions, systematic geographical biases could be detected. These biases are mainly latitudedependent, with variations close to 0.5 cm between Jason-1 and Jason-2, and 1 cm between TOPEX and Jason-1. Correcting these regional and systematic sea level differences (see the SL_cci Validation Report, Regional SSH bias corrections between altimetry missions, 2012), led us to better combine together these 3 altimetry missions and therefore better estimate the long-term sea level evolution at regional scales. The impact of these corrections on regional MSL trends plotted in Figure 5 from
1993 to 2010 is close to ±0.3 mm/yr, with large hemispheric dependence.

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231 4. New CCI-based sea level records

232 Sea level products were generated using the new altimeter corrections described in section 3. The 233 same procedure was adopted as for the SSALTO DUACS (Segment Sol Multimission Altimetrie et 234 Orbitographie, Data Unification and Altimeter Combination System) system (Dibarboure et al., 2011). 235 After calculating the along-track sea level for each of the 7 missions (TOPEX/Poseidon, Jason-1, 236 Jason-2, ERS-1, ERS-2, Envisat and Geosat Follow-on) over the [1993,2010] period, the main steps 237 consisted of: combining all missions together, reducing the orbit and the long wavelength errors, 238 computing the gridded sea level anomalies using an objective analysis approach (Ducet et al, 2000; 239 Le Traon et al, 2003), and generating mean sea level products (e.g., GMSL time series, gridded sea 240 level time series, etc.) dedicated for climate studies. The SL_cci products are monthly grids time 241 series with a spatial resolution of 0.25° degrees using a rectangular projection. The GMSL time series 242 (also at monthly interval) is based on the geographical averaging over the oceanic domain observed 243 by the altimetry data (82°S to 82°N) of the gridded data. Additional products (called indicators) are 244 provided, e.g., GMSL trend, regional MSL trends, amplitudes and phases of the main periodic signals 245 (annual, semi-annual), etc.

Access to the SL_cci products can be obtained by sending an email at the following address: <u>info-</u> sealevel@esa-sealevel-cci.org. The Product User Guide (PUG, 2013) and Product Specification Document (PSD, 2013) provide further details.

249 Comparisons between the SL_cci product and the AVISO-2010 products (Dibarboure et al., 2011) were performed by applying the formal validation protocol described above (section 2). Concerning 250 251 the GMSL trend, similar values were obtained for both time series: 3.2 mm/yr over the 1993-2010 252 time span. At the interannual time scale, (highlighted by calculating the difference between the two 253 GMSL time series (Figure 6, top panel), small differences in the range 1-2 mm or lower are noticed, 254 except for 1994 where a 4 mm jump is observed. This jump is due to an anomalous value of the 255 AVISO-2010 products caused by an inadequate merging of the TOPEX data with the ERS-1 data of the 256 non-repetitive geodetic phase (Pujol et al., 2014). The most impressive result is obtained by 257 separating the ERS-1/ERS-2/Envisat and TOPEX/Jason-1/Jason-2 global GMSL time series using 258 alternately the old and new altimeter corrections (Figure 7): the trend difference between both time 259 series is now close to 0.6 mm/yr from 1993 to 2010 instead of about 1.5 mm/yr previously. This improved consistency does not have a direct impact on the GMSL trend, which only depends on the
 TOPEX/Jason-1/Jason-2 missions. However, this provides increased confidence in the long-term
 GMSL time series.

263 Looking at the regional sea level trend differences (Figure 6, bottom panel), large geographically 264 correlated structures are observed. Their amplitude is in the ±2 mm/yr range. They primarily result 265 from the new orbit solutions (hemispheric effects), the new ERA-interim atmospheric fields (at high 266 latitudes), the new wet tropospheric correction, and the geographical biases arising when linking 267 altimetry missions together. Comparing with in-situ measurements (tide gauges and Argo-based 268 steric sea level) indicates a better consistency at the regional scale with the new SL cci data (see 269 SL_cci Product Validation Internal Report - PVIR, 2013). It is more difficult to detect any 270 improvement at short spatial scales, because either the spatial or temporal sampling of in-situ 271 measurements is not good enough or because the error generated by the collocation method 272 between the in-situ and altimetry data is larger than the target signal (Couhert et al., 2014). We also 273 examined the periodic (annual and semi-annual) sea level signals. We found differences in the order 274 of 5 mm on average for the amplitude of the annual signal. In some regions (the tropics), the 275 differences can reach 1 cm. Whilst we think that the new seasonal signal is improved compared to 276 the AVISO-2010 products, it is not possible to demonstrate this through any independent validation 277 diagnoses. Indeed, comparisons with the in-situ measurements are not accurate enough to observe 278 such signals.

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5. Validation of the temporal and spatial variations of global sea level:

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The SL_cci products delivered at the end of Phase 1 are currently under validation and evaluation.Two different approaches have been developed:

- (1) Assessment of the accuracy of the SL_cci products through their use in ocean reanalyses and
 Earth system models
- 285 (2) Assessment of the global sea level budget
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In approach (1), the accuracy of the SL_cci data is evaluated by quantifying the model performances and robustness (compared to the use of using a reference sea level data set, e.g., AVISO standard data) in representing a number of physical processes (e.g., the sea level drop associated with the 2011 La Niña, the Indonesian through flow, changes in the Arctic circulation, effects of monsoon on sea level, regional sea level fingerprint due to wind stress, steric sea level trend patterns, etc.). Approach (2) consists of comparing the SL_cci GMSL and variability to (i) other GMSL, and (ii) the sum of the climatic and non-climatic components estimated independently (changes in thermal expansion, glacier and ice sheet mass balance and land water storage).

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296 5.1 Assessment based on numerical ocean models

297 Ocean model simulations are an effective way of translating wind and heat fluxes information into 298 sea level variations, thus providing independent verification of their contribution to sea level. Sea 299 level from ocean-only simulations at different resolutions (1° degree, ¼° of degree) has been 300 contrasted with along-track data and with gridded (filtered and merged) sea level maps from AVISO 301 (Dibarboure et al., 2011) and SL_cci. The statistics of the comparison (correlation, rms error, 302 differences in trends) were similar when using AVISO and SL_cci data. Differences between models 303 and any observed estimations were much larger than the differences between observational 304 products. The spatial patterns of these differences were suggestive of model error. For instance, 305 small scale sea level variability is much larger in observed products than in models, which is 306 consistent with insufficient resolution in the models. In contrast the low frequency and large scale 307 variability is more obvious better resolved in models. The large scale patterns of interannual 308 variability and trends are consistent between models and observations, but differences exist 309 associated with the precise location of strong current systems, which models struggle to capture. 310 This information is in itself interesting, and suggests that a large part of the sea level variability is of 311 dynamic nature, associated with changes in the wind-driven circulation. Both AVISO and SL_cci were 312 useful to detect improvements in ocean model simulations due to the increased resolution.

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In the Arctic Ocean the SL_cci reprocessed data reveals some distinct features of the elevated trend in sea level rise, notably: in the Beaufort Sea, in the Norwegian Sea, in the Sub-Polar gyre, and in the North East Atlantic south of the Iceland-Faroe ridge. The Beaufort Sea rise of about 6.5-7 mm/year has also been reported by Morison et al., (2011) and Laxon et al., (2012), while the elevated feature of around 6-7 mm/year, as detected in the SL_cci field in the Lofoten Basin of the Norwegian Sea, compares rather well with the trend recovered from in-situ hydrographic observations.

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A first look at the three GCMs (General Circulation Model), NorESM (Norwegian Earth System Model), Hadley and IPSL (Institut Pierre-Simon Laplace), reveals large individual differences in the trend of sea level change, both regarding the overall trend as well as in its regional characteristic changes. The contributions to these simulated changes include the regional variability of the steric and the mass components, while there is no account of the GIA. In comparison to the SL_cci sea level change the NorESM simulations (1° resolution) yield the best agreement both in the Sub-Polar gyre, in the northeast Atlantic Ocean south of the Iceland-Faroe ridge, in the Lofoten basin of the Norwegian Sea and in the Beaufort Gyre. This inter comparison of the SL_cci trends with the trends derived from the three GCMs can therefore provide evidence for how realistic the model simulations are with respect to the regional variability of the water masses (steric height contribution) and variability, spreading and accumulation of freshwater discharges from melting ice sheets and glaciers (mass changes).

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In summary, as it was to be expected from the beginning, even ocean-only simulations are not able to identify the incremental improvement of SL_cci versus its predecessor. Nevertheless, this validation exercise has shown that the SL_cci is a robust dataset for ocean and climate models validation, and can discern verification metrics.

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339 5.2 Assessment based on ocean data assimilation

340 Data assimilation methods can be very effective methods to test the quality of the input data. This 341 approach was used here to evaluate the SL cci products, either by direct assimilation of the product 342 as an ocean synthesis (active mode) or by simple comparison with a reference state (passive mode), 343 obtained by a forced ocean-model combined with in-situ observations, and even other sea level 344 observations. In this way, the ocean synthesis, containing information both from the model forced 345 with realistic atmospheric state and observations, should have less error than an ocean-model 346 simulation alone. The passive comparison can be done a-posteriori (by comparing ocean reanalyses 347 with SL_cci), or during the assimilation process, by contrasting, at the appropriate location and time, 348 the along track altimeter altimetry data with the estimate given by an ocean model that assimilates 349 in-situ temperature and salinity.

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351 In a first step, sea surface height fields available from the GECCO2 assimilation approach (Köhl, 2014) 352 were compared to the AVISO products as well as to the SL_cci product, respectively. Of these two, 353 the AVISO product was used to constrain the model, but not the SL_cci product. The comparison was 354 performed to investigate whether the new SL_cci product is closer to the GECCO2 ocean reanalysis 355 product, which is constrained by most of the available global data sets, than the previous AVISO data set, a test that would highlight a better consistency of the new SSH data with ocean dynamics and 356 357 other ECV information. The comparisons have been performed separately for the ERS (ERS-1, ERS-2 358 and ENVISAT) and the TOPEX/Poseidon satellite-series (TOPEX/Poseidon, Jason-1 and Jason-2). Figure 8 shows the ratio (RMS_AVISO/RMS_SL_cci) of the RMS differences between the GECCO model and 359 360 the satellite time series of ERS-1, ERS-2 and ENVISAT for AVISO (RMS_AVISO) and SL_cci (RMS SL cci) in percent improvement at model resolution. Red indicates improvements of the SL cci 361

compared to the AVISO data set and blue degradation. Remarkable are the improvements in the 362 363 north Atlantic, in the Indian Ocean through flow and in many parts of the ocean. The regions where 364 SL_cci shows less skill compared to AVISO are the ones where the GECCO2 solution has adapted very 365 well to AVISO and at the same time where the STD of the datasets are very small, indicating a small 366 signal to noise ratio in these regions. Therefore, the model might have adapted to the not as good 367 AVISO data and thus gives less skill in comparison to the improved SL_cci dataset. The improved 368 regions (red colors) cover 62.8 % of the ocean area that had valid data for the comparison, leaving 369 37.2 % of the ocean area that has degraded (blue colors). Further, when averaging the ratio of 370 RMS_AVISO/RMS_SL_cci globally, weighted by the area of each grid point, a global mean 371 improvement of 0.91 % can be seen from the analysis on the model grid. This could demonstrate 372 that the SL_cci has been improved in many regions.

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374 Both AVISO and SL_cci sea levels have also been compared with the sea level from the ORAS4 ocean 375 reanalyses (Balmaseda et al., 2013), which assimilate in situ temperature, salinity and AVISO data 376 along track altimeter. Time series of standard area-averaged climate indices have been used to gain 377 insight on the differences between the AVISO and SL cci products. Figure 9 shows a time series of 378 the 12-month running mean sea level anomaly differences (respect AVISO for SL_cci (red) and ORAS4 379 (blue)). In the Eastern Pacific (5N-5S, 130W-90W left panel) both ORAS4 and SL_cci show a positive 380 offset with respect to AVISO data after 2005 (from 2005 onwards the ocean state in ORAS4 is 381 relatively well constrained by Argo). In addition, SL_cci and ORAS4 data consistently show stronger local maxima associated with El Nino 1997. The precursor of this El Niño is visible in the Western 382 383 Pacific slightly earlier, and it is also more pronounced in SL cci and ORAS4 than in AVISO (not shown). 384 The right panel of Figure 9 shows the equivalent time series for the Southern Indian Ocean (30°S-385 70°S, 20°E-150°E), where both ORAS4 and SL_cci consistently show a negative tendency with respect 386 to AVISO, suggesting that AVISO overestimates the sea level rise in this area. The differences in 387 trends between SL_cci and AVISO shown in these time series are similar to those shown in Figure 6 388 (bottom). The variability of the ORAS4 reanalysis agrees better with the SL_cci product than with 389 AVISO.

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391 5.3 Comparison of the SL_cci GMSL time series with other GMSL products

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We constructed a GMSL time series by geographically averaging the SL_cci gridded data between 66°S and 66°N. A simple cosine of latitude weighting was applied to the data. As no glacial isostatic adjustment (GIA) correction was applied to the gridded data, we added the usual +0.3 mm/yr GIA 396 trend from the SL_cci GMSL (as usually done by other processing groups). We further compared the 397 SL_cci GMSL with altimetry-based GMSL time series computed by different processing groups: 398 (AVISO, University of Colorado (CU), NOAA (National Oceanic and Atmospheric Administration), GSFC (Goddard Space Flight Center) and CSIRO (Australia's Commonwealth Scientific and Industrial 399 400 Research Organisation). The results are shown in Figure 10 (left panel). In terms of trends, all curves 401 are very really similar to each other and trend differences (<0.2 mm/yr) are fully covered by the 402 formal error on the trend computation. However, it is interesting to note that all sea level curves 403 differ significantly (by several mm) over an interannual time scale. This is illustrated in Figure 10 404 (right panel). This is particularly noticeable during the TOPEX/Poseidon period (1993-2001), with a 405 significant big departure of the CSIRO GMSL from other curves. The detrended SL_cci GMSL is in general close to the AVISO GMSL, although slight differences are noticed at the end of the study 406 407 period.

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409 5.4 Comparison of the SL_cci GMSL with steric and ocean mass components (sea level closure 410 budget); Interannual time scale.

411 GMSL change is a combination of ocean mass and steric (thermal expansion) changes. We compared 412 the GMSL computed from the SL_cci gridded product with the sum of steric and mass components 413 over the Argo and GRACE (Gravity Recovery and Climate Experiment) operating period (since ~2005). 414 Argo-based steric data used for this comparison is based on that processed by Karina von 415 Schuckmann (von Schuckmann and Le Traon, 2011). Ocean mass has been estimated using the RL05 416 data from the GRACE project (average of the three products: CSR, JPL and GFZ, Chambers et al., 417 2012). The GRACE and steric data have been averaged over the 66°S and 66°N domain. Figure 11 418 compares three GMSL products (AVISO, CU and SL cci) with the sum of steric and mass contributions 419 over 2005-2010. Error bars of the sum 'steric plus mass' time series are not shown for clarity. They 420 are estimated to within +/- 2 mm for individual monthly values. The mean trend over the study 421 period (2005-2010) has been removed. The three GMSLs present similar variations and show 422 reasonably good agreement with the sum of the components. Although small differences exist, the 423 best agreement is found for the SL cci GMSL. Correlation coefficients between the sum 'steric plus 424 mass' component and GMSL time series have also been computed. The highest correlation (of 0.65) 425 is found with the SL cci GMSL.

The results presented above are first attempts to validate the SL_cci products. We find some differences both in terms of global mean and regional variability with the standard products. Preliminary comparisons with the sum of the climate contributions (the sea level budget closure budget approach) suggest that the CCI product fits better the sum of the climatic components. However, this result is not robust considering the large uncertainties affecting the steric and mass components. Further work is needed on that matter, using different steric and ocean mass products
with assessed uncertainties. For instance, the steric height from ocean reanalyses can also be used
for global sea level budget closure (Balmaseda et al., 2013). This will be a topic for the CCI phase 2
activities.

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437 6. Error budget of sea level

Although improvements were made, the SL_cci products still contain remaining errors at different time scales. In order to inform users about these errors, we have established an error budget dedicated to the main spatio-temporal scales (i.e., global and regional, long-term - 5-10 years or more -, interannual - <5 years - and seasonal (see Table 2)). For each of these, an error was determined and compared to the sea level Climate User requirements (GCOS, 2011) which have been updated in the framework of the Sea Level CCI project (Sea Level CCI User Requirement Document -URD-, 2013).

445 Regarding the GMSL trend, an uncertainty of 0.5 mm/yr was estimated over the whole altimetry era 446 (1993-2010). This uncertainty is reduced by 0.1 mm/yr compared to the previous data based on 447 AVISO-2010 standards over 1993-2008 (Ablain et al., 2009). While small, this reduction is mainly due 448 a 2-year longer record as well as to the homogenization of the altimetry corrections between all the 449 missions. The main source of the error remains the radiometer wet tropospheric correction with a 450 drift uncertainty in the range of 0.2 - 0.3 mm/yr (Legeais et al., 2014). To a lesser extent, the orbit 451 error (Couhert et al., 2014) and the altimeter parameters (range, sigma-0, SWH) instabilities (Ablain 452 et al., 2012) also add additional uncertainty, of the order of 0.1 mm/yr. Notice that for these two 453 corrections, the uncertainties are higher in the first altimetry decade (1993-2002) where 454 TOPEX/Poseidon, ERS-1 and ERS-2 measurements display stronger errors (Ablain et al., 2013). 455 Furthermore, imperfect links between TOPEX-A and TOPEX-B (February 1999), TOPEX-B and Jason-1 456 (April 2003), Jason-1 and Jason-2 (October 2008) lead to the errors of 2 mm, 1 mm and 0.5 mm 457 respectively (Ablain et al., 2009). They cause a GMSL trend error of about 0.15 mm/yr over the 1993-458 2010 period. Although the SL cci project work has led to significant improvements, the remaining uncertainty of 0.5 mm/yr on the GMSL trend remains 0.2 mm/yr higher than the GCOS requirements 459 460 (of 0.3 mm/yr, see GCOS, 2011).

All sources of errors described above have also had an impact at the interannual time scale (< 5 years). Recent studies (Henry et al., 2013) have also shown that the methodology applied to calculate sea level is particularly sensitive for the interannual scales (Henry et al., 2014). We estimated that the 464 methodology uncertainty is on average ~2 mm over a 1-year period. Although improvements have 465 has been made, this level of error is still 1.5 mm higher than the GCOS requirement of (0.5 mm). This 466 may have consequences on the sea level closure budget studies at the interannual time scale. For the 467 annual signal, the amplitude error was estimated to be <1 mm. Knowing that the annual amplitude of 468 the GMSL is of the order of 9 mm, we can consider this error is low. Notice that no requirement has 469 yet been defined by GCOS for the periodic signals (at global and regional scales).

470 At the regional scale, the regional trend uncertainty is of the order of 2-3 mm/yr. Although the orbit 471 error has been significantly reduced for this spatial scale, it remains the main source of the error (in 472 the range of 1-2 mm/yr; Couhert et al., 2014) with large spatial patterns at hemispheric scale. The 473 Earth gravity field model errors explain an important part of these uncertainties (Rudenko et al., 474 2014). Furthermore, errors are higher in first decade (1993-2002) where the Earth gravity field 475 models are less accurate due to the unavailability of the Gravity Recovery and Climate Experiment 476 (GRACE) data before 2002. Additional errors are still observed, e.g., for the radiometer-based wet 477 tropospheric correction in tropical areas, other atmospheric corrections in high latitudes, and high frequency corrections in coastal areas. The combined errors give rise to an uncertainty of 0.5-1.5 478 479 mm/yr. Finally, the 2-3 mm/yr uncertainty on regional sea level trends remains a significant error 480 compared to the 1 mm/yr GCOS requirement, even if this project has led to a 0.5 to 1.5 mm/yr 481 reduction (Figure 6).

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7. Conclusions and perspectives

484 Several groups (AVISO, University of Colorado, CSIRO, JPL (Jet Propulsion Laboratory), etc...) are 485 currently processing satellite altimetry data to provide sea level products to user for climate applications. Within the SL_cci project, we have continued to improve the multi-mission sea level 486 487 products over the altimetry era (1993-2010) through the development and computation of new 488 corrections listed in Table 1. As far as possible, we have homogenized these corrections between all 489 the missions in order to reduce the sources of discrepancies. Thanks to our formal validation 490 protocol, we have been able to select the best corrections and algorithms applied in the sea level 491 calculation. We have produced new sea level products and additional indicators over the 1993-2010 492 period. The SL_cci products exhibit improvements of different importance for climate studies. Some of them are substantial for instance for the estimation of the regional sea level trends, with an error 493 494 reduction of 0.5-1.5 mm/yr with large correlated spatial patterns. In parallel, the uncertainties of 495 altimetry sea level have been better characterized and the sea level user requirements refined for 496 climate applications.

497 The validation exercise has demonstrated that the existence of an additional good quality sea level 498 record has value in itself. Firstly, it clearly shows that the AVISO and SL_cci altimeter-derived sea 499 level gridded products are robust (small uncertainty compared with the model error), and able to 500 identify model improvements. Therefore they are a suitable data set to define metrics in the 501 validation of ocean and climate models. SL_cci can be treated as an independent data set for 502 verification. It has been used in the recent inter-comparison of ocean reanalyses ORAIP (Balmaseda 503 et al., 2014, Hernandez et al., 2014). Preliminary results show that the SL cci is closer to the 504 ensemble mean of ocean reanalyses (a robust estimator) than its predecessor AVISO, and suggest 505 that some ocean reanalyses that assimilate AVISO may over-fit the altimeter data. Model outputs 506 using ocean assimilation techniques also provide independent sea level estimations that can be used 507 to validate the SL_cci. Results obtained in the frame of the SL_cci project show that the low 508 frequency variability and trends of SL_cci agree better with ocean data assimilation estimators than 509 with AVISO, especially in the Southern Ocean, the Eastern Pacific and coastal areas.

However, while a lot of improvements have been made, the user requirements are not yet reached. Remaining uncertainties are still 0.2 mm/yr and 1-2 mm/yr higher than the GCOS requirements for the GMSL trend and regional trends respectively. Similarly, the sea level error over a 1-year period is about 2 mm on average instead of the required 0.5 mm. Therefore it is still necessary to continue to improve the sea-level time series to better understand key scientific issues, as raised in the abstract.. Several ways of improvements have already been identified and will be implemented during phase 2 of SL cci project (January 2014 to December 2016).

517 For example, we plan to extend the sea level time series beyond 2010 using the same sea-level 518 corrections. By the end of year 2014, the current CCI_SL release will be extended until 2013 519 (included). And each subsequent year, we will extend the time series by 1 year. Additional 520 improvements will be implemented; in particular, new orbit solutions, use of new atmospheric 521 reanalyses based on ERA-Clim project (Dee et al., 2014), new ocean tides, new radiometer-based wet 522 troposphere corrections with improved long-term stability, etc. Furthermore, several level-2 523 altimetry data reprocessing activities are already planned by space agencies (CNES, NASA, ESA) for 524 Jason-1, TOPEX/Poseidon, Envisat and ERS missions, allowing us to benefit from homogenized data 525 both for instrumental parameters and geophysical corrections. In addition, we intend to account for 526 new altimeter missions already in orbit (CryoSat-2, SARAL/Altika) or to be launched in the near future 527 (Jason-3, Sentinel-3). They are all relevant to extend the sea level time series with the same level of accuracy, and to improve coastal and high latitude areas which are of great interest for climate 528 529 studies. Dedicated analyses will be performed in the Arctic region in order to improve sea level 530 estimates nearby or under sea ice where no data is currently available. In parallel, we will continue to refine the user requirements, further developing the link with users and space agencies. This will include a quantification of the requirements for accuracy and long-term stability for climate-quality observations of sea level in the coastal zone, a key area for climate change. We also would like to refine the budget error with the new measurements and the new corrections. Lastly and with the idea to continuously answer to the user needs, we will produce by the end of 2016, a new improved sea level time series covering the 1993-2015 period.

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	CL sei Validation Demont Manusing Matheada available at http://www.aca.acale.ol
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767 - 768 769	SL_cci Validation Report, High latitudes areas, available at: <u>http://www.esa-sealevel-cci.org/webfm_send/185</u> (last access : 20 June 2014)
770 - 771 772 773	SL_cci Validation Report, Coastal areas, available at: <u>http://www.esa-sealevel-</u> cci.org/webfm_send/186 (last access : 20 June 2014)
774	

- Figure 1: Definition of the temporal and spatial scales (on left panel) and the indicator value table (on
- right panel) allowing the impact characterization in sea level of new SL_cci corrections in comparison
- with corrections defined as reference (AVISO-2010).
- 778

Impact of new SL-CCI in SSH calculation in comparison with AVISO							
standards				Definition of the indicator value			
Climate Applications	Climate pplications Temporal Scales For 1 mission (Envisat, ERS, Jason, T/P,)			Significant impact	Low impact	No impact detected	
	Long-term evolution (trend)	ΝΟ ΙΜΡΑCΤ	Ţ	Trend >0.15 mm/yr	Trend>0.05 mm/yr	Trend< 0.05 mm/yr	
Global Mean Sea Level	Inter annual signals (> 1 year)	LOW IMPACT	4	Amplitude > 0.5 mm	Amplitude > 0.2 mm	Amplitude< 0.2 mm	
	Annual and semi- annual Signals	STRONG IMPACT	4	Amplitude> 1 mm	Amplitude > 0.2 mm	Amplitude< 0.2 mm	
Regional Mean Sea Level	Long-term evolution (trend)			Trend > 0.5 mm/yr	Trend>0.1 mm/yr	Trend< 0.1 mm/yr	
	Annual and semi- annual Signals			Amplitude > 5 mm	Amplitude > 0.5 mm	Amplitude< 0.5 mm	
Mesoscale	Signals < 2 months		Ţ	Crossovers Variance differences > 1 cm ²	Crossovers Variance differences > 0.2 cm ²	Crossovers Variance differences < 0.2 cm²	

- 781 Figure 2: Evolution of the sea level error reduction applying the new Dynamical Atmospheric and dry
- troposphere corrections derived from ERA-Interim reanalyses instead of operational ECMWF fields
- 783 (top) and impact on sea level regional trends (bottom).









- 787 Figure 3: Evolution of the error reduction versus the coastal distance applying the new GPD wet
- troposphere corrections instead of the reference radiometer-based corrections used in AVISO-2010.





- 792 Figure 4: Impact of the new orbit solutions on the regional sea level trends for ERS-2 (Reaper
- combined versus DEOS DGM-E04 orbit solutions), Envisat, Jason-1 and Jason-2 (CNES GDR-D versus
 CNES GDR-C orbit solutions).



-2





797 Figure 5 : MSL trend differences from 1993 to 2010 between sea level maps without and with

regional bias corrections for TOPEX/Jason-1 and Jason-1/Jason-2

799



Figure 6: GMSL (top panel) and regional sea level (bottom panel) differences between the SL_cci
(release 1.1) and AVISO products (release 2010).



- 808 Figure 7: GMSL time series separating ERS-1/ERS-2/Envisat and TOPEX/Jason-1/Jason-2 altimeter
- 809 missions using alternatively the old (AVISO-2010 standards) on left and new altimeter correction
- 810 (SL_cci) on right.



815 Figure 8: Ratio of the RMS differences RMS_AVISO and RMS_SL_cci between the GECCO model and

the satellite time series of ERS-1, ERS-2 and ENVISAT in percent improvement.



- 823 Figure 9: Differences (m) in the sea level time evolution (12 month running mean) respect the AVISO
- 824 product of SL_cci (red) and ORAS4 (blue) . Left: Eastern Equatorial Pacific (5N-5S, 130W-90-W). Right:
- 825 Southern Indian Ocean(30S-70S, 20E-150E). The differences in trends between SL_cci and AVISO are
- 826 confirmed by ORAS4. In the Eastern Pacific, both ORAS4 and SL_cci have stronger ENSO signature
- 827 than AVISO.



- 831 Figure 10: GMSL based on multi-mission satellite altimetry data processed by different groups
- 832 (including SL_cci project). Left/right panel : with/without the global mean trend.





- 836 Figure 11: Sum of steric and ocean mass component based on Argo and Grace data (see text section
- 5.4) (green curve) over the Jan. 2005-Dec. 2010 time period and different GMSL products (left
- panels). Right panel: difference between the GMSL products and sum of components.







- boxes indicate that the AVISO standards (release 2010) have been applied.

Corrections	ERS-1	ERS-2	Envisat	Jason-1	Jason-2	T/P	GFO
Orbit	Reaper com (Rudenko e	ibined orbit et al., 2012)	GDR-D CI	NES (Couhert et			
Instrumental correction			New PTR Correction (Garcia and Roca, 2010)				
Sea State Bias			V2.1 release		GDR-D release		
Wet Troposphere	GPD corrections (Fernandes et al., 2010, 2014)						
Dry Troposphere	ERA-interim based (Carrere et al., 2014)						
Dynamical atmospherical corrections	ERA-interim based (Carrere et al., 2014)						
Ocean tide	GOT 4.8 (Ray et al., 2013)						
Mean Sea Surface	DTU 2010 (Andersen et al., 2010)						

850 Table 2: Error budget of SL_cci products for the main climate scales

Spatial Scales	Temporal Scales	Altimetry Errors	User requirements	
	Long-term evolution (> 10 years)	< 0.5 mm/yr	0.3 mm/yr	
Global MSL	Interannual signals (< 5 years)	< 2 mm over 1 year	0.5 mm over 1 year	
	Annual signals	< 1 mm	Not defined	
Regional MSL	Long-term evolution (> 10 years)	< 3 mm/yr	1 mm/yr	
	Annual signals	< 1 cm	Not defined	